

1 **Differences in methanogenic pathways and methanogenic communities in**
2 **paddy soils under three typical cropping modes**

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16 **Key Points:**

- 17 • The CH₄ production potential, methanogenic pathways, and communities differ in three
18 cropping mode soils.
- 19 • Paddy soil properties, mainly soil pH and soil texture regulate methanogenesis by changing
20 methanogenic communities.

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23 Abstract

24 Microbial methane (CH₄) production varies among different cropping modes, which has important
25 implications for how to reduce CH₄ emissions from paddy fields. However, little is known about the
26 values of anaerobically produced δ¹³CH₄, methanogenic pathways, and their dominant communities in
27 different paddy soils. Through anaerobic incubation experiments and the stable carbon isotope with
28 fluoromethane inhibitor method, CH₄ production potential (MPP), the relative contribution of
29 acetoclastic methanogenesis (*f*_{ac}), and the abundance and community composition of methanogens in
30 paddy soils were measured under three typical cropping modes (Rice-Wheat, RW; Rice-Fallow, RF;
31 Double-Rice, DR) in China. The results showed that MPP was 30.7 μg CH₄ g⁻¹ d⁻¹ in DR soil, 57%
32 and 66% higher than that in RW and RF soils, respectively, possibly due to the lower pH and higher
33 abundance of *mcrA* gene. Moreover, RF soil had the highest produced δ¹³CH₄ value (-43.9‰) and the
34 lowest produced δ¹³CO₂ value (-26.3‰). Based on the carbon isotope fractionations associated with
35 H₂/CO₂-dependent methanogenesis (1.049–1.062), the values of *f*_{ac} estimated in RF soil (80–98%)
36 were much higher than that in RW (39–60%) and DR (52–75%) soils. It could be supported by that the
37 *Methanosarcina* (acetoclastic methanogens) were dominant in RF soil while *Methanosarcina* and
38 *Methanobacterium* (hydrogenotrophic methanogens) dominated in RW and DR soils. Redundancy
39 analysis revealed that the community structure of methanogens was significantly affected by soil pH,
40 indicating that the differences in methanogenic pathways under the three typical cropping modes
41 might be caused by the changes in community composition driven by soil pH. The findings suggest
42 that soil pH-induced methanogenic abundance and community composition drive paddy MPP and
43 methanogenic pathways, which would provide important insights into the CH₄ reduction in paddies.

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45 Plain Language Summary

46 In paddy soils, the microbial methanogenesis and its mediated CH₄ production potential are various
47 due to their various rice-based cropping modes. However, the methanogenic pathways, microbial
48 mechanisms, and their responses to key influencing factors under different cropping modes in China
49 are still poorly documented. We investigated the differences in pathways of CH₄ production,

50 methanogenic communities, and their responses to the corresponding soil properties under three
51 typical cropping modes, i.e. Rice-Wheat, Rice-Fallow, and Double-Rice rotation systems. Our results
52 demonstrated that there were significant differences in acetoclastic methanogenesis and dominant
53 methanogenic communities in the three cropping mode paddy soils, which could mainly be caused by
54 soil pH. This study provides a new perspective and further understanding of the methanogenic
55 pathways and their microbial mechanisms in different cropping modes.

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57 **Keywords**

58 Rice-based cropping modes; Methanogenic pathway; Methanogenic microbial community; Soil pH;

59 Carbon isotopic fractionation

60 **1. Introduction**

61 As the second-most powerful greenhouse gas after carbon dioxide (CO₂), methane (CH₄) has
62 received great global concern to address climate change (Liu et al., 2022). Rice fields are important
63 anthropogenic sources of worldwide CH₄ emissions and are estimated at 24–39 Tg yr⁻¹, accounting for
64 12–21% of the global agriculture emission per year (Saunio et al., 2020). Except for India (about 7.4
65 Tg yr⁻¹), the largest CH₄ emissions are found in China (about 6.2 Tg yr⁻¹), which contributes to about
66 21% of the total global budget from paddy fields (Carlson et al., 2016). Therefore, paddy fields in
67 China play a critical role in the global carbon cycle (Qi et al., 2021), and exploring the mechanism of
68 CH₄ production and CH₄ emission reduction from paddy fields is an important part of the current
69 carbon neutrality in China.

70 CH₄ emissions from rice fields are the net effects of three processes: CH₄ production, oxidation, and
71 transport from the soil into the atmosphere (Cai et al., 2009), of which CH₄ production is the basic
72 prerequisite for CH₄ emission (Le Mer and Roger, 2001). Methanogens use CH₄ precursors to produce
73 CH₄ under strictly anaerobic conditions (Conrad, 2007). The acetate and CO₂/H₂ were the two major
74 precursors of methanogenesis in paddy ecosystems, and the corresponding methanogenic pathways of
75 paddy soils were acetoclastic and hydrogenotrophic methanogenesis, respectively (Glissmann and
76 Conrad, 2000; Ji et al., 2018b). The acetoclastic and hydrogenotrophic methanogenic pathways would
77 fractionate organic material of a similar signature in ways, resulting in different signatures of product
78 CH₄ (Ji et al., 2018a; Whiticar et al., 1986). This difference in isotopic fractionation could in principle
79 be used to estimate the relative contribution of the two methanogenic pathways to total CH₄
80 production, which was determined by specific inhibition of acetoclastic methanogenesis with methyl
81 fluoride (CH₃F) (Conrad and Klose, 1999; Glissmann and Conrad, 2000).

82 China has a large rice planting area with wide distribution and diverse cropping modes, mainly
83 including three cropping modes of Rice-Wheat (RW), Rice-Follow (RF), and Double-Rice (DR).
84 Different cropping modes with various water management and fertilization conditions lead to diverse
85 soil properties (e.g. soil organic matter, soil pH, and soil organic acid, etc.) in paddy soils, changing
86 the supply of methanogenic substances, the community structure, and composition of methanogens,
87 then probably affecting CH₄ production and methanogenic pathways (Fu et al., 2021; Sun et al., 2018;

88 Wang et al., 1993; Yang et al., 2021). The RF paddy is mainly distributed in the hilly mountainous
89 areas of southwest China, which is flooded all year round resulting in higher CH₄ emissions than other
90 paddy fields (Cai et al., 2000; Mei et al., 1998). The DR paddy was mainly found in the central region
91 south of the Yangtze River, the Pearl River basin, and the Hainan Province. In general, these regions
92 have relatively more precipitations, higher temperatures, and longer duration of rain, which provide
93 favorable conditions for CH₄ production and emission (Mei et al., 1998). However, most of the
94 previous studies on CH₄ production and methanogenic pathway in Chinese paddy soils have focused
95 on the RW cropping mode (Zhang et al., 2011; Zhang et al., 2012), and little is known about RF and
96 DR soils with high CH₄ emissions and large emission reduction potential (Cai et al., 2000; Chen et al.,
97 2013). Furthermore, paddy soils under different cropping modes have various methanogenic
98 communities due to their specific soil conditions, which leads to changes in methanogenic pathways
99 (Jiang et al., 2022; Zhang et al., 2017). However, there is no clear information so far on the targeted
100 comparative study about the microbiological mechanisms between methanogenesis pathways and
101 environmental factors in paddy soils under different cropping modes. Therefore, it is of great guiding
102 significance to carry out the above research in-depth to take appropriate technical measures of
103 microbial regulation for emission reduction.

104 In this study, we hypothesized that the various soil properties in different cropping modes directly
105 or indirectly affect the abundance, diversity, and community composition of methanogens, thereby
106 regulating the CH₄ production and the methanogenic pathway. Therefore, a microcosm experiment
107 was conducted to determine the CH₄ production potential (MPP), carbon isotope composition ($\delta^{13}\text{C}$) of
108 CH₄ production, and the dynamics of communities and composition of methanogens in paddy soils of
109 three cropping modes (RW, RF, and DR) in China. DOC, acetate, soil pH, and other soil
110 physicochemical properties were measured to probe the environmental factors of the methanogenic
111 pathway with different cropping modes. The prime objectives of this research were: 1) to explore MPP
112 and the related key influencing factors in paddy soils with different cropping modes, 2) to reveal the
113 relevant composition and diversity of methanogenic communities in these paddy soils, and 3) to
114 investigate the CH₄ production pathways and the corresponding microbial mechanisms in paddy soils
115 with different cropping modes.

116 **2. Materials and methods**

117 **2.1 Soil sample**

118 The soils were sampled at the mature stage of rice from paddy fields under three typical Chinese
119 cropping modes: RW [located in Jurong City, Jiangsu Province, East China (31°57' N, 119°10' E)], RF
120 [located in Jianyang City, Sichuan Province, Southwest China (30°39' N, 104°55' E)], and DR [located
121 in Yingtan City, Jiangxi Province, Southeast China (28°24' N, 117°03' E)]. Soil samples were collected
122 following Soil Agro-Chemical Analyses procedures (Lu, 2000), for each model, cores (0–20 cm) were
123 taken from three representative rice fields according to the distribution method. Ten fresh soil samples
124 were then pooled to form a composite sample. Each soil was sieved (< 2 mm) and split into two
125 subsamples. One subsample was air-dried and stored for incubation studies at room temperature and
126 the other for soil physicochemical properties analysis stored at 4°C.

127 **2.2 Soil incubation**

128 The incubation procedure was the same as described by Ji et al. (2018a). Soil slurries of three
129 cropping modes (RW, RF, and DR) were prepared by mixing 20 g of air-dried soil with 20 mL of
130 deionized, sterile, and anoxic water at a ratio of 1:1. Pre-incubation of the slurries occurred in 120 mL
131 serum bottles to activate soil activity, closed with butyl rubber stoppers at 25 °C for 3 days. After pre-
132 incubation, the slurries were flushed with N₂ consecutively six times to purge the air in the mixture of
133 residual O₂ and CH₄ (Cai et al., 2009). Then, they have been incubated anaerobically in the dark at
134 25 °C for 34 days in the absence (no CH₃F addition, treatment control) and the presence of CH₃F [the
135 gas headspace of bottles was supplemented with 1% (treatment 1%) and 2% (treatment 2%)] in
136 triplicates. Notably, acetoclastic methanogens have been described to be inhibited at much lower
137 concentrations of CH₃F than hydrogenotrophic methanogens (Conrad and Klose, 1999). Therefore,
138 when CH₃F is added at a low concentration (0.5–2%), it is generally considered to be an inhibitor of
139 acetoclastic methanogenesis (Conrad and Klose, 1999; Ji et al., 2018a; b).

140 **2.3 Sample collection and analyses**

141 Gas samples were collected every three to four days during the 34-day incubation. The CH₄ and
142 CO₂ concentrations were analyzed with a gas chromatograph (GC) (Agilent 7890B, USA) equipped
143 with a flame ionization detector (FID). After 34 days of incubation, gas samples were collected to

144 analyze $\delta^{13}\text{C}$ of CH_4 and CO_2 . The stable carbon isotopes of CH_4 and CO_2 were analyzed using the
145 continuous flow technique on a Finnigan MAT 253 Plus Isotope Ratio Mass Spectrometer (Thermo
146 Fisher Scientific, Waltham, USA).

147 Soil properties were determined following Soil Agro-Chemical Analyses procedures (Lu, 2000).
148 Soil pH was measured in a 1:2.5 (v/v) ratio of soil to water (deionized water). Total soil organic carbon
149 (SOC) was analyzed by wet digestion with $\text{H}_2\text{SO}_4\text{-K}_2\text{Cr}_2\text{O}_7$. Active Fe (Fe(III)) and Mn (Mn(IV)) were
150 extracted with $\text{H}_2\text{C}_2\text{O}_4\text{-(NH}_4)_2\text{C}_2\text{O}_4$, and were determined by Inductively Coupled Plasma Mass
151 Spectrometer (Nexion2000, America). NO_3^- and SO_4^{2-} were extracted in a 2 M KCl solution at a
152 soil/water ratio of 1:5, and measured using an ion chromatograph (ICS-5000+, America). Soil cation
153 exchange capacity (CEC) was determined after extraction with 1 mol L^{-1} ammonium acetate. Total
154 nitrogen (TN) was analyzed by an elemental analyzer (Vario MAX).

155 At the beginning and end of all incubation, soil samples were also collected and a portion of them
156 was analyzed DOC and acetic acid content by a total organic carbon/total nitrogen analyzer (Multi NC
157 3100) and by high-pressure liquid chromatography (HPLC) (ELSD/UV, Agilent HPLC 1260),
158 respectively. The other portion was freeze-dried and extracted for total DNA using a FastDNA[®] SPIN
159 Kit for Soil (MP Biomedicals LLC, USA). DNA concentration and quality were determined using a
160 NanoDrop spectrophotometer (Thermo Fisher Scientific, MA, USA). The abundances of *mcrA* gene
161 were measured by Majorbio (Shanghai, China) via Real-Time qPCR on an Applied Biosystems (ABI)
162 7300 Real-Time PCR system (Thermo Fisher Scientific, MA, USA).

163 The primer pairs *mlas-mod/mcrA-rev* were used to quantify the copy numbers of *mcrA* gene (Angel
164 et al., 2011). The PCR mixture consisted of 10 μL of ChamQ SYBR Color qPCR Master Mix (2 \times), 0.8
165 μL of each primer, 0.4 μL of ROX Reference Dye 1 (50 \times), and 2 μL of DNA template, brought to a
166 final volume of 6 μL with sterile water. The primers *MLfF* and *MLfR* were used to amplify the *mcrA*
167 gene fragments targeting the 469–490 bps region. Then a library was constructed by pooling equal
168 amounts of individual barcoded amplicons of *mcrA* and sequenced on an ILLUMINA MISEQ PE250
169 system using a 2 \times 300 cycle combination mode by Majorbio (Shanghai, China).

170 2.4 Bioinformatics analysis

171 After the sequencing was completed, the raw bacterial sequences were processed using the QIIME
172 pipeline (Caporaso et al., 2010). All reads were subjected to quality control, de novo chimera filtering,
173 singleton filtering, and operational taxonomic units (OTUs) clustering according to the UPARSE
174 pipeline. The effective sequences were clustered into OTUs using a 97% identity threshold, and the
175 chimeras were detected and deleted via USEARCH (Edgar et al., 2011). Sequences that cannot be
176 classified into any known group were considered unclassified, and groups with < 1% average
177 proportion were merged into the “others” taxa. All sequences have been deposited in the NCBI
178 sequence Read Archive (SRA) database under the accession number PRJNA842891.

179 2.5 Calculations

180 2.5.1 CH₄ production potential and soil oxidation capacity

181 CH₄ production potential (MPP) under anaerobic incubation was determined, which was calculated
182 by the following equation (Zhang et al., 2011; Zhang et al., 2010):

$$183 \quad P = dc/dt \times V_H/W_S \times M_r/M_V \times 273/(273+25) \quad (1)$$

184 where P is MPP ($\mu\text{g g}^{-1} \text{d}^{-1}$), dc/dt is the rate of CH₄ accumulation ($\mu\text{L L}^{-1} \text{d}^{-1}$), V_H is the headspace
185 volume of the serum bottle (L), W_S is dry soil weight (g), M_r is the relative molecular mass of CH₄ (g),
186 M_V is the gas volume of an ideal gas (L).

187 Soil oxidation capacity (OXC) was calculated using a modification of the equation given by Zhang
188 et al. (2009b):

$$189 \quad \text{OXC} = 5[\text{NO}_3^-] + 2[\text{Mn(IV)}] + [\text{Fe(III)}] + 8[\text{SO}_4^{2-}] \quad (2)$$

190 where brackets denote molar concentrations (mol kg^{-1}).

191 2.5.2 The relative contribution of acetoclastic methanogenesis

192 The relative contribution of CO₂/H₂ and CH₃COOH to CH₄ production in paddy soils was
193 determined by applying the stable carbon isotope method (Sugimoto and Wada, 1993), which assumed
194 that total CH₄ production was equal to the sum of acetate fermentation (CH_{4(ac)}) and CO₂/H₂ reduction
195 (CH_{4(CO₂)}):

$$196 \quad \text{CH}_4 = \text{CH}_{4(\text{ac})} + \text{CH}_{4(\text{CO}_2)} \quad (3)$$

197 The contribution of acetate to total CH₄ production (f_{ac}) was calculated by the following (Sugimoto

198 and Wada, 1993):

$$199 \quad f_{ac} = \text{CH}_4(\text{ac}) / [\text{CH}_4(\text{ac}) + \text{CH}_4(\text{CO}_2)] \times 100\% \quad (4)$$

200 The $\delta^{13}\text{CH}_4$ of the evolved CH_4 primarily depended on the relative contribution of the two main
201 methanogenic pathways, according to the isotopic mass balance (Tyler et al., 1997):

$$202 \quad \delta^{13}\text{CH}_4 = \delta^{13}\text{CH}_4(\text{ac}) \times f_{ac} + \delta^{13}\text{CH}_4(\text{CO}_2) \times (1 - f_{ac}) \quad (5)$$

203 where $\delta^{13}\text{CH}_4(\text{ac})$ refers to $\delta^{13}\text{C}$ values of CH_4 produced by acetate; $\delta^{13}\text{CH}_4(\text{ac})$ is assumed a fixed value
204 (Bilek et al., 1999; Nakagawa et al., 2002; Sugimoto and Wada, 1993; Tyler et al., 1997) or varies
205 within a certain range (Sugimoto and Wada, 1993), which assumed to be -37% and -43% (Zhang et
206 al., 2012; Zhang et al., 2016).

207 $\delta^{13}\text{CH}_4(\text{CO}_2)$ referred to $\delta^{13}\text{C}$ values of CH_4 produced by CO_2/H_2 reduction, which was calculated by
208 the following equation (Sugimoto and Wada, 1993):

$$209 \quad \delta^{13}\text{CH}_4(\text{CO}_2) = (\delta^{13}\text{CO}_2 + 1000) / \alpha_{(\text{CO}_2/\text{CH}_4)} - 1000 \quad (6)$$

210 where $\delta^{13}\text{CH}_4(\text{CO}_2)$ is calculated from $\delta^{13}\text{C}$ (‰) and $\alpha_{(\text{CO}_2/\text{CH}_4)}$ for CO_2 obtained from anaerobic
211 incubation of soils without the addition of CH_3F (Fig. S1). The carbon isotope fractionation factor
212 during H_2/CO_2 reduction ($\alpha_{(\text{CO}_2/\text{CH}_4)}$) for methanogenesis was defined by Nakagawa et al. (2002):

$$213 \quad \alpha_{(\text{CO}_2/\text{CH}_4)} = (\delta^{13}\text{CO}_2 + 1000) / [\delta^{13}\text{CH}_4(\text{CO}_2) + 1000] \quad (7)$$

214 where $\delta^{13}\text{CH}_4(\text{CO}_2)$ is the $\delta^{13}\text{C}$ of CH_4 obtained by anaerobic incubation of soil with the addition of
215 CH_3F (Fig. S1).

216 **2.6 Statistical analysis**

217 The correlations between different parameters were assessed using Spearman's correlation analysis.
218 Differences in MPP, $\delta^{13}\text{C}$ of CH_4 production, *mcrA* gene abundance, and the relative abundance of
219 dominant methanogens among treatments were analyzed using the one-way analysis of variance
220 (ANOVA). The alpha diversity of methanogens was evaluated by Shannon index and Chao 1
221 estimator. Redundancy analysis (RDA) was used to study the relationship between methanogenic
222 community composition and environmental factors. Analysis of similarities (ANOSIM) was
223 performed to determine the differences between and within groups, and the Bray-Curtis dissimilarity
224 matrix was used to perform nonmetric multidimensional scaling analyses (NMDS). Network analysis

225 about methanogens associations based on the Spearman correlation was performed using Microeco
226 bioinformatics cloud (<https://bioincloud.tech/>).

227 **3. Results**

228 **3.1 Soil properties**

229 The physicochemical properties of three cropping mode paddy soils were shown in Table 1. RF soil
230 was alkaline clay soil (pH = 7.70), and RW and DR soils were acidic silt and sandy soils, respectively.
231 The lowest soil pH (4.89) was detected in DR soil. RF soil possessed higher SOC, TN, clay, Mn(IV),
232 NO_3^- , SO_4^{2-} , and CEC content, while it showed much lower sand and Fe(III) content than that of the
233 other two soils. DR soil possessed little clay, Mn(IV), SO_4^{2-} , OXC, and CEC content, and RW soil
234 possessed little NO_3^- content. Spearman correlation coefficients indicated that soil pH was positively
235 correlated with CEC, and soil clay content ($p < 0.01$, Fig. 1).

236 **3.2 The cumulative concentration of CH_4 production and MPP**

237 The cumulative concentration of CH_4 production increased rapidly in the early stage of incubation,
238 then gradually stabilized in the later stage of incubation (Fig. 2A–C). However, it varied in paddy soils
239 with different cropping modes, with the maximum and minimum peak values of CH_4 production
240 cumulative concentrations in DR ($150\,460\ \mu\text{mol mol}^{-1}$) and RF ($80\,493\ \mu\text{mol mol}^{-1}$), respectively.
241 Compared to the control treatment, the addition of 1% and 2% CH_3F resulted in a pronounced partial
242 inhibition of CH_4 production, significantly reducing CH_4 production cumulative concentration (Fig.
243 2A–C). DR soil had the highest MPP at $30.7\ \mu\text{g g}^{-1}\ \text{d}^{-1}$. Compared with DR soils, the MPP of RW and
244 RF soils was lower by 57% and 66%, respectively ($p < 0.05$, Fig. 2D). The MPP of different cropping
245 modes with 1% and 2% CH_3F treatment ranged from 0.98 to $7.51\ \mu\text{g g}^{-1}\ \text{d}^{-1}$. Moreover, MPP
246 negatively correlated with soil pH and soil clay content ($p < 0.01$, Fig. 1).

247 **3.3 Soil DOC and acetate concentration**

248 The contents of DOC and acetate before incubation in the three cropping modes were 181.1 to 443.4
249 mg kg^{-1} and 28.24 to $157.3\ \text{mg kg}^{-1}$, respectively (Fig. 3). In addition, the DOC and acetate contents
250 decreased significantly after incubation ($p < 0.05$). The highest DOC content was observed in DR soil
251 ($255\ \text{mg kg}^{-1}$), which was 51% and 17% higher than that in RW and RF soils, respectively ($p < 0.05$,
252 Fig. 3A). The DOC contents of all cropping modes increased with the CH_3F addition ($p < 0.05$). The

253 acetate content ranged from 5.68 to 7.19 mg kg⁻¹, with the highest content in DR and the order by
254 DR > RW > RF ($p > 0.05$, Fig. 3B). Similar to DOC contents, the acetate contents of all cropping
255 mode soils increased with the CH₃F addition ($p < 0.05$).

256 3.4 The $\delta^{13}\text{CH}_4$, $\delta^{13}\text{CO}_2$, $\alpha_{(\text{CO}_2/\text{CH}_4)}$, and f_{ac} values

257 After anaerobic incubation, the values of $\delta^{13}\text{CH}_4$ in paddy soils under different cropping modes
258 ranged from -54.4‰ to -43.9‰, and the values of $\delta^{13}\text{CO}_2$ ranged from -26.3‰ to -18.7‰ (Table 2).
259 RF soil had the highest $\delta^{13}\text{CH}_4$ value, which was more positive than RW and DR soils by 10.5‰ and
260 7.8‰, respectively ($p < 0.05$). However, it exhibited a much lower $\delta^{13}\text{CO}_2$ value in RF soil, being
261 more negative than RW and DR by 7.6‰ and 5.2‰, respectively ($p < 0.05$). Both $\delta^{13}\text{CH}_4$ and $\delta^{13}\text{CO}_2$
262 values in the addition of CH₃F were far lower than that of the control treatment ($p < 0.05$), with values
263 of -82.3‰ to -68.0‰ and -28.2‰ to -21.1‰, respectively.

264 The values of $\alpha_{(\text{CO}_2/\text{CH}_4)}$ in RW, RF, and DR soils ranged from 1.049 to 1.050 and from 1.057 to
265 1.062 in the addition of 1% and 2% CH₃F, respectively (Table 2). Compared to DR soil, the $\alpha_{(\text{CO}_2/\text{CH}_4)}$
266 values of RW and RF soils were significantly lower ($p < 0.05$). The f_{ac} values for RW, RF, and DR
267 soils ranged from 39% to 60%, from 80% to 98%, and from 52% to 75% when $\delta^{13}\text{CH}_{4(\text{ac})}$ values were
268 -37‰ to -43‰, respectively. There were significant differences in f_{ac} depending on the cropping
269 mode, and f_{ac} values for RF soil were much larger than those for RW and DR soils ($p < 0.05$).

270 3.5 The abundance and community composition of methanogens

271 RF soil had the lowest *mcrA* gene abundance among the three cropping modes (Fig. 4). Compared
272 to the RF soil, the abundance of *mcrA* gene in RW and DR soils increased by 29% and 40% ($p < 0.05$),
273 respectively. The addition of CH₃F significantly reduced *mcrA* gene abundance by 51–90% ($p < 0.05$).
274 The abundance of *mcrA* gene had a strong positive correlation with the MPP, while significantly
275 negatively correlated with soil pH, CEC, and clay content ($p < 0.001$, Fig. 1).

276 The number of OTUs in a single sample ranged from 168 to 275 (Table 3). Further, the Shannon
277 indices and Chao 1 estimators were calculated to evaluate the richness and diversity of methanogens in
278 soils under different cropping modes. A significant difference in Shannon indices was observed in RF
279 soil compared with RW and DR soils ($p < 0.05$). The number of OTUs and Chao 1 estimator in RW

280 soil was significantly higher than that in RW and DR soils ($p < 0.05$). Correlation analysis showed that
281 the number of OTUs and Shannon index negatively correlated with the MPP ($p < 0.01$), whereas the
282 Chao 1 estimator positively correlated with f_{ac} ($p < 0.05$, Fig. 1).

283 The NMDS based on Bray-Curtis distances indicated that there was a significant variation in the
284 community structure of methanogens in paddy soils under the three different cropping modes (Fig. S2,
285 $p < 0.001$). To more completely interpret the composition of the methanogenic community under three
286 cropping modes, the relative abundance of methanogens at which was given as the percentage of
287 different methanogens in total methanogens provided (Figs. 5 and S3). No matter at order, family, and
288 genus levels, the methanogens in different cropping mode soil samples varied in community
289 composition, but each cropping mode with different CH_3F addition possessed similar community
290 composition. At the order level (Fig. S3A), the dominant methanogens in paddy soils under different
291 cropping modes mainly included *Methanosarcinales* (15–91%), *Methanobacteriales* (3–38%),
292 *unclassified_p_Euryarchaeota* (4–21%), *norank_p_Euryarchaeota* (1–22%), *Methanocallales* (1–4%),
293 *Methanomassiliicocales* (1–3%), as well as *Methanmicrobiales* (1–2%). At the family level (Fig.
294 S3B), five known methanogen communities dominated in paddy soils under different cropping modes:
295 *Methanosarcinaceae*, *Methanobacteriaceae*, *Methanocellaceae*, *Methanomassiliicoccaceae*, and
296 *Methanotrichaceae*. The composition of methanogens differed significantly at the genus level under
297 different cropping modes. For instance, *Methanosarcina* was the dominant methanogen in RW and RF
298 soils, while *Methanobacterium* was in DR soils (Fig. 5). Furthermore, the relative abundance of
299 *Methanosarcina* was negatively correlated with MPP and *mcrA* gene abundance, while positively
300 correlated with soil pH (Fig. S4). The relative abundance of *Methanobacterium* positively correlated
301 with MPP and *mcrA* gene abundance ($p < 0.05$), but negatively correlated with soil pH ($p < 0.05$).

302 The bio-plot of RDA analyses illustrated that the eight environmental factors explained 84.11% of
303 the cumulative variance for the first principal component, and more importantly, the soil pH had
304 significant effects on the variation of methanogenic community structure (Fig. 6). Network analysis
305 showed, compared with RW and DR soils, the number of lines of methanogens strongly increased in
306 RF soil (Fig. 7, Table S2), which meant that the interaction between methanogens in RF soil was
307 closer than the other two soils. In addition, *Methanosarcinales* had the highest relative abundance in all

308 paddy soils and were negatively associated with most methanogens, especially in RW and RF soils.

309 **4. Discussion**

310 **4.1 Effect on MPP**

311 MPP in double-season rice fields (RW and DR) was significantly higher than in single-season rice
312 (RF) fields (Fig. 2D). The possible reason was that the abundance of methanogens in RW and DR
313 soils was higher than that in RF soil (Fig. 4). Soil organic matter and soil organic acids are the
314 important precursors of methanogenesis for consumption by methanogens (Ding and Cai, 2002;
315 Glissmann and Conrad, 2002). They are mainly derived from carbon input, e.g., root stubble and the
316 content of root exudate that is affected by cropping modes (Jiang et al., 2022). In theory, both DR and
317 RW were planted with crops harvested twice a year (Zhang et al., 2017), resulting in more readily
318 available root stubble and root exudate than RF soil. As such, the large carbon sources provided an
319 abundant substrate for CH₄ production, resulting in higher MPP in DR and RW soils (Fig. 2D). More
320 importantly, DR was characterized by flooding in both early and late rice seasons, while RW was in a
321 dry-wet alternation of winter wheat and summer rice (Zhang et al., 2017). It was reported that flooded
322 fields in the previous rice season would increase CH₄ production in the second season (Zhang et al.,
323 2011), which supported the MPP in DR soil was higher than that in RW soil (Fig. 2D). Therefore, CH₄
324 emissions in RW soil (He et al., 2022; Zhang et al., 2017) was often observed to be much lower than
325 those in DR soil (Wang et al., 2018; Zhong et al., 2021).

326 In addition to the different straw incorporation and water management, such cropping modes
327 inevitably lead to significant differences in soil physicochemical properties, like as soil pH, soil
328 texture, and so on (Table 1), which could affect CH₄ production by regulating methanogens (Wang,
329 2015). We found that soil pH was negatively correlated with *mcrA* gene abundance and the relative
330 abundance of most methanogens (Figs. 1 and S4), suggesting that soil pH probably affects MPP by
331 regulating the abundance and community composition of methanogens (Dubey et al., 2013).
332 Generally, different methanogens have certain differences in soil pH preference, and the optimal pH
333 for *Methanosarcinales* and *Methanobacteriales* is from 4.16 to 4.38 and from 3.64 to 4.04,
334 respectively (Galand et al., 2003; Yavitt et al., 2006). Thus, in alkaline RF soil, *Methanosarcinales* and
335 *Methanobacteriales* were also the dominant methanogens (Fig. S3A), and the abundance of *mcrA* gene

336 was far lower than that in acid RW and DR soils (Fig. 4), which might be one of the reasons for the
337 lower MPP of alkaline RF soil.

338 Soil texture might be another influencing factor for the MPP. First, clay soil has a strong retention
339 effect on the organic matter resulting in a low supply of organic substrate for methanogens, reducing
340 CH₄ production and emission (Kim et al., 2018; Mitra et al., 2002). Indeed, the MPP was negatively
341 correlated with soil clay content, and the highest soil clay content in RF had the lowest MPP (Figs. 1
342 and 2). Second, Mitra et al. (2002) showed that soil CEC affected different redox indices by stepwise
343 multiple regression analyses, and then had a significant effect on MPP. Soil CEC also affects CH₄
344 oxidation by mediating the NH₄⁺ behavior. Low soil CEC leads to a weaker soil buffering capacity and
345 a lower NH₄⁺ concentration, which has a strong inhibitory effect on the microbial oxidation of CH₄
346 (De Visscher et al., 1998). Therefore, these reasons might explain the negative correlation between soil
347 CEC and MPP in this study (Fig. 1). Third, the higher OXC values usually maintain higher levels of
348 redox potential (Eh) in paddy soil (Zhang et al., 2009b), which does not conducive to CH₄ production
349 (Zhang et al., 2015). The OXC of RW and RF soils was significantly higher than that of DR soil (Table
350 1) further verifying the lower MPP of RW and RF soils than DR soil in this study (Fig. 2D).

351 **4.2 Effect on methanogenic pathway**

352 In consistency with our hypothesis, f_{ac} values were different under various cropping modes (Table
353 2), which might due to the various substrate content among different cropping modes. Indeed, in
354 paddy fields, acetate was the most important substrate for the acetoclastic methanogens pathway (Ji et
355 al., 2018b). In our study, both DOC and acetate contents in RW soil were lowest, and the availability
356 of substrates for acetoclastic methanogens resulted in the lowest f_{ac} value (Fig. 3, Table 2).
357 Furthermore, the content of acetate in RF soil was the highest among the three soils, and RF soils had
358 the highest SOC (Table 1) with large number carbon, which also could provide rich methanogenic
359 substrates for acetoclastic methanogens (Luo et al., 2022).

360 In this study, the acetoclastic methanogen *Methanosaeta* was the dominant methanogen in RF soil,
361 and the f_{ac} value of RF alkaline soil was highest relative to those of RW and DR acidic soils (Table 2).
362 Lee et al. (2014) and Yang et al. (2021) reported that *Methanosaeta* was the dominant methanogen in
363 soils with soil pH of 6.0 to 8.0. Additionally, Liu and Ding (2011) also found that hydrogenotrophic

364 methanogens were dominant under acidic conditions. Thus, the f_{ac} decreased in acidic DR soil
365 dominated by hydrogenotrophic methanogens, while increased in alkaline RF soil dominated by
366 acetoclastic methanogens (Table 2). Our results were consistent with previous studies and further
367 demonstrated that the f_{ac} was affected by soil pH by influencing the compositions of methanogens (Lee
368 et al., 2014; Liu and Ding, 2011).

369 Based on the results of previous studies in Thailand, the USA, Japan, Italy, and China, f_{ac} values of
370 paddy soils varied with different cropping modes, ranging from 10% to 108% (Table S1). Our results
371 showed that f_{ac} values of the three cropping modes were all within the range of these studies, but there
372 were some differences between the f_{ac} values of our study and previous studies under the
373 corresponding cropping modes (Table 2). For instance, compared with the corresponding DR mode
374 (Yingtian) in China, the f_{ac} values in Thailand paddy were higher (Table S1), which may be due to the
375 enhanced activity of methanogens and the promoted methanogenic pathway as affected by higher
376 temperature in Thailand (Fey and Conrad, 2000). In addition, soil texture could change the microbial
377 community structure and function (Zhang et al., 2007), which was probably related to the differences
378 in soil organic carbon and soil carbon-nitrogen ratio in different soil grain size components (Zhou et
379 al., 2002). Xiao et al. (2019) reported that *Methanosarcina* had high activity in clay soil, which might
380 be the reason that f_{ac} values were highest in clay RF soil (Table 2). Thus, differences in soil texture
381 under the same cropping mode might also contribute to differences in methanogenic pathways
382 (Allison and Prosser, 1991; Qian and Cai, 2010).

383 Generally, $\delta^{13}\text{CH}_{4(\text{ac})}$ and $\alpha_{(\text{CO}_2/\text{CH}_4)}$ values have a strong influence on f_{ac} values (Zhang et al., 2009a).
384 Numerous studies have shown that $\delta^{13}\text{CH}_{4(\text{ac})}$ values measured in paddy soils are usually in the range
385 of -43‰ to -37‰ (Conrad et al., 2002; Fey et al., 2004; Krüger et al., 2002; Nakagawa et al., 2002),
386 and the $\alpha_{(\text{CO}_2/\text{CH}_4)}$ value was usually assumed as a fixed value (Bilek et al., 1999; Conrad et al., 2002;
387 Nakagawa et al., 2002). However, the $\alpha_{(\text{CO}_2/\text{CH}_4)}$ value in paddy soils varied greatly with the differences
388 in soil properties and soil carbon isotope in different cropping modes (Zhang et al., 2012; Zhang et al.,
389 2009a). In this study, the $\alpha_{(\text{CO}_2/\text{CH}_4)}$ value in paddy soil under different cropping modes was calculated
390 using Eq. (7) by the CH_3F inhibitor method (Fig. S1), ranging from 1.049 to 1.062 (Table 2), which

391 was consistent with previous studies (Zhang et al., 2016). Therefore, our study provided a more detailed
392 reference for soil $\alpha_{(\text{CO}_2/\text{CH}_4)}$ value in three typical cropping modes (RW, RF, and DR) in China. The f_{ac}
393 values of RW, RF and DR were from 27% to 36%, from 78% to 96%, and from 44% to 57%,
394 respectively, when the commonly $\alpha_{(\text{CO}_2/\text{CH}_4)}$ (1.045) selected to quantify the methanogenic pathways
395 of different cropping modes (Fey et al., 2004). The results indicate that the relative contribution of
396 acetate to total methanogenesis was mainly dependent on cropping modes rather than on the values of
397 $\alpha_{(\text{CO}_2/\text{CH}_4)}$ cited.

398 4.3 Effect on methanogenic communities

399 A higher abundance of *mcrA* gene was observed in DR mode than that in RW and RF modes (Fig.
400 4). Previous studies have shown that DR mode had less soil carbon loss and higher SOC stock than RF
401 and RF modes (Cha-un et al., 2017; Sun et al., 2019). Therefore, the higher abundance of *mcrA* gene
402 in DR soil than in the other two soils, which possibly driven by nutrient availability due to the
403 different soil properties among various cropping modes (Jiang et al., 2022). In theory, DR mode, a
404 biannual rice cultivation mode, had higher residual carbon (stubble and root), organic acid remaining,
405 and more favorable anaerobic conditions for methanogens in the soil. It will inevitably lead to the
406 release of more liable carbon and acetate by residual decomposition (Jiang et al., 2022), thus
407 potentially supplying sufficient available substrates for methanogen and increasing the abundance of
408 *mcrA* gene. Furthermore, RF mode had less fertilizer application and less stubble compared with RW
409 and DR modes, which may lead to fewer types and quantities of methanogenic substrates, and then
410 affect the low diversity and abundance of methanogens (You et al., 2022), which is also confirmed by
411 the low Shannon index and Chao 1 estimator observed in our study (Table 3).

412 The methanogenic community compositions varied among the three cropping modes, the dominant
413 methanogen of RW was *Methanosarcina*, while the dominant methanogen of RF and DR soil were
414 *Methanosarcina* and *Methanobacterium* (Fig. 5). The detected dominantly methanogens in the soils
415 coincided with previous research in rice fields (Jiang et al., 2022; Wang et al., 2021). Among the
416 different cropping modes, unique geographical features determined that the soil pH of the RF soil was
417 significantly higher than that of RW and RF soils (Table 1). This indicated that different types of

418 methanogens might have different preferences for soil pH, with acetoclastic methanogens being more
419 active under weak acidic and alkaline conditions and hydrogenotrophic methanogens being more
420 active under acidic conditions (Kotsyurbenko et al., 2007). In this study, the relative abundance of
421 acetoclastic methanogen (*Methaosarcina*) in RF alkaline soil was higher than that in RW and DR
422 acidic soils, while the relative abundance of hydrogenotrophic methanogen (*Methanobacterium*) in
423 RW and DR soils was higher than that in RF soil (Fig. 5). The findings indicated that differences in
424 soil pH driven by cropping modes might be the underlying cause of methanogenic communities.

425 Different cropping modes have different water and fertilizer management and climatic conditions,
426 which change the composition of methanogens and also affected the interaction of methanogens (Gu et
427 al., 2022). Network analysis showed the RF soil had the most lines, indicating that the interaction of
428 methanogens in both RW and DR soils was weaker than that in RF soil (Fig. 7, Table S2). The reason
429 might be due to the anaerobic environment conducive to the retention of methanogens activities and
430 the growth of methanogens in flooded RF soil (Xu et al., 2020; Zhang et al., 2017). The methanogens
431 population could increase in flooded RF soil (Bhullar et al., 2013; Pavlostathis and Giraldo, 1991),
432 thus forming a more complex microbial network. In addition, since only one crop is harvested in RF
433 soil, N fertilizer application is much less than that in DR and RW soils in which two crops are
434 harvested. This must theoretically lead to less root exudate. In other words, RF soil has less
435 methanogenic substrate than both DR and RW soils. Interestingly, we found that RF soil had a higher
436 SOC content (Table 1), but had relatively lower available organic carbon dominated by the unique clay
437 characteristics (Ding and Cai, 2003). These results probably lead to an insufficient supply of
438 methanogenic substrates for RF soils, which also explained the lowest MPP of RF soils among the
439 three cropping modes (Fig. 2). Meanwhile, highly active *Methaosarcina* consumed large amounts of
440 methanogenic substrates, resulting in a complex antagonistic interaction with other methanogens (Fig.
441 7).

442 The effect of CH_3F as an inhibitor on CH_4 oxidation and production in rice fields was first reported
443 by Frenzel and Bosse (1996). Subsequently, Conrad and Klose (1999) systematically investigated the
444 effects of different concentrations on CH_4 production and proposed that it inhibits only the acetoclastic
445 methanogenesis in a certain concentration range, which could be well used for the study of the

446 methanogenesis pathway. However, how CH₃F inhibited acetoclastic methanogenesis by affecting
447 methanogens had not been observed so far. We observed that the addition of CH₃F resulted in a strong
448 decrease in the abundance of *mcrA* gene by one order of magnitude (Fig. 4) and in the relative
449 abundance of acetoclastic methanogen (Fig. 5). The finding confirmed that CH₃F addition inhibited
450 methanogenesis mainly by reducing the abundance of acetoclastic methanogen (Conrad and Klose,
451 1999; Ji et al., 2018a). However, 1% CH₃F addition significantly changed the relative abundances of
452 *Methanosarcinales* and *Methanobacterials* in RF and DR soils, but not in RW soils (Fig. S3).
453 Additionally, a 2% CH₃F addition significantly changed the relative abundances of *Methanosarcinales*
454 and *Methanobacterials* in RW soil. The response of soil methanogens composition to CH₃F varied
455 among cropping modes and the sensitivity of soil to CH₃F differed among cropping modes. These
456 phenomena demonstrated that the difference in the inhibitory effect of CH₃F addition probably
457 attributes to the difference in acetoclastic methanogens under different cropping modes.

458 **5. Conclusion**

459 Our findings suggest that the potential of CH₄ production, the pathway of methanogenesis, and the
460 composition of methanogenic community were different under the three typical rice cropping modes
461 in China. The carbon isotope fractionation during hydrogenotrophic methanogenesis was 1.049–1.062,
462 making the acetoclastic methanogenesis in paddy soils range from 39% to 98%. The CH₄ production
463 potential was largely controlled by soil pH and *mcrA* gene abundance while the relative contribution
464 of acetoclastic methanogenesis was probably affected by soil pH and methanogenic community.
465 Moreover, it was found that CH₄ production potential was positively correlated with the abundance of
466 *mcrA* gene and *Methanobacterium* (hydrogenotrophic methanogens), respectively. More importantly,
467 soil pH was negatively correlated with the abundance of *mcrA* gene while positively related to
468 *Methanosarcina* (acetoclastic methanogens). The results indicated that both CH₄ production potential
469 and acetoclastic methanogenesis would mainly be regulated by the abundance and community of
470 methanogens driven by soil pH under the different rice cropping modes.

471

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476 **Conflict of Interest**

477 The authors declare no conflict of interest relevant to this study.

478 **Authors' contributions**

479 Wanyu Shen wrote the main manuscript text, Yang Ji and Guangbin Zhang modified the manuscript,
480 Qiong Huang and Xiaoli Zhu assisted with data analysis, Jing Ma and Hua Xu supervised the data. All
481 authors read and approved the manuscript.

482 **Data Availability Statement**

483 Research data are not publicly shared, please contact corresponding author for enquiries.

484

485 **Reference**

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681

682 **Table 1** The physicochemical properties of paddy soils in the three typical rice cropping modes in

683 China. Data are presented as the mean \pm SD (n = 3).

684

685 **Table 2** Values of $\delta^{13}\text{CH}_4$, $\delta^{13}\text{CO}_2$, $\alpha_{(\text{CO}_2/\text{CH}_4)}$, $\delta^{13}\text{CH}_4(\text{CO}_2)$, and f_{ac} in three cropping mode paddy soils.

686 Data are presented as the mean \pm SD (n = 3). The 1% and 2% indicate the proportions of CH_3F
687 addition. Different uppercase and lowercase letters denote significant differences among different
688 cropping modes and different proportions of CH_3F added treatment at $p < 0.05$, respectively. ^{a, b, c}

689 These values were calculated under different proportions of CH_3F addition with Equations 7, 6, and 5,
690 respectively.

691

692 **Table 3** Alpha diversity of *mcrA* gene in paddy soils under different cropping modes. Data are
693 presented as the mean \pm SD (n = 3). Different uppercase and lowercase letters denote significant
694 differences among different cropping modes and different proportions of CH_3F added treatment at $p <$
695 0.05, respectively. The 1% and 2% indicate the proportions of CH_3F addition.

696

697 **Fig. 1** The heatmap of correlations among soil physicochemical properties, soil oxidation capacity
698 (OXC), CH_4 production potential (MPP), and the diversity and abundance of *mcrA* gene in paddy soils
699 under different cropping modes. Red circles represent the positive correlation, blue circles represent
700 the negative correlation, and the size of the solid circle represents the size of the correlation
701 coefficient. *, **, and *** represent significant differences between each two variables at $p < 0.05$, $p <$
702 0.01, and $p < 0.001$, respectively.

703

704 **Fig. 2** The cumulative concentration of CH_4 production (A, RW; B, RF; and C, DR) and MPP (D) in
705 paddy soils under different cropping modes. Data are presented as the mean \pm SD (n = 3). Different
706 uppercase and lowercase letters denote significant differences among different cropping modes and
707 different treatments at $p < 0.05$, respectively. The 1% and 2% indicate the proportions of CH_3F
708 addition.

709

710 **Fig. 3** The contents of DOC (A) and acetate (B) in paddy soils under different cropping modes. Data

711 are presented as the mean \pm SD (n = 3). Different uppercase and lowercase letters denote significant
712 differences among different cropping modes and different proportions of CH₃F added treatment at $p <$
713 0.05, respectively. The 1% and 2% indicate the proportions of CH₃F addition, BI represents the before
714 the induction experiment.

715

716 **Fig. 4** The abundance of *mcrA* gene in paddy soils under different cropping modes. Data are presented
717 as the mean \pm SD (n = 3). Different uppercase and lowercase letters denote significant differences
718 among different cropping modes and different proportions of CH₃F added treatment at $p <$ 0.05,
719 respectively. The 1% and 2% indicate the proportions of CH₃F addition.

720

721 **Fig. 5** Relative abundance of methanogens under different cropping modes at the genus level. Data are
722 presented as the mean \pm SD (n = 3). The 1% and 2% indicate the proportions of CH₃F addition.

723

724 **Fig. 6** Redundancy analysis (RDA) ordination plots showing the relationship between the community
725 structure of methanogenic and environmental factors in paddy soils under different cropping modes at
726 genus level. The 1% and 2% indicate the proportions of CH₃F addition.

727

728 **Fig. 7** Network analysis of the correlation of methanogens in paddy soils under different cropping
729 modes (A, RW; B, RF; and C, DR). Each circle indicates a different individual. The node size
730 represents the node degree (a larger size indicates a higher degree). The color of the lines represents
731 positive (red) or negative (dark cyan) correlation. Only significant correlations ($p <$ 0.05) are shown.

732

Figure 1.

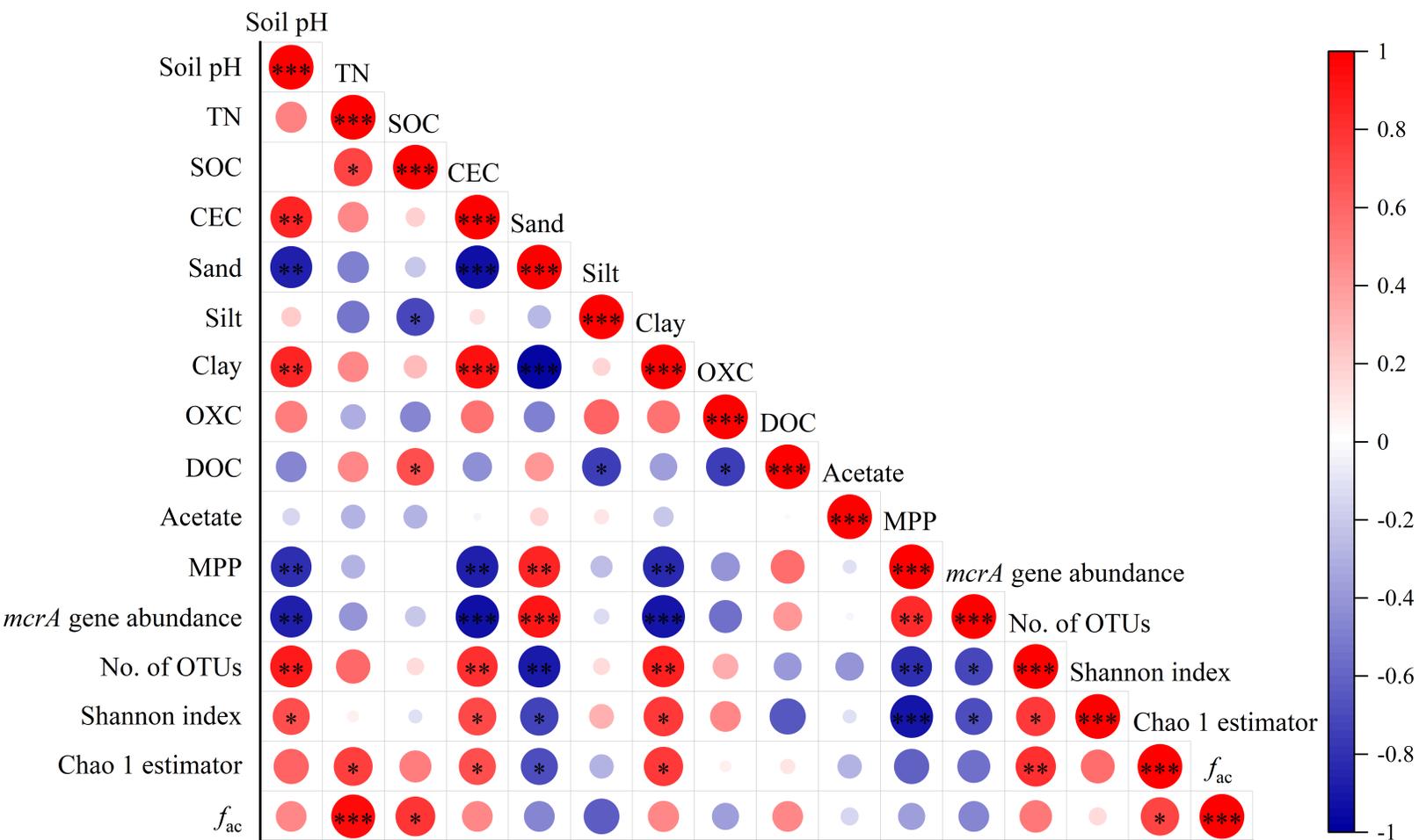


Figure 2.

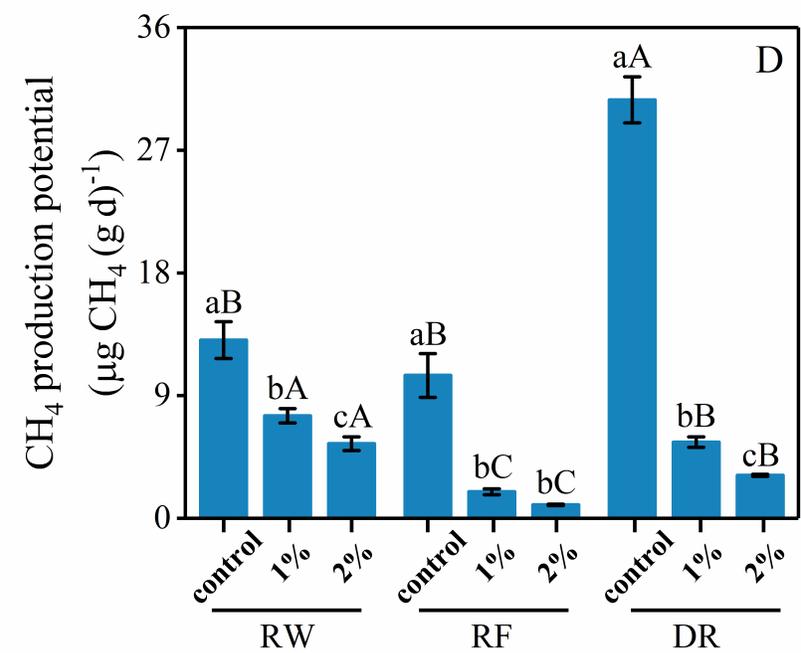
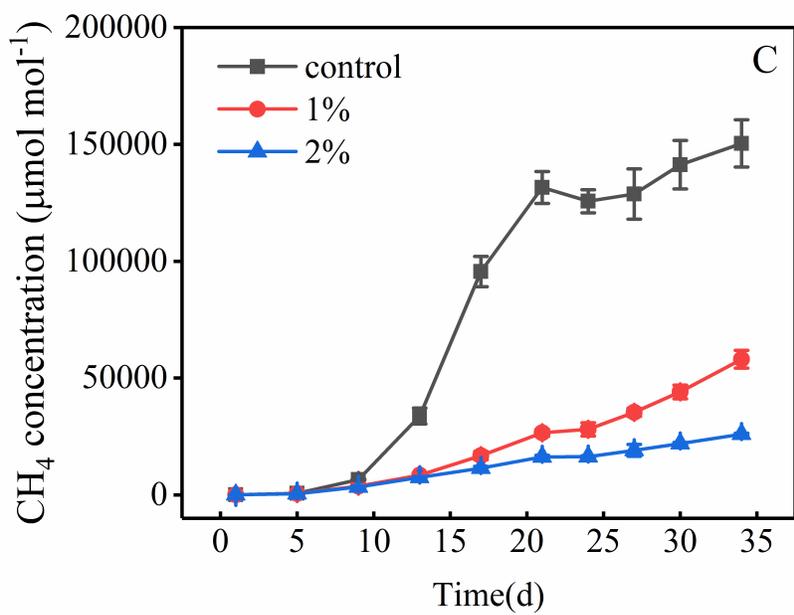
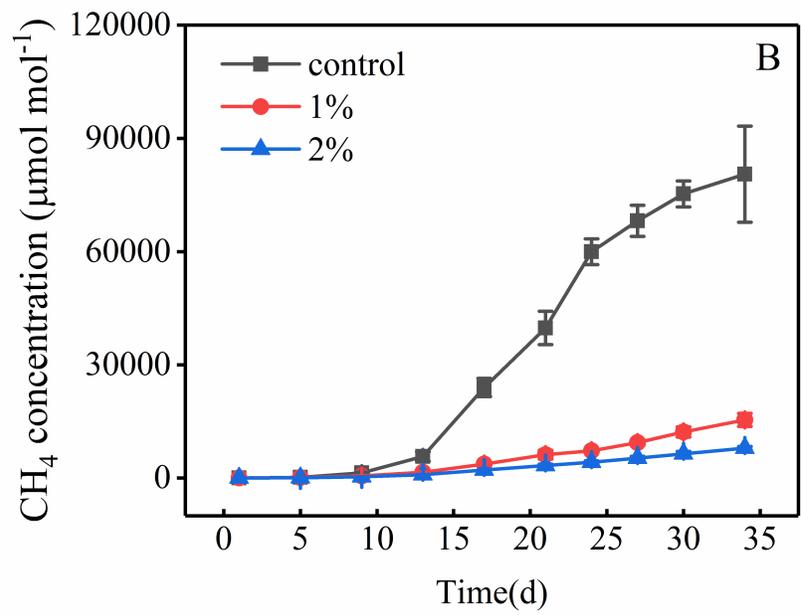
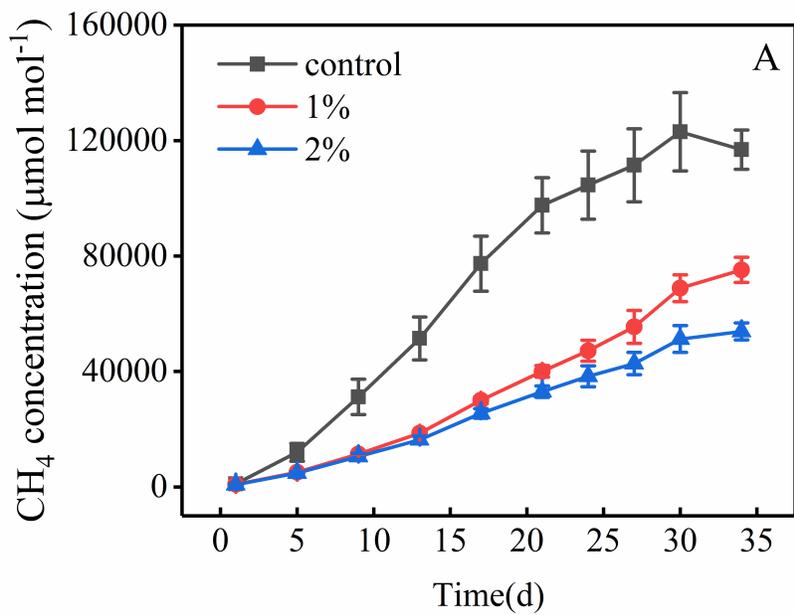


Figure 3.

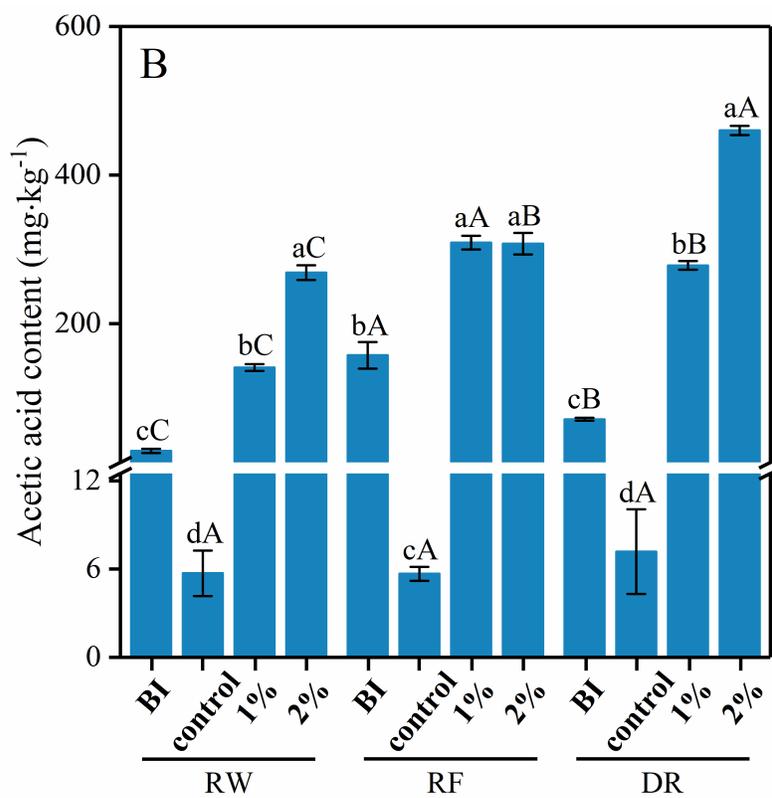
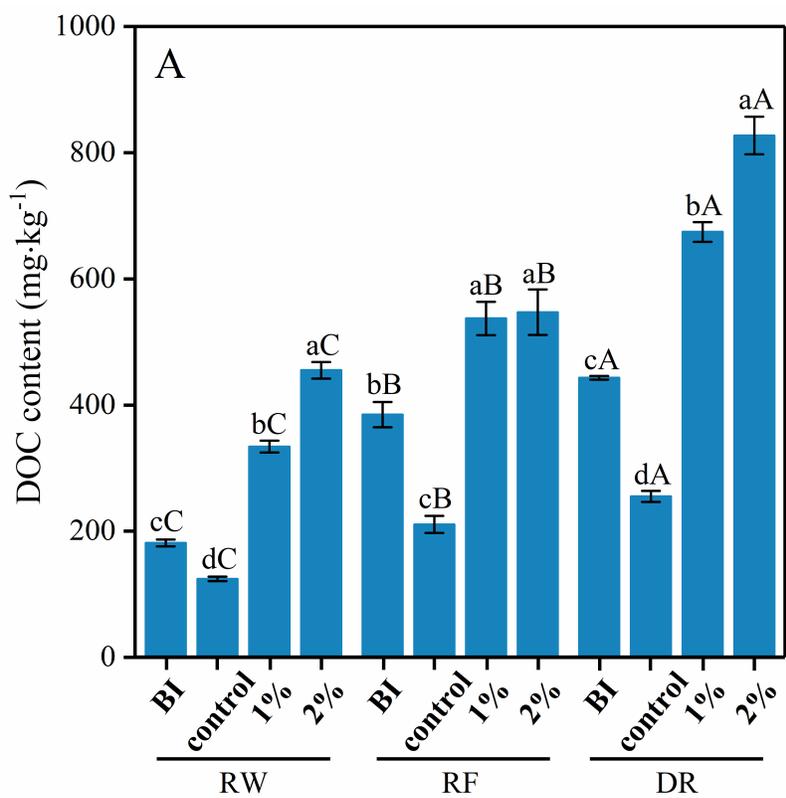


Figure 4.

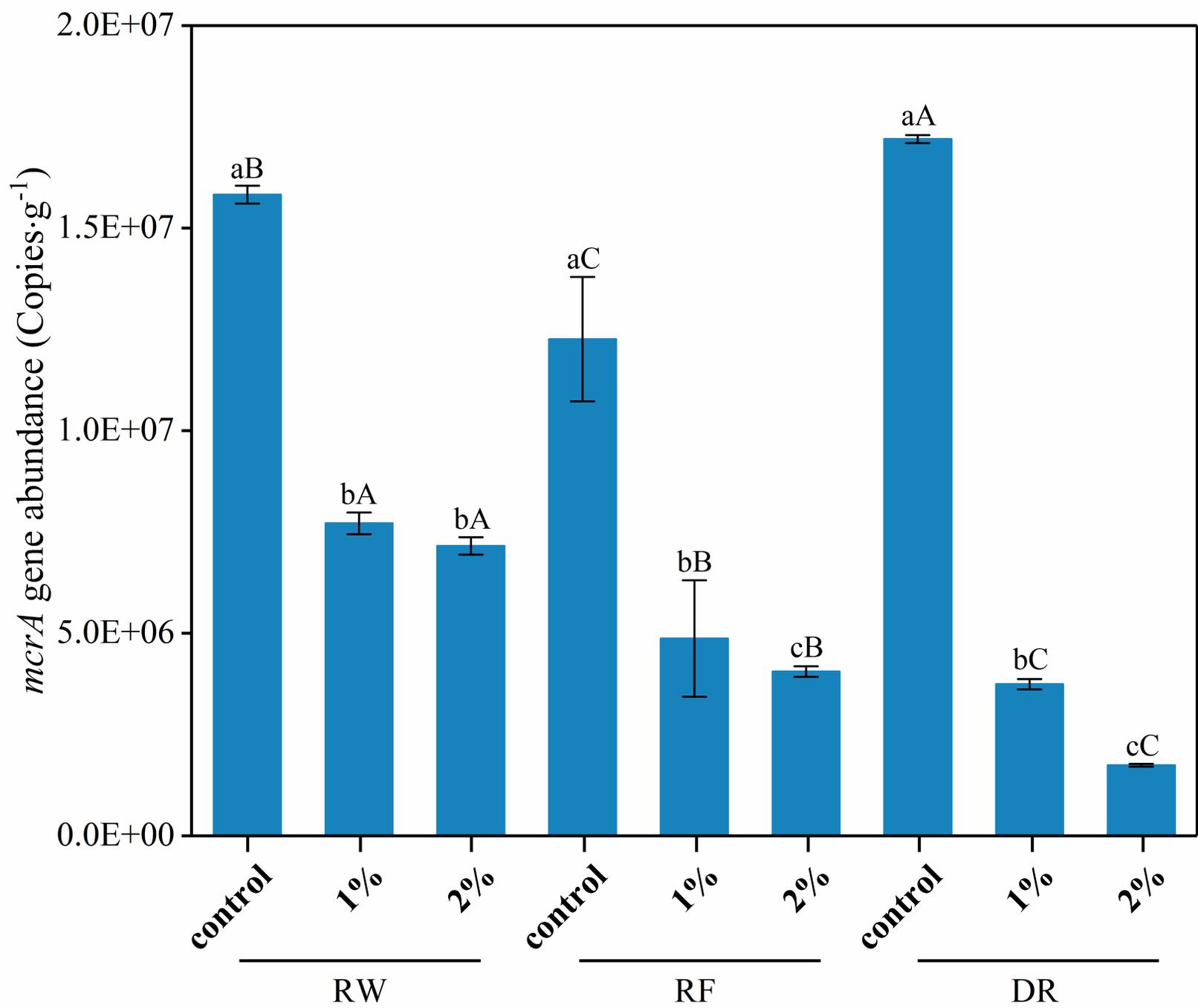


Figure 5.

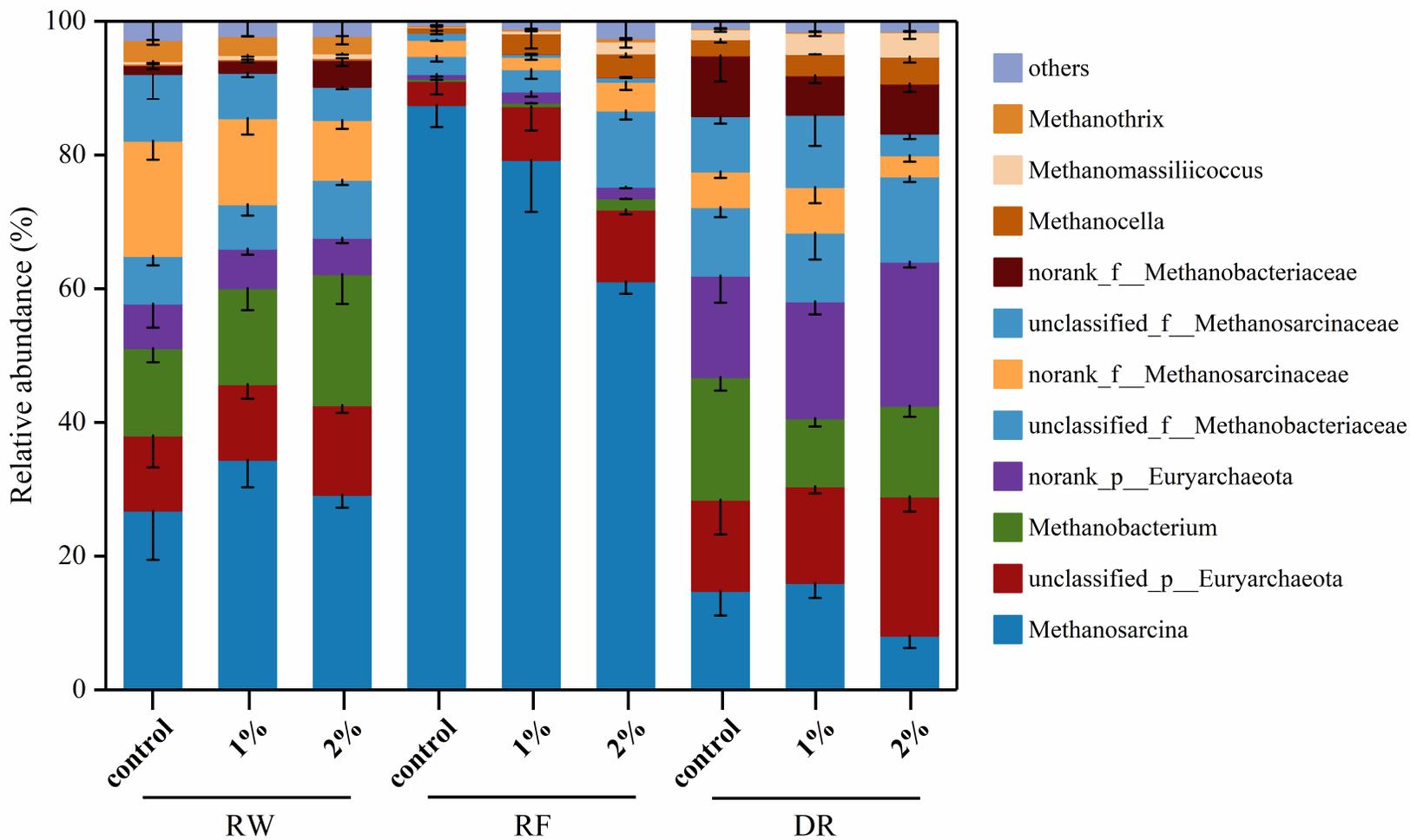


Figure 6.

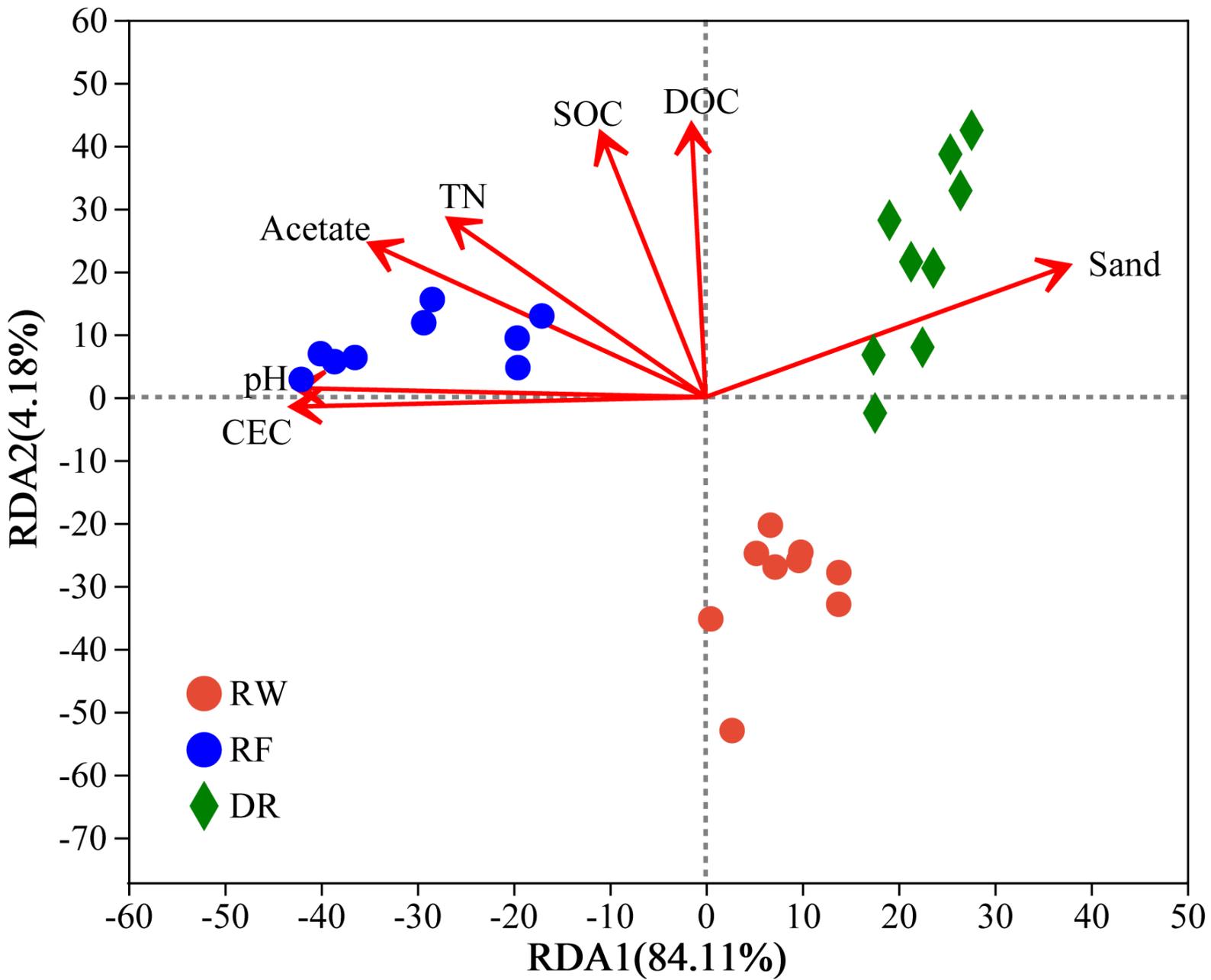


Figure 7.

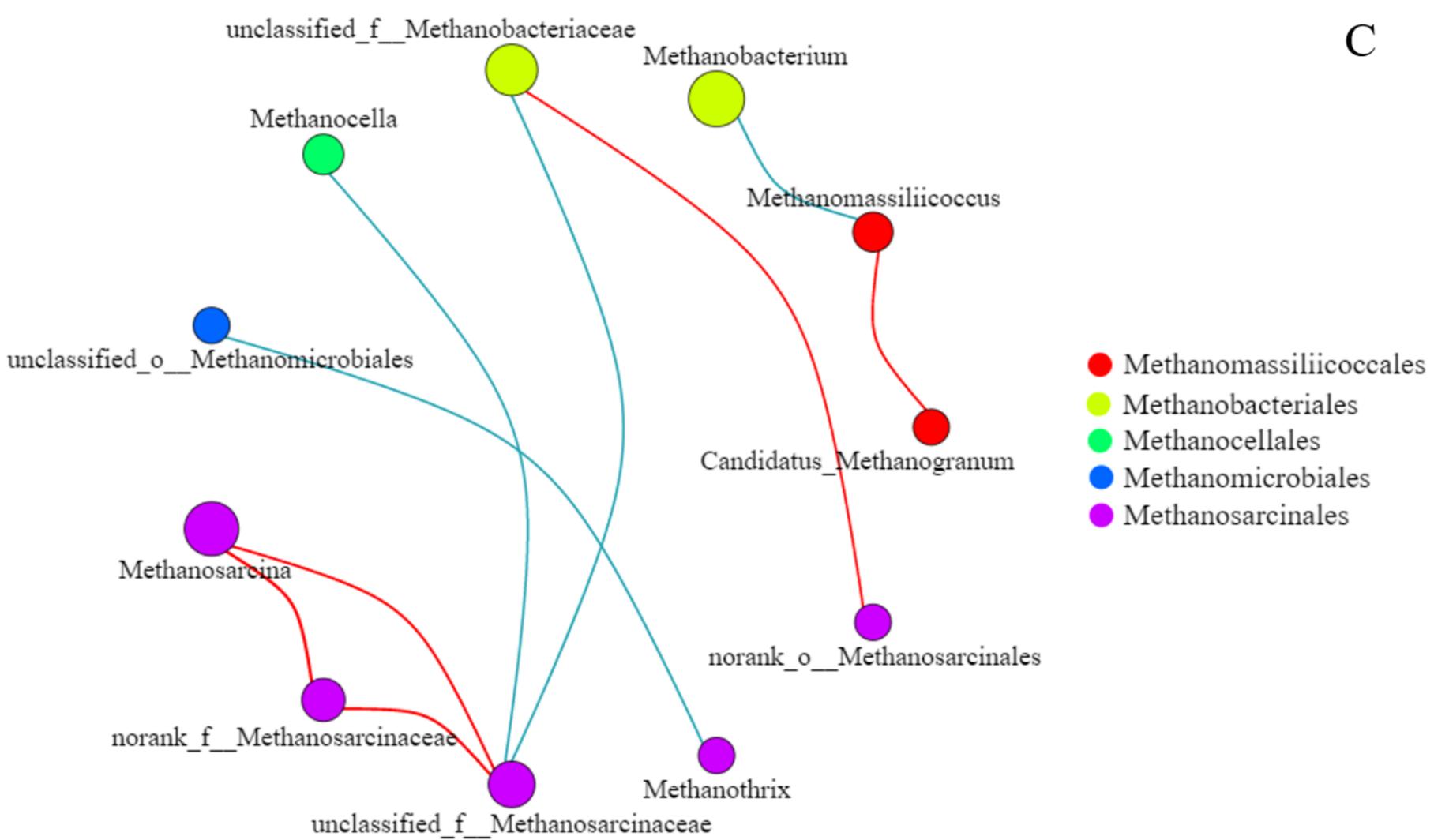
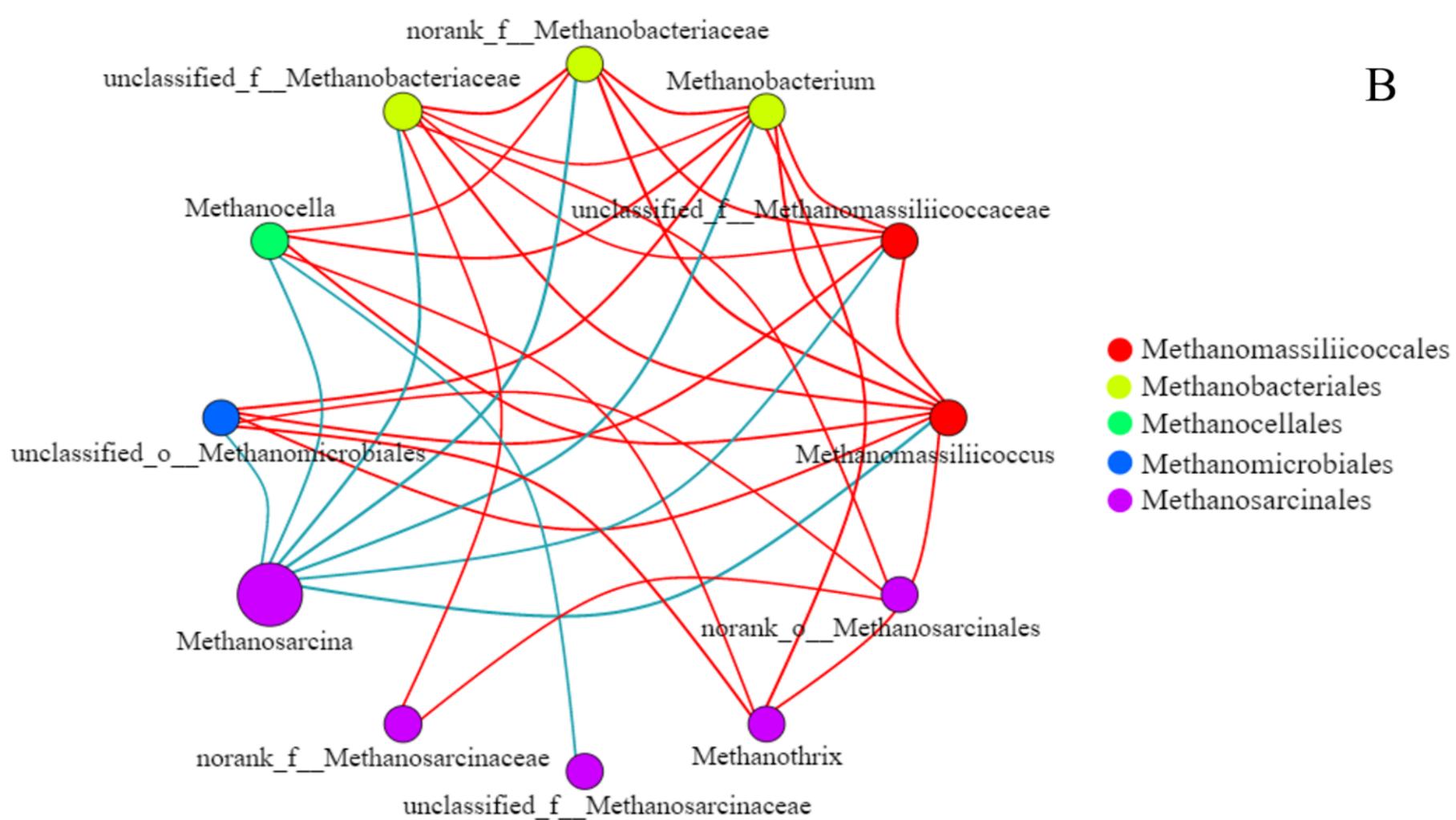
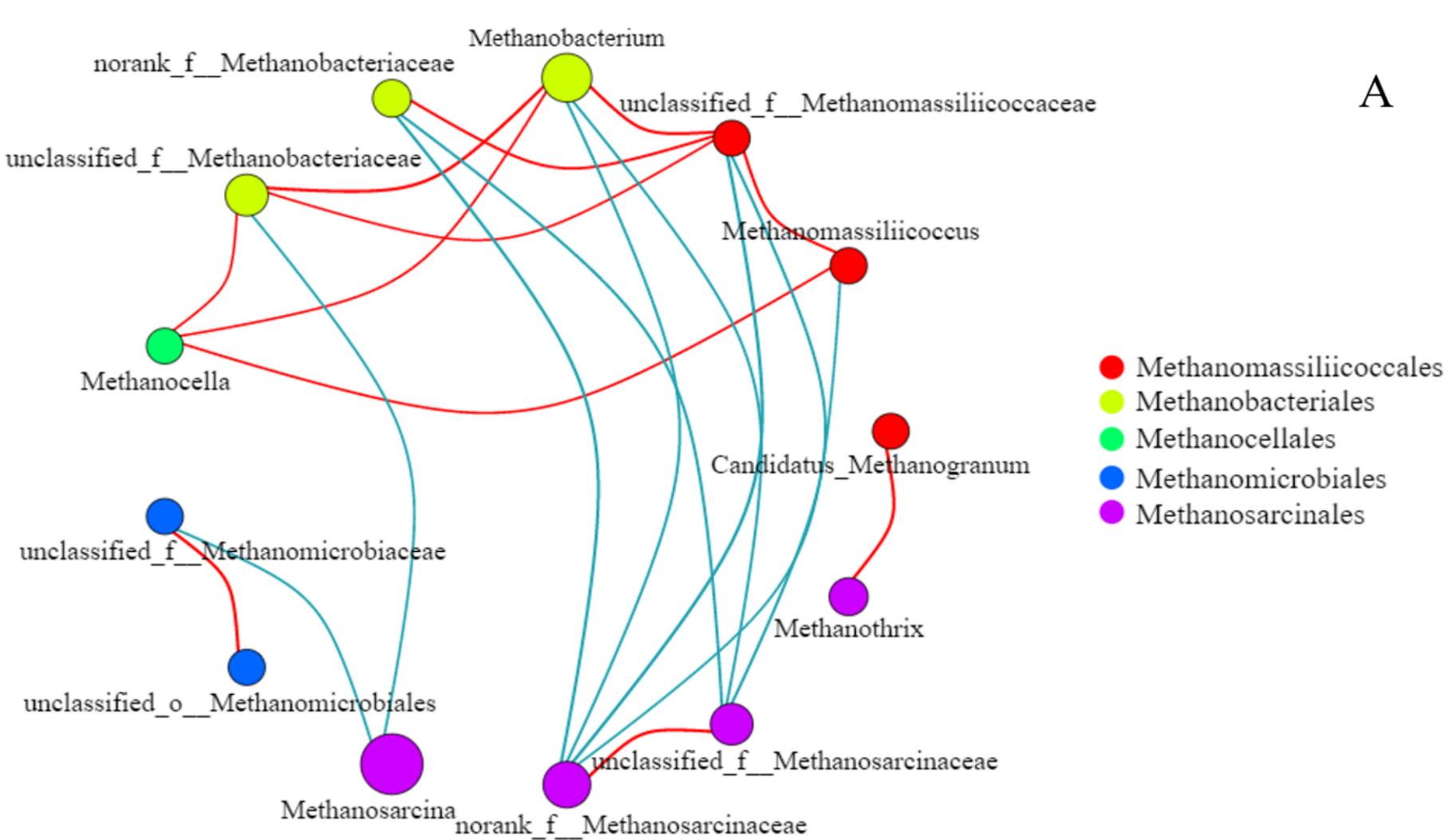


Table 1 The physicochemical properties of paddy soils in the three typical rice cropping modes in China. Data are presented as the mean \pm SD (n = 3)

Soils	SOC (g kg ⁻¹)	TN (g kg ⁻¹)	pH	$\delta^{13}\text{C}$ (‰)	Percentage of a particle of different size (%)			Fe (III) (mg kg ⁻¹)	Mn (IV) (mg kg ⁻¹)	NO ₃ ⁻ (mg kg ⁻¹)	SO ₄ ²⁻ (mg kg ⁻¹)	OXC (mol kg ⁻¹)	CEC (cmol kg ⁻¹)
					Sand	Silt	Clay						
RW	12.8 \pm 0.1	1.25 \pm 0.1	5.59 \pm 0.1	-26.2 \pm 0.2	36.9 \pm 0.4	39.8 \pm 0.3	23.3 \pm 0.6	4560 \pm 380	180 \pm 20	0.03 \pm 0.01	79.6 \pm 5.3	0.095 \pm 0.007	12.2 \pm 0.1
RF	19.8 \pm 0.3	2.31 \pm 0.4	7.70 \pm 0.3	-21.2 \pm 0.1	31.8 \pm 0.6	31.4 \pm 0.1	36.8 \pm 0.6	2940 \pm 60	250 \pm 20	1.01 \pm 0.02	292 \pm 9.2	0.086 \pm 0.002	22.7 \pm 0.4
DR	19.7 \pm 0.1	1.80 \pm 0.1	4.89 \pm 0.1	-28.8 \pm 0.2	50.5 \pm 1.1	31.3 \pm 1.0	18.2 \pm 0.6	3350 \pm 280	40 \pm 10	0.33 \pm 0.07	32.2 \pm 5.3	0.064 \pm 0.005	7.48 \pm 0.2

Table 2 Values of $\delta^{13}\text{CH}_4$, $\delta^{13}\text{CO}_2$, $\alpha_{(\text{CO}_2/\text{CH}_4)}$, $\delta^{13}\text{CH}_4(\text{CO}_2)$, and f_{ac} in three cropping mode paddy soils

Cropping mode	$\delta^{13}\text{CH}_4$ (‰)			$\delta^{13}\text{CO}_2$ (‰)			$\alpha_{(\text{CO}_2/\text{CH}_4)}$ ^a		$\delta^{13}\text{CH}_4(\text{CO}_2)$ ^b		f_{ac} (%) ^c		
	control	1%	2%	control	1%	2%	1%	2%	1%	2%	$\delta^{13}\text{CH}_4(\text{ac})$	1%	2%
RW	-54.4±0.2aC	-68.0±0.7bA	-74.5±1.2cA	-18.7±0.5aA	-21.1±0.2bA	-21.7±0.2bA	1.050±0.001bA	1.057±0.001aB	-65.7±0.4aA	-71.7±0.4bA	-37%	39±1.2eB	50±0.9eA
											-43%	50±1.4dB	60±1.0dA
RF	-43.9±1.2aA	-73.7±1.2bC	-81.3±0.3cB	-26.3±0.6aC	-28.2±0.2bC	-28.0±0.3bC	1.049±0.002bA	1.058±0.001aB	-71.8±0.6aC	-79.8±0.6bC	-37%	80±3.8bA	84±3.1bA
											-43%	97±4.3aA	98±3.4aA
DR	-51.7±0.5aB	-71.2±1.2bB	-82.3±0.4cB	-21.1±0.6aB	-25.1±0.1bB	-25.5±0.2bB	1.050±0.001bA	1.062±0.001aA	-67.5±0.6aB	-78.2±0.6bB	-37%	52±2.6dB	64±1.8dA
											-43%	64±3.0cB	75±1.9cA

Data are presented as the mean ± SD (n = 3). The 1% and 2% indicate the proportions of CH₃F addition. Different uppercase and lowercase letters denote significant differences among different cropping modes and different proportions of CH₃F added treatment at $p < 0.05$, respectively. ^{a, b, c} These values were calculated under different proportions of CH₃F addition with Equations 7, 6, and 5, respectively

Table 3 Alpha diversity of *mcrA* gene in paddy soils under different cropping modes

Cropping modes	Treatment	Effective sequence	No. of OTUs	Shannon index	Chao 1 estimator
RW	control	8145	275±26.2aA	4.01±0.21aA	350±16.7aA
	1%	8145	260±15.1aA	3.89±0.07aA	319±2.95bA
	2%	8145	266±5.69aA	3.82±0.20aAB	330±13.5abA
RF	control	8145	172±9.29bB	3.04±0.06bB	235±30.5aB
	1%	8145	179±1.00bB	2.99±0.28bB	219±15.1aB
	2%	8145	220±8.62aB	3.65±0.05aB	260±19.5aB
DR	control	8145	193±16.2aB	3.98±0.04bA	233±31.2aB
	1%	8145	185±6.66abB	4.10±0.05aA	200±14.0abB
	2%	8145	168±3.06bC	3.93±0.05bA	187±6.57bC

Data are presented as the mean ± SD (n = 3). Different uppercase and lowercase letters denote significant differences among different cropping modes and different proportions of CH₃F added treatment at $p < 0.05$, respectively. The 1% and 2% indicate the proportions of CH₃F addition