



PUBLICATIONS

Global Biogeochemical Cycles

Supporting Information for

Suppression of nitrogen deposition on global forest soil CH₄ uptake depends on nitrogen status

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36 **Introduction**

37 The uploaded Data Set S1 (CH₄_exp dataset in main text) was used to derive the response
38 factors of soil CH₄ flux to N input in global forests; Data Set S2 (CH₄_obs dataset in main text)
39 was used to estimate the soil CH₄ fluxes in global forests; Data Sets S3–S7 were used to classify
40 the N-limited and N-saturated forests on global level; Data Set S8 contains environmental
41 factors (MAT, MAP, soil texture, etc.) for global estimations; Data Set S9 contains global forest
42 soil CH₄ budgets reported in previous studies. The data analysis process and produced figures
43 can be replicated with the uploaded R script (Code S1).

44

45 **Text S1. Nitrogen saturation status of global forests indicated by sensitivity of soil N₂O**
46 **emission to N deposition.**

47 Globally, human-induced increase in atmospheric N deposition is changing forests from a
48 nitrogen-limited to nitrogen-saturated status. In N-limited forests, plants and microbes utilize N
49 conservatively for a lower proportion of input N to be leaked from tight N cycling processes
50 (Chapman et al., 2006; Van Der Heijden et al., 2008). However, when forests become N-
51 saturated, input N exceeds the N demand of plants and microbes, leading to excessive
52 utilization of N, and thus, the N cycle becomes more open (Hietz et al., 2011). Therefore, a
53 higher proportion of input N is lost via leaching or gaseous emission (Aber et al., 1989). This
54 implies that increased gaseous N emissions (N₂O, NO, and N₂) per unit of N deposition (i.e.,
55 higher sensitivity of gaseous N emissions to N deposition) may indicate forests reaching N
56 saturation (Aber et al., 1998). Coincidentally, studies have measured nitrous oxide (N₂O)
57 greenhouse gas emissions under different N input levels since the 1980s in global forests, using
58 a controlled experiment design and standard sampling method (Holland et al., 1999). The
59 accumulated experimental data provide an opportunity to quantify the sensitivity of N₂O
60 emissions to N deposition in various forests, and indicate the N limitation or saturation status of
61 global forests.

62
63 *Gathering data*

64 To quantify the sensitivity of soil N₂O emissions to N deposition (s_N), we compiled soil N₂O
65 emission data observed in N addition experiments conducted in global forests. On 03/30/2022,
66 we searched for papers and theses published before 01/01/2022 from the Web of Science Core
67 Collection database (www.webofscience.com) and China National Knowledge Infrastructure
68 Theses and Dissertations Database (<https://oversea.cnki.net/kns?dbcode=CDMD>), using the
69 following keywords: "forest" AND "greenhouse gas" OR "N₂O" OR "nitrous oxide". The retrieved
70 7422 papers and 718 theses were then refined manually based on the following criteria: (i)
71 experimental N addition was conducted in forest ecosystem; (ii) literature recorded the location,
72 time, and dose of the experiment(s); (iii) soil N₂O flux was observed in experimental sites and
73 measured using gas chromatograph technique (Holland et al., 1999). As a result, the compiled
74 "N₂O_exp" dataset (Data Set S3) contained 553 observations from 102 sites worldwide (Fig. S7).

75 Similarly, we compiled data on the soil N₂O emission rates of global forests observed under
76 natural conditions. We refined from the same papers and theses as above, using a different set
77 of criteria: (i) no nutrients, including N, were artificially added to the forest site so the site only
78 received naturally deposited N; (ii) literature recorded the location, and time of flux
79 measurement; (iii) soil N₂O flux was observed in the field and measured using gas
80 chromatograph technique (Holland et al., 1999). The compiled "N₂O_obs" dataset (Data Set S4)
81 contained 246 observations from 140 sites worldwide (Fig. S7).

82 In addition, we compiled data on total N loss (N leaching and gaseous N emission
83 combined), N leaching, and change in soil N pool, from N addition experiments in global forests.
84 We searched in the aforementioned databases using the following keywords: "forest" AND
85 "nitrogen addition" OR "fert*" AND "nitrogen loss" OR "nitrogen leaching" OR "nitrogen
86 budget". Retrieved 2693 papers and theses were then refined based on the following criteria: (i)
87 literature recorded the location, time, and dose of experimental N addition in forests; (ii) total N
88 loss rate, N leaching rate, or change rate of soil N pool was observed or estimated in the

89 experiments. The compiled "Ncycle_exp" dataset (Data Set S5) contained 169 observations from
90 37 sites (Fig. S7).

91 To analyze the relationship between s_N and N saturation status, we compiled data on field-
92 observed N-limited and N-saturated forests indicated by N leaching. On 10/31/2022, we
93 searched for literature in the aforementioned databases using the following keywords: "forest"
94 AND "leaching" AND "nitrogen limit*" OR "nitrogen saturat*". Retrieved 823 papers and theses
95 were then refined based on the following criteria: (i) literature recorded whether the forest was
96 N-limited or N-saturated, and its location; (ii) literature used nitrogen leaching as an indicator of
97 N limitation or saturation status. The compiled "Nleach" dataset (Data Set S6) contains 136
98 observations from 92 sites worldwide (Fig. S6). We also used data on field-observed N-limited
99 and N-saturated ecosystems indicated by plant growth response to N input ("NuLi" dataset;
100 Data Set S7) from a published database by Du et al. (2020). It covers 106 sites worldwide, 65 of
101 which are forest sites (Fig. S6).

102 Moreover, we extracted auxiliary information from the literature on environmental factors
103 (including mean annual temperature, MAT; mean annual precipitation, MAP; mean annual N
104 deposition rate, N_{depo} ; coefficients of temporal variation, MAT.cv, MAP.cv, and $N_{depo}.cv$; soil sand
105 content, soil clay content, and other soil properties) for the forest sites in the datasets. However,
106 the literature did not provide the necessary auxiliary information for all sites; therefore, spatial
107 datasets were used to fill in the missing data based on the location of the sites. Global
108 temperature and precipitation datasets were obtained from the Climatic Research Unit,
109 University of East Anglia (https://crudata.uea.ac.uk/cru/data/hrg/cru_ts_4.03/). The soil C:N ratio
110 was obtained from a published database (Shangguan et al., 2014). Other soil properties were
111 obtained from the HWSD dataset ([https://www.fao.org/soils-portal/data-hub/soil-maps-and-](https://www.fao.org/soils-portal/data-hub/soil-maps-and-databases/harmonized-world-soil-database-v12/en/)
112 [databases/harmonized-world-soil-database-v12/en/](https://www.fao.org/soils-portal/data-hub/soil-maps-and-databases/harmonized-world-soil-database-v12/en/)). N deposition rate and forest cover data
113 were from published databases (Ackerman et al., 2019; Liu et al., 2020). The forest biome map
114 was derived from the Global Forest Monitoring project (Hansen et al., 2010).

115

116 *Quantifying the sensitivity of soil N₂O emissions to N deposition*

117 Under low N input, the soil N₂O emission rate responds almost linearly to N input, whereas
118 high N input may induce non-linear responses (Aber et al., 1998; D.-G. Kim et al., 2013). High N
119 input may change ecosystem properties, leading to a deviation from the natural response of
120 ecosystems to environmental change. Therefore, we used a linear model (Eq. S1) to define and
121 quantify the sensitivity (s_N) of soil N₂O emissions to N deposition (or low N input), for s_N to
122 reflect ecosystem properties (i.e., N saturation status).

$$123 \quad R_{N_2O} = s_N \times N_{depo} + R_0 \quad (\text{Eq. S1})$$

124 where R_{N_2O} is the soil N₂O emission rate (kgN₂O-N ha⁻¹ yr⁻¹), N_{depo} is the atmospheric N
125 deposition rate (kg N ha⁻¹ yr⁻¹), s_N is the sensitivity of soil N₂O emission to N deposition,
126 quantified as soil N₂O emission per unit of low N input (kgN₂O-N kgN⁻¹), and R_0 is the
127 background soil N₂O emission rate when there is no N deposition or artificial N addition
128 (kgN₂O-N ha⁻¹ yr⁻¹).

129 A segmented regression analysis on N₂O_exp dataset showed that there is one change
130 point in the linear relationship between N input rate and R_{N_2O} , which is 174.70 ± 19.73 kgN ha⁻¹
131 yr⁻¹. That is in line with change points estimated or used in previous studies (Bouwman et al.,
132 2002; Hoben et al., 2011; M. Lu et al., 2022; McSwiney & Robertson, 2005; Shcherbak et al.,

133 2014). Conservatively, experimental data with N addition rates not exceeding $150 \text{ kg N ha}^{-1} \text{ yr}^{-1}$
134 were used as “low N input” data in further analysis. The N deposition rates in global forests were
135 lower than the level (Ackerman et al., 2019). For all the low-N input data in the $\text{N}_2\text{O_exp}$ dataset,
136 we aggregated them to $0.5^\circ \times 0.5^\circ$ grids based on their coordinates to match the spatial
137 resolution with environmental factors and reduce random errors in sampling. A linear model
138 (Model: $R_{\text{N}_2\text{O}} \sim \text{N input rate}$) was built for each grid with low-N input data. The slope of the
139 linear model was the estimated s_{N} of the grid (Table S3).

140 Based on the estimated s_{N} of all grids and the corresponding environmental factors, we
141 built a generalized linear model to simulate s_{N} (Table S4). In addition, another generalized linear
142 model was built to simulate R_0 .

143 To validate s_{N} , we firstly used the modeled s_{N} , together with the modeled R_0 and N_{depo}
144 datasets, to estimate $R_{\text{N}_2\text{O}}$ (Eq. S1). The estimated $R_{\text{N}_2\text{O}}$ values were compared with $R_{\text{N}_2\text{O}}$
145 observations ($\text{N}_2\text{O_obs}$ dataset) and indirectly validated the intermediate variable s_{N} (Fig. S8). In
146 addition, s_{N} was validated using a second approach. The sensitivity of N loss to N input (c_1), the
147 sensitivity of N leaching to N input (c_2), and the end-product ratio of nitrification and
148 denitrification processes (c_3) were either derived from the Ncycle_exp dataset or extracted from
149 the literature; s_{N} was then calculated from these parameters (Eq. S2).

$$s_{\text{N}} = c_3 \times (c_1 - c_2) \quad (\text{Eq. S2})$$

151 The limited observations allowed us to calculate s_{N} on a biome scale (Fig. S7), which was
152 then compared with the biome-mean value of the modeled s_{N} to validate it. The good
153 agreement also validated the modeled s_{N} ($r = 0.998$).

154 *Determining N saturation status of global forests using s_{N}*

155 We tested whether s_{N} can distinguish between N-limited and N-saturated forests using
156 data from forests having field-observed N saturation status data. First, we combined Nleach and
157 NuLi datasets to enlarge the sample size and derive a universal classification. Excluding three
158 duplicate sites in both datasets, the combined dataset had 154 sites with field-observed N
159 saturation status (86 N-limited and 68 N-saturated sites).

161 We modeled the s_{N} of the 154 sites using environmental factors (Table S4). We then
162 analyzed the s_{N} of N-limited and N-saturated forests and verified if there were significant
163 differences on the global and biome scales. In Western Europe, North America, and East Asia,
164 where there were abundant sites, we also compared the s_{N} of forests with N-limited or N-
165 saturated status on a regional scale. The mean s_{N} was significantly different on global and
166 regional scales ($p < 0.001$; Fig. S9), proving that s_{N} can indicate N limitation or saturation status in
167 forests.

168 Then we calculated an optimal threshold for s_{N} using data from 154 sites with field-
169 observed N saturation status and s_{N} information. The bootstrap method accounted for the
170 different sample sizes of N-limited and N-saturated sites (Davison & Hinkley, 1997). Specifically,
171 from the 154 sites, we randomly sampled 10 N-limited and 10 N-saturated sites and selected a
172 cutoff value for their s_{N} at a precision of $0.0001 \text{ kgN}_2\text{O-N kgN}^{-1}$. Sites in which s_{N} were above the
173 cutoff value were classified as “N-saturated,” and the rest were classified as “N-limited.” The
174 classified N saturation status of the sites was compared with field observations to determine the
175 accuracy of the classification, which was calculated as the proportion of sites accurately classified
176 into the same category as that observed. All possible cutoff values were tested, and the one with

177 the highest classification accuracy was the “optimal” cutoff value. Random sampling and
178 detection of optimal cutoff values were repeated 5000 times, during which some optimal cutoff
179 values were detected more frequently than others. The optimal threshold for s_N in all samples
180 was the most frequently detected optimal cutoff value, which was $0.0143 \text{ kgN}_2\text{O-N kgN}^{-1}$.

181 The N saturation status of global forests was determined based on the optimal threshold.
182 Forests with s_N above the threshold were classified as N-saturated, and the rest were classified
183 as N-limited. The accuracy of the classification was higher than 70% on global and regional
184 scales (Fig. S6). Based on the classification, we produced a rasterized map of N-limited and N-
185 saturated forests ($0.5^\circ \times 0.5^\circ$ resolution) in ArcGIS (ESRI, 2011).

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188 **Text S2. Inferring the variation of methane production and oxidation rates from the**
189 **variation of observed methane fluxes**

190 Soil CH₄ flux observed on the soil-air interface is codetermined by methane production
191 (methanogenesis) and oxidation rates (Eq. S3). However, it has been difficult to disentangle the
192 responses of methane production and oxidation to N input, because of the limited ability to
193 separately observe methanogenesis and methane oxidation processes in the field. Here, we
194 inferred the variation of methane production and oxidation rates from the variation of observed
195 methane fluxes. This could further support the “three stage” hypothesis we proposed.

$$196 \quad R_{CH_4} = R_{CH_4_{prod}} - R_{CH_4_{oxid}} \quad (\text{Eq. S3})$$

197 where R_{CH_4} is the observed soil CH₄ flux (kg ha⁻¹ yr⁻¹), positive R_{CH_4} value means methane
198 emission, whereas negative R_{CH_4} value means methane uptake; $R_{CH_4_{prod}}$ is methane production
199 rate (kg ha⁻¹ yr⁻¹); $R_{CH_4_{oxid}}$ is methane oxidation rate (kg ha⁻¹ yr⁻¹).

200 The change in methane production and oxidation rates could hardly be inversely calculated
201 from the R_{CH_4} values. Here, we inferred the change in $R_{CH_4_{prod}}$ and $R_{CH_4_{oxid}}$ by analyzing the mean
202 values and standard deviations of R_{CH_4} .

203 Firstly, the standard deviation of R_{CH_4} could be calculated from that of $R_{CH_4_{prod}}$ and $R_{CH_4_{oxid}}$
204 (not considering the interaction between $R_{CH_4_{prod}}$ and $R_{CH_4_{oxid}}$; Eq. S4).

$$205 \quad SD(R_{CH_4}) = \sqrt{SD(R_{CH_4_{prod}})^2 + SD(R_{CH_4_{oxid}})^2} \quad (\text{Eq. S4})$$

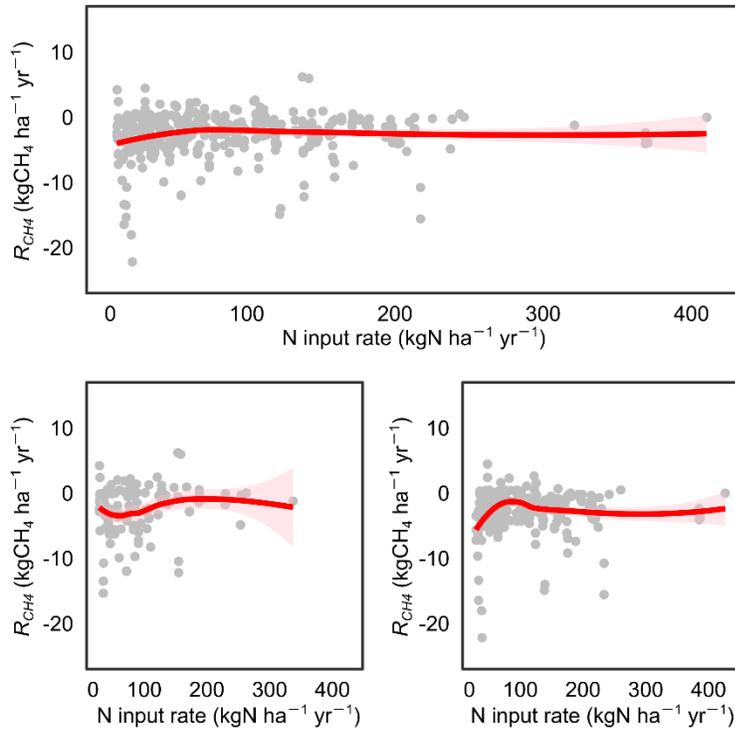
206 Usually, when the expected value of a variable becomes higher, its observations will be
207 more dispersed. This is because the random errors in the observations are often proportional to
208 their values. That is to say, statistical dispersion of $R_{CH_4_{prod}}$ and $R_{CH_4_{oxid}}$ (as indicated by their
209 standard deviations) should be positively related to their mean values.

210 Therefore, the decrease in the standard deviation of R_{CH_4} under high N input (Fig. S2) may
211 result from: (1) $R_{CH_4_{prod}}$ decreased under high N input, and $R_{CH_4_{oxid}}$ didn't change or slightly
212 increased; (2) $R_{CH_4_{oxid}}$ decreased under high N input, and $R_{CH_4_{prod}}$ didn't change or slightly
213 increased; (3) both $R_{CH_4_{prod}}$ and $R_{CH_4_{oxid}}$ decreased under high N input.

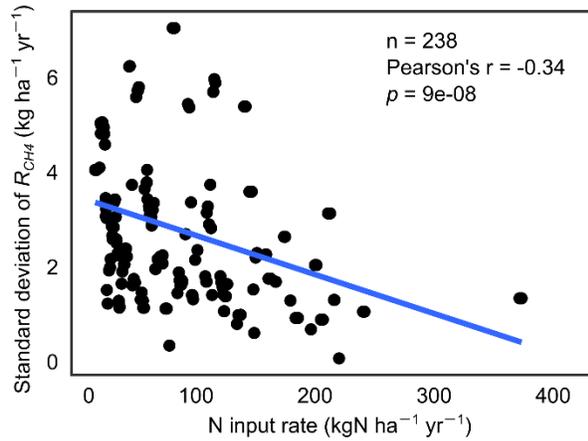
214 Meanwhile, we observed that the mean values of R_{CH_4} remained nearly unchanged under
215 high N input (Fig. S1a), which may result from: (i) both $R_{CH_4_{prod}}$ and $R_{CH_4_{oxid}}$ increased under high
216 N input; (ii) both $R_{CH_4_{prod}}$ and $R_{CH_4_{oxid}}$ decreased under high N input; (iii) both $R_{CH_4_{prod}}$ and $R_{CH_4_{oxid}}$
217 remained constant under high N input.

218 Combining the two evidences (standard deviation and mean values of R_{CH_4} under high N
219 input), it can be inferred that only hypotheses (3) and (ii) can be true at the same time. That is,
220 both $R_{CH_4_{prod}}$ and $R_{CH_4_{oxid}}$ decreased under high N input.

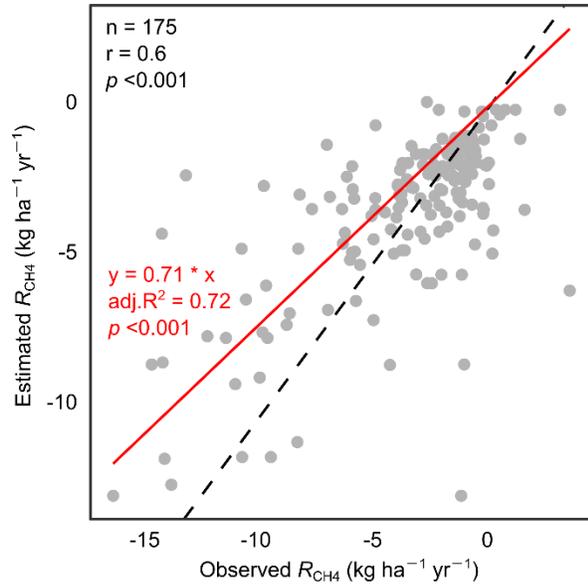
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 223 **Fig. S1.** Locally weighed regression (“LOWESS”) model on soil CH₄ emission rate and N input
 224 rate. (a) Using all observations compiled from global N addition experiments, the N input rates
 225 of which were no greater than 400 kgN ha⁻¹ yr⁻¹ (n = 448). The few but variable observations on
 226 soil CH₄ fluxes at sites where N input rates were above 400 kgN ha⁻¹ yr⁻¹ (n = 17) were not used
 227 in further analysis. (b) LOWESS model constructed using data from N-limited sites and also
 228 where N addition experiments have been conducted for no more than 3 years when CH₄
 229 emissions were observed (n = 131); (c) LOWESS model constructed using data from N-saturated
 230 forests, or data from sites where N addition experiments have been conducted for more than 3
 231 years before observing the CH₄ fluxes (n = 317). Pink shadings represent the standard errors of
 232 the fitted models.
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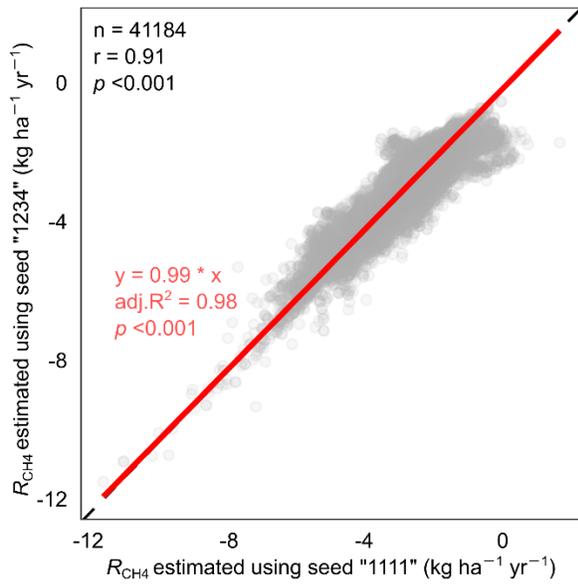
234
 235 **Fig. S2.** Standard deviation of soil methane flux (R_{CH_4}) was negatively correlated to N input rate.
 236 Data corresponding to N input levels above $400 \text{ kgN ha}^{-1} \text{yr}^{-1}$ were not included in this analysis,
 237 because the very limited observations may not sufficiently reveal the statistical distribution of
 238 R_{CH_4} . There were 238 unique N input rates that was no greater than $400 \text{ kgN ha}^{-1} \text{yr}^{-1}$. In practice,
 239 standard deviation was calculated for R_{CH_4} corresponding to each N input rate, and N input rates
 240 less than $2 \text{ kgN ha}^{-1} \text{yr}^{-1}$ in difference (e.g., standard deviation of R_{CH_4} corresponding to 5 kgN
 241 $\text{ha}^{-1} \text{yr}^{-1}$ was calculated using observations whose N input rates were within the range of 3 to 7
 242 $\text{kgN ha}^{-1} \text{yr}^{-1}$). That was to make sure that there were sufficient observations for each N input
 243 level.
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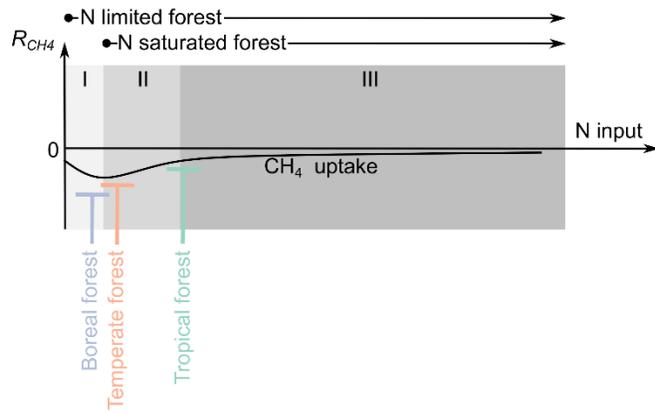
Fig. S3. Comparing observed soil CH₄ flux (R_{CH_4}) in testing dataset with that estimated using averaged outputs from 1,000 random forest regression models. The red line and font indicate the fitted linear model on estimated and observed R_{CH_4} values.

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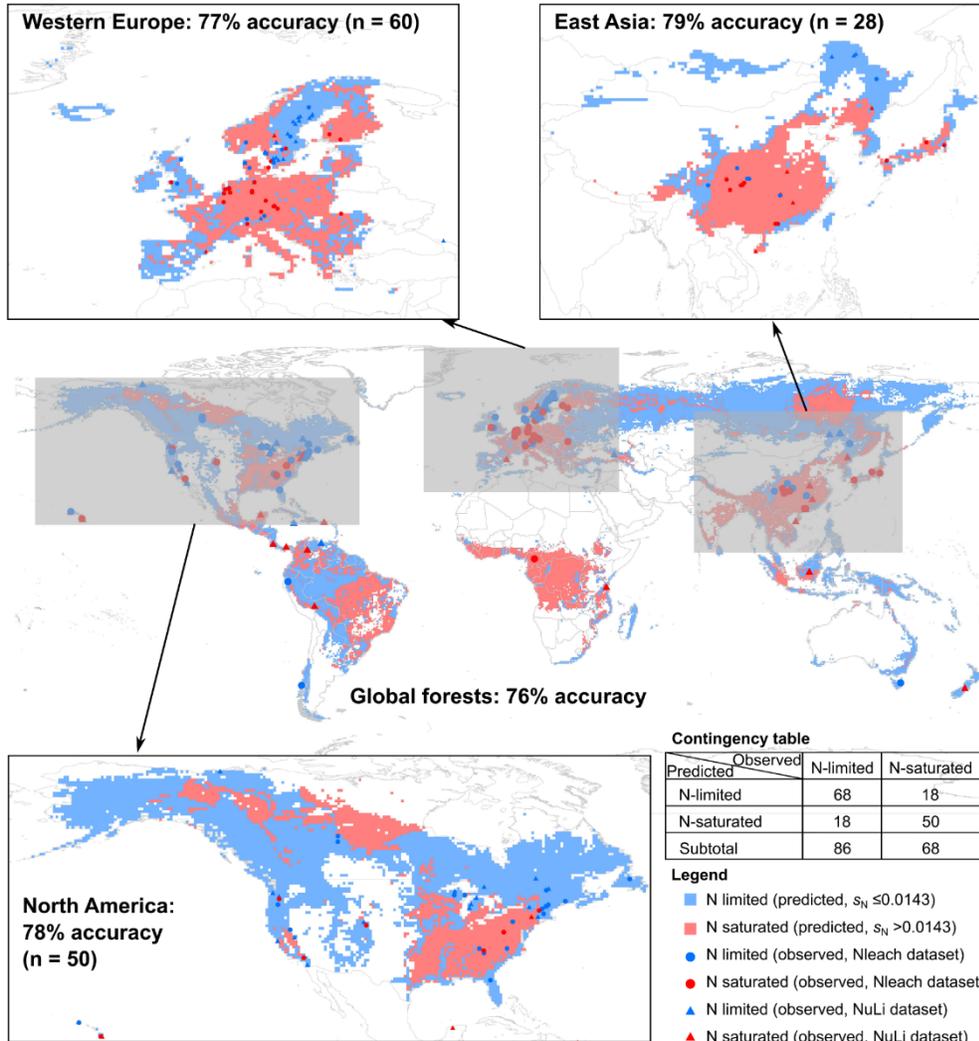


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Fig. S4. Comparing soil CH₄ flux (R_{CH_4}) estimated from different models built out of different training datasets. The sampling of observations to form a training (or testing) dataset was randomized by using different “seeds”. Each seed corresponds to a determined set of samples, and different seeds lead to different samples. In this study, we randomly used seeds “1111” and “1234” for sampling. This analysis was to ensure that the estimated R_{CH_4} values were not dependent on which data were used for training and testing the models, so that the derived spatial pattern of R_{CH_4} was robust on grid level.

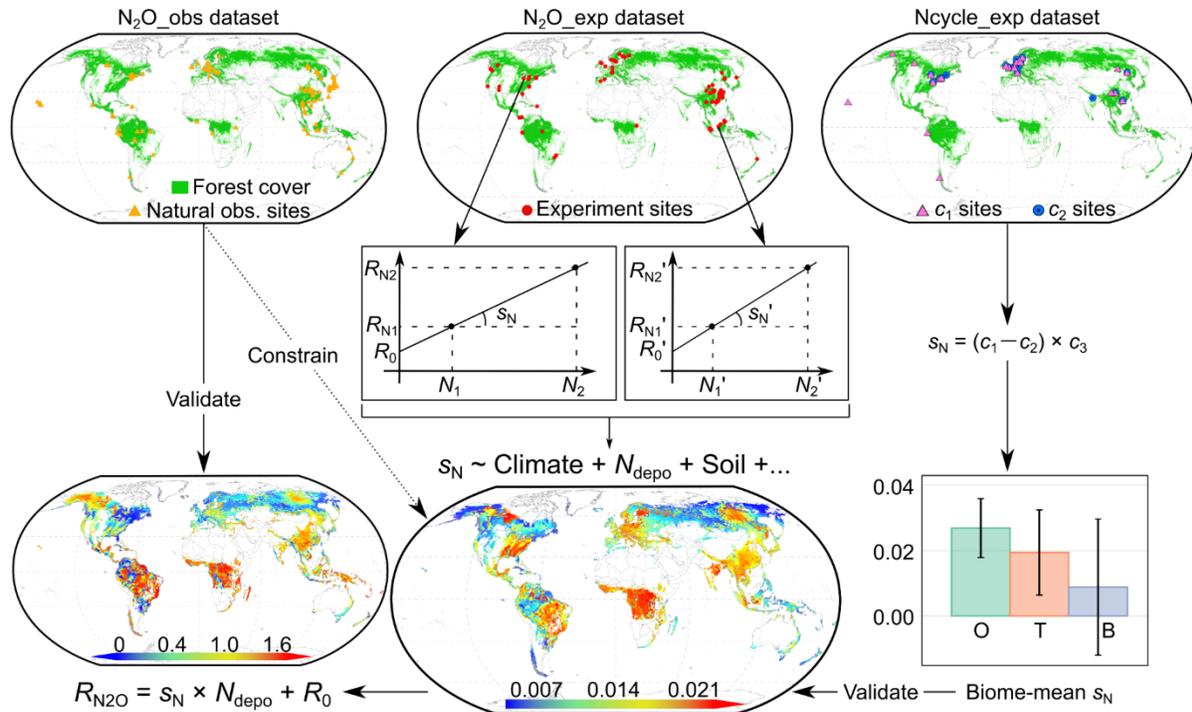


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 262 **Fig. S5.** Various forests are at different "stages" (in the stimulating-suppressing-weakening
 263 "three stages" framework), in accordance with the overall effects of N deposition on soil CH₄
 264 fluxes in the forests.
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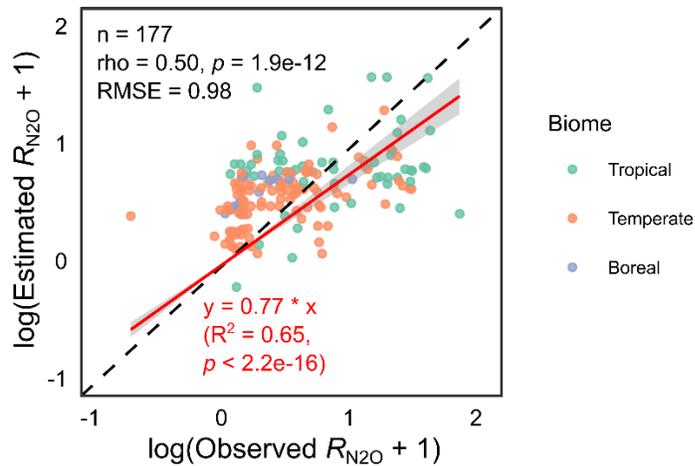


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Fig. S6. Classified N-limited and N-saturated forests based on the sensitivity of soil N₂O emission to N deposition (s_N) compared with field-observed N limitation or saturation status, with extra details in regions where field-observations were more abundant.

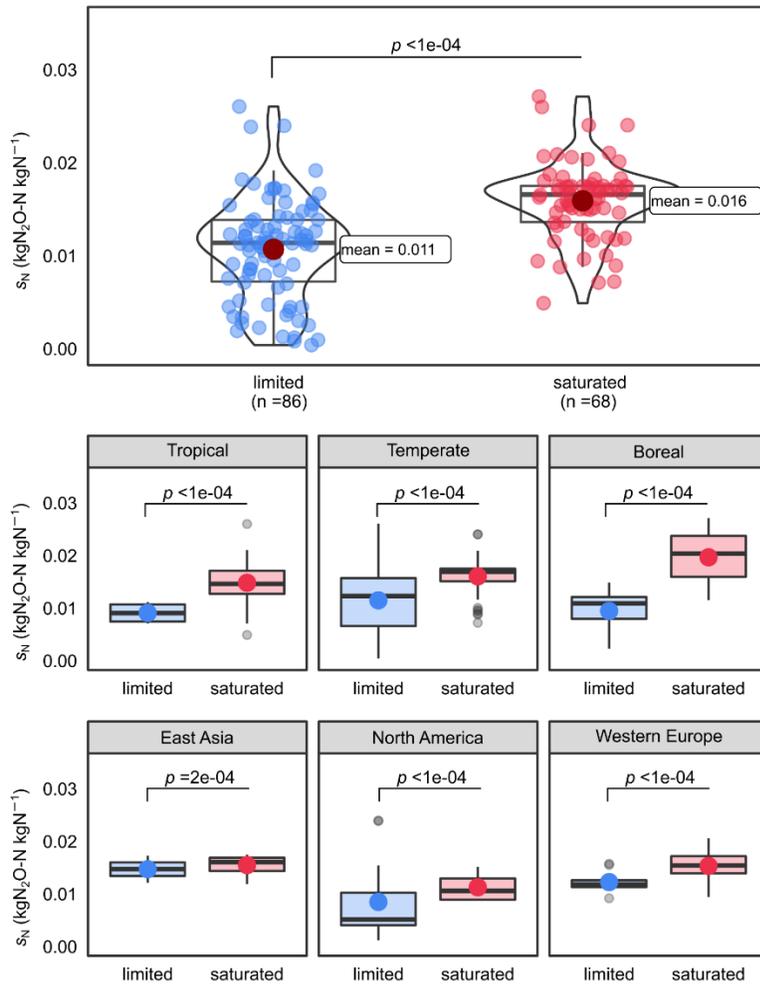


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 272 **Fig. S7.** Workflow illustrating the quantification and validation of the sensitivity of soil N₂O
 273 emission to N deposition (s_N) of global forests. N_1 and N_2 are different rates of low N input, and
 274 R_{N1} and R_{N2} are the corresponding soil N₂O emission rates. N_{depo} : N deposition rate ($\text{kgN ha}^{-1} \text{yr}^{-1}$);
 275 R_0 : background soil N₂O emission rate ($\text{kgN}_2\text{O-N ha}^{-1} \text{yr}^{-1}$); c_1 : sensitivity of N loss to N
 276 deposition (kgN kgN^{-1}); c_2 : sensitivity of N leaching to N deposition (kgN kgN^{-1}); c_3 : ratio of N₂O
 277 to other gaseous end-products from nitrification and denitrification processes ($\text{kgN}_2\text{O-N kgN}^{-1}$).
 278 O: Tropical; T: Temperate; B: Boreal.
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Fig. S8. Comparing estimated and observed soil N₂O emission rates (R_{N_2O}). Observations were aggregated to 0.5°×0.5° grids to match with the spatial resolution of the environmental factors. Each point represents a grid-year. Points of different colors represent grid-years in different biomes. Dashed black line is the 1:1 line. The red line and fonts show a linear regression model on estimated and observed R_{N_2O} .



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Fig. S9. Comparing the sensitivity of soil N₂O emission to N deposition (s_N) of N-limited and N-saturated forests on global and regional scales.

291 **Table S1.** Parameters of segmented linear regression models on soil CH₄ flux (R_{CH4}) and N input
 292 rate.

No.	Model ($R_{CH4} \sim$ N input rate)	Parameters
①	$y = -0.037*x - 2.45$	$n = 53, R^2 = 0.01, p = 0.44$
②	$y = 0.045*x - 5.75$	$n = 49, R^2 = 0.06, p = 0.09$
③	$y = -0.004*x - 0.73$	$n = 29, R^2 = 0.00, p = 0.80$
④	$y = 0.096*x - 5.28$	$n = 121, R^2 = 0.10, p = 0.0003$
⑤	$y = -0.006*x - 1.53$	$n = 196, R^2 = 0.03, p = 0.02$

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295 **Table S2.** Parameters of the constructed random forest regression models.

Model	$R_{CH_4} \sim \text{MAT} + \text{MAT.cv} + \text{MAP} + \text{MAP.cv} + N_{depo} + N_{depo.cv} + \text{Sand} + \text{Clay} + s_N$
	mtry
Parameters	ntree
	Number of runs

296 R_{CH_4} : soil CH₄ emission rate; MAT: mean annual temperature; MAP: mean annual precipitation; N_{depo} : mean
 297 annual N deposition; Sand: soil sand content; Clay: soil Clay content; MAT.cv, MAP.cv and $N_{depo.cv}$ are the
 298 corresponding coefficients of temporal variation; s_N : sensitivity of soil N₂O emission to N deposition,
 299 which indicates soil N limitation or saturation status. The predictors were selected based on mechanistic
 300 relevance and data availability.
 301

302 **Table S3.** Linear models on soil N₂O emission rate (R_{N_2O}) and N input rate (model: $R_{N_2O} \sim$ N input rate) built with low N input data (N
303 addition rate ≤ 150 kgN ha⁻¹ yr⁻¹) from global forest experiment sites, and the derived sensitivity (s_N) of soil N₂O emission to N
304 deposition and background N₂O emission rate (R_0).

N o.	Longitude range	Latitude range	Biome	s_N	R_0	n	adj.R ²	p value	References
1	(19,19.5)	(64,64.5)	Boreal	0.002	0.045	2	NA	NA	(Rutting et al., 2021)
2	(30.5,31)	(62.5,63)	Boreal	0.025	5.132	4	0.14	0.347	(Regina et al., 1998)
3	(22.5,23)	(62,62.5)	Boreal	0.013	0.538	2	NA	NA	(Ojanen et al., 2019)
4	(8,8.5)	(58.5,59)	Boreal	0.026	0.343	6	0.57	0.052	(Sitaula et al., 1995a, 1995b)
5	(-3.5,-3)	(55.5,56)	Temperate	0.02	0.258	6	0.18	0.224	(U. M. Skiba et al., 1998)
6	(-3,-2.5)	(55.5,56)	Temperate	0.006*	-0.009	6	0.73	0.019	(U. Skiba et al., 1999; U. M. Skiba et al., 1998)
7	(1.5,2)	(52.5,53)	Temperate	0.004	0.233	2	NA	NA	(U. M. Skiba et al., 1998)
8	(9.5,10)	(51.5,52)	Temperate	0.042*	0.51	10	0.48	0.015	(Borken et al., 2002; Brumme & Beese, 1992; Marife D Corre et al., 2003)
9	(128.5,129)	(47,47.5)	Boreal	0.015	0.777	11	0.02	0.300	(He, 2015; L. Song et al., 2017; Tian et al., 2018)
10	(8.5,9)	(47,47.5)	Temperate	0.003	-0.062	4	0.63	0.134	(Krause et al., 2013)
11	(-80.5,-80)	(43.5,44)	Temperate	0.009	1.374*	4	0.79	0.073	(Lutes et al., 2016)
12	(-72.5,-72)	(43,43.5)	Temperate	0.012	-0.216	2	NA	NA	(M. S. Castro et al., 1992)
13	(141,141.5)	(43,43.5)	Temperate	0.025	1.647	2	NA	NA	(Y. S. Kim et al., 2012)
14	(-72.5,-72)	(42.5,43)	Temperate	0.001	0.074	6	0.05	0.323	(Richard D. Bowden et al., 1991)
15	(128,128.5)	(42,42.5)	Temperate	0.01	0.67	2	NA	NA	(Geng et al., 2017)
16	(127.5,128)	(41.5,42)	Temperate	0.029	2.287	13	0.11	0.141	(Bai et al., 2014; Cheng et al., 2016; B. Peng et al., 2021)
17	(-80.5,-80)	(41.5,42)	Temperate	0.003	0.217	2	NA	NA	(R. D. Bowden et al., 2000)
18	(-4,-3.5)	(40,40.5)	Temperate	0.001*	0.026*	4	0.95	0.017	(Lafuente et al., 2020)
19	(112,112.5)	(36.5,37)	Temperate	0.056	2.754	3	0.98	0.068	(H. Yu, 2019)
20	(111,111.5)	(31.5,32)	Temperate	0.013**	0.483	27	0.28	0.003	(Zhaolan Lin, 2013; Zhaolan Lin et al., 2012; R. Wang, 2012; Xu et al., 2017)
21	(110,110.5)	(31.5,32)	Temperate	0.023	-0.31	4	0.54	0.166	(Pan, 2013)
22	(120.5,121)	(30.5,31)	Temperate	0.017	1.135	4	0.51	0.181	(Tu & Zhang, 2018)
23	(119.5,120)	(30,30.5)	Temperate	0.003	1.238***	16	0.01	0.308	(X. Chen, 2014; X. Chen et al., 2014; Ziwen Lin, 2019; X. Z. Song et al., 2020; Z. Wang, 2014)
24	(120,120.5)	(30,30.5)	Temperate	0.012**	0.834*	12	0.64	0.001	(J. Zhang, 2013; J. Zhang et al., 2013)
25	(106.5,107)	(29.5,30)	Temperate	0.025*	0.875*	3	1	0.018	(Xie et al., 2018)

26	(115.5,116)	(29.5,30)	Temperate	0.012	2.025	6	0.14	0.248	(C. Li et al., 2019)
27	(116.5,117)	(28,28.5)	Temperate	0.013	0.16	2	NA	NA	(Fan et al., 2020)
28	(118,118.5)	(27,27.5)	Tropical	0.015	1.948	9	0.12	0.190	(S. Chen, 2012)
29	(115,115.5)	(26.5,27)	Tropical	0.026***	-0.092	54	0.47	<0.001	(Dang, 2015; X. Li, 2017; X. Y. Li et al., 2015; Sun & Zhang, 2015; J. Wang, 2016; L. Wang, 2015; L. Wang et al., 2016; Y. Wang, 2015; Y. S. Wang et al., 2016; L. Zhang, 2013)
30	(117,117.5)	(26,26.5)	Tropical	0.007	0.5	3	0.55	0.313	(Wu, 2018)
31	(118,118.5)	(25.5,26)	Tropical	0.012	0.601	4	0.33	0.257	(Yuan, 2016)
32	(113,113.5)	(23.5,24)	Tropical	0.014	-0.226	3	0.77	0.220	(Cai, 2013)
33	(112.5,113)	(23,23.5)	Tropical	0.027*	0.19	22	0.15	0.041	(H. Chen et al., 2016; Gao et al., 2017; Mo et al., 2006)
34	(112.5,113)	(22.5,23)	Tropical	0.004	1.919***	14	0.11	0.129	(W. Zhang et al., 2014)
35	(106.5,107)	(22,22.5)	Tropical	0.012***	-0.038	8	0.84	0.001	(Hong, 2015)
36	(107,107.5)	(22,22.5)	Tropical	0.043*	-0.089	10	0.51	0.013	(R. Li et al., 2014, 2015; Yang, 2015; Kai Zhang et al., 2015)
37	(107.5,108)	(22,22.5)	Tropical	0.007**	0.589**	4	0.98	0.007	(K. Zhang et al., 2017)
38	(101,101.5)	(21.5,22)	Tropical	0.037	2.101*	9	0.18	0.144	(Yan, 2006; Zhou et al., 2016)
39	(110.5,111)	(21,21.5)	Tropical	0.018	3.195	3	0.69	0.256	(F. M. Wang et al., 2014)
40	(-80,-79.5)	(9,9.5)	Tropical	0.021**	0.674	8	0.71	0.005	(M. D. Corre et al., 2014; Koehler et al., 2009)
41	(-82.5,-82)	(8.5,9)	Tropical	0.019	1.063	8	0.32	0.083	(M. D. Corre et al., 2014; Koehler et al., 2009)
42	(116.5,117)	(6,6.5)	Tropical	0.007**	0.517**	10	0.61	0.005	(Hall et al., 2004)
43	(31.5,32)	(1.5,2)	Tropical	0.018***	1.756***	4	1	0.001	(Tamale et al., 2021)
44	(102,102.5)	(-1.5,-1)	Tropical	0.022**	0.919*	7	0.84	0.002	(Aini et al., 2015)
45	(-79.5,-79)	(-4,-3.5)	Tropical	0.005	0.135	3	0.44	0.356	(Muller et al., 2015)
46	(-79,-78.5)	(-4.5,-4)	Tropical	0.006	0.471	3	0.5	0.333	(Muller et al., 2015)
47	(-79.5,-79)	(-4.5,-4)	Tropical	0.006	-0.11	3	0.95	0.106	(Muller et al., 2015)

305 * $p < 0.05$; ** $p < 0.01$; *** $p < 0.001$; NA, not applicable

306 **Table S4.** Generalized linear models on environmental factors and the sensitivity (s_N) of
 307 soil N₂O emission to N deposition and the background N₂O emission rate (R_0).

	Estimate	SE	t	p
Refined model on s_N[†] (Deviance explained = 91.1%, n=46)				
Clay	4.77E-03	1.83E-03	2.605	0.013*
Sand	3.15E-03	9.20E-04	3.419	0.001**
log(N_{depo})	2.01E-02	1.14E-02	1.769	0.085
Clay × log($N_{depo.cv}$)	2.13E-03	9.35E-04	2.282	0.028*
Sand × log($N_{depo.cv}$)	1.17E-03	3.82E-04	3.056	0.004**
Clay × Sand	-1.90E-04	6.94E-05	-2.735	0.009**
Clay × Sand × log($N_{depo.cv}$)	-1.14E-04	3.66E-05	-3.112	0.003**
Refined model on R_0[‡] (Deviance explained = 43.2%, n = 45)				
log($N_{depo.cv}$)	1.99E-01	9.56E-02	2.084	0.043*
MAT × Sand × Clay	3.04E-06	5.99E-07	5.072	0.000***
MAP × MAP.cv × log(N_{depo})	-8.31E-04	2.91E-04	-2.854	0.007**

308 MAT: mean annual temperature; MAP: mean annual precipitation; N_{depo} : mean annual N
 309 deposition; Sand: soil sand content; Clay: soil clay content.
 310 [†] $s_N \sim (\text{Clay} + \text{Sand} + \log(N_{depo}) + \text{Clay} \times \log(N_{depo.cv}) + \text{Sand} \times \log(N_{depo.cv}) + \text{Clay} \times \text{Sand} + \text{Clay}$
 311 $\times \text{Sand} \times \log(N_{depo.cv}))^2$
 312 [‡] $R_0 \sim \text{EXP}(\log(N_{depo.cv}) + \text{MAT} \times \text{Sand} \times \text{Clay} + \text{MAP} \times \text{MAP.cv} \times \log(N_{depo})) - 0.5$
 313 * $p < 0.05$; ** $p < 0.01$; *** $p < 0.001$
 314

315 **Data Set S1. (separate file)**
316 Compiled dataset on soil CH₄ flux from N addition experiments in global forests
317 (CH₄_exp dataset in main text).
318
319 **Data Set S2. (separate file)**
320 Compiled data on soil CH₄ flux under natural conditions in global forests (CH₄_obs
321 dataset in main text).
322
323 **Data Set S3. (separate file)**
324 Compiled dataset on soil N₂O emission rate from N addition experiments in global
325 forests (N₂O_exp dataset in Text S1).
326
327 **Data Set S4. (separate file)**
328 Compiled data on soil N₂O emission rate under natural conditions in global forests
329 (N₂O_obs dataset in Text S1).
330
331 **Data Set S5. (separate file)**
332 Compiled dataset on N loss rate, N leaching rate and change rate of soil N pool from N
333 addition experiments in global forests (Ncycle_exp dataset in Text S1).
334
335 **Data Set S6. (separate file)**
336 Compiled dataset on global forest N saturation status (limited or saturated) indicated by
337 N leaching rate (Nleach dataset in Text S1).
338
339 **Data Set S7. (separate file)**
340 An existing dataset from Du et al. (2020) on global forest N saturation status (limited or
341 saturated) indicated by plant growth response to N input (NuLi dataset in Text S1).
342
343 **Data Set S8. (separate file)**
344 Data on environmental factors (MAT, MAP, N deposition rate, etc.) in global forests,
345 extracted from spatial datasets mentioned in Methods section.
346
347 **Data Set S9. (separate file)**
348 Global forest soil methane budgets estimated in previous studies.
349
350 **Code S1. (separate file)**
351 R code script used to carry out the data analysis processes, and produce the figures.
352
353
354