

1 **Observed Changes in Interannual Precipitation Variability in the United**
2 **States**

3 Ryan D. Harp^{1,2}, Daniel E. Horton^{1,2}

4
5 ¹Institute for Sustainability and Energy at Northwestern, Northwestern University, Evanston, IL

6 ²Department of Earth and Planetary Sciences, Northwestern University, Evanston, IL

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15 **Key Points**

- 16 • We find widespread robust changes in two measures of interannual precipitation
17 variability across the United States
- 18 • We detect robust increases (decreases) in annual precipitation and wet day frequency
19 across the central and eastern (western) United States
- 20 • We explore the interaction of changes in precipitation frequency and wet day
21 precipitation intensity on interannual variability

22 **Abstract**

23

24 Characterizing robust changes in precipitation patterns over time is critical for water resource
25 management and agricultural planning. Here, we explore observed trends in interannual
26 precipitation variability using a suite of metrics that describe changes in precipitation patterns
27 over time. We analyze daily *in-situ* Global Historical Climatology Network precipitation data
28 from 1970 to present over seventeen internally consistent sub-national United States domains
29 using the regional Mann-Kendall trend test. We find increasing trends in annual precipitation
30 and wet day frequency for most of the central and eastern U.S., but decreasing trends in the
31 western U.S. Importantly, we also identify widespread significant trends in interannual
32 precipitation variability with increasing variability in the southeast, decreasing variability in the
33 far west, and mixed signals in the Rocky Mountains and north-central U.S. Our results provide
34 important context for water resource management and a new observational standard for climate
35 model performance assessments.

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38 **Plain Language Summary**

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40 While many studies have examined how annual precipitation and precipitation frequency have
41 changed, few examine the variability, or consistency, of year-over-year precipitation. We test for
42 these trends in daily observations from the Global Historical Climatology Network within
43 seventeen regions within the U.S. We find changes in yearly precipitation variability for most

44 regions, though results in the central U.S. are mixed. We also identify rising annual
45 precipitation and precipitation frequency for the central and eastern U.S. and falling annual
46 precipitation and frequency for the western U.S. Our results are important for agriculture and
47 water resource management and can be compared against climate models to determine how
48 well they reproduce our findings.

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52 **Keywords**

53 precipitation, interannual variability, precipitation variability, GHCN, NEON, NCA

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58 **Introduction**

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60 Precipitation patterns are shifting globally due to climate change (Douville et al., 2021). These
61 changes are broadly driven by increased moisture availability due to rising temperatures (i.e.,
62 the Clausius-Clapeyron relationship) and shifts in atmospheric circulation patterns (e.g.,
63 poleward expansion of the Hadley cell; Polade et al., 2014), and are constrained by Earth's
64 energy budget (Pendergrass and Hartmann, 2014). Observationally-based historical studies and
65 model-based future projections of precipitation commonly characterize changes across metrics
66 like annual mean, wet day frequency, and measures of extremes. However, constraining the
67 temporal variability of precipitation changes, using metrics such as interannual variability, is
68 important to inform a number of societally-impactful realms. Interannual variability of
69 precipitation describes the degree of consistency in year-over-year precipitation amounts:
70 higher variability equates to greater irregularity about the annual mean, which brings increased
71 challenges to fields like water resource management. Greater precipitation variability has been
72 shown to reduce crop yields, including for staples like corn and rice (Shortridge, 2019; Rowhani
73 et al., 2011), as well as decrease a grazing area's ability to support livestock (Sloat et al., 2018).
74 Hydrologically, shifts in interannual precipitation variability may also be driving increased
75 variability in Great Lake water levels (Gronewold et al., 2021) and water quality via increased
76 agricultural runoff (Loecke et al., 2017). Despite the importance of interannual variability,
77 summary assessments like the U.S. National Climate Assessment have not yet included

78 characterizations of its changes, instead focusing on mean and extreme precipitation (Easterling
79 et al., 2017). Here, to better constrain historical changes in the year over year distribution of
80 precipitation across the U.S., we examine shifts in observed interannual precipitation
81 variability, as well as annual precipitation amounts and precipitation frequency – two metrics
82 useful for understanding and explaining observed changes in interannual precipitation
83 variability.

84

85 *How is interannual precipitation variability projected to change?*

86 Global climate models project that interannual precipitation variability will increase with rising
87 greenhouse gas concentrations (Boer, 2009; Polade et al., 2014; Berg and Hall, 2015). Increases in
88 the interannual variability of precipitation of 3 to 5%/K are projected globally, with 4 to 5%/K
89 projected over land (Pendergrass et al., 2017; Wood et al., 2021; Chou and Lan, 2012), though
90 some projections estimate smaller increases (He and Li, 2018). He and Li (2018) explain that the
91 drivers of changes in interannual precipitation variability vary spatially; the increase of mean
92 state specific humidity leads to an increase in variability over areas of climatological ascent.
93 Conversely, variability increases in areas of climatological descent are driven primarily by
94 changes in mean state precipitation. Good et al. (2016) further tie interannual precipitation
95 variability to wet season length, individual rainfall event intensity, and variability in interstorm
96 wait times.

97

98 A number of studies, while not U.S.-focused, have used global climate models to project
99 changes in interannual precipitation variability over the U.S. Wood et al. (2021) used initial

100 perturbation large ensemble projections and found increasing interannual variability over the
101 U.S. by the end of the century, particularly in the winter months, under the RCP8.5 emissions
102 scenario. Polade (2014) revealed a similar widespread, though slight, increase in interannual
103 variability across the U.S. (up to 4%), with a hotspot of increased variability over the southwest.
104 This finding is replicated by Berg and Hall (2015) for California using a suite of RCP8.5-driven
105 CMIP5 models. Similarly, Swain et al. (2018) projected increases in annual precipitation
106 “whiplash events” (i.e., sub-20th percentile annual precipitation followed by super-80th
107 percentile in the next year) for California. Chou and Lan (2012) find an expanding range of
108 projected annual precipitation under the A1B emissions scenario over the U.S. midwest,
109 northeast, and northwest, driven by rising maximum annual totals. These findings are mirrored
110 by Pendergrass et al. (2017), who noted a similar spatial pattern using projections from the
111 RCP8.5 emissions scenario.

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113 Despite numerous model projections of interannual precipitation variability change, there
114 remains a dearth of observation-based analyses on the topic. Recently, Zhang et al. (2021)
115 identified increases in the coefficient of variation of precipitation for regions of the
116 southwestern and central United States using *in situ* observations from 1976-2019, however we
117 are aware of no other U.S.-focused analyses. To address this deficiency of observation-based
118 analyses and produce an observational standard for model studies, we explore changes in
119 interannual variability and relevant precipitation metrics throughout the U.S. using a full
120 complement of *in situ* measurements.

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Methods

To characterize interannual precipitation variability in the U.S. we use daily *in-situ* station data from the Global Historical Climatology Network Daily (GHCN-D). The National Centers for Environmental Information curate the GHCN-D database, which consists of observations from over 80,000 stations from 180 countries and territories, including the most complete collection of U.S. daily data available (Menne et al., 2012). GHCN-D observations have a sensitivity of 0.1 mm and are subjected to a sequence of quality control tests to identify climatological outliers, duplicate data, and other inconsistencies (Durre et al., 2010). We subject available U.S. station records to additional constraints to determine station observations with sufficient length and completeness for trend analysis. Specifically, we require station records to consist of 90% or more complete station-years to qualify, where a station-year must contain 90% or more of all possible daily records within the year to be considered complete. These screenings filtered our set of available U.S. stations from 63,571 to 2,542 (using a 1970 start year); domain summary statistics of station availability are shown in Table S7. To overcome some of the limitations of individual station statistics, such as internal variability (e.g., Fischer et al., 2013), we center our analysis on regional trends and utilize the set of established domains determined by the National Ecological Observatory Network (NEON). These twenty domains were created to possess internally homogeneous climates but remain distinct across-domains, as determined by a multi-variable analysis using nine climate variables (Keller et al., 2008; Schimel et al., 2011). As labeled in Figure 1a, we use the seventeen domains that are predominantly within the

144 contiguous United States. We also perform our analysis for U.S. National Climate Assessment
145 regions (Easterling et al., 2017) with results included within the Supporting Information.

146 We employ the regional Mann-Kendall trend tests to identify trends in precipitation
147 data at the NEON-domain level. The Mann-Kendall trend test is nonparametric and determines
148 if a trend exists in the data regardless of underlying distribution or linearity (Mann, 1945;
149 Kendall, 1975). As the Mann-Kendall trend test relies on the rank of values in place of actual
150 values, it is less subject to outliers and is suitable for detecting robust trends in hydrological
151 time series (Hamed, 2008). It is commonly used in studies assessing trends of precipitation over
152 time (e.g., Zhang et al., 2021; Roque-Malo and Kumar, 2017). The regional Mann-Kendall trend
153 test determines if a significant trend emerges across the collection of time series within a region
154 (Helsel and Frans, 2006). We use the Theil-Sen slope estimator, which is similar in its
155 underlying design to the Mann-Kendall trend test, to determine the slope of identified trends
156 (Sen, 1968; Theil, 1950).

157 We focus our analysis on four precipitation metrics: changes in interannual precipitation
158 variability, interannual coefficient of variation (a.k.a. relative interannual variability), annual
159 precipitation, and annual wet day frequency, where a wet day is defined as a station-day
160 observing 1 mm or more of precipitation (a threshold common in precipitation analyses; e.g.,
161 Vaittinada Ayar and Mailhot, 2021; Ye, 2018; Zolina et al., 2013; Giorgi et al., 2019). Collectively,
162 these four variables either directly characterize interannual variability, or provide crucial
163 information to explain shifts in interannual variability. The relationships between these
164 variables are described in the discussion section and illustrated in the Supporting Information.

165 Here, we define interannual variability as the standard deviation in annual precipitation
166 over a moving 11-year window. We use an 11-year window to limit the influence of known
167 dominant modes of interannual climate variability (e.g., ENSO), though a sensitivity analysis
168 reveals generally stable results for five to fifteen-year moving windows (Table S1-S2). We
169 similarly determine the coefficient of variation by dividing the standard deviation by the mean
170 annual precipitation over the concurrent 11-year moving window. The coefficient of variation is
171 often used as a measure of precipitation variability as it removes the strong dependence of
172 precipitation variability on the mean itself (e.g., Giorgi et al., 2019). For ease of understanding,
173 we will refer to the coefficient of variation as the relative interannual variability for the
174 remainder of this article.

175 In addition to performing a sensitivity analysis on the width of the moving window, we
176 analyzed the stability of precipitation trends across time periods by choosing different starting
177 dates in 10 year increments, such that calculations are performed every ten years from 1920
178 through 1980. We present findings using a 1970 starting date as it provides a balance of
179 widespread station availability and length of observation record, but highlight discrepancies we
180 identify within the sensitivity analysis in the discussion section. The full results of the
181 sensitivity analysis presented in the Supporting Information (Tables S1-S8).

182

183

184 **Results**

185

186 To properly inform changes in interannual variability, we must also assess changes in annual
187 precipitation and precipitation frequency over our domain. We find statistically significant ($p <$
188 0.05) increases in annual mean precipitation for all domains east of the Rocky Mountains. These
189 increases in annual precipitation range from 5.2-23.7 mm/decade (0.4-2.5%/decade; excluding
190 the Atlantic Neotropical domain), with larger increases for a subset of central and eastern
191 domains (Northeast, Great Lakes, Prairie Peninsula, Appalachians and Cumberland Plateau)
192 ranging from 18.4-23.7 mm/decade (1.6-2.5 %/decade) (Figures 1a and s, Tables S9-S10). We
193 identify statistically significant negative trends in annual precipitation over the western U.S.
194 between -9.6 to -2.7 mm/decade (-2.4 to -0.6 %/decade; excluding the non-significant Northern
195 Rockies domain). Spatial patterns in changes in annual wet day frequency largely mirror
196 changes in annual mean precipitation, with some additional non-significant domains (Figure
197 1b). We observe statistically significant increases in wet day frequency for northern domains
198 east of the Rocky Mountains, and statistically significant decreases for most western domains,
199 as well as the Southern Plains and Southeast domains. Changes in wet day frequency range
200 from -1.0 to 0.9 wet days/decade (-3.3 to 1.0%/decade) with the greatest increases generally
201 located in the most northern and southern domains (Figures 1b and 2, Tables S9-S10).

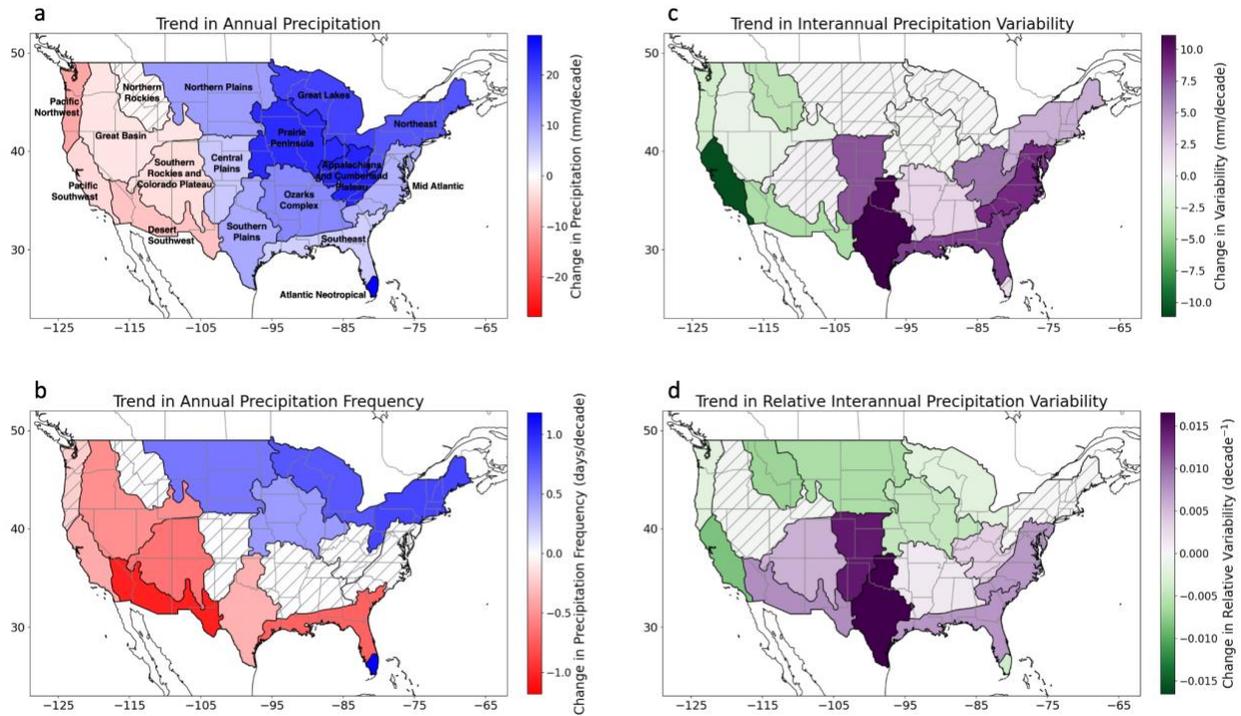
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203 Given robust changes in observed annual precipitation, it is important to determine if these
204 changes have been equitably distributed. To assess precipitation distribution, we quantify two
205 metrics of year-over-year variability. We identify statistically significant trends in both the
206 interannual variability and relative interannual variability of precipitation for most NEON
207 domains in the United States (Figures 1c-d, 2, Tables S9-S10). Changes in interannual variability

208 range from -1.1 to 2.0 mm/decade (-4.4 to 9.5%/decade), with changes not reaching statistical
209 significance for five domains, predominantly in the north central U.S. Generally speaking,
210 interannual variability is decreasing in domain clusters in the western U.S. (five domains), and
211 increasing in the south central and northeastern U.S. (Figure 1c; seven domains). We observe
212 broadly similar spatial patterns in trends of relative interannual variability, though seven
213 domains switched from significant to non-significant trends or vice versa. Additionally, the
214 direction of change in the Desert Southwest domain switched from significantly negative to
215 significantly positive (Figures 1c-d, 3). We explore this discrepancy in the discussion section.
216 Collectively, trends in relative interannual variability range from -3.2 to 9.6%/decade with
217 statistically significant changes occurring in all but two domains (Great Basin and Northeast).
218 Results for U.S. NCA regions reveal similar spatial patterns and can be found in the Supporting
219 Information (Figures S1-S2, Tables S11-S12).

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223 *Figure 1: Domain Trends in Various Precipitation Metrics. (a) Map of changes in annual precipitation*

224 *for each NEON domain within the contiguous U.S. Red-blue fill indicates domain-level trends in annual*

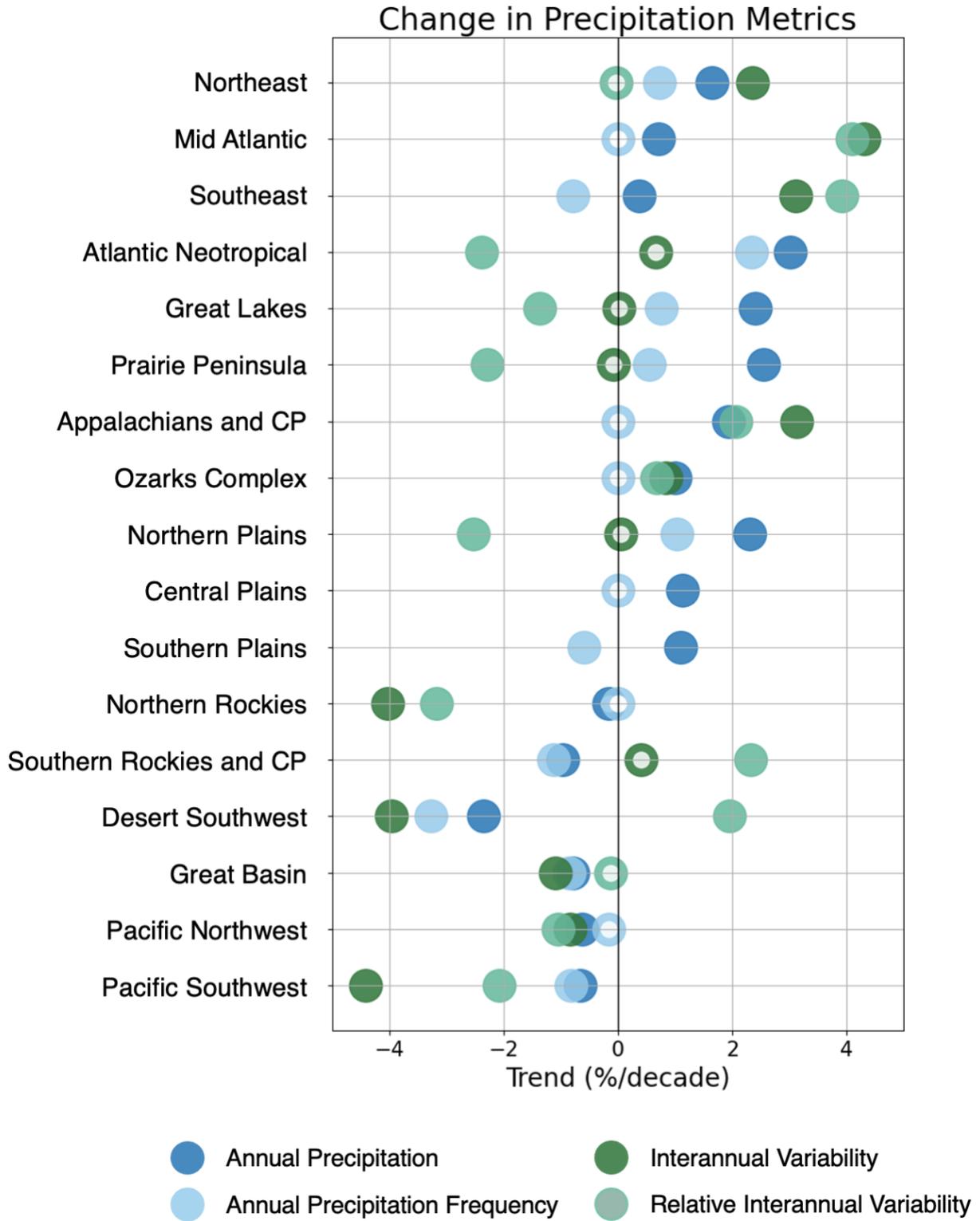
225 *precipitation in mm/decade (dark grey borders). Hatching indicates domain trends of zero or those not*

226 *reaching statistical significance. (b) Same as (a) but for annual precipitation frequency and units of*

227 *days/decade. (c) Same as (a) but for interannual precipitation variability with purple-green fill and units*

228 *of mm/decade. (d) Same as (c) but for relative interannual precipitation variability and units of decade⁻¹.*

229



230

231 *Figure 2: Domain Trends in Annual Precipitation Metrics. Trends in annual precipitation (dark blue),*

232 *annual precipitation frequency (light blue), interannual precipitation variability (dark green), and*

233 *relative interannual precipitation variability (light green) for each domain. Trends are normalized against*
234 *the mean value within each domain to produce trends in percent change/decade. Non-filled circles*
235 *indicate non-significant domain-trends ($p < 0.05$). Note outlying trends in both metrics of interannual*
236 *variability for the Central and Southern Plains are not displayed.*

237

238

239 **Discussion**

240

241 Broadly, our analysis of precipitation trends in the United States reveals increasing interannual
242 variability for the south-central and eastern U.S., decreasing interannual variability for the
243 Pacific coast, and mixed trends in the north-central and Rocky Mountain portions of the U.S.,
244 depending on the variability metric of interest. These changes are side-by-side with rising
245 annual precipitation and wet day frequency over the central and eastern United States, with
246 falling trends in the western United States.

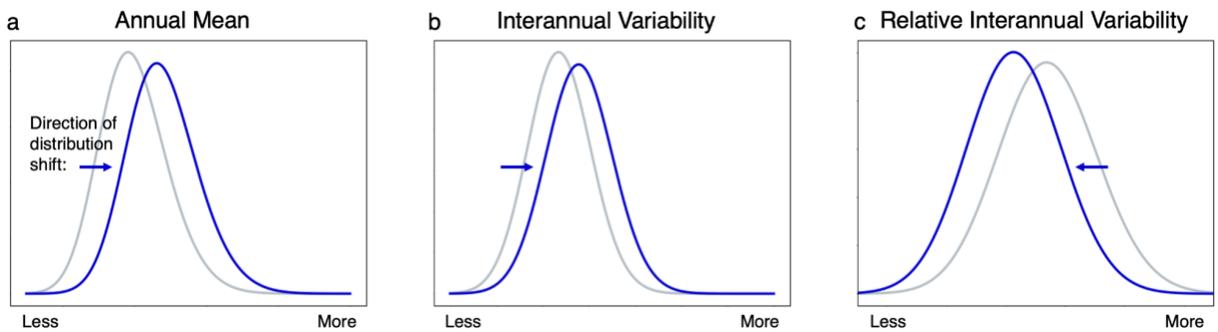
247 One result of particular interest is the finding that interannual variability *increased* but
248 relative interannual variability *decreased* at a statistically significant level in the Desert
249 Southwest. In addition, regardless of directionality, the trends in variability differed across
250 metrics by one percent or more for eight domains. We explain this between-metric discrepancy
251 through an examination of the components which influence interannual variability.

252

253 *Interannual variability vs relative interannual variability*

254 Together, changes in frequency and daily precipitation intensity drive changes in interannual
255 and relative interannual precipitation variability. We demonstrate the underlying principles of
256 these interactions using theoretical examples in Figures 3 and S3-S5. These examples apply 10%
257 increases in precipitation frequency and intensity along with three different possible
258 transformations of the underlying precipitation distribution – (1) a uniform increase at all
259 intensities (Figures 3 and S3), (2) increases in the higher intensities (Figure S4), and (3) increases
260 in the lower and medium intensities (Figure S5) – to demonstrate the interplay between annual
261 variability metrics and wet day frequency and intensity.

262



263

264 *Figure 3: Response of Annual Mean Precipitation and Interannual Variability of Precipitation to*
265 *Changes in Wet Day Frequency. (a) Initial probability distribution function (light grey) of annual*
266 *precipitation based on Great Lakes domain precipitation intensity distribution. Projected probability*
267 *distribution function (blue) after incorporating 10% increase in wet day frequency. (b) Same as (a) but*
268 *for interannual variability of precipitation. (c) Same as (a) but for relative interannual variability of*
269 *precipitation. Figure is replicated within the full combination of changes in Figures S3-S5.*

270

271 Holding intensity constant (1), an increase in wet day frequency leads to an increase in
272 interannual variability but a *decrease* in relative interannual variability (Figure 3 and Figures
273 S3b-c, S4b-c, S5b-c). As would be expected, an increase in wet day frequency produces an
274 increase in annual precipitation totals (Figures S3a, S4a, S5a). This rise in mean state leads to a
275 corresponding increase in interannual variability as larger annual totals provide greater
276 flexibility for interannual fluctuations. However, when accounting for the shift in baseline,
277 *relative* interannual variability decreases. As wet day frequency rises, the contribution of
278 extreme events toward annual totals is reduced, along with the likelihood that a given year of
279 precipitation will be unduly influenced by extreme outlier events. Consequently, year-over-year
280 annual precipitation totals become more consistent with more frequent precipitation. This
281 scenario can be seen in reverse for the Desert Southwest domain: interannual variability
282 decreases and relative interannual variability increases (Harp and Horton (*in review*) found no
283 shift in underlying precipitation intensities for the Desert Southwest). A similar, abbreviated
284 discussion of precipitation frequency on interannual variability can be found in Polade et al.
285 (2014).

286 The impacts of shifts in wet day precipitation intensity are more nuanced (compare
287 across Figures S3-S5). Generally, increases in mean wet day precipitation intensity alone (2) will
288 lead to increases in interannual variability, however, the standard deviation of the underlying
289 wet day precipitation intensity distribution has critical impacts on relative interannual
290 variability (Figures S3d-f, S4d-f, S5d-f). For example, if the standard deviation of wet day
291 precipitation intensity does not change, then an increase in the mean wet day precipitation
292 intensity leads to negligible impacts on relative interannual variability (Figure S3f). However,

293 an increase in standard deviation leads to an increase in relative interannual variability and vice
294 versa (Figures S4f, S5f). Ultimately, changing interannual variability is a byproduct of changes
295 in wet day frequency and the underlying precipitation distribution – both the change in mean
296 and standard deviation of the distribution are important – which can combine to produce
297 differential impacts on interannual variability and relative interannual variability. This is
298 illustrated by observed changes over the Northeast domain. Here, both wet day frequency and
299 intensity increase (Harp and Horton, *in review*) leading to a 2.4% rise in interannual variability
300 but no change in relative interannual variability, mirroring the hypothetical shown in Figures
301 S4h-i.

302

303 *Highlighted Domains: Northern Rockies, Central/Southern Plains*

304 As discussed above, generally speaking, the paths to shifting interannual precipitation
305 variability are driven by a combination of changes in precipitation frequency or the underlying
306 precipitation intensity distribution. Intriguingly, the Northern Rockies domain displays
307 decreases in both interannual variability and relative interannual variability despite observing
308 no statistically significant change in annual precipitation or wet day frequency. One potential
309 explanation for these discrepancies is that underlying trends exist in one or more of these
310 variables that do not rise to the level of statistical significance based on the data analyzed here.
311 It is also possible that there are shifts in circulation patterns or storm tracks which are persistent
312 *within* years but vary *between* years, such as a shift in atmospheric river frequency tied to modes
313 of climate variability. For instance, shifts in El Niño Southern Oscillation teleconnection patterns
314 could explain increased interannual variability in precipitation metrics, despite no long-term

315 trends in precipitation or the underlying intensity. We leave this as an avenue for future
316 research.

317 A second pair of notable domains are the Central and Southern Plains. These two
318 domains have the most substantial changes in both interannual variability (6.2% and 9.5%,
319 respectively) and relative interannual variability (6.1% and 9.6%, respectively) in either
320 direction despite modest changes in annual precipitation and wet day frequency. These changes
321 are likely driven by strong shifts in the underlying distribution of precipitation intensity toward
322 heavier rainfall (Harp and Horton, *in review*) with increases in mean wet day intensity of 4.6%
323 and 8%, respectively.

324

325 *Comparison with earlier literature*

326 Our results on changes in observed annual precipitation largely mirror earlier findings
327 from the fourth National Climate Assessment (Easterling et al., 2017) with subtle differences
328 over the southeastern and northwestern U.S. Additionally, we find similar trends in wet day
329 frequency as earlier *in-situ*, station-based observational studies such as Pal et al. (2013), though
330 there is some discrepancy in findings over the western U.S. Despite a similar observation-driven
331 and interannual variability-focused methodology, we identify differences between our findings
332 and those of Zhang et al. (2021), where a similar methodology was applied to observations in
333 NEON domains in the western U.S. Specifically, within the domains of overlap in our studies,
334 we find statistically significant changes in relative interannual variability for all domains except
335 for the Great Basin, while Zhang et al. find statistically significant changes in just three domains.
336 The identified trends for these three domains do, however, agree with our results. We similarly

337 find significant results across more domains for annual precipitation and wet day frequency
338 than Zhang et al., though the directions of any identified trends nearly perfectly overlap. These
339 discrepancies may be a byproduct of methodological decisions. For example, despite also using
340 GHCN-D data, Zhang et al. focus their analysis on the period from 1976-2019 and use a shorter
341 moving window (five years) for calculation of relative interannual variability, though our
342 sensitivity analysis did not reveal strong window width dependency.

343 While an imperfect comparison, we also compare our results of observed interannual
344 variability with a suite of studies using high emission scenario model projections to determine if
345 trends emerging in historical observations mirror future estimates. Our findings of increasing
346 interannual variability of precipitation in the midwest and northeast match those of Chou and
347 Lan (2012) and Pendergrass et al. (2017), though we disagree over the sign of change in the
348 northwest U.S. Both Chou and Lan (2012) and Pendergrass et al. (2017) attribute rising
349 interannual precipitation variability to greater moisture availability connected with increasing
350 temperatures. The spatial patterns in interannual variability shifts we identify also differ from
351 the generally uniform nationwide-increases projected by Wood et al. (2021) and Polade (2014),
352 particularly in the western U.S. Similarly, our findings of falling interannual variability in
353 California disagree with the modeled increases presented in multiple studies Berg and Hall
354 (2015) and Swain et al. (2018), though both studies do not predict an emergence of signal until
355 the middle 21st century. It should be noted that while our study examines changes in
356 interannual variability over a period of rapidly increasing greenhouse gas concentrations and
357 subsequent climate impacts, unlike the above studies, we do not explicitly examine the effects of
358 climate change on interannual variability.

359

360 *Limitations and Sensitivity Analysis Implications*

361 There are potential limitations of our study, beginning with an underlying assumption
362 that stations within NEON domains are relatively homogeneous. While NEON domains were
363 created to possess internally consistent climates, within-domain variability may exist and
364 inconsistent station availability may influence domain-level findings. The quantity of qualifying
365 stations varies also between domains and can impact the reliability of results, this is especially
366 true for the Atlantic Neotropical domain with only six qualifying stations.

367 Our sensitivity analysis revealed two domain clusters with start year-dependent results,
368 in agreement with Kunkel (2003) which describes the importance of length of record for
369 analysis, and notes that shorter time series may exhibit different trends than a greater length of
370 record for a similar location. First, the direction of interannual variability trends over the three
371 Plains domains and the Ozarks Complex between a 1950 and 1960 start date. Similarly, results
372 for the western U.S. show a distinct shift in precipitation trends between a 1950 or earlier start
373 date and a 1960 or 1970 start date. This shift occurs in trends for all metrics and across at least
374 half of the western NEON domains (Tables S3-S6). Thus, while we have focused on results
375 using a 1970 start date and 11-year moving window, we highlight that this combination should
376 not be considered definitive. We further include results of analysis based on a 1950 start date in
377 the Supporting Information (Figures S6-S9, Tables S13-S16). Finally, it should again be noted
378 that although we examine trends in precipitation through a period of time of increasing
379 greenhouse gas emissions and resultant climate impacts, we do not claim to directly attribute
380 changes to anthropogenic climate change.

381

382

383 **Conclusion**

384 We use curated daily *in situ* precipitation measurements from the GHCN to examine
385 domain-level trends in annual precipitation metrics, with a focus on interannual variability. We
386 identify rises in annual precipitation in the central and eastern U.S. and declines in the western
387 U.S. Trends in wet day frequency broadly mirror those of annual precipitation. We also reveal
388 significant trends in interannual precipitation variability and relative precipitation variability
389 across the United States, though with some differences in within-domain trends depending on
390 the variability metric of interest. Broadly, we find an increase in precipitation variability across
391 both metrics for the southeastern U.S., a decrease along the west coast, and mixed signals in the
392 central U.S. These findings have important implications for understanding the impact of
393 changing precipitation variability on agriculture and water resource planning. The full
394 complement of our results can be compared against climate model projections to inform climate
395 model analyses across the full spectrum of precipitation metrics. Finally, we recommend that
396 future studies carefully consider how interannual precipitation variability is characterized (i.e.,
397 interannual variability vs relative interannual variability) and any subsequent implications.

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400

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409 Information Technology.

410

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412 **Open Research and Availability Statement**

413 The National Centers for Environmental Information hosts publicly available Global Historical
414 Climatology Network Daily data at [https://www.ncei.noaa.gov/products/land-based-](https://www.ncei.noaa.gov/products/land-based-station/global-historical-climatology-network-daily)
415 [station/global-historical-climatology-network-daily](https://www.ncei.noaa.gov/products/land-based-station/global-historical-climatology-network-daily). Code developed by the authors to conduct
416 the analysis and produce the figures within this study is available at
417 [github.com/ryandharp/Observed_Changes_in_Interannual_Precipitation_Variability_in_the_U](https://github.com/ryandharp/Observed_Changes_in_Interannual_Precipitation_Variability_in_the_United_States)
418 [nited_States](https://github.com/ryandharp/Observed_Changes_in_Interannual_Precipitation_Variability_in_the_United_States). Analysis code will be placed and archived on Zenodo upon completion of the peer
419 review process, at which time the finalized link to archive, DOI, and data citation will be
420 included in this statement.

421

422

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