

1 **Physical-chemical properties of non-soluble particles in a**
2 **hailstone collected in Argentina**

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9 **Key Points:**

- 10 • Hailstones near Cordoba, Argentina, show diverse carbonaceous particles, with the
11 largest particles located at center of the hailstone.
12 • Silicate and carbonaceous particles coexist within the same hailstone layers, orig-
13 inating from soil and agricultural sources.
14 • Anthropogenic activities contribute to heavy metals in hailstones, emphasizing the
15 need to consider urban influences in aerosol modeling.

Abstract

This study presents a novel analysis of a hailstone collected near Cordoba, Argentina, quantifying the composition, size distribution, and potential sources of non-soluble particles contained within. The hailstone contained diverse particles, with sizes ranging from 1.9 to 150.3 μm , primarily carbonaceous, including in the center, suggesting a possible biological and geological influence on hail formation. Silicate particles were distributed throughout the hailstone, likely from eroded soil and agricultural activities. Finally, salts were detected in the outer layers of the hailstone and may have originated from the nearby salt lake. This study highlights the regional influence of various land use types on hail formation and growth and points to the potential impacts of natural and anthropogenic factors on hailstone composition.

Plain Language Summary

In this study, scientists investigated hailstones collected in Argentina to understand the particles trapped inside them and their potential impact on hail formation. Hailstorms cause significant damage and economic losses, and understanding how hailstones form is crucial. A unique method was used in this project to analyze the hailstones without melting them, revealing the size, composition, and origin of particles within an individual hailstone.

The hailstone contained various types of particles, mostly carbon-based, possibly from biological and geological sources. Silicate particles, originating from eroded soil and agricultural activities, were also present, suggesting that local winds carried these particles into the clouds where hail formed.

The research highlights the importance of local environmental factors, such as land use, in influencing the composition of hailstones. The findings contribute to our understanding of the complex processes involved in hail formation and shed light on the potential impacts of both natural and human-related factors on hailstone composition. This study opens the door for further research on hailstones collected in different environmental conditions, providing valuable insights for future studies and potential applications in weather forecasting and risk assessment.

1 Introduction

Falling hailstones are destructive natural phenomena contributing to billion-dollar disasters in the U.S. (Changnon, 2008; Sander et al., 2013; Allen et al., 2017; Kumjian et al., 2019) and significant agricultural losses globally (Calori et al., 2016; Streifeneder et al., 2023). Understanding hail formation, including its environmental controls and variations within storm modes, has been limited owing to challenges in validating remotely sensed proxies (Cecil & Blankenship, 2012; Bang & Cecil, 2019; Bruick et al., 2019) and studying hail growth globally (e.g., Allen et al. 2020). Aerosols ingested into convective cloud updrafts are known to serve as hail embryos via heterogeneous nucleation, either through frozen cloud drops initially forming on cloud condensation nuclei (CCN) or rimed ice crystals forming on non-soluble ice nucleating particles (INP); however, the effects of CCN and INP on hail formation remain largely inconclusive from modeling studies (Fan et al., 2013; Lebo & Morrison, 2014).

Previous research has used collected hailstones to examine the role of environmental aerosols, CCN, and INPs in hail formation and found links to local land use. For example, biological ice nuclei were found in hailstone embryos in the U.S. Rocky Mountains (Michaud et al., 2014), while hailstones collected in Slovenia (Šantl-Temkiv et al., 2013) and the triple border region of Paraguay, Brazil, and Argentina (Beal et al., 2021) noted signatures of the respective regions' soil. Others point to anthropogenic markers

64 through the presence of microplastics (Kozjek et al., 2023), highlighting the implications
65 of human activity on hail formation. The analysis techniques in those studies all required
66 melting the hailstones, removing information on particle size distribution or composi-
67 tion with respect to the hailstone embryo, necessarily neglecting non-soluble particles,
68 or both. This present work, using a new method (Bernal Ayala, Rowe, Arena, Nachlas,
69 & Asar, 2024) to address several limitations of previous work, aims to gain a more com-
70 prehensive understanding of hailstones’ composition and inferred microphysical processes
71 through analysis of non-soluble particles contained within hailstones.

72 This study analyzes a 4 *cm* hailstone collected near Cordoba, Argentina (Fig. 1;
73 lime green star), an area known for its intense hail-producing storms (Zipser et al., 2006;
74 Cecil & Blankenship, 2012; Rasmussen et al., 2014; Bernal Ayala et al., 2022) and the
75 focus of recent collaborative field campaigns: the 2018-19 Cloud, Aerosol, and Complex
76 Terrain Interactions (CACTI) (Varble, 2021) and the 2018 Remote Sensing of Electrifi-
77 cation, Lightning, and Mesoscale/Microscale Processes with Adaptive Ground Obser-
78 vations (RELAMPAGO) (Nesbitt et al., 2021). A survey of INPs collected at the sur-
79 face over the Sierras de Córdoba (SDC; Fig. 1) during CACTI (Testa et al., 2021) found
80 that diverse plant communities in the region release high amounts of biological particles
81 and intensively farmed areas contribute organic soil dust, indicating their potential role
82 in hail formation via INPs. The hailstone in this analysis formed from an isolated su-
83 percell that developed at 1700 UTC on 8 February 2018 on the northern section of the
84 SDC (red star in Fig. 1) that did not feature large-scale environmental conditions typi-
85 cal of widespread convection in this region (Rasmussen & Houze, 2016), thus suggest-
86 ing a local, regional influence on storm formation (Bernal Ayala et al., 2022). We, there-
87 fore, hypothesize a contribution from both biological and soil-derived particulates to hail
88 formation and growth in this event. This study is the first to characterize the elemen-
89 tal composition of particles in their original spatial context within the hailstone interior.
90 Using a novel application of microscopy-based methods for hailstone analysis (Bernal Ay-
91 ala, Rowe, Arena, Nachlas, & Asar, 2024), this study addresses the following questions:
92 1) What is the particle size and spatial distribution of non-soluble particles trapped within
93 the hailstone sample? 2) What are the major element compositions of the non-soluble
94 particles from different regions of the hailstone? 3) Which land use regions are poten-
95 tial sources of the non-soluble particles identified in the hailstone?

96 2 Data and Methods

97 2.1 Hailstone Collection and Preparation

98 The 8 February hailstone was collected by residents of Villa Carlos Paz (VCP; lime
99 green star, Fig.1) through contact with Dr. Lucia Arena at the Facultad de Matemática,
100 Astronomía, Física y Computación at the Universidad Nacional de Córdoba “FAMAF-
101 UNC” and the COSECHEROS Program (*Cosecheros de granizo Córdoba*, 2022). Col-
102 lected hailstones were processed at the FAMAF-UNC’s subzero facility, Laura Levi At-
103 mospheric Physics Laboratory. Once cut, hailstones were sublimated and coated with
104 a specific polyvinyl formal (FORMVAR) solution diluted in ethylene dichloride to trap
105 the particles within the hailstone layers. More details on the hailstone collection and prepa-
106 ration are provided in Bernal Ayala, Rowe, Arena, Nachlas, and Asar (2024).

107 2.2 Microscopy analysis and elemental characterization

108 The unique microscopy technique for analysis is detailed in Bernal Ayala, Rowe,
109 Arena, Nachlas, and Asar (2024) and summarized here. The technique is based on the
110 method of separating insoluble particles by adapted sublimation (Arena, 2023), which
111 involves covering a thin sheet of hail, approximately 1*mm* thick, with a 1% FORMVAR
112 solution to promote sublimation at below-zero temperatures to capture particles beneath
113 the plastic film. Once the ice has sublimated, the particles can be studied at room tem-

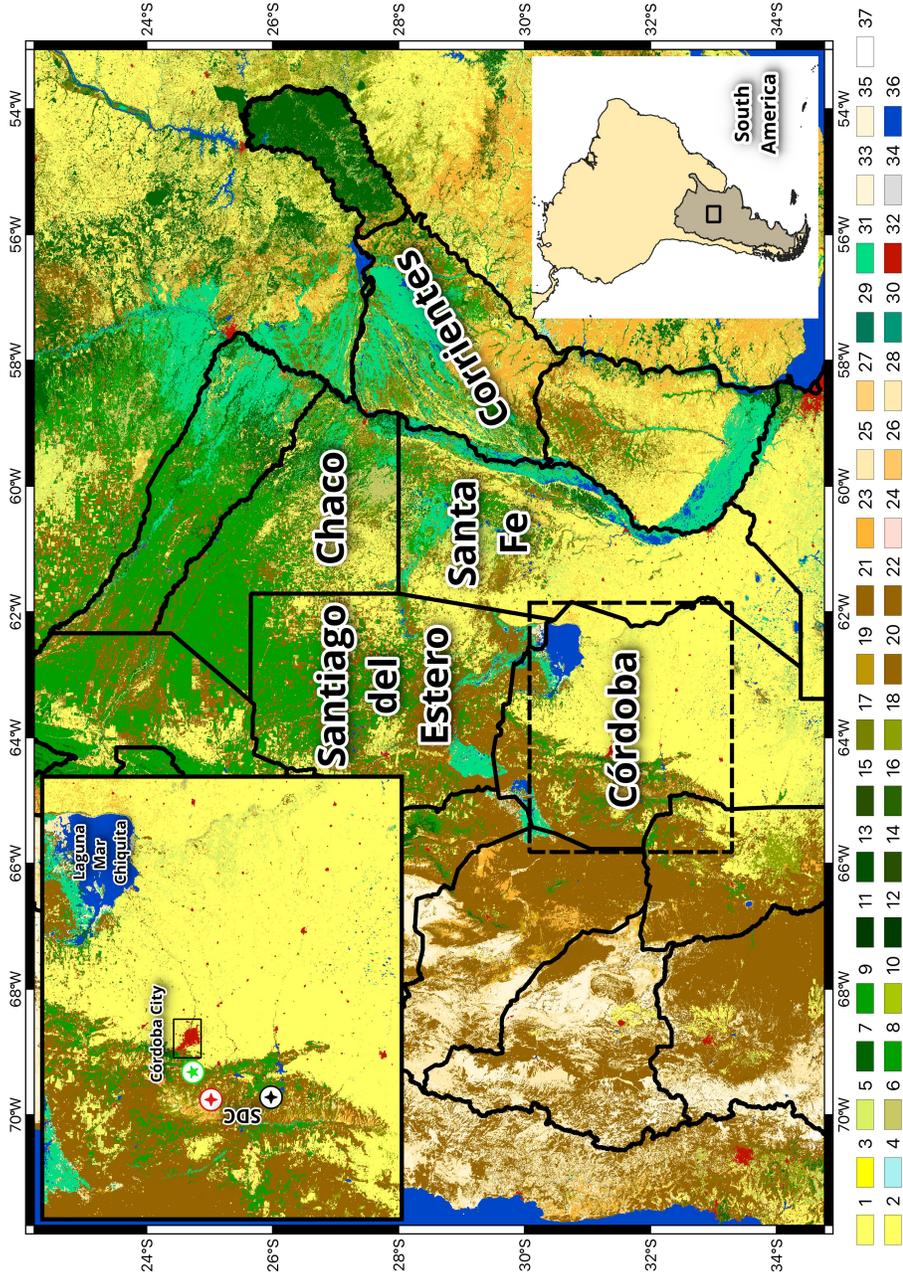


Figure 1. Map of northern Argentina covering an area shown in the black box in the lower right panel, including the Córdoba study area and nearby provinces. Within the dashed box covering northern Córdoba province is an inset (top left) highlighting Córdoba City to the east of the Sierras de Córdoba (SDC) and points of interest for this analysis: the site of the CACTI experiment observations in Villa Yacanto (black star), the hail collection location (lime green star), and the initiation point of the hail-producing cell (red star). Color fill represents the C3S Land Cover map available through the C3S Climate Data Store (CDS): 1-cropland rainfed, 2-cropland rainfed, 3-cropland irrigated, 4-cropland irrigated, 5-mosaic cropland, 6-mosaic natural vegetation, 6-tree broadleaved evergreen closed to open, 7-tree broadleaved deciduous closed to open, 8-tree broadleaved deciduous closed, 9-tree broadleaved deciduous open, 10-tree needleleaved evergreen closed to open, 11-tree needleleaved evergreen closed, 12-tree needleleaved evergreen open, 13-tree needleleaved deciduous closed to open, 14-tree needleleaved deciduous closed, 15-tree needleleaved deciduous open, 16-tree mixed, 17-mosaic tree and shrub, 18-mosaic herbaceous, 19-shrubland, 20-shrubland evergreen, 21-shrubland deciduous, 22-grassland, 23-lichens and mosses, 24-sparse vegetation, 25-sparse tree, 26-sparse shrub, 27-sparse herbaceous, 28-tree cover flooded fresh or brackish water, 29-tree cover flooded saline water, 30-shrub or herbaceous cover flooded, 31-urban, 32-bare areas, 33-bare areas consolidated, 34-bare areas unconsolidated, 35-water, 36-snow and ice.

perature. This novel approach applies microscopy techniques more frequently used in other disciplines, such as Geology (Hurley et al., 2021; Han et al., 2022), to uniquely provide information on non-soluble particle physical and chemical characteristics with respect to the hailstone center. A 2-D cross-section of the hail sample was created using an OLYMPUS LEXT OLS4000 Confocal Laser Scanning Microscope (CLSM), using the embryo as a reference and then scanning the sample from both ends. Within this area, sectors were selected for further magnification to identify individual particles manually with respect to the hailstone embryo. ProfilMOnline was used to calculate particle size. This detailed approach provided particle size distribution for 76 identified particles within this 4 cm hailstone, including particles as small as $1\ \mu\text{m}$ within individual sectors of the 2-D cross-section.

The sublimated hailstone sample was gold-coated and analyzed using a ZEISS FEG-SEM Sigma Scanning Electron Microscope (SEM-EDS) with backscatter and secondary electron images created at 15kV and 8.5mm working distance. Manual SEM particle examination confirmed 73 of the original 76 particles from CLSM analysis for further physical-chemical analysis. To avoid interference with the sample substrate, EDS spectra were acquired from the center of each particle using a single-point analysis technique to determine the elemental composition of individual particles. The following elements were identified, with abundance $\geq 1\%$ by weight (C, N, O, F, Na, Mg, Al, Si, S, Cl, K, Ca, Ti, Cr, Fe, Ni, Zn, Br, Mo), with elements $\geq 10\%$ considered predominant. Finally, a large area elemental cross-section was acquired at $0.8\ \mu\text{m}$ resolution across a diameter of the equatorial plane to evaluate the elemental distribution throughout the hailstone sample, providing insights into the overall composition distribution from the center (e.g., embryo or nuclei).

2.3 Air Mass Trajectories

The NOAA Air Resources Laboratory’s Hybrid Single-Particle Lagrangian Integrated Trajectory model (HYSPLIT)(Stein et al., 2015) was used to generate a 24-hour air mass back-trajectory using European Environment Agency Reanalysis data sets (ERA5) (Hersbach et al., 2020), with a vertical resolution of 37 pressure levels: 25hPa from 1025 through 750hPa , 50hPa from 750 through 300hPa , and 25hPa from 275 through 100hPa . These trajectories were initiated at 1700 UTC with hourly intervals, starting at unique model heights of 100 (first pressure level above surface), 500 and 1000 (boundary-layer variability), and (d) 1500 (low-level jet level) meters AGL from the convective core coordinates $[-64.75,-31.59]$ identified at initiation time using channel 11 ($8.4\ \mu\text{m}$) from the geostationary satellite GOES-16 (Bernal Ayala et al., 2022). A trajectory matrix with a 7×5 grid with 0.3 spacing was also processed from the initial point and time for five days to investigate any path variations for an air parcel before convective initiation. The area covered by all back-trajectories was divided into grid cells with dimensions of 0.28 longitude and 0.28 latitude (i.e., ERA-5 horizontal resolution). Each trajectory occurrence in each grid cell was normalized based on the time spent over each grid, including trajectory endpoints for all the heights (Ashbaugh et al., 1985). Residence-time coefficient pixels were then compared with the 22-class C3S Climate Data Store Land Cover classification gridded map from 2023 (e.g., Fig. 1). This approach provides insight into the highest probability of a specific land use being a source region for the non-soluble particles analyzed in this study.

3 Results

3.1 Particle Size and Elemental Composition of Individual Particles

The CLSM analysis of the 73 individual particles revealed maximum particle sizes ranging from 1.9 to $150.3\ \mu\text{m}$ with a precision uncertainty of $\pm 0.2\ \mu\text{m}$ (Fig. 2) and an average particle size of $40\ \mu\text{m}$. These sizes were larger on average than particles found

164 in the CACTI INP study (Testa et al., 2021) in Villa Yacanto, Argentina (black star in
165 Fig. 1), which analyzed ground-based INP measurements from instrumentation capa-
166 ble of measuring sizes only up to 20 μm . While particles of up to 100 μm are likely to
167 remain suspended in the atmosphere for up to 2 days (Jaenicke, 1978; Bakan et al., 1987),
168 the larger particles observed here, up to 150 μm , suggest that intense local winds were
169 required to loft those particles into the cloud. Overall, the greatest size variability is ob-
170 served in particles trapped in the center of the hailstone.

171 To evaluate if those size differences are owing to particle composition, the 73 iden-
172 tified particles were characterized based on elemental weight percentages from the SEM-
173 EDS analysis (Laskin et al., 2012) using the statistical-based Orange’s k-means charac-
174 terization method with silhouette scores (Demšar et al., 2013). Based on the outcomes
175 derived from the particle classification output of the k-means analysis, alongside the aware-
176 ness that EDS alone is incapable of distinguishing between biological and non-biological
177 species, we systematically categorized the particles into five general groups: Carbona-
178 ceous (e.g., Fig. 3, A), Carbonaceous with heavy metals, Silicates (e.g., Fig. 3, B), Sil-
179 icates with heavy metals, and Salts (e.g., Fig. 3, C). Particles included in the Carbona-
180 ceous group contained C greater than 10 % and greater than Cl and Si weight percent-
181 ages. Particles in the Silicates group had Si greater than 10 % and greater than C and
182 Cl. Particles containing Cl greater than 10 % and greater than C and Si were catego-
183 rized in the Salts group.

184 This classification scheme revealed that the particles in the hailstone were primar-
185 ily carbonaceous (66%; Fig. 2), containing carbon either in the inorganic, organic, or
186 elemental form that can be from anthropogenic, biological, or geological origin. Of the
187 carbonaceous particles, 19% contained heavier metals such as Br, Fe, Ni, Mo, Ti, and
188 Cr, with weight percentages individually greater than 1%. Regardless of whether these
189 carbonaceous particles contained heavy metals, they were spread throughout the hail-
190 stone with no apparent higher concentration in any one layer relative to the center. The
191 carbonaceous particles also contained the largest particle size with the highest size vari-
192 ability ranging from 5 to 150 μm with no preferential size distribution throughout the
193 sample. The second most prevalent group was particles categorized in the Silicates group.
194 This group composed 22% of all particles with size variability ranging from 2 to 55 μm
195 (Fig. 2). Similar to the carbonaceous group, silicate particles were spread through the
196 sample, and no apparent maximum concentration or size was observed in layers with re-
197 spect to the center. A subsidiary category to Silicates was Silicates with heavier met-
198 als, which only composed 4% of the sample, with sizes from 11 to 45 μm . These sam-
199 ples had at least 10% of Silicon as particles in the Silicates group. Still, these particles
200 contained Br and Fe with weight percentages individually greater than 1% and were found
201 in the outer layers of the hailstone. Lastly, Salt particles compose 8% of the entire sam-
202 ple. Salt particle sizes range from 8 to 29 μm (Fig. 2), primarily concentrated in the outer
203 layers of the hailstone. The full 2-D chemical cross section (Fig. 4) confirmed the pres-
204 ence of C throughout the sample. Si was challenging to distinguish due to glass inter-
205 ference. Finally, high Cl concentrations indicated the presence of salt particles in the hail-
206 stone’s outer layers (Fig. 4 bottom), as revealed in the analysis of randomly selected in-
207 dividual particles, but also near the center of the hailstone (Fig. 4 top).

208 3.2 Possible Source Regions

209 The analysis of the 24-hour HYSPLIT back trajectory (Fig. 5, A, top) sheds light
210 on the possible origins of the particles in the hailstone sample with distinct patterns at
211 different altitude levels. Levels 100 and 500 m AGL come initially from the northwest
212 but then loop over the northeastern part of the SDC, likely owing to the upslope flow
213 of surface winds from the north and northeast (Bernal Ayala et al., 2022). This upslope
214 flow is supported by the terrain and height change of the wind, moving upslope supported
215 by the terrain analysis seen in Fig. 5, A (bottom figure). The 1000 and 1500 m trajec-

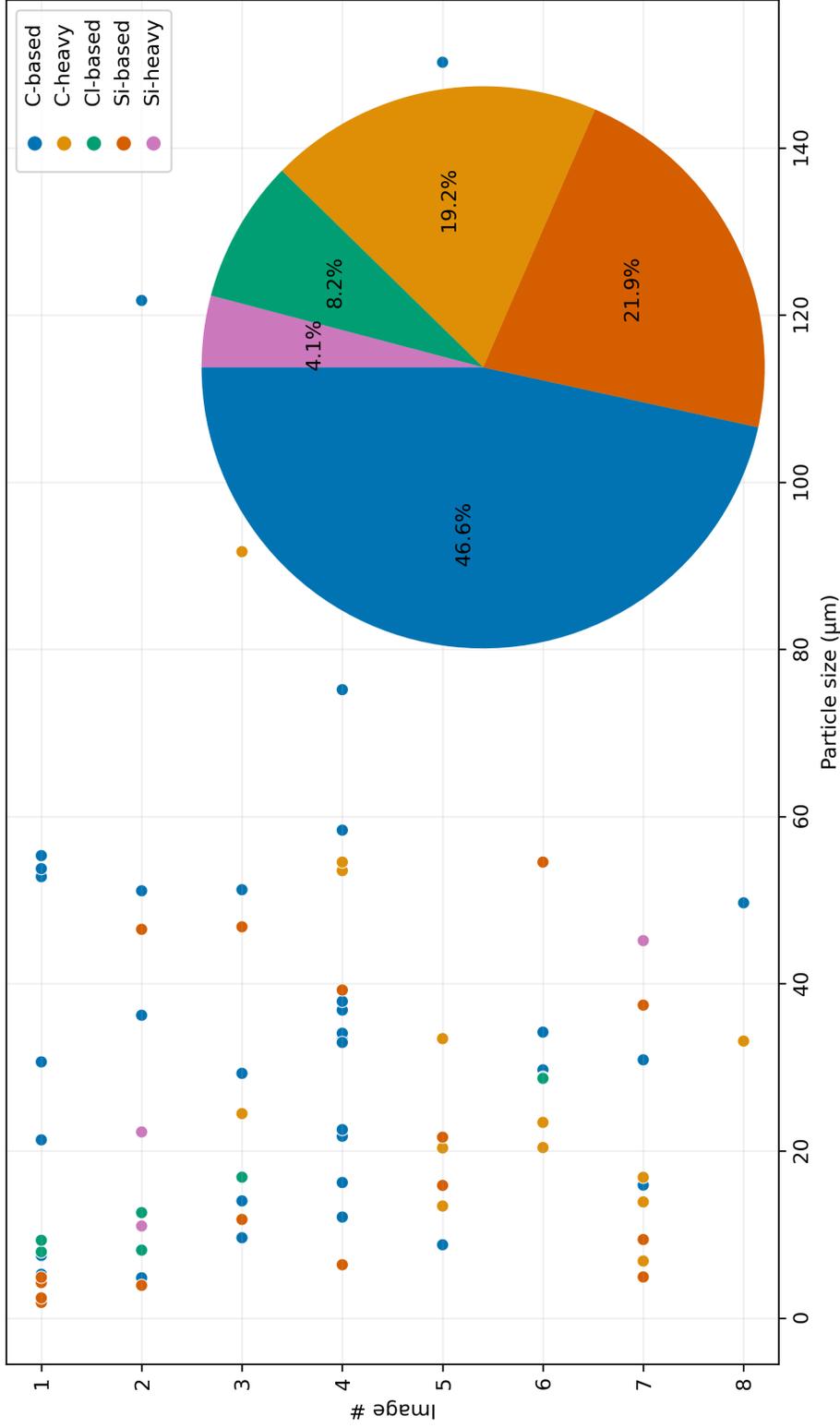


Figure 2. Particle size distribution graph (X-axis in μm) through the cross-section of the hailstone along a diameter of the equatorial plane (Y-axis in μm), with 0 and 8 representing the outermost layers of the hailstone, and images 4 and 5 depicting the center of the hailstone. The pie chart insert illustrates the elemental composition percentages of the 76 identified particles within the 4cm hailstone with colors corresponding to those used in the particle size distribution plot.

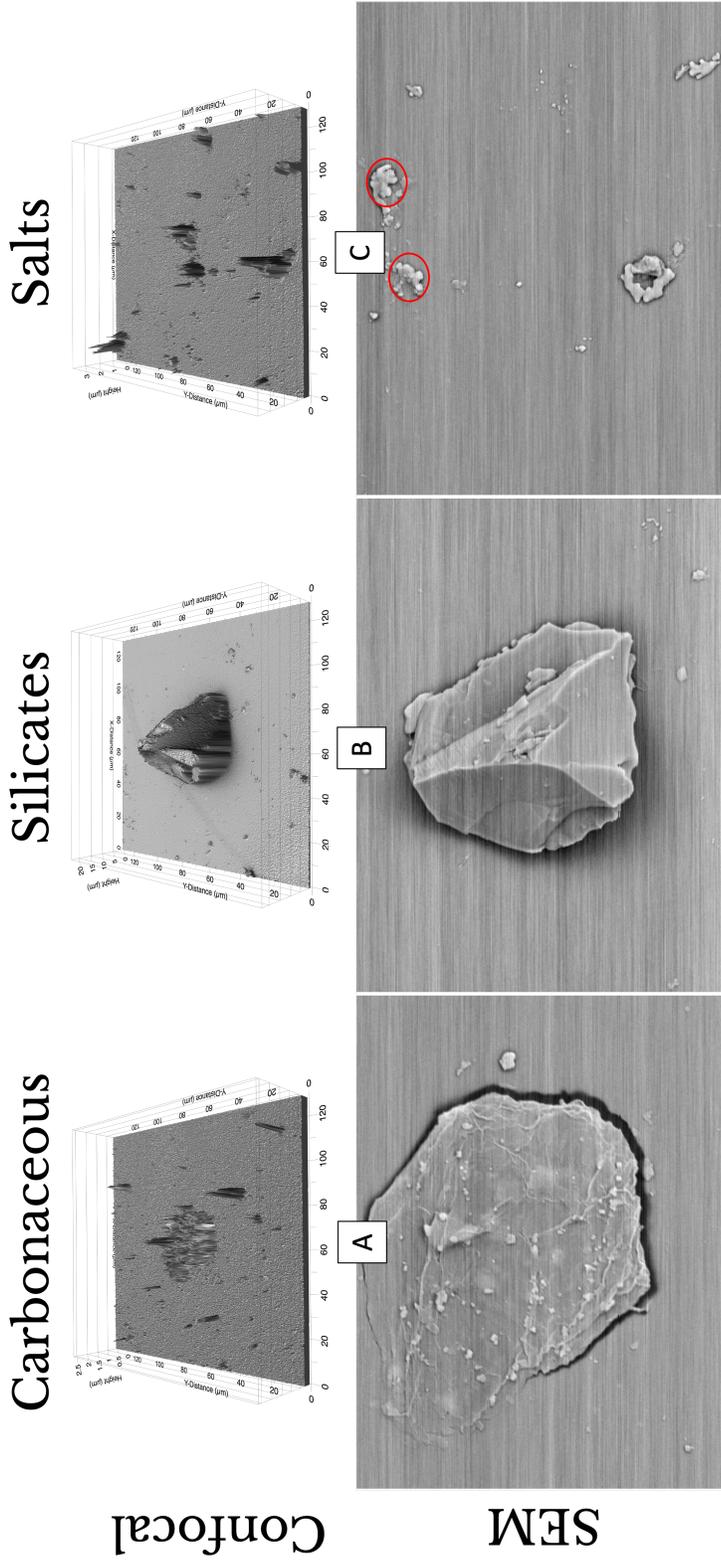


Figure 3. Examples from each of the primary particle categories are presented, along with their ProFilmOnline topographical output obtained through a Confocal Laser Scanning Microscope (CLSM, top), accompanied by their corresponding Scanning Electron Microscopy (SEM, bottom) image.

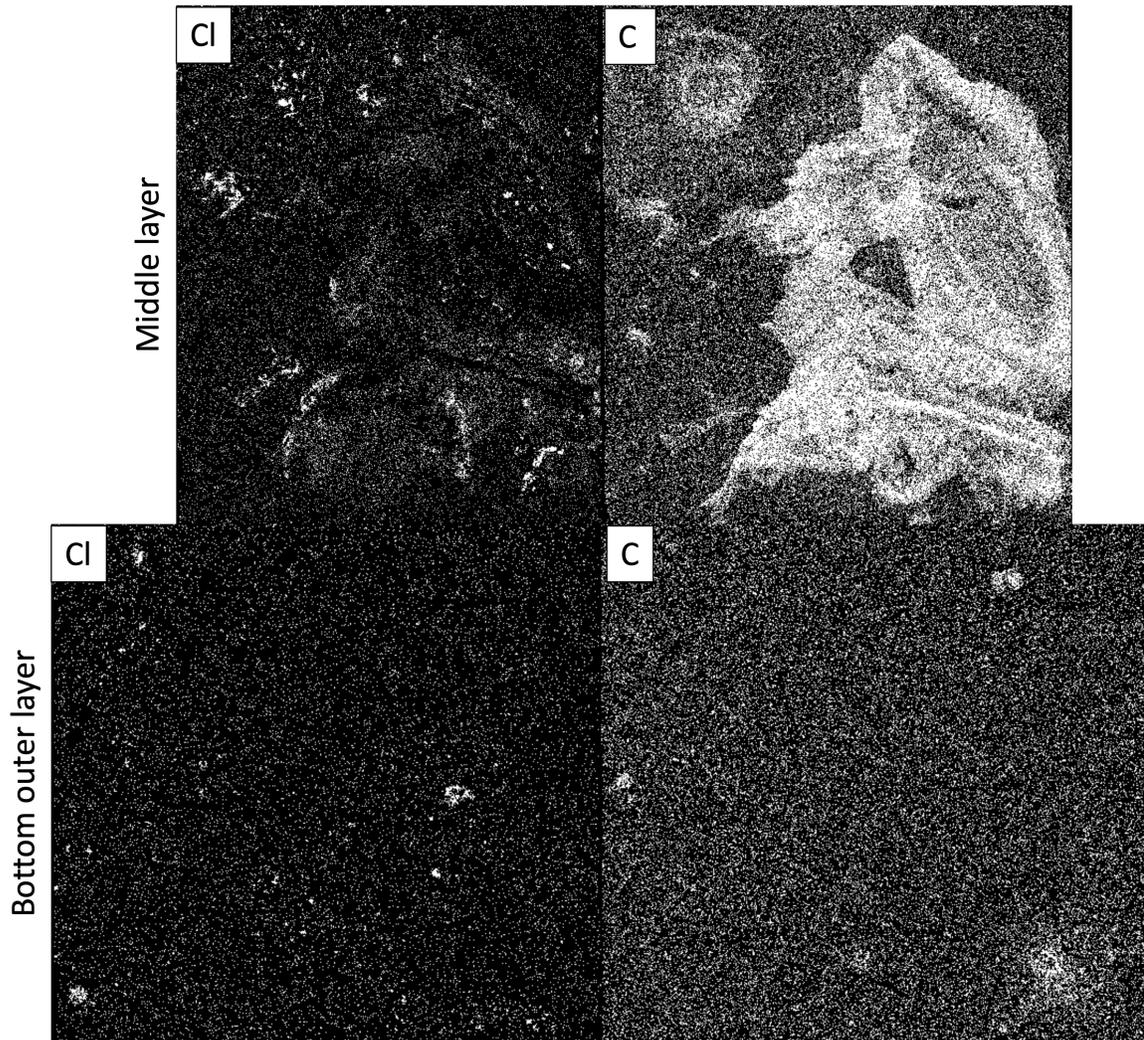


Figure 4. 2-D EDS cross-section of the complete hailstone sample processed in Fiji-ImageJ V2.14.0. This cross-section illustrates the intensity of elements Cl and C, with brighter spots indicating higher concentrations. Middle layers are positioned between images 4 and 5 in Figure 2, while the outer layers are situated between images 7 and 8.

216 tories initially come from the north and northwest, then show a slight curvature north
217 of the Cordoba Province and later from the northeast. The residence time coefficients
218 calculated for 5-day back-trajectories (Fig. 5, B) similarly highlight the regional influ-
219 ence. Under the assumption that the particles arriving at the location where the hail-
220 producing storm initiated are more likely emitted from regions where the air masses spent
221 more time (Yadav et al., 2021; Ren et al., 2021; Testa et al., 2021), the grid cells show-
222 ing the high residence-time coefficients are considered potentially principal sources for
223 the particles found in the hailstone. The regions with the highest residence-time coef-
224 ficients (orange, yellow, and red in Fig. 5, B) are generally located over the SDC, Cor-
225 doba City, Argentina’s largest natural salt lake (Laguna Mar Chiquita), and Provinces
226 such as Santiago del Estero, Chaco, Santa Fe, and Corrientes (Fig. 1). These results re-
227 veal that discernible sources only within Argentina’s geographical limits account for pos-
228 sible non-soluble particulate sources in the analyzed hailstone sample, supporting Bernal Ay-
229 ala et al. (2022)’s conclusion of local environmental factors responsible for this case. Fur-
230 thermore, in comparing the residence-time coefficient pixels from the 5-day back trajec-
231 tory with the C3S Land Cover map (Fig. 5, B), we find the most predominant land uses
232 (Fig. 5, C) showing high residence-time coefficient areas are Shrublands, Croplands, ar-
233 eas covered by Mixed Vegetation, Urban areas (mostly Cordoba city), and areas with
234 a body of water (the Salt Lake), consistent with the findings of carbonaceous, mineral,
235 and salt particles, as well as heavy metals, throughout the hailstone.

236 4 Discussion

237 Our results show, for the first time, the sizes, composition, and distribution with
238 respect to the stone’s center of non-soluble particles trapped in a hailstone in Argentina.
239 This analysis reveals the presence of carbonaceous particles within the embryo of the hail-
240 stone, consistent with the prevalence of biological particles in this region (Testa et al.,
241 2021) and the presence of carbonaceous particles near the center of the hailstones in other
242 parts of the world (Šantl-Temkiv et al., 2013; Michaud et al., 2014; Beal et al., 2021).
243 Additionally, observations of carbonaceous particles up to $150\ \mu\text{m}$ in diameter suggest
244 the influence of strong local winds, which effectively suspended these larger particles for
245 a sufficient time, enabling them to be entrained within the updraft. These results rein-
246 force the important role of biological and geological particles in understanding hail for-
247 mation in deep convective clouds globally.

248 Silicates were also detected near the hailstone’s center with particle sizes ranging
249 from 7 to $39\ \mu\text{m}$, overlapping with the carbonaceous particle size ranges. These silicates
250 can be attributed to eroded particles from the surrounding mountains, specifically K-
251 feldspar commonly found in agricultural topsoil, as well as soil dust particles resulting
252 from agricultural activities in this region (Testa et al., 2021). The identification of larger
253 silicate particles in the center of the hailstone suggests that they were sufficiently large
254 to nucleate at temperatures similar to those of the carbonaceous particles, aligning with
255 the findings of Beal et al. (2021). Among the mineral particles, K-feldspar has previously
256 been identified as the most ice-active component for promoting ice nucleation (Kiselev
257 et al., 2017). When compared to pure mineral standards, the coexistence of Na, Cl, Mg,
258 Ca, and various elements within the silicate particles in our sample suggests the aggre-
259 gation of particles with multiple mineral phases. This finding aligns with the notion that
260 when silicates become suspended in the air, they can potentially aggregate with other
261 atmospheric particles, resulting in the coagulation of compositions that enhance their
262 nucleation potential. This effect is particularly pronounced at higher temperatures, as
263 Pruppacher and Klett (1980) indicated.

264 Regarding the outer layer of the hailstone, we observed a consistent size distribu-
265 tion encompassing various groups of particle composition, including carbonaceous par-
266 ticles that were evenly distributed throughout the sample. Within this outer layer, we
267 detected smaller silicate particles that may have been too small to act as INPs during

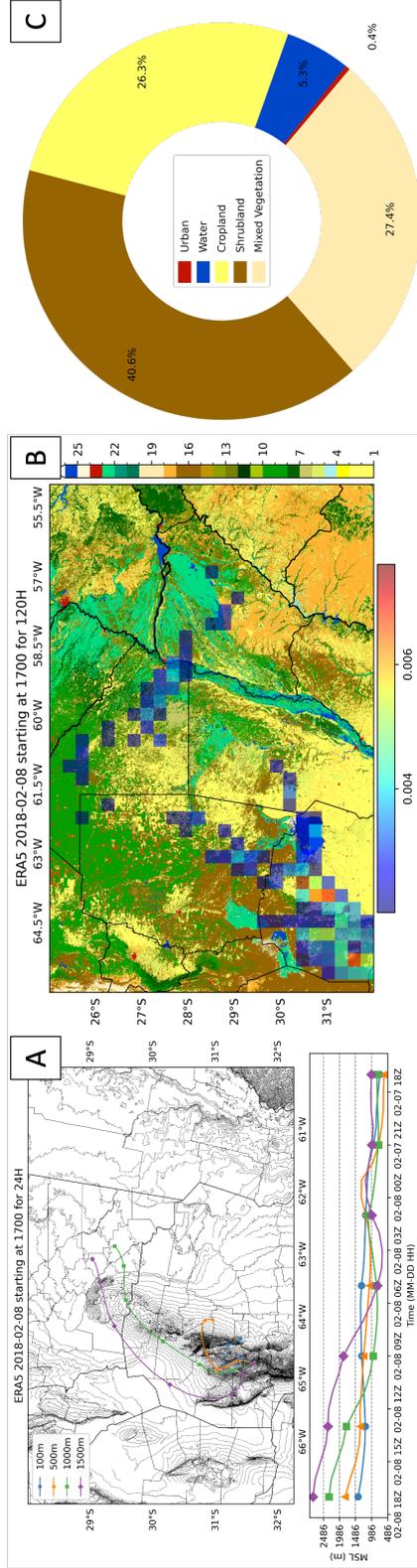


Figure 5. Figure A (top) shows 24-hour HYSPLIT back trajectories color-coded by height (in meters), as the legend defines. The bottom part of Figure A displays terrain analysis in meters above sea level for the same time periods as the back trajectories. Figure B shows residence time coefficients calculated for 5-day back trajectories overlaid on the C3S Land Cover map. The land cover colors correspond to a subset of categories shown in Figure 1, specifically: 1/2/3-Cropland rain-fed, 4-Cropland irrigated, 5-Mosaic cropland, 6-Mosaic natural vegetation, 7-Evergreen broadleaved, 8/9/10-Deciduous broadleaved, 11-Evergreen needleleaved, 12-Mixed trees, 13-Mosaic tree/shrub, 14-Mosaic herbaceous, 15/16-Shrubland, 17-Grassland, 18-Sparse vegetation, 19-Sparse herbaceous, 20-Fresh water flooded tree cover, 21- Saline water flooded tree cover, 22-Flooded shrub/herbaceous cover, 23-Urban, 24-Bare areas, 25-Water, 26-Snow and ice. Figure C displays the predominant land uses within the residence time coefficient pixels observed in Figure B. Similar land uses were grouped into the following categories: Urban (23), Water (25,26), Cropland(1,2,3,4,5), Shrubland(15,16,17), Mixed Vegetation (6,7,8,9,10,11,12,13,14,18,19). Flooded Vegetation (20,21,22) and bare areas (24) were not included.

268 the formation of the hailstone embryo. Still, they could have served as CCN and accu-
269 cumulated onto the hailstone during phases when the temperature allowed for the accu-
270 mulation of supercooled liquid water, thereby continuing the growth of the hailstone lay-
271 ers. Additionally, we detected salts in the hailstone, which likely originated from par-
272 ticles acting as CCN. This result suggests that the salts, potentially originating from the
273 salt lake northeast of the SDC, could mix with other atmospheric particles, potentially
274 leading to an increased presence of insoluble CCNs during hailstone growth.

275 The SEM-EDX analysis revealed that 12 of the 73 analyzed particles contained heav-
276 ier metals such as Br, Fe, Ni, Ti, Mo, and Cr, most combined with silicates and biolog-
277 ical particles. Anthropogenic sources may contribute to the presence of heavier metals
278 in the particles (Beal et al., 2021), which has implications for modeling aerosol interac-
279 tions. Therefore, considering the influence of anthropogenic activities in nearby urban
280 areas is essential, as they may transport heavier metals that may mix with other silicates
281 and carbon-based particles suspended in the atmosphere. Additionally, they can be ab-
282 sorbed by surrounding vegetation, which can subsequently re-emit them back into the
283 atmosphere (Pruppacher & Klett, 1980). These metals later become available to mix with
284 other atmospheric particles, potentially affecting ice nucleation temperatures within the
285 cloud.

286 5 Conclusions

287 Using a novel technique described in Bernal Ayala, Rowe, Arena, Nachlas, and Asar
288 (2024), this study provides first-of-their-kind insights into the composition, size distri-
289 bution, and potential sources of non-soluble particles within a hailstone collected near
290 Cordoba, Argentina. The results indicate that the hailstone contains diverse particles,
291 with carbonaceous particles dominating the hailstone sample, including the embryo, with
292 silicate and salt particles distributed throughout the sample. The presence of carbona-
293 ceous particles suggests a strong influence of biological and geological sources in hail for-
294 mation, while silicates originate from eroded soil particles and agricultural activities. The
295 particle size distribution, particularly the detection of larger particles up to 150 μm , points
296 to the role of local winds in lofting particles into the convective cloud, where they could
297 participate in nucleation and growth processes.

298 This study highlights the regional influence of various land use types, including shrub-
299 lands, croplands, urban areas, and bodies of water, specifically a salt lake. It supports
300 the idea that local environmental factors play a significant role in the particles in the hail-
301 stone. Overall, this study enhances our understanding of the complex interplay between
302 land-use sources emitting atmospheric particles, local environmental conditions, and hail-
303 stone formation processes. The findings contribute to ongoing research in hail formation
304 and offer insights into the potential impacts of natural and anthropogenic factors on the
305 elemental chemical composition of hailstones. Future work will analyze additional hail-
306 stones collected in this region under different environmental conditions.

307 6 Open Research Section

308 The physical (CLSM) and chemical (EDS) data of hailstones are stored in an Ex-
309 cel sheet, which has been uploaded and is accessible through Zenodo (Bernal Ayala, Rowe,
310 Arena, & Nachlas, 2024). We obtained the ERA5 reanalysis data from Hersbach et al.
311 (2023) and the land cover classification gridded map from the Copernicus Climate Change
312 Service, Climate Data Store Service (2019). This ERA5 data was input to the National
313 Oceanographic and Atmospheric Administration’s Hybrid Single-Particle Lagrangian In-
314 tegrated Trajectory (HYSPLIT) model (version 5.3.0) (available at <https://www.ready.noaa.gov/HYSPLIT.php>)
315 to produce residence time coefficient plots based on the air mass trajectories.

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