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Advantages of Inter-calibration for Geostationary Satellite Sensors onboard Twin Satellites

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Key Points:

- Inter-calibration based on ray-matching between geostationary satellite sensors onboard twin satellites.
- Implementation of collocation process in the visible wavelengths and the application of weighted statistics for ensuring collocation stability.
- The novel advantages of the GEO-GEO inter-calibration approach for future applications

19 **Abstract**

20 To address the increasing demand for diurnal information on trace gases and aerosols, a series of
21 geostationary (GEO) satellite programs called GEO-constellation have been initiated, with the
22 launch of the Geostationary Environment Monitoring Spectrometer (GEMS) onboard
23 Geostationary Korea Multi-Purpose Satellite 2B (GK2B). To assess the sensor performance of
24 GEMS in orbit, the current work suggests employing an inter-calibration methodology involving
25 the Advanced Meteorological Imager (AMI) aboard its twin satellite, GK2A. Twin satellites
26 have a significant advantage in obtaining collocation datasets across diverse spatiotemporal
27 conditions, enabling rigorous collocation criteria effectively reducing mismatch uncertainty. The
28 collocation results present robust correlation coefficients over 0.98, revealing the current
29 calibration characteristics of the sensors. This research emphasizes the advantages of the GEO-
30 GEO inter-calibration, particularly the capability of analyzing spatial and temporal dependencies.
31 These findings confirm the mutual benefit of utilizing the sensors in similar configurations,
32 highlighting their importance for future satellite monitoring endeavors.

33 **Plain Language Summary**

34 Understanding the daily changes in air pollutants is crucial for grasping how these substances
35 move and disperse in the air, aiding efforts to reduce pollution. In this regard, satellites have a
36 distinct advantage in observation owing to their wide spatial coverage at regular intervals. GEMS
37 is one of the geostationary sensors providing such information for the Asia-Pacific region nearly
38 7-8 times a day. To evaluate the reliability of GEMS, this study proposes an inter-calibration
39 method by comparing GEMS observations with those of AMI aboard GK2A. These satellites
40 have a unique advantage as they fly close to each other, observing the Earth with matched optical
41 viewing paths. Scenes simultaneously observed by AMI and GEMS exhibit strong agreement,
42 thus revealing the inherent observation characteristics of each sensor. These findings confirm the
43 mutual benefit of utilizing the sensors in similar configurations for satellite monitoring during
44 the operation.

45 **1 Introduction**

46 A global network of geostationary (GEO) satellites for air quality monitoring is to be
47 established soon, including the Geostationary Environment Monitoring Spectrometer (GEMS),
48 Tropospheric Emissions: Monitoring of Pollution (TEMPO) and Sentinel-4 over East Asia,
49 Northern America, and Europe, respectively. These GEO satellite sensors measure specific areas
50 multiple times a day, which gives a significant advantage in providing diurnal information on
51 trace gases and aerosol properties (J. Kim et al., 2020; Zoogman et al., 2017). The observations
52 obtained from the GEO constellation will be crucial for monitoring the long-range transport of
53 air pollutants and changes in pollutant levels throughout the day. However, the effective
54 utilization of the measurements relies on the capability to ensure consistency in measurements
55 across sensors having varying designs, specifications, and calibration processes. This
56 underscores the need for implementing measures to monitor measurement quality throughout
57 their operational period.

58 In this regard, inter-calibration has been an effective measure for post launch calibration
59 (Chander et al., 2013). Especially for GEO sensors, ray-matching with low Earth orbit (LEO)
60 satellite sensors has been widely applied, considering that LEO sensors cover different field of
61 regards (FOR) of various GEO imagers (Doelling et al., 2016; Jiang et al., 2009; Minnis et al.,

62 2002a, 2002b). Ray-matching between GEO-GEO sensors, however, has not been fully
63 investigated as the FORs of GEO sensors cover distinct regions, like in the case of GEMS and
64 TEMPO. The characteristic has led to the development of inter-calibration for the GEO sensors
65 through intermediary means (Chander et al., 2013), such as radiative transfer models (Alsweiss
66 et al., 2015), transfer measurements (L. Wang et al., 2009), or numerical weather prediction
67 models (S. J. Lee & Ahn, 2021; Saunders et al., 2013).

68 However, there is a unique circumstance where direct GEO-GEO comparison becomes
69 feasible especially when the sub-nadir satellite longitudes of the spacecrafts are exceptionally
70 close. The onboard sensors in this condition observe Earth scenes with coincident optical
71 viewing paths, which greatly reduces collocation mismatch. Luckily, GEMS onboard the
72 Geostationary Korea Multi-Purpose Satellite-2B (GK2B) has a corresponding pair satisfying the
73 condition, the Advance Meteorological Imager (AMI) onboard its twin satellite, GK2A. It is
74 facilitated because the two satellites are positioned within a 0.05° control box centered around
75 128.2°E longitude.

76 Collocation between the sensors is straightforward and has several advantages such as:
77 (1) the huge number of collocated samples; (2) full coverage of solar and viewing angles; and (3)
78 wide spatiotemporal coverage. With the advantages, here we focus on optimizing the collocation
79 conditions and further clarify the current calibration issues of the sensors through spatial and
80 temporal analyses. If possible, GEMS and AMI could serve as useful sources for understanding
81 the observation characteristics of the sensors during their operational periods, in terms of relative
82 standards. Additionally, the GEO-GEO comparison can serve as a high-quality collocation
83 reference for GEO-LEO collocation, given the strictest collocation criteria.

84 **2 Data and Methods**

85 2.1 Sensor specification

86 2.1.1 GEMS

87 GEMS scans the Asia-Pacific region (5°S - 45°N , 75 - 145°E) in the east-west direction by
88 moving a scan mirror, while maintaining a fixed north-south field of view of about 7.78° (Choi et
89 al., 2019; J. Kim et al., 2020). GEMS is designed to measure spectral radiances ranging from 300
90 to 500 nm (Level 1B) with a spectral resolution of better than 0.6 nm. To monitor and calibrate
91 the Level 1B products, two transmissive solar diffusers and light emitting diode (LED) are
92 deployed as the on-board calibration system. Alongside this system, monitoring and calibration
93 methods have been devised and performed with various statistical approaches since the
94 completion of the in-orbit test in October 2020 (Kang et al., 2020, 2022; Y. Lee et al., 2020).

95 The assessment has revealed a significant calibration issue in GEMS solar observations,
96 specifically associated with spatial dependence along the north-south direction. The dependence
97 is attributed to the unresolved azimuth angle dependence in the bidirectional transmittance
98 distribution function (BTDF) of the solar diffusers (Kang et al., 2023). Theoretically, the
99 dependence should specifically impact solar irradiance but not Earth radiance. This distinction
100 originates from the shared optical paths within the instrument for both solar and Earth
101 observation modes, except for the solar diffusers. To confirm the cause of the dependence, it is
102 necessary to ensure that the Earth's radiance is not affected by this. However, radiances
103 observing various Earth scenes make it challenging to discern the dependence originating from

104 the BTDF. In this situation, inter-calibration with AMI emerges as a useful means to evaluate the
105 issue, given that the collocation datasets can perfectly cancel out the scene variabilities. To
106 validate this, we employ the upgraded version (provisional) of GEMS solar irradiance and
107 evaluate the BTDF correction (refer to Section 3.1) along with addressing other calibration
108 concerns.

109 2.1.2 AMI

110 AMI has six visible/near infrared and ten infrared channels for continuous atmospheric
111 monitoring. The first visible channel at a central wavelength of 470 nm (Ch01) is employed for
112 the comparison since the spectral response function (SRF) of the channel is fully encompassed
113 by the GEMS observations. To obtain surface information, an infrared channel centered at 10.5
114 μm (window channel, Ch13) is collocated together with Ch01. This study employs full-disk
115 observations taken at ten-minute intervals for Ch01 and Ch13, with spatial resolutions of 1 and 2
116 km, respectively. For visible and near-infrared channels, the National Meteorological Satellite
117 Center (NMSC) has applied multiple calibration techniques such as solar calibration with a solar
118 diffuser, vicarious calibration using the radiative transfer model (RTM) over the Australian
119 deserts, and ray-matching with the Moderate Resolution Imaging Spectroradiometer (MODIS)
120 onboard Terra and Aqua and the Visible Infrared Imaging Radiometer Suite (VIIRS) onboard
121 NOAA-20 and Suomi-National Polar-orbiting Partnership (Suomi-NPP). The calibration results
122 (D. Kim et al., 2021) as well as the related information is well organized in the following website
123 (URL: <http://210.125.45.71/enhome/html/gsics/vicariousIntroGK2A.do>).

124 2.2 Ray-matching

125 Ray-matching had been developed to relate measured information of GEO and LEO
126 sensors in early days for the Clouds and the Earth's Radiant Energy System (CERES) project
127 (Minnis and Harrison, 1984; Minnis et al., 1991). Afterwards, the concept of relating the GEO-
128 LEO measurements has been employed to calibrate GEO imagers by setting a well-calibrated
129 LEO sensor as a reference (Minnis et al., 2002a, 2008; Doelling et al., 2016; Xiong et al., 2020).
130 Targeting specific scenes such as deep convective clouds (DCC) has been further developed as it
131 can reduce the bidirectional reflectance distribution function (BRDF) effects of the clouds (Hu et
132 al., 2004; Bhatt et al., 2017). Setting a particular target contributes to effectively detecting
133 sensor-specific signals, while it reduces the number of datasets as a trade-off. To mitigate this, a
134 statistical approach has been developed especially for the DCC calibration, and it has proven
135 effective in detecting sensor degradation (Doelling et al., 2013).

136 In this study, we take advantage of ray-matching between AMI and GEMS, observing
137 spatiotemporally matched scenes, all with coinciding optical viewing paths without requiring
138 additional treatment. When empirically checking the position vectors of GK2A and GK2B over
139 the course of a year, the angles between the position vectors in the Earth-centered, Earth-fixed
140 (ECEF) coordinates vary up to 0.06° . This variation introduces viewing angle differences
141 approximately on the order of $0.01\text{-}0.1^\circ$, which is practically small compared to the current ray-
142 matching condition between GEO-LEO, mostly 1% for the viewing zenith angle (VZA)
143 differences (D. Kim et al., 2021). The advantage guarantees huge collocation datasets,
144 facilitating further data filtering and a more rigorous spatiotemporal match, as detailed in the
145 following sections.

146 2.2.1 Spatial, temporal, and spectral matching

147 Spatial averaging is applied to both sensor observations with a grid size of 0.1° . The
 148 spatial averaging can reduce random noise originated from natural variability and image
 149 navigation and registration (INR) uncertainty. In addition to the averaging, standard deviation is
 150 also derived for each grid to characterize scene inhomogeneity (Doelling et al., 2013). This
 151 inhomogeneity is closely related to collocation error and is further employed as weights in the
 152 weighted linear regression (refer to Section 2.3).

153 Temporal matching is performed by calculating the observation time difference for each
 154 grid. GEMS provides approximately 700 scan images for 30-minutes observation. During the
 155 GEMS observation, several AMI full-disk images are taken and pieced together with the
 156 temporal match threshold ($\Delta t < 5$ minutes). The time matching inherently aligns solar zenith
 157 angle (SZA) and solar azimuth angle (SAA) between the sensors without additional treatment.

158 Finally, the hyperspectral observations of GEMS are convolved with the AMI SRF for
 159 spectral matching. The convolved radiance and irradiance of GEMS are normalized by the sum
 160 of the SRF. The bidirectional reflectance is calculated after the convolution process as follows:

$$161 \quad (1) R = \frac{\pi I}{\mu_0 F}$$

162 where F , I , R and μ_0 denotes the measured solar irradiance, Earth radiance, reflectance, and the
 163 cosine of SZA for the optical path length correction.

164 2.2.2 Filtering test

165 After the collocation process, the grids consisting of land and sun-glint pixels are
 166 screened out. Land scenes have higher natural variability increasing variances in radiance biases
 167 between AMI and GEMS. Sun-glint scenes corresponding to specular reflection seem to be more
 168 vulnerable to temporal mismatches, which could add a systematic bias depending on the
 169 observation time difference. The SZA and VZA are also limited to 60° , because the grids with
 170 larger zenith angles may have high collocation error caused by longer optical path lengths
 171 (Sterckx et al., 2013). Similarly, the longitudes are also constrained to values over 100°E ,
 172 because the number of GEMS pixels within a grid box is very small ($<4-5$ pixels) in the far-
 173 western region in the GEMS FOR. The number of collocation datasets even after these filtering
 174 conditions is around 20,000-40,000 every hour, while the number varies with times and seasons.

175 2.3 Weighted linear regression

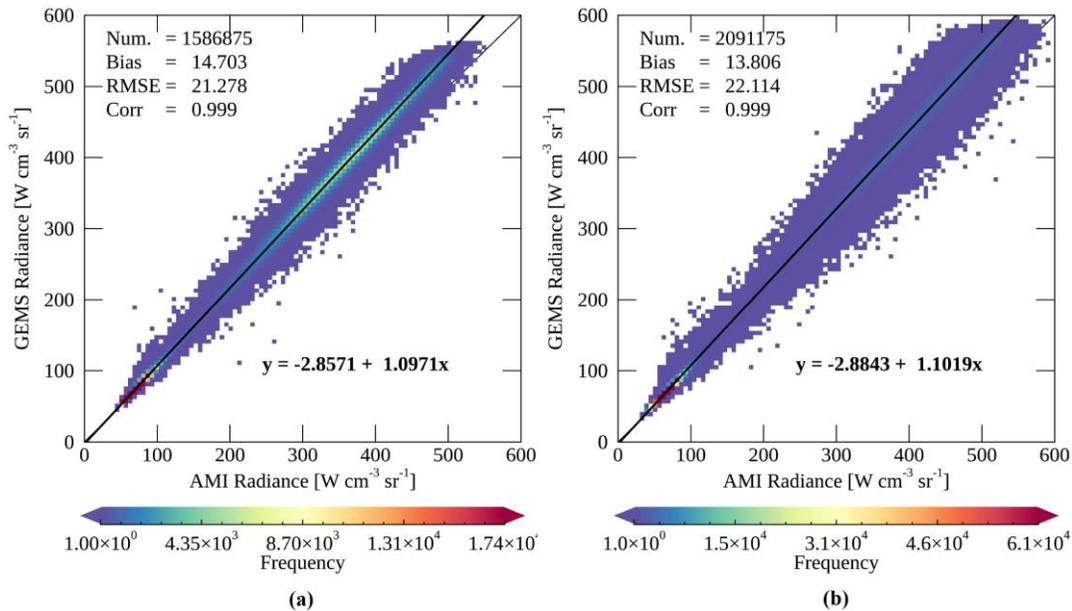
176 To obtain stable signals from the scenes ranging from dark ocean to very bright clouds,
 177 we employ weighted linear regression. The standard deviation mentioned in Section 2.2.1 serves
 178 as a weight for each collocated grid, increasing the contribution from highly homogeneous
 179 scenes in the statistics. The method has been widely applied to collect spatially homogeneous
 180 cloud tops (Doelling et al., 2013). However, the standard deviation tends to increase with
 181 brighter scenes, thereby reducing the contribution of the cloud scenes. To ensure equal
 182 contribution regardless of signal levels, the standard deviation of each grid is divided by the
 183 mean (Hewison, 2013), resulting in the relative uncertainty (v_i and w_i) as follows:

$$184 \quad (2) v_i = u(X_i)^{-1} = \left(\frac{\sigma(X_i)}{m(X_i)} \right)^{-1} \quad \text{and} \quad w_i = u(Y_i)^{-1} = \left(\frac{\sigma(Y_i)}{m(Y_i)} \right)^{-1}$$

185 where σ and m indicates the standard deviation and the mean of reflectances for each
 186 grid, respectively. The subscript i indicates the grid index in the collocation datasets and X_i and
 187 Y_i are the AMI and GEMS observations within the i -th grid. The weighted linear regression
 188 shows more stable results when accounting for the uncertainties associated with X_i and Y_i
 189 together. Empirically, the weighted regression presents a stable trend compared to non-weighted
 190 regression especially when the interval becomes shorter such as on a daily or hourly basis. The
 191 applied regression method with the weights is the generalized distance regression (GDR)
 192 introduced by the following document (ISO/TS 28037, 2010).

193 3 Results and Discussion

194 AMI and GEMS observations are collected over two years from November 2020 to May
 195 2023 every hour during daytime. The datasets presented in Section 3.1 particularly undergo
 196 screening based on the spatial inhomogeneity condition ($u(X_i) < 5\%$ and $u(Y_i) < 5\%$). Figure 1
 197 presents the collocation datasets measured in January and July 2021, showing the regression
 198 slopes larger than unity. The positive bias of GEMS will be discussed in detail in Section 3.2.1.
 199 The correlation presents a good agreement with the coefficients over 0.98 regardless of
 200 observation times and seasons. The GEMS radiances over $592.35 \text{ W cm}^{-3} \text{ sr}^{-1}$ are affected by
 201 saturation, resulting in the limiting feature for higher signals. Even after applying the scene
 202 homogeneity condition, the number of collocation datasets exceeds 1,500,000 every month. With
 203 the datasets, the following sections will present spatial and temporal analysis results between
 204 AMI and GEMS.

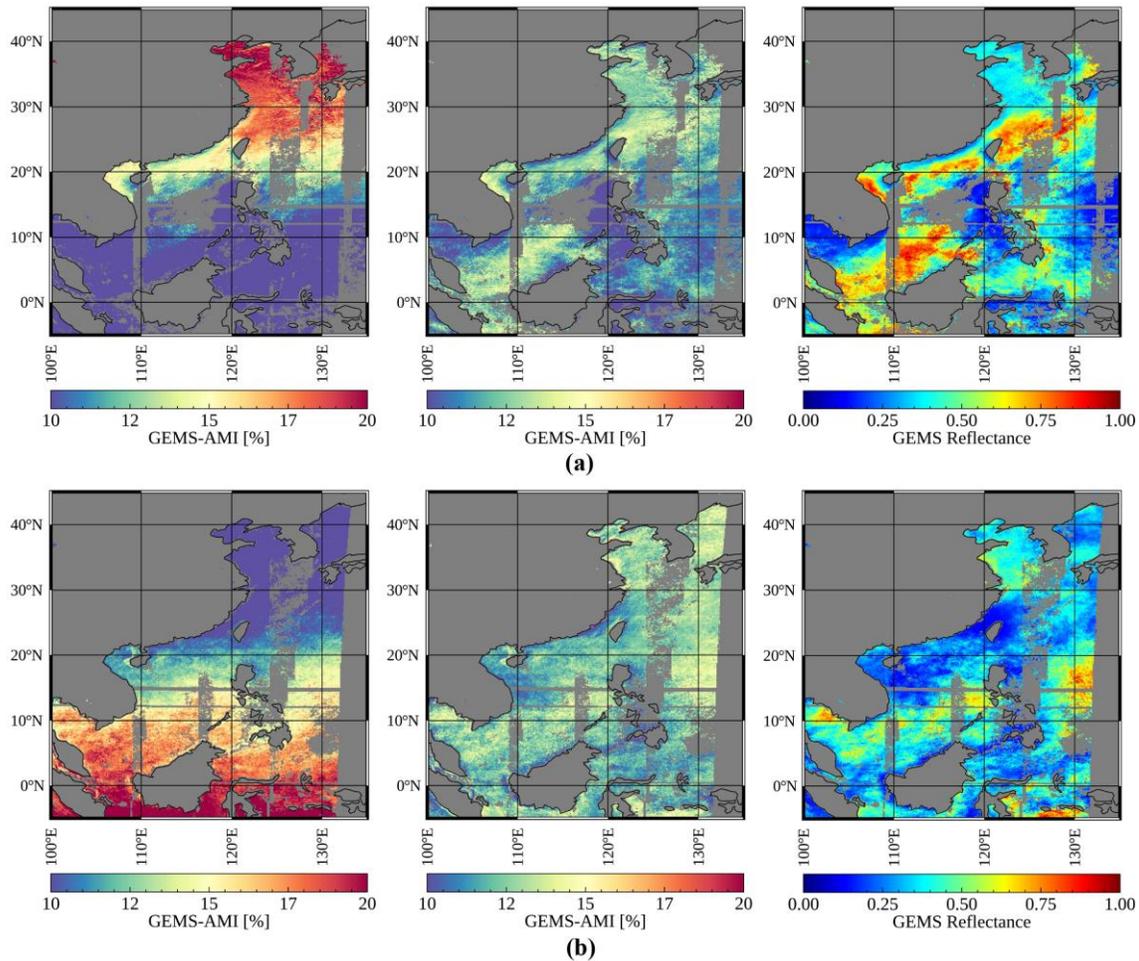


205 **Figure 1.** Scatter density plots of AMI and GEMS radiances measured in (a) January and (b) July 2021 (01-06
 206 UTC) with statistics based on the weighted linear regression. The color bar indicates the number of collocation
 207 datasets after binning with the bin size of $1 \text{ W cm}^{-3} \text{ sr}^{-1}$.
 208

209 3.1 Spatial analysis

210 3.1.1 North-south spatial dependence

211 As mentioned earlier, a primary issue regarding GEMS Level 1B products was the north-
 212 south spatial dependence in solar irradiance, which is expected to show up in Earth reflectance.
 213 To verify this, mean biases of reflectance between AMI and GEMS are calculated with the
 214 collocation datasets measured in January and July 2021. Invalid grids including bad pixels
 215 around 10-15 °N latitudes are filtered out during the collocation process (Y. Lee et al., 2023).
 216 The north-south spatial dependence has seasonal variation, and it is clearly shown in Figure 2
 217 (the first column) with the reversed pattern for different seasons. The BTDF correction updated
 218 by Kang et al. (2023) has greatly improved GEMS solar irradiance and, consequently, Earth
 219 reflectance. The second column in Figure 2 demonstrates the improvement with consistent trends
 220 across all latitudes. The update has effectively removed the north-south dependence, reducing
 221 biases in solar irradiance from approximately 20% to within 4%. Regarding reflectance, the
 222 biases converge to 15% across all latitudes.



223 **Figure 2.** The mean biases in percentage between AMI and GEMS reflectances before (the first column) and after
 224 (the second column) the GEMS BTDF update. The third column presents mean reflectances of GEMS for each grid.
 225 The collocation datasets are the observations in (a) January and (b) July 2021 (01-06 UTC).
 226

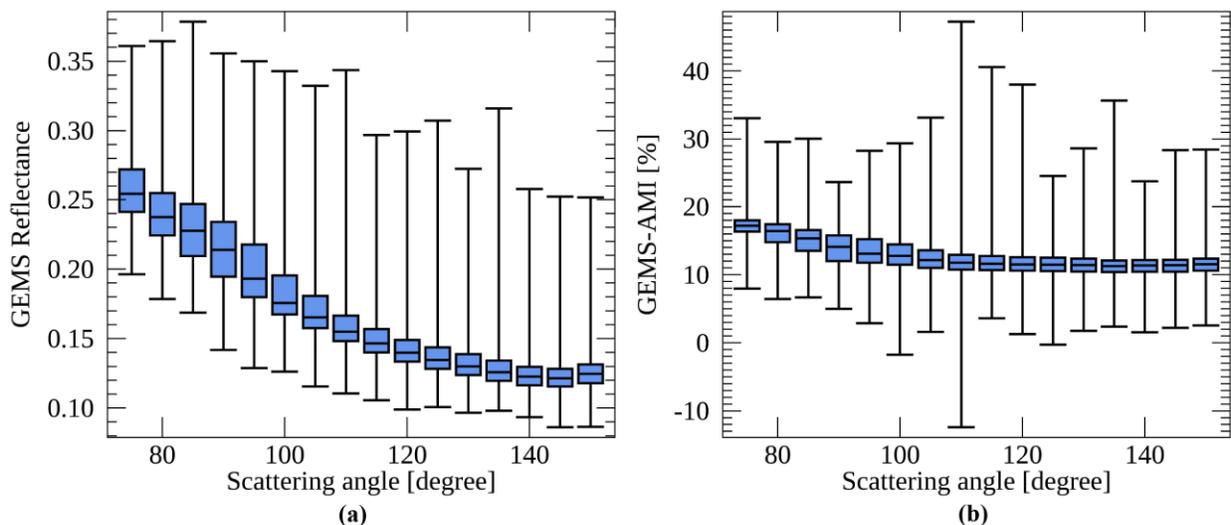
227 3.1.2 Signal dependence

228 Another finding in Figure 2 is the appearance of signal dependence after removing the
 229 north-south spatial dependence. Particularly for lower reflectances (below 0.4), the positive
 230 biases rapidly increase, as the signal levels increase. The dependence associated with lower
 231 reflectances is further investigated with the scattering angle as follows:

$$232 \quad (3) \cos\Theta = -\cos\theta_0\cos\theta - \sin\theta_0\sin\theta\cos(\phi_0 - \phi)$$

233 where Θ , θ_0 , θ , ϕ_0 and ϕ represent the scattering angle, SZA, VZA, SAA, and viewing azimuth
 234 angle (VAA), respectively. In the visible spectral domain, Rayleigh scattering predominantly
 235 accounts for about 80% of the top of the atmosphere (TOA) reflectances, particularly over dark
 236 ocean (Sterckx et al., 2013; M. Wang, 2016). The scattering angle computed in Equation (3)
 237 determines the intensity of Rayleigh scattering, as the input parameter for the Rayleigh scattering
 238 phase function. The intensity generally increases when the direction approaches forward or
 239 backward scattering, corresponding to the angles of 0° or 180° , respectively.

240 Figure 3 depicts the scattering angle dependence for the collocation datasets having AMI
 241 reflectance lower than 0.3 and the brightness temperatures (Ch13) greater than 280 K. The
 242 selected datasets are grouped by scattering angles with the angle interval of 5° . In Figure 3a, the
 243 dependence on scattering angle is evidently clear, with the reflectance distributions decreasing
 244 for higher scattering angles. This is because the light observed at the TOA has been more
 245 scattered within the atmosphere under backscattering conditions. It should be noted that the
 246 effects of reflected light from ocean surfaces may exist, albeit with a smaller impact. The
 247 distributions having lower scattering angles exhibit higher mean biases as shown in Figure 3b,
 248 presenting consistent results with the signal dependence. This indicates that the signal and
 249 scattering angle dependencies might be intertwined, particularly for lower radiances. The
 250 potential cause of these dependencies is still unclear, but it needs to be clarified in further
 251 research since the scattering angle dependence could be mis-interpreted as temporal dependence.
 252 In other words, variations in scattering angles, influenced by the solar position, could lead to
 253 seasonal or diurnal fluctuations.
 254



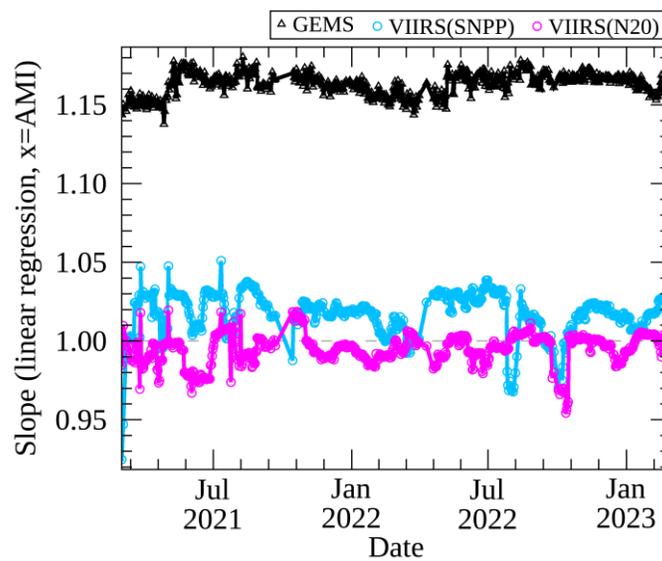
255 **Figure 3.** The box plots for dark ocean scenes of (a) GEMS reflectances and (b) the biases with AMI grouped by
 256 scattering angles with a class width of 5° .
 257

258 3.2 Temporal analysis

259 The previous section analyzes dependencies based on geolocation variables, as
 260 collocation datasets cover a wide range of spatiotemporal conditions. This section focuses on
 261 temporal aspects of the collocation datasets collected at various observation intervals. For
 262 temporal analysis, the weighted regression is used due to the stable trends exhibited by the
 263 indicator.

264 3.2.1 Radiometric calibration

265 Figure 4 presents the regression slopes of AMI and GEMS, compared with the slopes
 266 between AMI and VIIRS (GEO-LEO) onboard Suomi-NPP and NOAA20 provided by NMSC.
 267 The results show that the regression slopes between AMI and GEMS are more stable compared
 268 to the GEO-LEO ray-matching results. As VIIRS is a LEO sensor, the collocation thresholds
 269 need to be more relaxed, which leads to a wider range of regression slopes over time.
 270 Nevertheless, the GEO-LEO regression slopes are closer to unity, which confirms that GEMS
 271 has a positive bias of approximately 15% in its signals especially beyond 450 nm, necessitating a
 272 correction in its radiometric calibration coefficients. Most GEMS Level 2 products utilize
 273 reflective spectral features for retrieval, such as employing differential optical absorption
 274 spectroscopy (DOAS) (Platt & Stutz, 2008). While overall systematic bias across spectra may
 275 have minimal impact on retrieval under such approaches, it still has the potential to influence
 276 specific retrieval processes relying on theoretical irradiance for radiance scaling (Cho et al.,
 277 2023).
 278



279 **Figure 4.** The regression slopes of reflectances derived from AMI with GEMS (black triangles) and VIIRS onboard
 280 Suomi-NPP and NOAA-20 (blue and magenta circles, respectively). The data is collected from March 2021 to
 281 February 2023 at daily intervals.
 282

283 3.2.2 Diurnal variation

284 The GEO-GEO inter-calibration offers unique insights into diurnal variation of measured
 285 signals, which are not as readily available through the GEO-LEO comparison. Table 1 presents
 286 the statistics of AMI and GEMS radiances measured in 2021 at different observation times (01-
 287 06 UTC) with correlation coefficient, regression slope, bias, and root mean square error (RMSE).

288 There is a decreasing trend in both regression slope and bias, which is similarly observed in
 289 reflectance (not shown) across different seasons and surface types (land or ocean). This indicates
 290 there is a consistent diurnal influence in the comparison results of AMI and GEMS. Although the
 291 cause of this temporal dependence remains unclear, it may stem from various factors, including
 292 angle dependence as discussed in Section 3.1.2. The results emphasize the potential of the GEO-
 293 GEO inter-calibration in enhancing our understanding of temporal fluctuations, thereby possibly
 294 refining inter-calibration methodologies in future research.

295

296 **Table 1.** Statistics of AMI and GEMS Radiances Measured in 2021 at Different observation Times Ranging from 1
 297 to 6 UTC

UTC	Correlation coefficient	Slope	Bias [W cm ⁻³ sr ⁻¹]	RMSE [W cm ⁻³ sr ⁻¹]
1 (10 KST)	0.987	1.107	13.3	25.8
2 (11 KST)	0.984	1.104	13.6	27.4
3 (12 KST)	0.981	1.101	13.6	28.6
4 (13 KST)	0.982	1.099	13.3	27.1
5 (14 KST)	0.985	1.099	12.4	25.8
6 (15 KST)	0.984	1.099	10.8	22.9

298 4 Conclusions

299 This research introduces an inter-calibration method, specifically utilizing ray-matching,
 300 to compare and monitor Level 1B products between instruments onboard twin satellites,
 301 GEMS/GK2B and AMI/GK2A. The close alignment of their sub-nadir positions offers a
 302 practical advantage, enabling matched optical viewing paths and generating extensive
 303 collocation datasets that cover a broad range of observation conditions.

304 With the datasets over two years, we conducted comprehensive spatial and temporal
 305 analyses between AMI and GEMS. The spatial analysis successfully addressed the calibration
 306 update of GEMS for solar irradiance and reflectance, though signal and scattering angle
 307 dependencies underscore the necessity for further analysis. Temporal analysis highlighted sensor
 308 drifts in regression slopes and radiometric calibration uncertainties in GEMS while emphasizing
 309 the significance of diurnal variation studies.

310 Our findings emphasize the critical role of GEO-GEO inter-calibration in enhancing our
 311 understanding of the measurement characteristics. Future investigations should focus on
 312 identifying the root causes of observed dependencies and biases, thereby advancing the
 313 effectiveness of inter-calibration techniques. This study aims to inform and facilitate future
 314 research endeavors, potentially aiding in the monitoring of sensors in similar configurations,
 315 such as the Flexible Combined Imager (FCI) and Sentinel-4.

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320

321 Open Research

322 The AMI Level 1B products are available at
 323 <https://datasvc.nmsc.kma.go.kr/datasvc/html/main/main.do?lang=en>. The AMI SRFs for all
 324 channels and land sea mask are also available at
 325 <https://datasvc.nmsc.kma.go.kr/datasvc/html/base/cmm/selectPage.do?page=static.software>. The
 326 GSICS inter-calibration coefficients between AMI and VIIRS were provided by NMSC, and the
 327 coefficients can be reproduced by applying the methodology described in D. Kim et al. (2021).
 328 The GEMS Level 1C products can be accessed through the SFTP service provided by the
 329 Environmental Satellite Center (ESC) of the National Institute of Environmental Research
 330 (NIER) (<https://nesc.nier.go.kr/en/html/cntnts/91/static/page.do>), following approval from the
 331 institute. The datasets are not publicly available so far due to the regulation of the institution for
 332 the Level 1C products, and the datasets are accessible to researchers only volunteering on the
 333 GEMS calibration and validation project. The MATLAB code for the weighted regression
 334 algorithm is freely available at <https://www.npl.co.uk/resources/software/iso-ts-28037-2010e>.
 335

336 References

- 337 Alsweiss, S. O., Jelenak, Z., Chang, P. S., Park, J. D., & Meyers, P. (2015). Inter-calibration Results of the Advanced
 338 Microwave Scanning Radiometer-2 Over Ocean. *IEEE Journal of Selected Topics in Applied Earth
 339 Observations and Remote Sensing*, 8(9), 4230–4238. <https://doi.org/10.1109/JSTARS.2014.2330980>
- 340 Chander, G., Hewison, T. J., Fox, N., Wu, X., Xiong, X., & Blackwell, W. J. (2013). Overview of intercalibration of
 341 satellite instruments. *IEEE Transactions on Geoscience and Remote Sensing*, 51(3), 1056–1080.
 342 <https://doi.org/10.1109/TGRS.2012.2228654>
- 343 Cho, Y., Kim, J., Go, S., Kim, M., Lee, S., Kim, M., et al. (2023). First Atmospheric Aerosol Monitoring Results
 344 from Geostationary Environment Monitoring Spectrometer (GEMS) over Asia. *Atmospheric Measurement
 345 Techniques Discussions*, 2023, 1–29. <https://doi.org/10.5194/amt-2023-221>
- 346 Choi, W. J., Moon, K.-J., Yoon, J., Cho, A., Kim, S., Lee, S., et al. (2019). Introducing the geostationary environment
 347 monitoring spectrometer. *Journal of Applied Remote Sensing*, 13(01), 1.
 348 <https://doi.org/10.1117/1.jrs.13.019901>
- 349 Doelling, D. R., Morstad, D., Scarino, B. R., Bhatt, R., & Gopalan, A. (2013). The Characterization of Deep
 350 Convective Clouds as an Invariant Calibration Target and as a Visible Calibration Technique. *IEEE
 351 Transactions on Geoscience and Remote Sensing*, 51(3), 1147–1159.
 352 <https://doi.org/10.1109/TGRS.2012.2225066>
- 353 Doelling, D. R., Haney, C. O., Scarino, B. R., Gopalan, A., & Bhatt, R. (2016). Improvements to the Geostationary
 354 Visible Imager Ray-Matching Calibration Algorithm for CERES Edition 4. *Journal of Atmospheric and
 355 Oceanic Technology*, 33(12), 2679–2698. <https://doi.org/10.1175/JTECH-D-16-0113.1>
- 356 Hewison, T. J. (2013). An Evaluation of the Uncertainty of the GSICS SEVIRI-IASI Intercalibration Products. *IEEE
 357 TRANSACTIONS ON GEOSCIENCE AND REMOTE SENSING*, 51(3).
 358 <https://doi.org/10.1109/TGRS.2012.2236330>
- 359 Hu, Y., Wielicki, B. A., Yang, P., Stackhouse, P. W., Lin, B., & Young, D. F. (2004). Application of Deep
 360 Convective Cloud Albedo Observation to Satellite-Based Study of the Terrestrial Atmosphere: Monitoring the
 361 Stability of Spaceborne Measurements and Assessing Absorption Anomaly. *IEEE TRANSACTIONS ON
 362 GEOSCIENCE AND REMOTE SENSING*, 42(11). <https://doi.org/10.1109/TGRS.2004.834765>
- 363 International Organization for Standardization Technical Specification (ISO/TS). (2010). ISO 28037:2010
 364 Determination and Use of Straight-line Calibration Functions. Geneva, Switzerland: International Organization
 365 for Standardization (ISO). Retrieved from <https://www.iso.org/standard/44473.html>
- 366 Jiang, G.-M., Yan, H., & Ma, L.-L. (2009). Intercalibration of SVISSR/FY-2C Infrared Channels Against
 367 MODIS/Terra and AIRS/Aqua Channels; Intercalibration of SVISSR/FY-2C Infrared Channels Against
 368 MODIS/Terra and AIRS/Aqua Channels. *IEEE Transactions on Geoscience and Remote Sensing*, 47(5).
 369 <https://doi.org/10.1109/TGRS.2008.2005200>

- 370 Kang, M., Ahn, M. H., Liu, X., Jeong, U., & Kim, J. (2020). Spectral calibration algorithm for the geostationary
371 environment monitoring spectrometer (Gems). *Remote Sensing*, 12(17), 1–17.
372 <https://doi.org/10.3390/rs12172846>
- 373 Kang, M., Ahn, M. H., Ko, D. H., Kim, J., Nicks, D., Eo, M., et al. (2022). Characteristics of the Spectral Response
374 Function of Geostationary Environment Monitoring Spectrometer Analyzed by Ground and In-Orbit
375 Measurements. *IEEE Transactions on Geoscience and Remote Sensing*, 60.
376 <https://doi.org/10.1109/TGRS.2021.3091677>
- 377 Kang, M., Ahn, M.-H., Lee, Y., Eo, M., & Kim, J. (2023, October 25). GEMS Performance and Lessons Learned
378 [PowerPoint slides]. Committee on Earth Observation Satellites (CEOS) AC-VC-19/ACSG Joint Meeting.
379 Retrieved from [https://ceos.org/document_management/Virtual_Constellations/AC-VC/Meetings/AC-VC-](https://ceos.org/document_management/Virtual_Constellations/AC-VC/Meetings/AC-VC-19/presentations/Presentations%20PDF/2.%20Wed%2025%20Oct%202023%20-%20TRACE%20GASES%20AND%20AEROSOLS%20AIR%20QUALITY/We-14_Kang_GEMS_v1.pdf)
380 [19/presentations/Presentations%20PDF/2.%20Wed%2025%20Oct%202023%20-](https://ceos.org/document_management/Virtual_Constellations/AC-VC/Meetings/AC-VC-19/presentations/Presentations%20PDF/2.%20Wed%2025%20Oct%202023%20-%20TRACE%20GASES%20AND%20AEROSOLS%20AIR%20QUALITY/We-14_Kang_GEMS_v1.pdf)
381 [%20TRACE%20GASES%20AND%20AEROSOLS%20AIR%20QUALITY/We-14_Kang_GEMS_v1.pdf](https://ceos.org/document_management/Virtual_Constellations/AC-VC/Meetings/AC-VC-19/presentations/Presentations%20PDF/2.%20Wed%2025%20Oct%202023%20-%20TRACE%20GASES%20AND%20AEROSOLS%20AIR%20QUALITY/We-14_Kang_GEMS_v1.pdf)
382 March 21, 2024
- 383 Kim, D., Gu, M., Oh, T.-H., Kim, E.-K., & Yang, H.-J. (2021). Introduction of the Advanced Meteorological Imager
384 of Geo-Kompsat-2a: In-Orbit Tests and Performance Validation. *Remote Sensing*, 13(7), 1303.
385 <https://doi.org/10.3390/rs13071303>
- 386 Kim, J., Jeong, U., Ahn, M. H., Kim, J. H., Park, R. J., Lee, H., et al. (2020). New era of air quality monitoring from
387 space: Geostationary environment monitoring spectrometer (GEMS). *Bulletin of the American Meteorological*
388 *Society*, 101(1), E1–E22. <https://doi.org/10.1175/BAMS-D-18-0013.1>
- 389 Lee, S. J., & Ahn, M.-H. (2021). Synergistic Benefits of Intercomparison Between Simulated and Measured
390 Radiances of Imagers Onboard Geostationary Satellites. *IEEE Transactions on Geoscience and Remote*
391 *Sensing*, 59(12), 10725–10737. <https://doi.org/10.1109/TGRS.2021.3054030>
- 392 Lee, Y., Ahn, M. H., & Kang, M. (2020). The new potential of deep convective clouds as a calibration target for a
393 geostationary UV/VIS hyperspectral spectrometer. *Remote Sensing*, 12(3). <https://doi.org/10.3390/rs12030446>
- 394 Lee, Y., Ahn, M.-H., Kang, M., & Eo, M. (2023). Spectral replacement using machine learning methods for
395 continuous mapping of the Geostationary Environment Monitoring Spectrometer (GEMS). *Atmospheric*
396 *Measurement Techniques*, 16(1), 153–168. <https://doi.org/10.5194/AMT-16-153-2023>
- 397 Minnis, P., & Harrison, E. F. (1984). Diurnal Variability of Regional Cloud and Clear-Sky Radiative Parameters
398 Derived from GOES Data. Part III: November 1978 Radiative Parameters. *Journal of Applied Meteorology and*
399 *Climatology*, 23(7), 1032–1051. [https://doi.org/10.1175/1520-0450\(1984\)023](https://doi.org/10.1175/1520-0450(1984)023)
- 400 Minnis, P., Young, D. F., & Harrison, E. F. (1991). Examination of the Relationship between Outgoing Infrared
401 Window and Total Longwave Fluxes Using Satellite Data. *Journal of Climate*, 4(11), 1114–1133.
402 [https://doi.org/10.1175/1520-0442\(1991\)004<1114:EOTRBO>2.0.CO;2](https://doi.org/10.1175/1520-0442(1991)004<1114:EOTRBO>2.0.CO;2)
- 403 Minnis, P., Nguyen, L., Doelling, D. R., Young, D. F., Miller, W. F., & Kratz, D. P. (2002a). Rapid Calibration of
404 Operational and Research Meteorological Satellite Imagers. Part I: Evaluation of Research Satellite Visible
405 Channels as References. *Journal of Atmospheric and Oceanic Technology*, 19(9), 1233–1249.
406 [https://doi.org/10.1175/1520-0426\(2002\)019<1233:RCOOAR>2.0.CO;2](https://doi.org/10.1175/1520-0426(2002)019<1233:RCOOAR>2.0.CO;2)
- 407 Minnis, P., Nguyen, L., Doelling, D. R., Young, D. F., Miller, W. F., & Kratz, D. P. (2002b). Rapid Calibration of
408 Operational and Research Meteorological Satellite Imagers. Part II: Comparison of Infrared Channels. *Journal*
409 *of Atmospheric and Oceanic Technology*, 19(9), 1250–1266. [https://doi.org/10.1175/1520-](https://doi.org/10.1175/1520-0426(2002)019<1250:RCOOAR>2.0.CO;2)
410 [0426\(2002\)019<1250:RCOOAR>2.0.CO;2](https://doi.org/10.1175/1520-0426(2002)019<1250:RCOOAR>2.0.CO;2)
- 411 Minnis, P., Doelling, D. R., Nguyen, L., Miller, W. F., & Chakrapani, V. (2008). Assessment of the Visible Channel
412 Calibrations of the VIRS on TRMM and MODIS on Aqua and Terra. *Journal of Atmospheric and Oceanic*
413 *Technology*, 25(3), 385–400. <https://doi.org/10.1175/2007JTECHA1021.1>
- 414 National Physical Laboratory. (2014, November 17). Determination and use of straight-line calibration functions
415 [Software]. NPL Resources. Retrieved from <https://www.npl.co.uk/resources/software/iso-ts-28037-2010e>.
- 416 Platt, U., & Stutz, J. (2008). Differential Absorption Spectroscopy. In *Differential Optical Absorption Spectroscopy:*
417 *Principles and Applications* (pp. 135–174). Berlin, Heidelberg: Springer. [https://doi.org/10.1007/978-3-540-](https://doi.org/10.1007/978-3-540-75776-4_6)
418 [75776-4_6](https://doi.org/10.1007/978-3-540-75776-4_6)
- 419 Saunders, R. W., Blackmore, T. A., Candy, B., Francis, P. N., & Hewison, T. J. (2013). Monitoring Satellite
420 Radiance Biases Using NWP Models. *IEEE Transactions on Geoscience and Remote Sensing*, 51(3), 1124–
421 1138. <https://doi.org/10.1109/TGRS.2012.2229283>
- 422 Sohn, B.-J., Ham, S.-H., & Yang, P. (2009). Possibility of the Visible-Channel Calibration Using Deep Convective
423 Clouds Overshooting the TTL. *Journal of Applied Meteorology and Climatology*, 48(11), 2271–2283.
424 <https://doi.org/10.1175/2009JAMC2197.1>

- 425 Sterckx, S., Livens, S., & Adriaensen, S. (2013). Rayleigh, Deep Convective Clouds, and Cross-Sensor Desert
426 Vicarious Calibration Validation for the PROBA-V Mission. *IEEE Transactions on Geoscience and Remote*
427 *Sensing*, 51(3), 1437–1452. <https://doi.org/10.1109/TGRS.2012.2236682>
- 428 Wang, L., Cao, C., & Goldberg, M. (2009). Intercalibration of GOES-11 and GOES-12 Water Vapor Channels with
429 MetOp IASI Hyperspectral Measurements. *Journal of Atmospheric and Oceanic Technology*, 26(9), 1843–
430 1855. <https://doi.org/10.1175/2009JTECHA1233.1>
- 431 Wang, M. (2016). Rayleigh radiance computations for satellite remote sensing: accounting for the effect of sensor
432 spectral response function. *Opt. Express*, 24(11), 12414–12429. <https://doi.org/10.1364/OE.24.012414>
- 433 Xiong, X., Angal, A., Chang, T., Chiang, K., Lei, N., Li, Y., et al. (2020). MODIS and VIIRS Calibration and
434 Characterization in Support of Producing Long-Term High-Quality Data Products. *Remote Sensing 2020*, Vol.
435 12, Page 3167, 12(19), 3167. <https://doi.org/10.3390/RS12193167>
- 436 Zoogman, P., Liu, X., Suleiman, R. M., Pennington, W. F., Flittner, D. E., Al-Saadi, J. A., et al. (2017). Tropospheric
437 emissions: Monitoring of pollution (TEMPO). *Journal of Quantitative Spectroscopy and Radiative Transfer*,
438 186, 17–39. <https://doi.org/https://doi.org/10.1016/j.jqsrt.2016.05.008>