

1 **Hypoxic Blackwater Events - Identifying High Risk Catchments in Estuaries**
2 **Now and Under Future Climate Scenarios**

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7 **Key Points:**

- 8 • A methodology is presented to differentiate estuarine catchments by their potential to
9 generate hypoxic blackwater.
- 10 • Local topographic and hydrodynamic conditions strongly influence the potential volume
11 of blackwater generated over a catchment.
- 12 • Climate change may increase the frequency and severity and change the distribution of
13 blackwater risk throughout an estuary.

14 **Abstract**

15 Hypoxic blackwater events occur worldwide, affecting inland and coastal waters. These
16 events have been exacerbated by man-made floodplain drainage, leading to large-scale fish
17 kills and ecological degradation. This paper presents a new method to identify estuarine
18 catchment areas that are most likely to generate hypoxic conditions. The method uses
19 established blackwater risk factors, including vegetation type, inundation extent and
20 duration, ground-truthed in eastern Australia. A catchment is at higher risk of blackwater
21 generation if (i) it is located where floodwaters are high and/or drainage is impeded, (ii) the
22 site topography includes an extensive, low-lying floodplain; and/or (iii) the land-use and
23 environmental characteristics have a high deoxygenation potential. Blackwater impacts
24 in an estuary are determined by the floodplain connectivity with the estuary, and the
25 discharge characteristics of the catchment drainage system. Where multiple, proximate
26 catchments have similar drainage conditions, compounding blackwater plumes can
27 overwhelm the assimilation capacity of the estuary. Climate change may significantly
28 increase the volume and frequency of blackwater events in estuarine environments as a
29 result of reduced drainage due to sea level rise, higher temperatures, and more intense and
30 sporadic rainfall events. It is recommended that management measures be introduced to
31 mitigate the effects of climate change and avoid further widespread hypoxic blackwater
32 events.

33 **1 Introduction**

34 Blackwater is the common name for dark coloured water characterised by low levels of
35 dissolved oxygen (DO) (Howitt, Baldwin, Rees, & Williams, 2007; Moore, 2006). Blackwaters
36 are typically associated with the presence of humic substances leached from decaying
37 vegetation (Coble, Koenig, Potter, Parham, & McDowell, 2019) or floodplain sediments
38 (Meyer, 1990; Rixen et al., 2008). This transfer of dissolved organic matter and nutrients
39 from riparian zones and floodplains to adjacent waterways is a natural part of the nutrient
40 cycle. However, excessive stimulation of microbial respiration, particularly via elevated
41 levels of organic matter and higher temperatures, can lead to hypoxic conditions, whereby
42 DO concentrations are reduced below 2mg/L (Rabalais, Turner, & Wiseman Jr., 2002;
43 Vithana, Sullivan, & Shepherd, 2019). Indeed, hypoxic blackwater plumes have been linked
44 to decreased aquatic ecosystem health, including mass fish kills, and detrimental impacts on

45 sessile flora and fauna (Hladyz, Watkins, Whitworth, & Baldwin, 2011; Pahor & Newton,
46 2013; Vaquer-Sunyer & Duarte, 2008). Blackwater may also affect the broader estuarine
47 environment through disrupted trophic levels, interrupted life cycles, reductions in suitable
48 habitat, overcrowding, and forced migration (Rabalais et al., 2002; Vaquer-Sunyer & Duarte,
49 2008).

50 Globally, hypoxia is frequently associated with the decomposition of autochthonous algal
51 blooms fuelled by the eutrophication of inland and coastal waterways, as exemplified by the
52 infamous dead zones of the Gulf of Mexico (Diaz, 2001; Górnaiak, 2017), the Baltic (Conley et
53 al., 2011) and Black Seas (Rabalais et al., 2002). However, hypoxic conditions can also
54 develop following direct, precipitous carbon loading of lakes, rivers or estuaries as a result
55 of extended catchment inundation, in what are commonly known as blackwater events
56 (Moore, 2006). Globally, flood-induced blackwater events are widespread, having been
57 reported on the Paraguay River, Brazil (Hamilton, Sippel, Calheiros, & Melack, 1997), the
58 Atchafalaya River, USA (Bonvillain et al., 2011; Pasco et al., 2016) and in Lake Filsø, Denmark
59 (Kragh, Martinsen, Kristensen, & Sand-Jensen, 2020). In Australia, blackwater events have
60 affected inland rivers in arid and temperate regions of the Murray-Darling Basin (King,
61 Tonkin, & Lieshcke, 2012; Whitworth, Baldwin, & Kerr, 2012) and the Edward-Wakool River
62 system (Hladyz et al., 2011); tropical waters in the Katherine (Townsend, Boland, & Wrigley,
63 1992) and Mary (Townsend & Edwards, 2003) Rivers of the Northern Territory; and coastal
64 estuaries, including the Hunter (Carney et al., 2015; Hitchcock, Westhorpe, Glamore, &
65 Mitrovic, 2021), Clarence (Johnston, Kroon, Slavich, Cibilic, & Bruce, 2003) and Richmond
66 (Walsh, Copeland, & Westlake, 2004) Rivers in northern New South Wales (NSW).

67 While eutrophic and blackwater hypoxic events may occur naturally, anthropogenic
68 activities appear to have escalated their frequency and magnitude (Carstensen, Andersen,
69 Gustafsson, & Conley, 2014; Kerr, Baldwin, & Whitworth, 2013; Paerl, 2006; Rabalais et al.,
70 2002; Wong et al., 2010). Dead zones have spread exponentially since the 1960s (Diaz &
71 Rosenberg, 2008), with scientifically confirmed accounts of hypoxia affecting over 500
72 catchments globally (Breitburg et al., 2018; Díaz & Rosenberg, 2011). This increase has been
73 attributed to anthropogenic impacts, including wastewater discharges (Breitburg et al.,
74 2018), modifications to floodplain hydrology following river regulation and water extraction
75 (Whitworth et al., 2013), and the construction of drainage and flood mitigation works

76 (Moore, 2006), with changes to vegetation and land-use contributing to increased nutrient
77 loads (Arellano et al., 2019; Conley et al., 2007; Godinho et al., 2019).

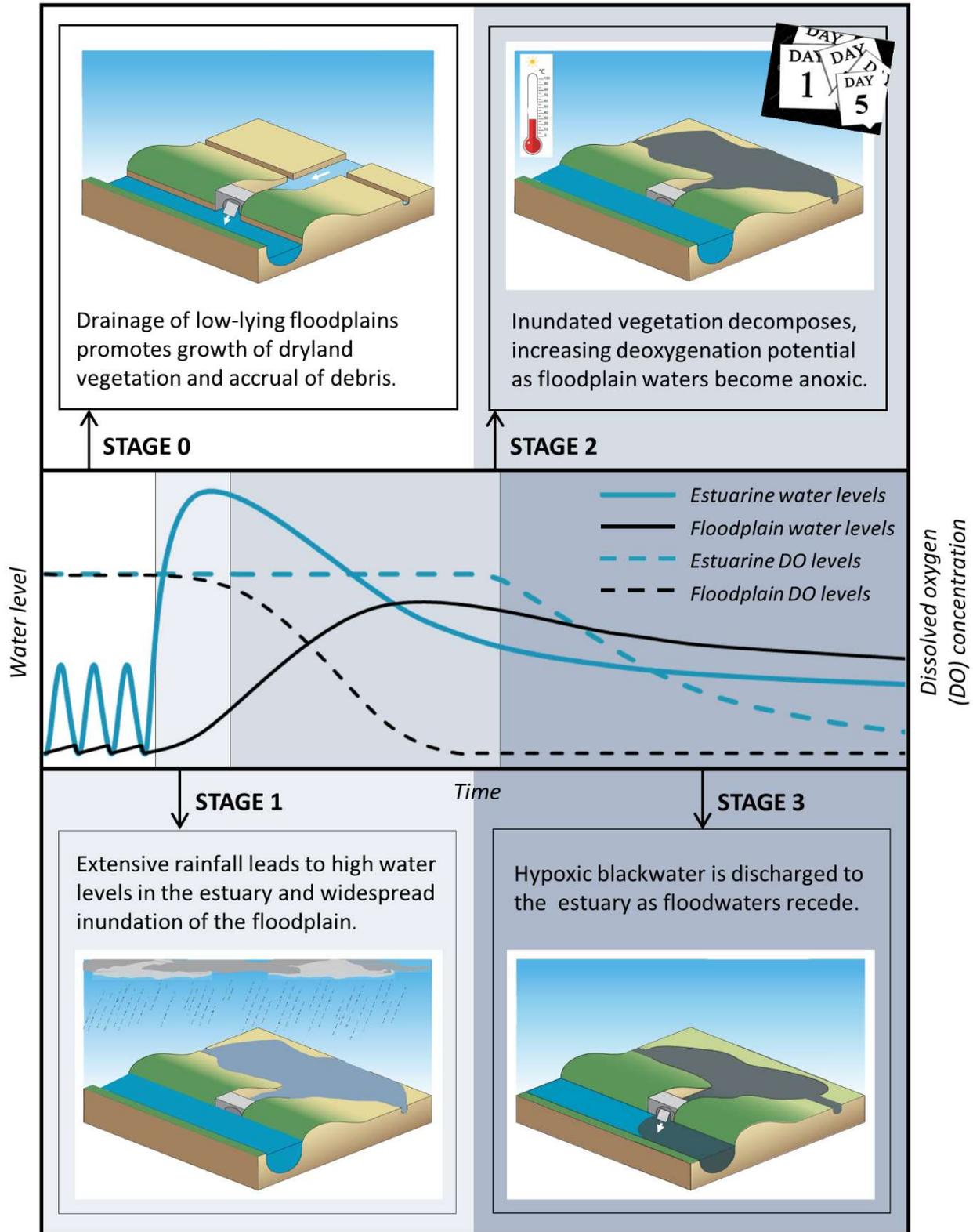
78 Efforts to mitigate the impacts of hypoxia have included nutrient reduction schemes,
79 primarily aimed at improving the quality of industrial and municipal wastewater discharges
80 (Conley, 2012), reducing the impacts of agricultural land-use practices (Tallis et al., 2019),
81 changes to river regulation procedures (Kerr et al., 2013; King et al., 2012; Watts et al.,
82 2018), and dilution/aeration techniques (Whitworth et al., 2013). However, effective
83 mitigation requires a thorough understanding of the mechanisms that contribute to the
84 formation and persistence of blackwater. To this aim, conceptual models have been
85 developed to improve blackwater management for the Atchafalaya River (Pasco et al.,
86 2016), Gulf of Mexico (Scavia et al., 2017) and San Francisco Bay (Cloern, Schraga, Nejad, &
87 Martin, 2020) in the United States, and the Murray-Darling River system in Australia
88 (Whitworth and Baldwin (2016)). For example, the Blackwater Risk Assessment Tool (BRAT)
89 developed by Howitt (2007) and Hladyz (2011) describes the carbon load generated over the
90 Murray-Darling floodplain via litter accumulation and decomposition rates, carbon leaching,
91 microbial degradation and respiration. The carbon load is then converted to a volumetric
92 biochemical oxygen demand (BOD) by considering the area, depth and duration of
93 inundation (Howitt et al., 2007; Whitworth & Baldwin, 2016). Other key variables are
94 temperature, re-aeration and dilution potential (Howitt et al., 2007). The Dissolved Oxygen-
95 Dissolved Organic Carbon (DODOC) model developed by Mosley, Wallace, Rahman, Roberts,
96 and Gibbs (2021) has subsequently integrated hydrologic capability with the BRAT ecological
97 assessment to better represent the complex interactions between catchment hydrology and
98 floodplain inundation (Gibbs, Wallace, & Mosley, 2022; Mosley et al., 2021). Consequently,
99 environmental flow releases in the regulated, lowland inland rivers of Australia are now
100 managed to avoid periods of high carbon production on floodplains (Saintilan, Kelleway,
101 Mazumder, Kobayashi, & Wen, 2021).

102 Many of the same risk factors that apply to the generation of blackwater over the
103 floodplains of inland rivers also apply to coastal floodplains. However, tidal flows present a
104 different hydrodynamic regime to that experienced in riverine environments, and within
105 many estuarine environments the retention and release of floodplain waters has been
106 heavily modified by extensive flood mitigation and drainage schemes. Over the course of

107 the 20th century, meandering channels with low hydraulic gradients and limited outlets have
108 been replaced by extensive networks of drainage channels that have breached the natural
109 hydraulic separation between an estuary and the adjacent coastal floodplains (Tulau, 2011).
110 Blackwaters that were once predominately retained within extensive tracts of freshwater
111 and tidal wetlands, allowing the carbon cycle to complete and the water column to become
112 re-aerated, are now discharged swiftly to the estuary (Johnston, Slavich, Sullivan, & Hirst,
113 2003).

114 These floodplain drainage schemes have been directly linked to blackwater events that
115 resulted in major fish kills throughout the northern estuaries of NSW (Walsh et al., 2004).
116 Such events closed the Macleay River to fishing for 3 months in 2001 (Walsh et al., 2004),
117 and the Richmond River for 4.5 months in 2001 (Steffe, Macbeth, & Murphy, 2007) and 2
118 months in 2008 (Wong et al., 2010). Consequently, analyses of blackwater events in the
119 Clarence (Johnston et al., 2003) and Richmond (Wong et al., 2010) Rivers of Australia have
120 identified three characteristic stages of an estuarine blackwater event (Figure 1):

- 121 1. Initially low-lying floodplain areas are inundated as floodwaters rise. The concentration
122 of dissolved organic carbon (DOC) in the water column starts to increase and microbial
123 metabolization of DOC rapidly reduces the DO concentration (Stage 1).
- 124 2. Floodplain waters become anoxic and the deoxygenation potential (DOP) increases as
125 the BOD continues to rise (Stage 2).
- 126 3. The assimilation capacity of the river decreases as the floodwaters recede and the high
127 oxygen demand of the blackwaters discharged from the floodplain drainage systems has
128 a substantial impact on the adjacent waterway (Stage 3). This stage may be exacerbated
129 by the drainage of acidic groundwaters from behind the flood mitigation structures.



130

131 **Figure 1.** Conceptual stages of a blackwater event in an estuary (graph adapted from

132 (Johnston et al., 2003))

133 The influence of constructed drainage systems and tidal floodgates on the magnitude and
134 frequency of hypoxic blackwater events is evident in each of the three stages identified in
135 Figure 1. Indeed, the construction of drainage schemes on estuarine floodplains has
136 increased the potential for blackwater generation for a number of reasons. First, enhanced
137 drainage promotes dry-land pastoral grasses and agricultural crops in favour of water-
138 tolerant vegetation species. These dry-land species are more susceptible to inundation and
139 can provide a highly labile source of carbon for microbial metabolism (Eyre, Kerr, &
140 Sullivan, 2006; Johnston et al., 2003). Second, the reduced frequency of floodplain
141 inundation increases the availability of carbon as organic debris accumulates between flood
142 events (Ning, Petrie, Gawne, Nielsen, & Rees, 2015; Pahor & Newton, 2013). Finally,
143 enhanced drainage intensifies the impact of blackwater by promoting the rapid discharge of
144 anoxic floodwaters with high DOP directly to the estuary and by increasing the volume of
145 blackwater that is transferred to the estuary (Wong et al., 2010).

146 Studies to date have focussed on the first two impacts of constructed drainage schemes;
147 developing a detailed understanding of the local mechanisms by which blackwater is
148 generated over estuarine floodplains. However, there have been limited investigations into
149 the hydrodynamic interactions between the estuary and floodplain. Indeed, variability of
150 rainfall and inundation across various drainage catchments within the floodplain has been
151 identified as a primary difficulty in prioritising areas for blackwater management (Moore,
152 2006). Estuarine water levels are fundamental to the retention of water on the floodplains
153 and the release of impounded water to the estuary. The former presents a limitation to the
154 volume of blackwater that may be generated on the floodplain, while the latter is a
155 determining factor regarding blackwater impacts upon the estuarine environment.

156 This paper presents a methodology to address this knowledge gap by incorporating the
157 hydrodynamic regime and the topographic constraints underlying the generation of hypoxic
158 blackwater in estuarine floodplains. The methodology quantifies the relative potential for
159 blackwater contribution from each drainage catchment within an estuary. Water quality
160 surveys from historic blackwater events in south-eastern Australia are used to explore the
161 validity of the blackwater risk assessment. The results provide insights into the susceptibility
162 of various catchments within an estuary to blackwater under current and future climate

163 conditions. It is anticipated that this approach may be used to optimise strategic monitoring
164 programmes and future management options.

165 **2 Methodology**

166 2.1 Blackwater risk factors

167 Differentiating the potential for blackwater generation between various catchments within
168 an estuary requires the identification and quantification of risk factors that may contribute
169 to a blackwater event. These include biological (carbon availability and microbial
170 metabolization), chemical (for example, inorganic reactions, acidity and salinity) and
171 physical (primarily temperature and pressure) mechanisms affecting the DOP of the
172 floodwaters. Further, the extent and duration of inundation over the catchment is critical to
173 the volume of blackwater that may be generated. An overview and the assessment of these
174 risk factors is presented in the following sections.

175 2.1.2 BOD and carbon availability

176 BOD is a critical factor contributing to changes in DO levels within a water column, with the
177 rate of oxidation assumed to be directly proportional to the BOD (Cox, 2003). BOD accounts
178 for both the chemical oxidation of inorganic cations such as Fe^{2+} , Mn^{2+} and S^{2-} (Johnston,
179 Slavich, & Hirst, 2005; Vithana et al., 2019; Wong et al., 2010), and the microbial
180 decomposition of biodegradable organic matter (Hladyz et al., 2011).

181 In coastal estuaries of southeast Australia, the impacts of chemical oxidation are associated
182 with acid sulfate soils (ASS) (Johnston et al., 2003). ASS are chemically stable when
183 undisturbed, however constructed drainage systems have exposed large floodplain areas to
184 oxygen, producing sulphuric acid and, via dissolution, metallic cations. Secondary oxidation
185 of these ions can create a significant oxygen demand within the water column (Johnston et
186 al., 2003; Lin, Wood, Haskins, Ryffel, & Lin, 2004) and the associated acidity has been
187 independently attributed to fish kills (Walsh et al., 2004). Analysis of the geochemical
188 signature of floodwaters in the main channel of the Clarence (Johnston et al., 2003) and
189 Richmond Rivers (Wong et al., 2010) identified the anaerobic decomposition
190 of floodplain vegetation as the primary process leading to generation of hypoxic blackwater
191 conditions. Further, mesocosm experiments in the Richmond (Eyre et al., 2006) and Edward-

192 Wakool (Hladyz et al., 2011) Rivers determined the BOD of a variety of vegetation types and
193 confirmed the potential for microbial decomposition of inundated floodplain vegetation to
194 be sufficient to trigger a blackwater event.

195 The rate at which vegetation decays and deoxygenates water during periods of prolonged
196 inundation differs depending on the vegetation type (Eyre et al., 2006; Johnston et al., 2005;
197 Whitworth & Baldwin, 2016), with DOP being a factor of, *inter alia*, the labile carbon
198 concentration (often measured as dissolved organic carbon) and temperature (Wong et al.,
199 2011). Labile carbon is the organic component that is readily bio-available, with a higher
200 lability associated with faster decomposition rates (Zhang et al., 2019). For example, when
201 examining the impacts of flooding observed in the Clarence River, Johnston et al. (2003)
202 attributed the relatively high oxygen demand from one catchment to the dominance of
203 labile dryland pasture species compared to another that was vegetated predominately by
204 recalcitrant *Melaleuca quinquenervia* forest. Thus, floodplains dominated by endemic
205 wetland plant species would be less likely to generate blackwater than those that have been
206 drained and revegetated with pasture grasses or crops that are less tolerant of inundation
207 (Vithana et al., 2019; Wong et al., 2011). Research determining the oxygen demand of
208 various vegetation, litter and soil types has subsequently been used to hypothesise the
209 relative risk and potential contribution of different land-uses to blackwater events (Liu,
210 Watts, Howitt, & McCasker, 2019).

211 2.1.2.1 Quantifying the DOP risk factor

212 It is difficult to establish strong links between land-use (for which spatial data is readily
213 available) and BOD (Amiri & Nakane, 2008). However, a review of Australian literature
214 identified five vegetation types typical of coastal floodplains for which experimental data
215 regarding DOP has been established. Experimental methods differ between the various
216 studies, making direct comparison difficult and there is limited data available regarding the
217 spatial distribution of vegetation on coastal floodplains. Consequently, a risk-based
218 approach linking vegetation types with comparative DOP was devised for this study, as
219 detailed in Table 1. The vegetation types and corresponding risk factors were then assigned
220 to each land-use identified in the 2017 (released in June 2020) Australian Land Use and
221 Management (ALUM) classification (DPIE, 2020). Full details are presented in the
222 Supplementary Material.

223 **Table 1.** Blackwater generation risk factors for typical vegetation types in coastal NSW

Vegetation Type	Water Tolerance		Comparative Deoxygenation Potential		Risk Factor**
	Rating	Score	Rating	Score	
Dryland grasses (e.g. pasture)	Low	3	High ^{a, c, e} Medium ^b	3	3
Forestry (other than tea tree) *	Low	3	High ^{b, c} Low ^d	3	3
Tea tree leaves*	Low	3	Low – Medium ^a Medium ^e	2	2
Sugar cane ⁺	Medium	2	Low – Medium ^a	2	2
Freshwater wetland grasses (e.g. Grey rush)	High	1	Low ^a Medium ^d	1	1

224 ^a Eyre et al. (2006)225 ^b Whitworth and Baldwin (2016)226 ^c Liu et al. (2019)227 ^d Johnston et al. (2005)228 ^e Southern Cross GeoScience (2019)229 * Low water tolerance is attributed to the presence of readily available leaf litter, rather
230 than the likelihood of plants dying.231 ⁺ Sugar cane is relatively tolerant to water, but the presence of waste after harvest increases
232 the DOP.

233 ** The risk factor is the average of the scores for water tolerance and comparative DOP.

234 2.1.3 Floodplain inundation characteristics

235 Flood mitigation and drainage systems have been implemented to increase floodplain
236 productivity by limiting floodplain inundation and increasing drainage efficiency. This
237 reduces the risk of blackwater generation during smaller, localised rainfall events, as
238 inundated catchments can drain freely when downstream waterways are not in flood.
239 However, regardless of the efficiency and scale of drainage infrastructure, floodplain
240 drainage is limited by the receiving (downstream) water level. Consequently, widespread
241 blackwater generation is more commonly associated with extensive flooding when
242 floodplain drainage is restricted by the rate of floodwater recession in the estuary.

243 2.1.3.1 Inundation duration

244 To determine the potential for blackwater generation in a floodplain, it is important to
245 quantify the inundation duration required to generate and sustain a blackwater event. This
246 will vary depending on catchment characteristics, seasonal and antecedent conditions, and
247 the unique hydrologic and hydraulic profile of each flood event.

248 The microbial metabolisation of highly labile carbon sources can rapidly deoxygenate a
249 water column. Organic compounds on the floodplains will start to decompose within hours
250 of inundation (Vithana et al., 2019; Wallace, Ganf, & Brookes, 2008), with the most labile
251 fractions leached of carbon within the first 24 hours. In-situ mesocosm experiments on the
252 floodplain of the Richmond River by Eyre et al. (2006) indicated that microbial
253 metabolisation of harvested sugar cane and dropped tea tree leaves can reduce the
254 dissolved oxygen levels in a 300 mm deep water column to 3 – 4 mg/L within 10 hours,
255 while slashed pasture grass can deoxygenate the same volume of river water almost
256 completely ($DO < 1$ mg/L) under the same conditions. Similar experiments have shown that
257 DO was reduced to near 0 mg/L over a period of two to three days for a variety of
258 vegetation types (Johnston et al., 2005; Liu et al., 2019; Vithana et al., 2019).

259 After prolonged immersion, living plants may start to die and more recalcitrant plants
260 decompose, contributing new carbon sources to the water column (Hladyz et al., 2011).
261 Additionally, experiments by Vithana et al. (2019) and Liu et al. (2019) suggest that both
262 DOC and BOD can continue to rise for more than two weeks after initial inundation.
263 Similarly, Wong et al. (2011) showed that chemical oxygen demand (COD) peaked

264 approximately 15 to 20 days after the flood peak in a backswamp on the Clarence River,
265 NSW. This suggests that the DOP is likely to persist for floodplain inundation durations of at
266 least two weeks.

267 During the 2001 floods in the Richmond, Clarence and Macleay Rivers, it was reported that
268 river water started to deoxygenate as floodplains began to drain, becoming completely
269 deoxygenated within 1 to 3 days, depending on site conditions (Walsh et al., 2004). Mass
270 deoxygenation of coastal estuaries in NSW is typically observed 4 to 6 days after the peak of
271 a flood event (Johnston et al., 2003; Southern Cross GeoScience, 2019; Wong et al., 2010).
272 In part, this reflects the recession of the flood hydrograph and the limited drainage from the
273 floodplain to the estuary during prolonged floods. However, similar observations were
274 made by Bonvillain et al. (2011) following flooding of the Atchafalaya River Basin (USA) in
275 September 2008, where hypoxic conditions were recorded within 3 days and extensive fish
276 kills occurred within 5 days.

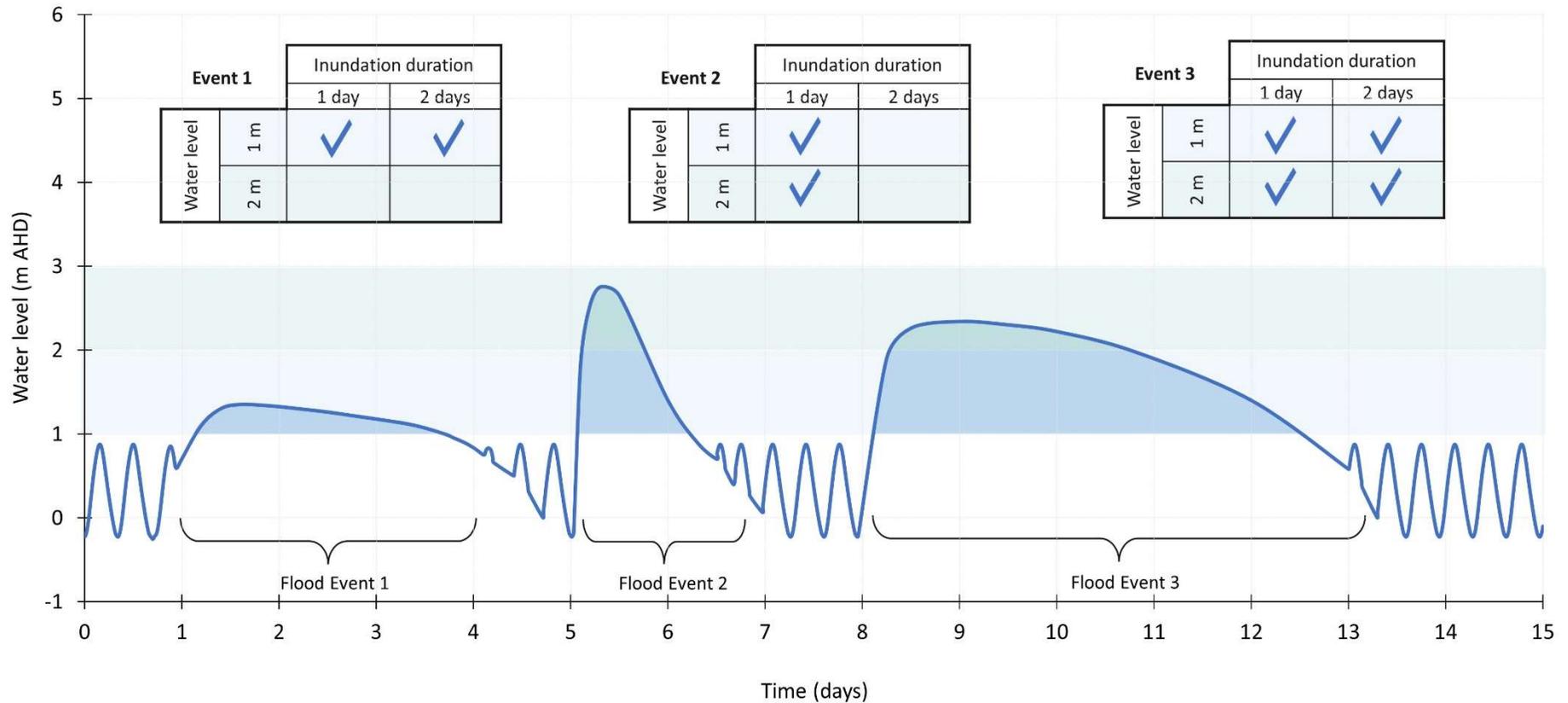
277 Both experimental evidence and recorded observations indicate that floodplain inundation
278 of less than 24 hours can generate blackwater during flood events. However, the limited
279 volume of blackwater generated during this short period is unlikely to have a significant
280 impact on receiving waters. Conversely, the longer that floodwaters are retained on a
281 floodplain, the greater the volume of water that can be deoxygenated (Hladysz et al., 2011),
282 the higher the DOP of the floodwaters (Eyre et al., 2006), and the greater the risk of a
283 blackwater event when those waters are discharged to the receiving estuary (Howitt et al.,
284 2007).

285 2.1.3.2 Inundation depth and extent

286 While there are no known long-term records of floodplain inundation within estuaries of
287 NSW, the Department of Planning and Environment maintains a network of water level
288 gauges throughout the estuarine channels (MHL, 2023). These gauges provide continuous
289 long-term, historic water level records which were adopted to represent the depth and
290 duration of floodplain inundation as onsite drainage is controlled by the downstream water
291 levels. The spatial extent of inundation across the floodplain was determined by
292 extrapolating these historic estuarine water levels using ground-truthed digital elevation
293 models (DEMs) with a 5 metre grid resolution (Geoscience Australia, 2018).

294 2.1.3.3 Inundation depth and duration matrices

295 Flood hydrographs are highly variable. The influence of the hydrograph shape on the
296 potential for blackwater generation is illustrated in Figure 2. For example, flood events 1
297 and 2 generate approximately the same volume of flow within the estuary, whereas the
298 higher peak water levels generated during event 2 have a shorter duration. Despite the
299 higher peak flood levels, for floodplain areas below 1m AHD (Australian Height Datum,
300 which equates approximately to mean sea level), the inundation conditions presented by
301 flood event 1 (Figure 2) are likely to generate more hypoxic blackwater than those of event
302 2, as inundation would be maintained over a longer period of time. Floodplain areas above
303 2m AHD would only generate blackwater during flood event 3 as they would not be
304 inundated during event 1 and there would be insufficient inundation duration during
305 event 2.



306

307 **Figure 2.** A conceptual hydrograph depicts three different flood events to highlight the blackwater potential risks for each event. At a
 308 floodplain elevation of 1m AHD, all floods contribute to the blackwater risk but for different inundation periods. At an elevation of 2m AHD,
 309 only flood events 2 and 3 would contribute blackwater. Flooding that persists over a broader area and a longer period (Flood Event 3) has a
 310 higher likelihood of producing larger volumes of blackwater.

311 Based on the conceptual model depicted in Figure 2, inundation risk matrices can be
312 established at each water level gauge by identifying the historic frequency at which various
313 combinations of inundation depth and duration have been exceeded. As such, the
314 inundation duration was calculated at 0.1m increments in water level between 0.1m and
315 5.0m AHD. This covers the range of water levels and topographic elevations typical of
316 estuarine floodplains. Additionally, as the tidal cycle will restrict discharges from the
317 floodplain drainage systems during flood events, the long-term average mean high water
318 (MHW) level (as documented by Fitzhenry, Alley, Hesse, and Couriel (2012)) was adopted as
319 the minimum floodplain inundation level at each gauge location.

320 Within the estuaries of NSW, recorded flooding events rarely exceed five days. As the critical
321 duration of inundation for the generation of blackwater may be achieved during floods with
322 a duration of less than one day, event durations between one and five days were adopted
323 for the inundation matrices. For each day a flood event exceeded any particular water level,
324 it was assumed to contribute to the blackwater risk for that duration. This ensured that
325 longer duration events contributed more blackwater risk than events of short duration,
326 although the matrix could be extended to incorporate additional flood durations for
327 estuaries regularly subject to longer floods.

328 To maximise the available data and account for any differences in the data record at each
329 gauge, the full data record was analysed for each water level monitoring station. These
330 results were then normalised by the length of the data record to calculate the average
331 recurrence interval of each incremental inundation level at each gauge location.

332 2.2 Aggregated blackwater risk assessment

333 For each catchment, a blackwater contribution factor was calculated using spatial analysis
334 tools within Geographic Information System (GIS) over a 5m square grid. This information
335 was used to determine the area inundated and the land-use risk factor associated with that
336 area for every 0.1m increment in water level (as illustrated in Figure 3). The blackwater
337 contribution factor corresponding to the water level was then calculated for all
338 combinations of inundation duration and frequency. A statistical mean of the factors
339 calculated from the matrix of inundation levels affecting each catchment was adopted as

340 the aggregated blackwater risk factor for that catchment. This factor was then used to rank
341 the catchments according to their relative potential to generate blackwater.

342 2.3 Study area

343 The methodology described in Section 2 was applied to seven major estuaries within NSW,
344 including (from north to south, as indicated in Figure 4):

- 345 • Tweed River;
- 346 • Richmond River;
- 347 • Clarence River;
- 348 • Macleay River;
- 349 • Hastings River;
- 350 • Manning River; and
- 351 • Shoalhaven River.

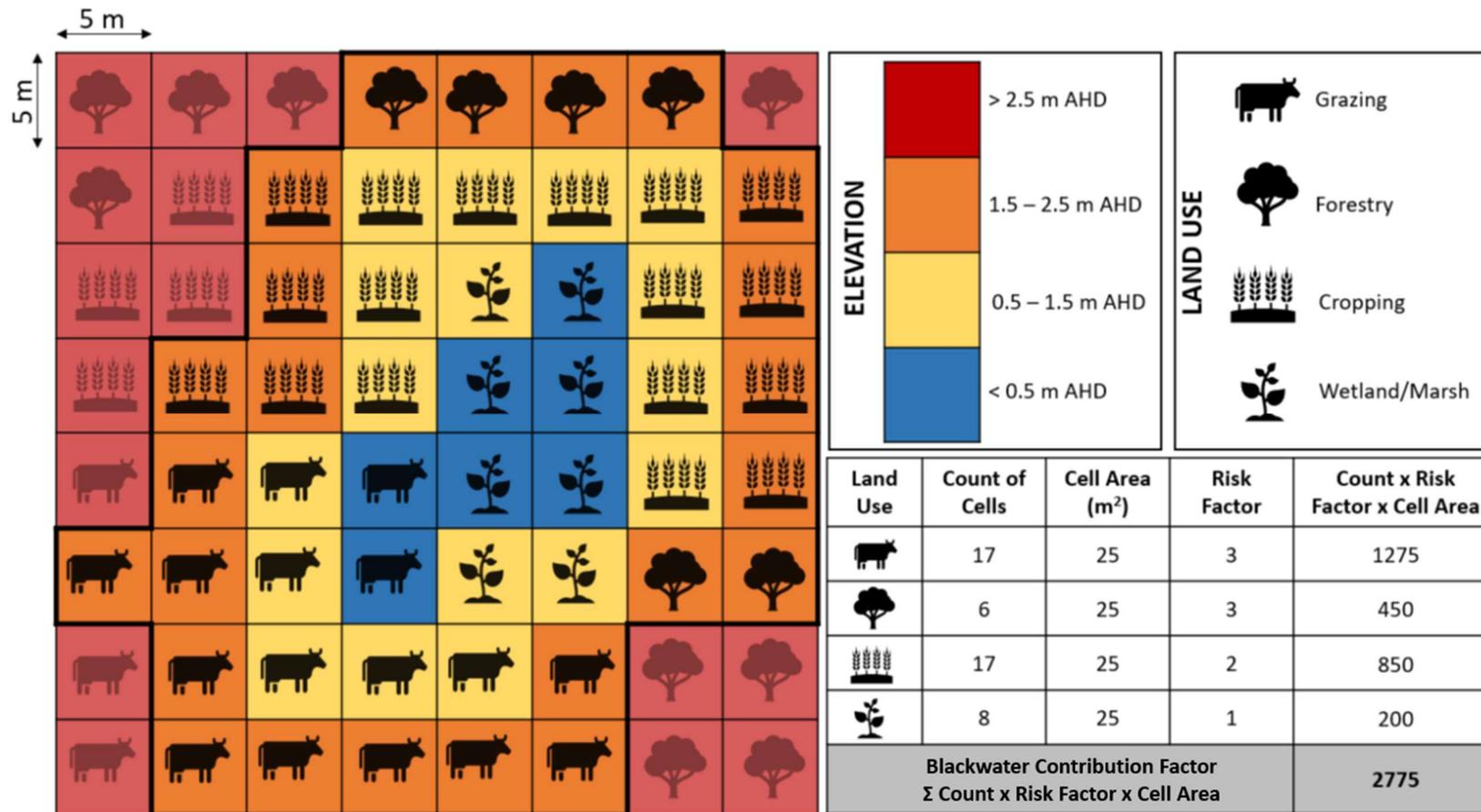
352 Each selected estuary has been previously identified as blackwater pollution sites resulting
353 from the clearing and drainage of coastal wetlands (Fletcher and Fisk, 2017). In particular,
354 the Richmond River has been the focus of detailed investigations into the causes of
355 blackwater events that resulted in mass fish kills.

356 Details of government water level gauges used in this assessment are identified in the
357 Supplementary Material, including gauge locations and historic flow distribution curves.

358 **3 Results**

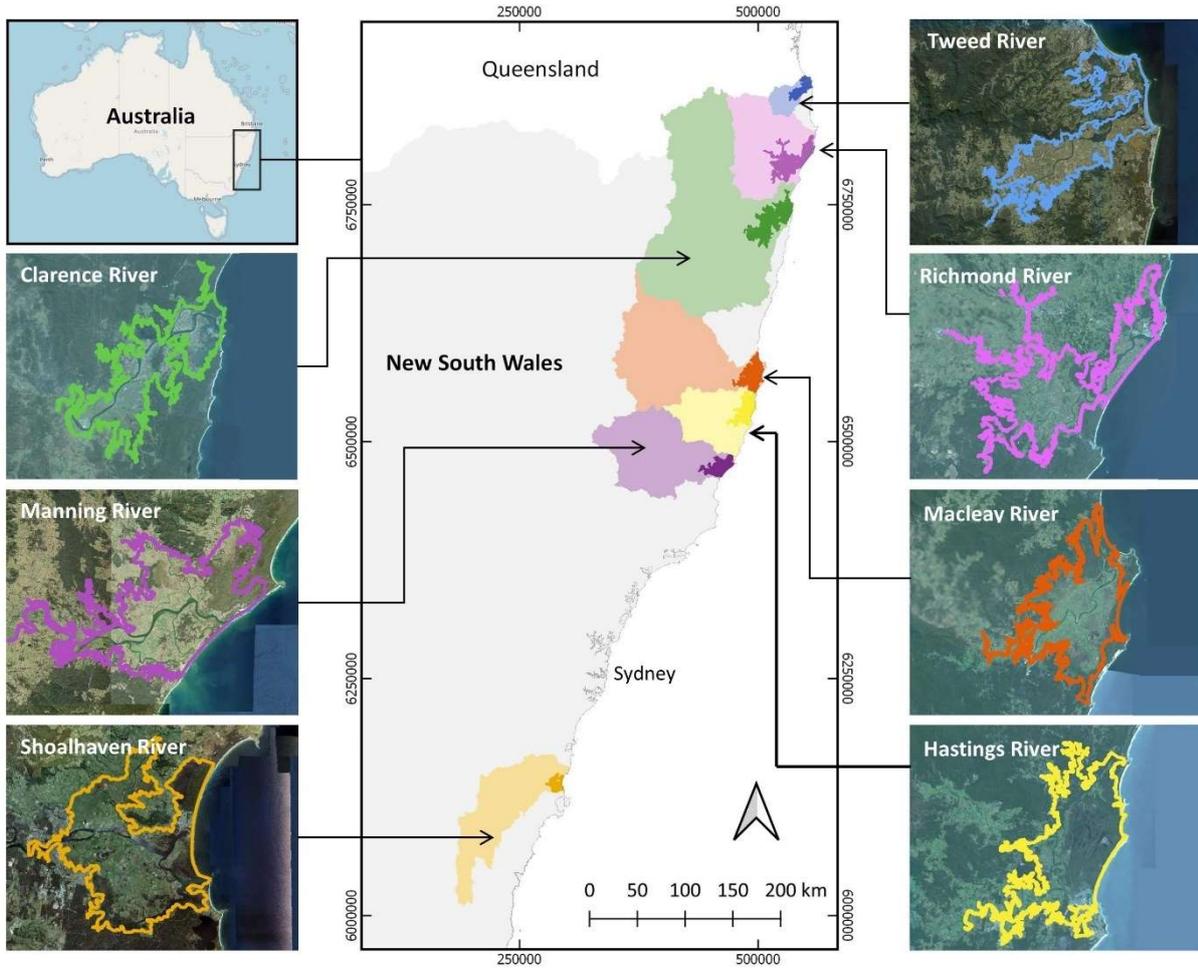
359 The calculated aggregated blackwater risk factors provide an objective, data-driven
360 evidence base to identify which catchments present the highest potential risk of generating
361 blackwater. The ranking of catchments based on blackwater risk can be used to inform and
362 prioritise floodplain management options. Ongoing and future monitoring programs may be
363 optimised and used to further validate and refine the assessment methodology.

364 The risk factors and catchment ranking within each estuary are tabulated with a statistical
365 analysis of the corresponding historic water levels in the Supplementary Material. Maps
366 indicating the distribution of blackwater risk are also provided, with sample results for the
367 Richmond River presented in Figure 5. These maps incorporate the median extent of

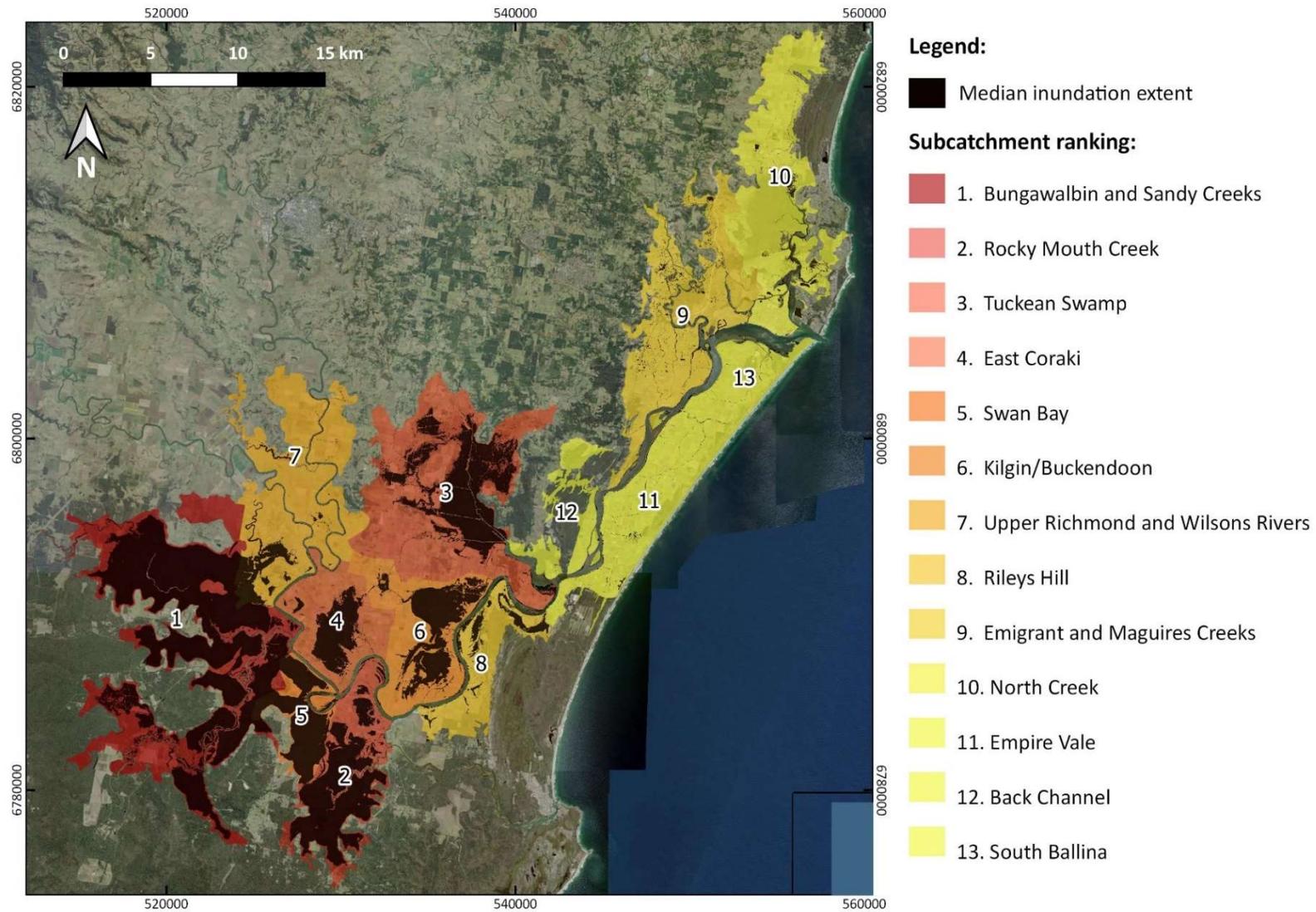


368

369 **Figure 3.** An example calculation of a catchment blackwater contribution factor for an inundation level of 2.5 m AHD. The coloured matrix
 370 depicts the land-use type and elevation. The number of each land-use cells multiplied by the area and risk factor summed for the catchment is
 371 the blackwater contribution factor for that catchment. Within an estuary, multiple catchments were calculated to rank priority risk areas.



372
373 **Figure 4.** Selected estuaries (shaded) and estuarine floodplain areas below 5m AHD (outlined)
374 defining the study area.



375

376 **Figure 5.** Ranking of catchments by aggregated blackwater risk factor and extent of inundation under median water levels for the Richmond
377 River estuary (maps of results for all estuaries within the study area are included in the Supplementary Material).

378 inundation corresponding to the historic water levels in the estuary to provide an indication
379 of the areal extent of potential blackwater generation within each catchment.

380 3.1 Validation of results

381 In the Richmond River the highest risk catchments were identified as Bungawalbyn and
382 Sandy Creeks (ranked 1), Rocky Mouth Creek (ranked 2), and the Tuckean Swamp (ranked
383 3). These results correlate well with several previous scientific investigations and help to
384 validate the assessment outcomes. Exceptionally low oxygen levels were previously
385 recorded at Bungawalbyn and Rocky Mouth Creeks following the February 2001 flood (Eyre
386 et al., 2006). Similarly, Wong et al. (2010) identified the Tuckean, Bungawalbyn, Sandy and
387 Rocky Mouth Creek catchments as primary sources of DOP following the January 2008 flood.
388 The same three catchments were attributed the highest risk of blackwater generation and
389 impact by an expert panel under the Richmond Estuary Ecosystem Health Monitoring
390 Strategy (Moore, 2006) and identified as the three top priority areas for blackwater
391 management within the Richmond River floodplain by a collaborative Australian Research
392 Council project (Southern Cross GeoScience, 2019). The latter study prioritised lower
393 Bungawalbyn Creek as the largest blackwater generator on the floodplain.

394 The validity of the blackwater risk methodology is strengthened by a review of water
395 quality, fish and crustacean mortality data undertaken by Walsh et al. (2004) following
396 major flooding in February and March 2001. This investigation also identified blackwater risk
397 areas within the Richmond River as Rocky Mouth Creek, Bungawalbyn Creek and Tuckean
398 Swamp. Additionally, Walsh et al. (2004) prioritised the Coldstream River (ranked 1 under
399 this study's methodology) and Everlasting Swamp (in the Sportsmans Creek catchment,
400 ranked 2) in the Clarence River, and Belmore Swamp (ranked 2), the Swan Pool (Kinchela
401 Creek, ranked 1) and Seven Oaks wetland (Collombatti-Clybucca, ranked 3) in the Macleay
402 River, further supporting this study's methodology.

403 3.2 Limitations of methodology

404 3.2.1 Land use, ambient and antecedent conditions

405 Due to the relative homogeneity of land-use across the study area, the catchment rankings
406 presented herein are highly influenced by the catchment size and inundation extent. The

407 Spearman's rank correlation between the aggregated blackwater risk factor and the area
408 flooded at the median inundation level varied from 0.82 to 0.97 throughout each estuary,
409 with the exception of the Shoalhaven River, where it was 0.55. Within the Shoalhaven River,
410 the largest discrepancy was within the Comerong Island catchment, where over 12% of the
411 land area is below the median inundation level but does not contribute to the blackwater
412 risk factor as it is a tidal wetland. Further refinement of the weighting of blackwater risk
413 factors may therefore be required in landscapes with more diverse land uses.

414 Where land-use is more critical in the assessment of blackwater risk, a weighting factor
415 scaled to BOD may provide greater differentiation than the direct rank-order weighting
416 adopted in this study. Scaled weightings may vary with time of inundation, as indicated by
417 the results of the mesocosm and inundation experiments on which the rank-order
418 weightings were based, (e.g. Liu et al. (2019) and Vithana et al. (2019)). The depth of
419 inundation may also influence the rate of deoxygenation as shallow waters are likely to be
420 warmer (increasing microbial metabolisation rates and reducing oxygen saturation levels)
421 and subject to photochemical deoxygenation processes (Southern Cross GeoScience, 2019).

422 Indeed, decomposition rates are strongly affected by environmental factors such as
423 temperature, solar radiation, salinity and acidity (Voß, Fernández, & Schäfer, 2015) and BOD
424 has been shown to respond to plant density, shadowing, soil mineralogy and light
425 transmission (Cox, 2003; Voß et al., 2015). Conversely, the bioavailability of organic matter
426 has generally been found to respond more directly to land management practices, such as
427 chemical and nutrient application, harvesting practices and stocking rates (Voß et al., 2015).
428 Thus the potential for blackwater generation over various catchment areas may
429 alternatively be differentiated by assessing variations in land management (both historic
430 and current) and intensity of use (Barlow, Christy, & Weeks, 2009; Buck, Niyogi, &
431 Townsend, 2004) rather than vegetation types or land-use categories, as the accumulation
432 of surface litter is likely to drive the majority of oxygen demand in many environments
433 (Mehring et al., 2014).

434 Similarly, seasonal changes and antecedent conditions will also impact the potential BOD. By
435 reviewing antecedent weather conditions in the Richmond River, Wong, Walsh, and Morris
436 (2018) found that fish kills were more common when the previous six months had been
437 drier than usual prior to the blackwater flooding event. Conversely, wetter than average

438 conditions are likely to reduce the amount of DOC that is available, thereby lowering the risk
439 of blackwater generation (Hladyz et al., 2011). Antecedent conditions will influence the
440 accumulation of organic matter and the bioavailability of carbon on the floodplain.

441 Vegetation stress due to drought, for example, may increase litterfall (Whitworth et al.,
442 2012), with the accumulation of organic debris increasing as the time between flood events
443 is extended (Wong et al., 2018; Xiong, 1997). The amount of carbon leached from organic
444 matter also reduces when the litter has been previously inundated.

445 Nevertheless, the spatial and temporal variability of environmental influences on BOD is
446 unlikely to affect the intrinsic risk across a catchment or estuary within the study area
447 presented in this assessment (refer to Supplementary Material). Indeed, current literature
448 suggests that various vegetation and land cover types all have the ability to deoxygenate a
449 waterbody (Kobayashi et al., 2009; Liu et al., 2019; O'Connell, Baldwin, Robertson, & Rees,
450 2000). Under equivalent environmental conditions, it would therefore be the extent and
451 duration of inundation throughout a catchment which provides the greatest risk differential
452 for blackwater generation.

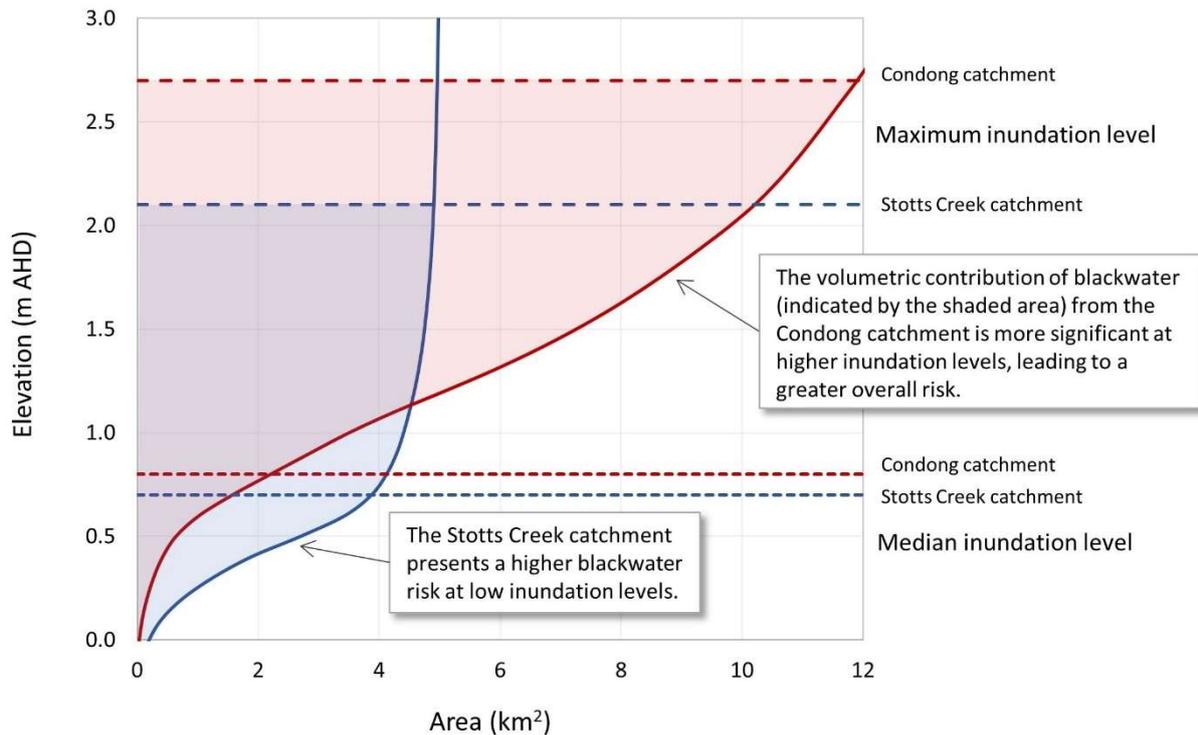
453 3.2.2 Water level data

454 The accuracy of the assessment with respect to inundation characteristics relies on the
455 suitability of the water level data available. In this regard, the results for the Shoalhaven
456 River may also have been adversely affected by the limited distribution of water level data
457 within major tributaries of the estuary. No gauges are located within the estuarine reaches
458 of Broughton Creek, while water levels on the Crookhaven River are only recorded at the
459 downstream limits of the estuary (refer to Supplementary Material). As the inundation
460 extents were estimated by a direct extrapolation of the estuarine water levels, it is
461 important to obtain a representative distribution of flood conditions throughout the
462 estuary. If the adopted water levels are lower than those typically experienced at any
463 catchment (as may occur when upstream water levels are poorly represented), the
464 inundated area, and corresponding aggregated blackwater risk factor, will be
465 underestimated.

466 4 Discussion**467 4.1 Floodplain inundation characteristics**

468 Results from the catchment rankings provide insights into the influence of estuarine
469 hydrodynamics on the blackwater risk profiles. For example, in the Tweed River the median
470 extent of inundation over the Stotts Creek catchment (3.9km²) exceeds that experienced in
471 the Condong catchment (2.2km²), yet Condong presents a higher overall blackwater risk
472 than Stotts Creek (Figure 5). This reflects the differences in catchment size and topography
473 as well as the water surface elevations.

474 As illustrated in Figure 6, Stotts Creek presents a relatively flat, low-lying floodplain below
475 the local median inundation level of 0.7m AHD. However, the topography rises steeply, with
476 limited additional catchment area contributing to potential blackwater generation at higher
477 water levels. Conversely, the Condong catchment would produce limited volumes of
478 blackwater until the surface water levels exceed 0.5m AHD. Once inundation levels reach
479 1.2m AHD, the area contributing to blackwater generation in the Condong catchment would
480 exceed the Stotts Creek catchment and the risk factor would increase accordingly. At higher
481 inundation levels (experienced during more significant, but less frequent rainfall events), the
482 large potential volumetric contribution of blackwater from the Condong catchment
483 outweighs that of the Stotts Creek catchment. In general, both greater topographic
484 exposure and higher inundation levels increase the overall blackwater risk presented by any
485 particular catchment to the receiving estuary.



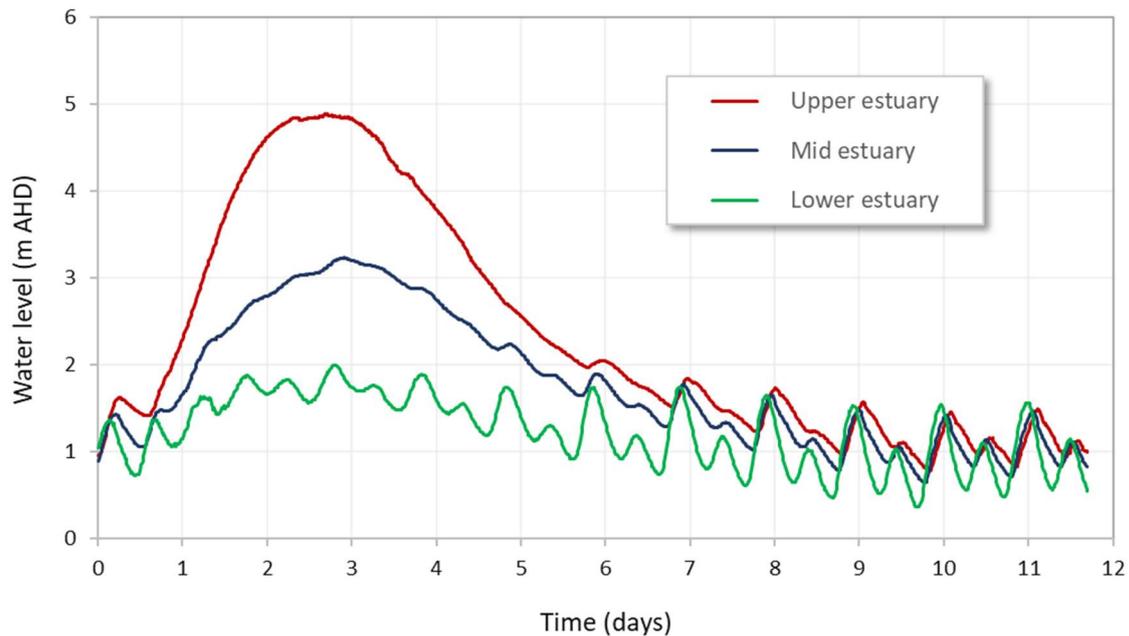
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487 **Figure 6.** Hypsographic graph of the Stotts Creek (blue) and Condong (red) catchments in
 488 the Tweed River estuary, illustrating the changing risk profiles based on inundation levels.

489 Comparing the Stotts Creek and Condong catchments (Figure 6) also highlights the
 490 sensitivity of the method to the duration and frequency of flood events. This was further
 491 investigated by modifying the inundation matrix to assess the aggregated blackwater risk
 492 factor. Removing the 1 and 2 day inundation durations from the aggregated blackwater risk
 493 assessment matrix reduced the overall inundation levels, as lower water levels are
 494 experienced during the rising and falling limbs of a flood hydrograph. Conversely, limiting
 495 the analysis to flood events with a recurrence interval of 2 to 5 years or 3 to 5 years typically
 496 increased the median inundation levels as fewer minor events were included in the
 497 statistics. Importantly, the relative risk remained consistent throughout this sensitivity
 498 analysis, with Spearman's rank correlation remaining above 0.9 in all estuaries except the
 499 Shoalhaven River (detailed in the Supplementary Material). Notably, the ranking of
 500 catchments within the Shoalhaven River was more sensitive to the influence of shorter
 501 duration events, although the correlation between the top five ranked catchments
 502 remained above 0.9 in all analyses.

503 The robustness of the assessment to variations in the flood duration and recurrence interval
504 reflects the consistency of the hydrodynamic response to flood events throughout each
505 estuary. As illustrated in Figure 7, peak flood levels are typically highest in the upper
506 estuary, where water levels rise rapidly and remain high due to restricted drainage from
507 elevated tailwaters in the mid- and lower portions of the estuary. Flood profiles in the lower
508 estuary are further moderated as the hydraulic energy and flow volumes are dispersed over
509 the low-lying floodplains. Additionally, in the lower estuary water levels are less affected by
510 flood flows as offshore waters can assimilate large flood volumes. These effects are
511 exemplified in the comparison between the Condong (located in the upper estuary and
512 subject to higher inundation levels) and Stotts Creek (mid-estuary) catchments (Figure 6).

513 Based on these spatial differences in drainage across an estuary, the free-draining lower
514 reaches of an estuary typically have a lower risk of blackwater generation. The
515 environmental impacts of blackwater discharged from these downstream catchments may
516 also be mitigated by high tidal flushing and reduced residence times due to their proximity
517 to the ocean (Johnston et al., 2003; Rabalais et al., 2002). To this aim, Eyre and Twigg (1997)
518 indicated that dissolved oxygen levels were higher in parts of the estuary with higher salinity
519 (or increased tidal flushing) for the first seven weeks after the flood event. In contrast, the
520 upper estuary is likely to have a higher blackwater risk as residence times may be prolonged
521 (i.e. any blackwater released would remain in the estuary for extended periods) and flood
522 levels tend to remain elevated for longer periods (i.e. any blackwater generated may
523 discharge into an estuary with less assimilation capacity).



524

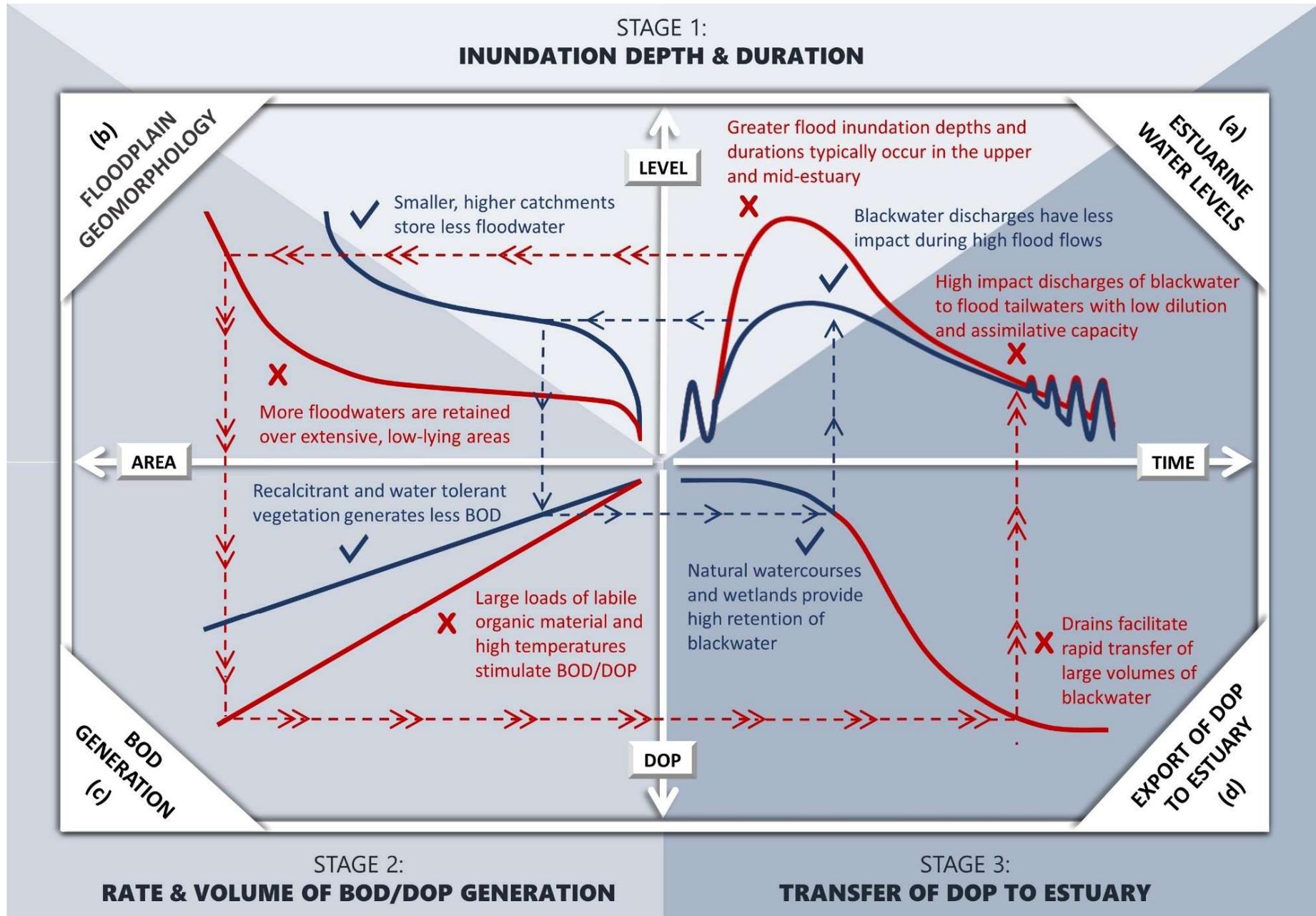
525 **Figure 7.** Spatial illustration of typical water level response to a flood in the lower, mid and
 526 upper reaches of an estuary from an event recorded in the Richmond River in June 2005.

527 The upper estuary was measured at Coraki (refer to Supplementary Material for gauge
 528 locations) and represents conditions for the Upper Richmond and Wilsons River catchments
 529 (Figure 5). The mid estuary was measured at Woodburn (Rocky Mouth Creek catchment)
 530 and the lower estuary at Wardell (Empire Vale and Back Channel catchments).

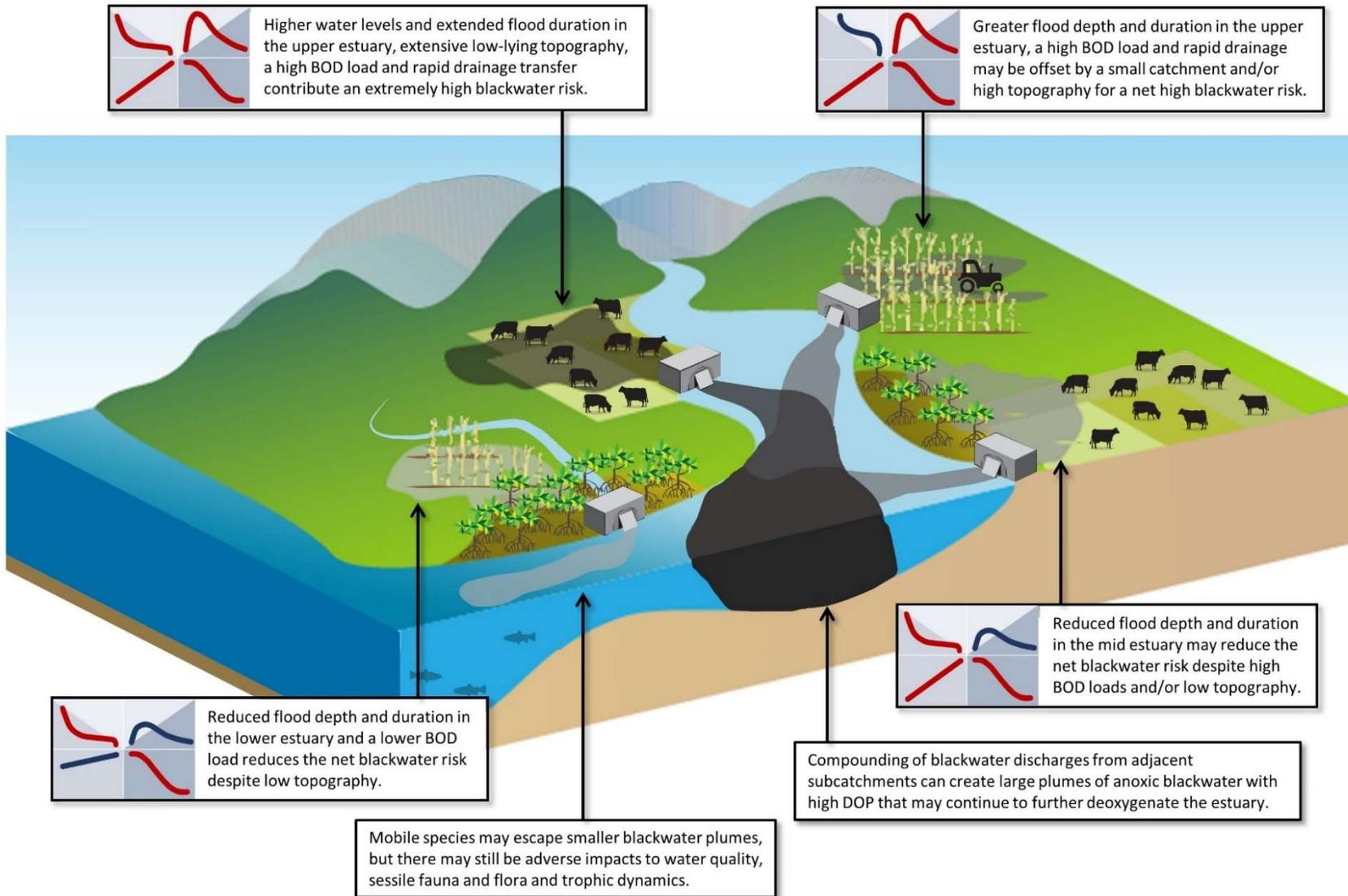
531 As inundation levels and inundation durations are often spatially correlated within an
 532 estuary, the risk factors for blackwater generation and impact may be simplified to the
 533 estuarine water levels, floodplain geomorphology, and the potential for BOD generation.
 534 These primary factors are conceptualised in Figure 8 for an idealised system. Based on these
 535 factors, a catchment may present a higher risk of blackwater generation and impact if:

- 536 • it is located in a portion of the estuary where floodwaters are maintained at higher
 537 levels and/or for longer durations (Figure 8(a));
- 538 • the topography contains an extensive, low-lying floodplain that can accommodate large
 539 volumes of floodwater (Figure 8(b)); and/or
- 540 • the land use and environmental characteristics have a higher potential to generate BOD
 541 (Figure 8(c)).

542 The environmental impact of blackwater in an estuary is then determined by the floodplain
543 connectivity with the estuary and the discharge characteristics of the catchment drainage
544 system (Figure 8(d)). Modern drainage systems facilitate the rapid transfer of DOP to the
545 estuary, with substantial volumes of blackwater discharged as the flood recedes (Figure
546 8(a)). These discharges can be particularly harmful when there is limited assimilative
547 capacity in the estuary. The impact from any individual catchment discharge is also affected
548 by discharges from nearby catchments as individual blackwater plumes can intermix. The
549 dilution capacity and potential environmental impacts from these blackwater plumes are
550 highlighted in Figure 9.



552 **Figure 8.** Conceptual diagram of the key risk factors for blackwater generation on an estuarine floodplain. The diagram should be read counter
553 clockwise from the top right corner. A high-risk scenario is realised by following the progression of a blackwater event via the red arrows from
554 Stages 1 to 3. Commencing in quadrant (a), high flood levels in the estuary lead to extensive inundation (b) during Stage 1. Increased
555 biochemical oxygen demand (BOD) generated at Stage 2 (c) will result in a high deoxygenation potential (DOP). The greatest impacts occur
556 when the blackwater discharges at Stage 3 (d) overwhelm the assimilative capacity of the receiving waters (a). An alternative low risk scenario
557 is proposed via the blue arrows.



559 **Figure 9.** Conceptual diagram of common blackwater risk factors and compounding discharges within an estuary. The risk from each drainage
560 catchment is described with reference to the four risk factors illustrated in Figure 8.

561 4.2 Spatial distribution of floodplain drainage

562 Impacts of compounding blackwater plumes are typified by the blackwater risk profiles of
563 the Clarence and Richmond Rivers. The inundation extent at the median inundation level for
564 the Clarence River (285 km²) is more than twice the Richmond River (137 km²). However,
565 the impact of blackwater events in the Richmond River has been more severe, with
566 extensive fish kills regularly reported after comparable flood events (Walsh et al., 2004).
567 This has previously been attributed to the substantially larger river discharges and
568 assimilation capacity of the Clarence River, which also provides opportunities for fish to seek
569 refuge in the less affected parts of the estuary (Walsh et al., 2004).

570 Results from the aggregated blackwater risk analysis indicate that the highest ranked sub-
571 catchments in the Clarence River discharge to different parts of the estuary. As such,
572 blackwater discharges are likely to impact the estuary at different stages of a flood event. In
573 comparison, the Richmond River has four of the five highest ranking sub-catchments
574 discharging into the estuary within a 20km reach. With such close proximity, these
575 individual plumes are likely to intermix, resulting in compounding impacts that are more
576 likely to overwhelm the assimilation capacity of the estuary. These potential compounding
577 impacts are illustrated in Figure 9.

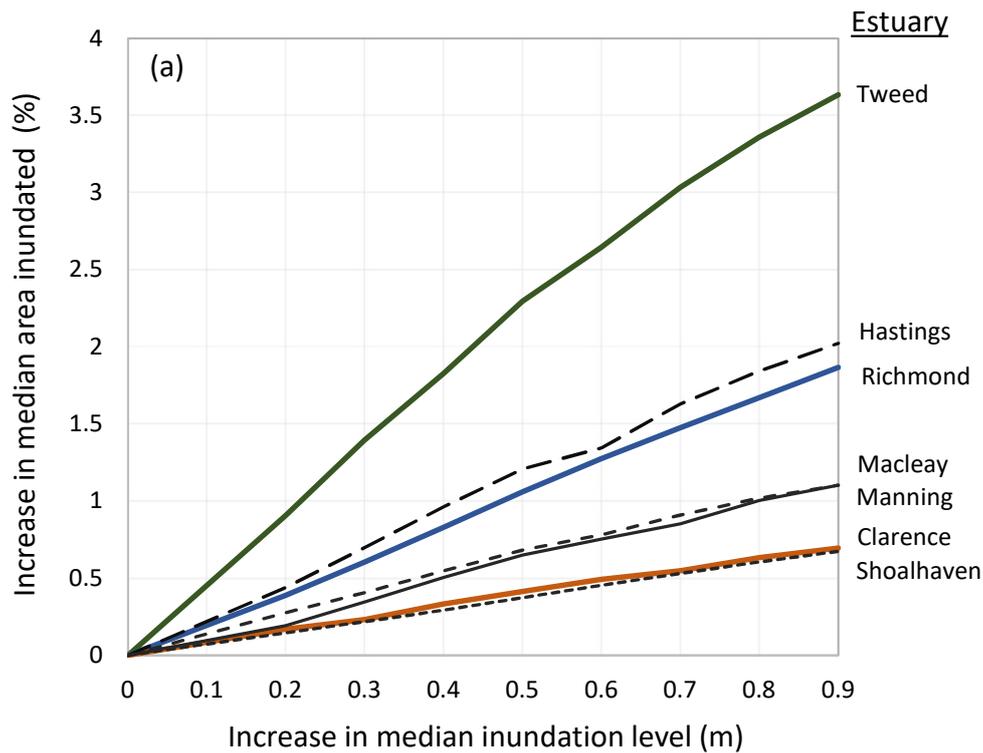
578 4.3 Climate change impacts

579 Investigating the sensitivity of spatial and temporal factors on blackwater generation also
580 provides insights into how climate change may influence future blackwater conditions.
581 Climate change has been predicted to increase temperatures (Masson-Delmotte et al.,
582 2021), and create more intense and sporadic rainfall conditions (Dey, Lewis, Arblaster, &
583 Abram, 2019). These impacts are expected to exacerbate hypoxic blackwater events
584 (Carstensen et al., 2014; Godinho et al., 2019; Koehn, 2022; Vaquer-Sunyer & Duarte, 2008),
585 with larger floods increasing the spatial and temporal inundation patterns, while warmer
586 conditions may increase the microbial reaction rates (Wong et al., 2010). Changes in flood
587 frequency and meteorological conditions may also encourage the accumulation of organic
588 matter between floods providing high BOD potential.

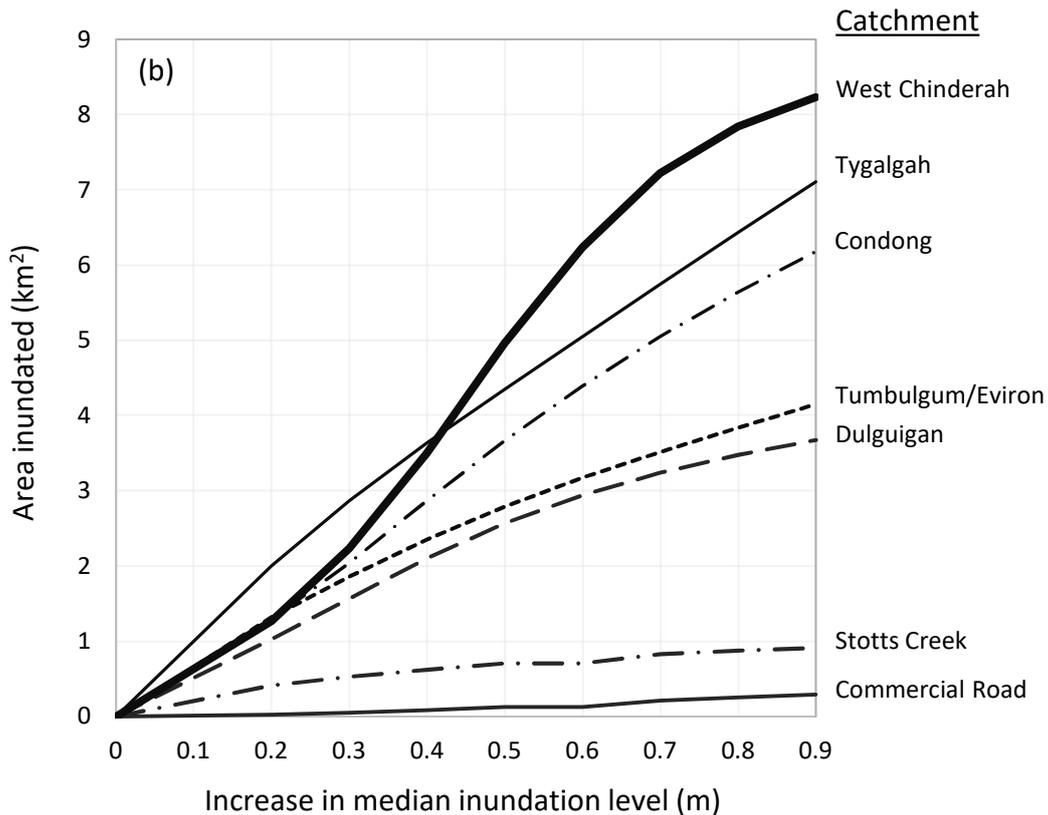
589 In estuarine environments these effects will be further compounded by sea level rise, which
590 is expected to increase standing water levels and reduce drainage (Waddington et al., 2022).

591 Higher downstream water levels may lead to increased spatial inundation, while reduced
 592 drainage will extend the temporal inundation. Overall, both factors will likely increase
 593 blackwater generation risk.

594 The potential for blackwater generation from sea level rise was investigated by examining
 595 the projected increase in area inundated for incremental increases in water levels across the
 596 floodplain sub-catchments. Unsurprisingly, the larger Clarence, Richmond and Macleay
 597 River systems, which already suffer the highest rate of fish mortality from blackwater
 598 events, would experience the greatest increases under rising sea levels. Normalising the
 599 increase in area inundated by the total catchment area (Figure 10(a)) indicates that the
 600 future risk of blackwater creation in the Tweed River is likely to be substantially higher than
 601 in other estuaries. This is primarily attributable to the topography of the Tweed River
 602 floodplain, where there is a greater increase in the area inundated by higher water levels.



603



604

605 **Figure 10.** The (a) increase in median area inundated after an incremental rise in water level
 606 expressed as a percentage of the total catchment area. Within any given estuary (the
 607 Tweed River is used as an example here) the sub-catchments most extensively inundated
 608 and the locations discharging the most blackwater may change (b).

609 This analysis can be further extended to identify which catchments within a particular
 610 estuary present an increased risk of blackwater generation under existing or future
 611 conditions. For example, while the West Chinderah catchment of the Tweed River estuary
 612 currently presents a relatively moderate risk, the potential for blackwater generation may
 613 increase exponentially once the median inundation level rises by approximately 0.3m
 614 (Figure 10(c)).

615 4.4 Management and mitigation of blackwater risk

616 The improved understanding of estuarine blackwater events, as provided by this
 617 assessment, enables a nuanced assessment of potential mitigation measures. In areas
 618 identified as low risk for blackwater generation, minor changes to land management may
 619 mitigate blackwater generation by reducing the bioavailability of carbon. For example, the
 620 removal of cuttings from slashed pastures has been suggested as an effective means of

621 reducing the DOP of grazing lands (Eyre et al., 2006). Alternatively, mechanical oxygenation
622 has been employed to mitigate blackwater impacts in the Swan and Canning Rivers of
623 Western Australia (Greenop, Lovatt, & Robb, 2001), although this is generally regarded as an
624 emergency response measure as it is substantially limited by the area that can be serviced
625 and operational costs (Baldwin, Boys, Rohlf, Ellis, & Pera, 2022).

626 As the potential for blackwater generation increases, the impact of effective mitigation
627 measures on current land uses is also likely to increase. Where the median inundation level
628 is relatively shallow, the depth and density of floodplain drains may be reduced to
629 encourage the growth of water-tolerant vegetation and facilitate reaeration of the water
630 column (Hamilton et al., 1997; Rixen et al., 2008). However, in areas subject to prolonged,
631 deep inundation, reaeration will be hampered by the lower ratio of the air-water interface
632 to flood volume and there is likely to be greater plant morbidity and contribution to BOD.
633 Such conditions are typified in the low energy backswamp environments, which may be best
634 suited to reinstatement into natural wetland systems. Indeed, tidal restoration and the
635 reinstatement of coastal and floodplain wetlands has been identified as the preferred
636 management measure for the mitigation of blackwater events in the coastal estuaries of
637 NSW (Moore, 2006).

638 Social and economic ramifications accompanying land-use change may be significant due to
639 the level of development throughout the floodplains (Moore, 2006;
640 Southern Cross GeoScience, 2019). Identification of areas at highest risk of backwater
641 generation (as discussed in this paper) is recommended to support trials and further
642 investigations into these options. This will ensure that decisions to reinstate wetland
643 systems are evidence-based and transparent, optimising water quality benefits and ongoing
644 floodplain productivity.

645 **5 Conclusion**

646 Hypoxic blackwater events occur worldwide and affect both inland and coastal waters. The
647 mechanisms underpinning these events are associated with the microbial metabolism of
648 carbon and the accumulation/discharge of deoxygenated water during and post flood
649 events. This paper presents a methodology developed to prioritise catchments within an
650 estuarine floodplain based on their potential to generate hypoxic blackwater and to identify

651 those catchments from which blackwater discharges are likely to have the most significant
652 impact on the estuary. Local topography and changes to the flood hydrograph as it
653 progresses along an estuary are shown to influence the extent and duration of inundation
654 over the floodplain. In turn, inundation characteristics are identified as critical factors in
655 determining the volume of blackwater generated and the DOP discharged to the receiving
656 waters. Concerns related to increased blackwater generation from climate change, and sea
657 level rise in particular, are highlighted.

658 It is anticipated that this research will be used to inform evidence-based decision making to
659 manage catchment risks to water quality. This may enable a strategic approach to future
660 investments in floodplain and estuarine research and management measures to optimise
661 economic, social and environmental outcomes.

662 **Acknowledgements**

663 The blackwater risk assessment methodology presented herein was developed with funding
664 from the New South Wales Marine Estate Management Strategy (MEMS). Katrina
665 Waddington was supported by a scholarship jointly funded by UNSW Sydney and MEMS.
666 The authors would like to thank Anna Blacka from UNSW Sydney for her assistance with the
667 preparation of figures. We also thank Danial Khojasteh for his constructive comments which
668 helped us to greatly improve this manuscript.

669 **Declaration**

670 The authors declare that they have no known competing financial interests or personal
671 relationships that could have appeared to influence the work reported in this paper.

672 **Data availability statement**

673 Water level data was sourced from the NSW Water Level Data Collection Program (MHL,
674 2023) available at <https://mhl.nsw.gov.au/Data-Level>. Land-use mapping was based on the
675 Australian Land Use and Management (ALUM) classification (DPIE, 2020). Climate (monthly
676 temperature) data presented in the Supplementary Material was downloaded from the
677 Climate Data Online service of the Australia Bureau of Meteorology website [Climate Data
678 Online - Map search \(bom.gov.au\)](#). Mapping was undertaken using the QGIS software, which
679 can be freely downloaded from [Discover QGIS](#). Digital elevation data (Geoscience Australia,

680 2018) was obtained from the National Elevation Data Framework spatial dataset [Elvis](https://elvis.fsdf.org.au)
681 [\(fsdf.org.au\)](https://elvis.fsdf.org.au) managed by Geoscience Australia [Digital Elevation Data | Geoscience Australia](https://digital.elevation.gov.au)
682 [\(ga.gov.au\)](https://digital.elevation.gov.au).

683 **References**

- 684 Amiri, B. J., & Nakane, K. (2008). Entire catchment and buffer zone approaches to modeling linkage
685 between river water quality and land cover—A case study of Yamaguchi Prefecture, Japan. *Chinese*
686 *Geographical Science*, *18*(1), 85-92. doi:10.1007/s11769-008-0085-6
- 687 Arellano, A., Bianchi, T. S., Osburn, C., D'Sa, E., Ward, N. D., Oviedo-Vargas, D., . . . Kurian, G. (2019).
688 Mechanisms of Organic Matter Export in Estuaries with Contrasting Carbon Sources. *Journal of*
689 *Geophysical Research: Biogeosciences*, *124*(10), 3168-3188.
- 690 Baldwin, D. S., Boys, C. A., Rohlf, A.-M., Ellis, I., & Pera, J. (2022). Field trials to determine the
691 efficacy of aerators to mitigate hypoxia in inland waterways. *Marine and Freshwater Research*, *73*(2),
692 211-222. doi:<https://doi.org/10.1071/MF20365>
- 693 Barlow, K., Christy, B., & Weeks, A. (2009). *Nutrient generation and transport at the catchment scale*.
694 Paper presented at the 18th World IMACS Congress and MODSIM09 International Congress on
695 Modelling and Simulation.
- 696 Bonvillain, C. P., Halloran, B. T., Boswell, K. M., Kelso, W. E., Harlan, A. R., & Rutherford, D. A. (2011).
697 Acute Physicochemical Effects in a Large River-Floodplain System Associated with the Passage of
698 Hurricane Gustav. *Wetlands*, *31*(5), 979. doi:10.1007/s13157-011-0213-4
- 699 Breitburg, D., Levin, L. A., Oschlies, A., Grégoire, M., Chavez, F. P., Conley, D. J., . . . Zhang, J. (2018).
700 Declining oxygen in the global ocean and coastal waters. *Science*, *359*(6371), eaam7240.
701 doi:doi:10.1126/science.aam7240
- 702 Buck, O., Niyogi, D. K., & Townsend, C. R. (2004). Scale-dependence of land use effects on water
703 quality of streams in agricultural catchments. *Environmental Pollution*, *130*(2), 287-299.
704 doi:<https://doi.org/10.1016/j.envpol.2003.10.018>
- 705 Carney, R. L., Mitrovic, S. M., Jeffries, T., Westhorpe, D., Curlevski, N., & Seymour, J. R. (2015). River
706 bacterioplankton community responses to a high inflow event. *Aquatic Microbial Ecology*, *75*(3),
707 187-205.
- 708 Carstensen, J., Andersen, J. H., Gustafsson, B. G., & Conley, D. J. (2014). Deoxygenation of the Baltic
709 Sea during the last century. *Proceedings of the National Academy of Sciences*, *111*(15), 5628-5633.
710 doi:doi:10.1073/pnas.1323156111

- 711 Cloern, J. E., Schraga, T. S., Nejad, E., & Martin, C. (2020). Nutrient Status of San Francisco Bay and Its
712 Management Implications. *Estuaries and Coasts*, 43(6), 1299-1317. doi:10.1007/s12237-020-00737-
713 w
- 714 Coble, A. A., Koenig, L. E., Potter, J. D., Parham, L. M., & McDowell, W. H. (2019). Homogenization of
715 dissolved organic matter within a river network occurs in the smallest headwaters. *Biogeochemistry*,
716 143(1), 85-104.
- 717 Conley, D. J. (2012). Save the Baltic Sea. *Nature*, 486(7404), 463-464. doi:10.1038/486463a
- 718 Conley, D. J., Carstensen, J., Ærtebjerg, G., Christensen, P. B., Dalsgaard, T., Hansen, J. L. S., &
719 Josefson, A. B. (2007). LONG-TERM CHANGES AND IMPACTS OF HYPOXIA IN DANISH COASTAL
720 WATERS. *Ecological Applications*, 17(sp5), S165-S184. doi:10.1890/05-0766.1
- 721 Conley, D. J., Carstensen, J., Aigars, J., Axe, P., Bonsdorff, E., Eremina, T., . . . Zillén, L. (2011). Hypoxia
722 Is Increasing in the Coastal Zone of the Baltic Sea. *Environmental Science & Technology*, 45(16),
723 6777-6783. doi:10.1021/es201212r
- 724 Cox, B. A. (2003). A review of currently available in-stream water-quality models and their
725 applicability for simulating dissolved oxygen in lowland rivers. *Science of The Total Environment*, 314-
726 316, 335-377. doi:[https://doi.org/10.1016/S0048-9697\(03\)00063-9](https://doi.org/10.1016/S0048-9697(03)00063-9)
- 727 Dey, R., Lewis, S. C., Arblaster, J. M., & Abram, N. J. (2019). A review of past and projected changes in
728 Australia's rainfall. *WIREs Climate Change*, 10(3), e577. doi:<https://doi.org/10.1002/wcc.577>
- 729 Diaz, R. J. (2001). Overview of Hypoxia around the World. *Journal of environmental quality*, 30(2),
730 275-281. doi:10.2134/jeq2001.302275x
- 731 Diaz, R. J., & Rosenberg, R. (2008). Spreading Dead Zones and Consequences for Marine Ecosystems.
732 *Science*, 321(5891), 926-929. doi:doi:10.1126/science.1156401
- 733 Díaz, R. J., & Rosenberg, R. (2011). Introduction to Environmental and Economic Consequences of
734 Hypoxia. *International Journal of Water Resources Development*, 27(1), 71-82.
735 doi:10.1080/07900627.2010.531379
- 736 DPIE. (2020). NSW Landuse 2017 Version 1.2. Retrieved from
737 <https://datasets.seed.nsw.gov.au/dataset/nsw-landuse-2017-v1p2-f0ed>
- 738 Eyre, B., & Twigg, C. (1997). Nutrient behaviour during post-flood recovery of the Richmond River
739 Estuary, Northern NSW, Australia. *Estuarine, Coastal and Shelf Science*, 44(3), 311-326.

- 740 Eyre, B. D., Kerr, G., & Sullivan, L. A. (2006). Deoxygenation potential of the Richmond River Estuary
741 floodplain, northern NSW, Australia. *River Research and Applications*, 22(9), 981-992.
742 doi:10.1002/rra.950
- 743 Fitzhenry, M., Alley, K., Hesse, S., & Couriel, E. (2012). *OEH NSW Tidal Planes Analysis: 1990 - 2010*
744 *Harmonic Analysis*. Retrieved from
745 Geoscience Australia. (2018). Retrieved from <https://elevation.fsdf.org.au/>
- 746 Gibbs, M. S., Wallace, T., & Mosley, L. M. (2022). Constraining organic matter composition and
747 dynamics as a dominant driver of hypoxic blackwater risk during river Murray floodplain inundation.
748 *Hydrological processes*, 36(3), e14529. doi:<https://doi.org/10.1002/hyp.14529>
- 749 Godinho, F. N., Segurado, P., Franco, A., Pinheiro, P., Pádua, J., Rivaes, R., & Ramos, P. (2019).
750 Factors related to fish kill events in Mediterranean reservoirs. *Water Research*, 158, 280-290.
751 doi:<https://doi.org/10.1016/j.watres.2019.04.027>
- 752 Górnjak, A. (2017). A new version of the Hydrochemical Dystrophy Index to evaluate dystrophy in
753 lakes. *Ecological Indicators*, 78, 566-573. doi:<https://doi.org/10.1016/j.ecolind.2017.03.030>
- 754 Greenop, B., Lovatt, K., & Robb, M. (2001). The use of artificial oxygenation to reduce nutrient
755 availability in the Canning River, Western Australia. *Water Sci Technol*, 43(9), 133-144.
- 756 Hamilton, S. K., Sippel, S. J., Calheiros, D. F., & Melack, J. M. (1997). An anoxic event and other
757 biogeochemical effects of the Pantanal wetland on the Paraguay River. *Limnology and*
758 *Oceanography*, 42(2), 257-272.
- 759 Hitchcock, J. N., Westhorpe, D., Glamore, W., & Mitrovic, S. (2021). Estuarine zooplankton responses
760 to flood pulses and a hypoxic blackwater event.
- 761 Hladyz, S., Watkins, S. C., Whitworth, K. L., & Baldwin, D. S. (2011). Flows and hypoxic blackwater
762 events in managed ephemeral river channels. *Journal of Hydrology*, 401(1), 117-125.
763 doi:<https://doi.org/10.1016/j.jhydrol.2011.02.014>
- 764 Howitt, J. A., Baldwin, D. S., Rees, G. N., & Williams, J. L. (2007). Modelling blackwater: Predicting
765 water quality during flooding of lowland river forests. *Ecological Modelling*, 203(3), 229-242.
766 doi:<https://doi.org/10.1016/j.ecolmodel.2006.11.017>
- 767 Johnston, S., Kroon, F., Slavich, P., Cibilic, A., & Bruce, A. (2003). Restoring the balance-guidelines for
768 managing floodgates and drainage systems on coastal floodplains.

- 769 Johnston, S. G., Slavich, P. G., & Hirst, P. (2005). Changes in surface water quality after inundation of
770 acid sulfate soils of different vegetation cover. *Soil Research*, *43*(1), 1-12.
771 doi:<https://doi.org/10.1071/SR04073>
- 772 Johnston, S. G., Slavich, P. G., Sullivan, L. A., & Hirst, P. (2003). Artificial drainage of floodwaters from
773 sulfidic backswamps: effects on deoxygenation in an Australian estuary. *Marine and Freshwater*
774 *Research*, *54*(6), 781-795. doi:<https://doi.org/10.1071/MF02016>
- 775 Kerr, J. L., Baldwin, D. S., & Whitworth, K. L. (2013). Options for managing hypoxic blackwater events
776 in river systems: A review. *Journal of Environmental Management*, *114*, 139-147.
777 doi:<https://doi.org/10.1016/j.jenvman.2012.10.013>
- 778 King, A. J., Tonkin, Z., & Lieshcke, J. (2012). Short-term effects of a prolonged blackwater event on
779 aquatic fauna in the Murray River, Australia: considerations for future events. *Marine and*
780 *Freshwater Research*, *63*(7), 576-586. doi:<https://doi.org/10.1071/MF11275>
- 781 Kobayashi, T., Ryder, D. S., Gordon, G., Shannon, I., Ingleton, T., Carpenter, M., & Jacobs, S. J. (2009).
782 Short-term response of nutrients, carbon and planktonic microbial communities to floodplain
783 wetland inundation. *Aquatic Ecology*, *43*(4), 843-858. doi:10.1007/s10452-008-9219-2
- 784 Koehn, J. D. (2022). Key steps to improve the assessment, evaluation and management of fish kills:
785 lessons from the Murray–Darling River system, Australia. *Marine and Freshwater Research*, *73*(2),
786 269-281. doi:<https://doi.org/10.1071/MF20375>
- 787 Kragh, T., Martinsen, K. T., Kristensen, E., & Sand-Jensen, K. (2020). From drought to flood: Sudden
788 carbon inflow causes whole-lake anoxia and massive fish kill in a large shallow lake. *Science of The*
789 *Total Environment*, *739*, 140072. doi:<https://doi.org/10.1016/j.scitotenv.2020.140072>
- 790 Lin, C., Wood, M., Haskins, P., Ryffel, T., & Lin, J. (2004). Controls on water acidification and de-
791 oxygenation in an estuarine waterway, eastern Australia. *Estuarine, Coastal and Shelf Science*, *61*(1),
792 55-63.
- 793 Liu, X., Watts, R. J., Howitt, J. A., & McCasker, N. (2019). Carbon and nutrient release from
794 experimental inundation of agricultural and forested floodplain soil and vegetation: influence of
795 floodplain land use on the development of hypoxic blackwater during floods. *Marine and Freshwater*
796 *Research*. doi:10.1071/mf18452
- 797 Masson-Delmotte, V., Zhai, P., Priani, A., Connors, S. L., Pean, C., Berger, S., . . . Zhou, B. (2021). *IPCC:*
798 *Climate Change 2021: The Physical Science Basis. Contribution of Working Group I to the Sixth*
799 *Assessment Report of the Intergovernmental Panel on Climate Change*. Retrieved from

- 800 Mehring, A. S., Kuehn, K. A., Tant, C. J., Pringle, C. M., Lowrance, R. R., & Vellidis, G. (2014).
801 Contribution of surface leaf-litter breakdown and forest composition to benthic oxygen demand and
802 ecosystem respiration in a South Georgia blackwater river. *Freshwater Science*, 33(2), 377-389.
803 doi:10.1086/675507
- 804 Meyer, J. L. (1990). A blackwater perspective on riverine ecosystems. *BioScience*, 40(9), 643-651.
- 805 MHL. (2023). New South Wales water level data collection program. Retrieved from
806 <https://mhl.nsw.gov.au/Data-Level>
- 807 Moore, A. (2006). Blackwater and fish kills in the Richmond River estuary. Defining the issues -
808 Assessing the risks - Providing management options. *Southern Cross University: Lismore, NSW,*
809 *Australia.*
- 810 Mosley, L. M., Wallace, T., Rahman, J., Roberts, T., & Gibbs, M. (2021). An integrated model to
811 predict and prevent hypoxia in floodplain-river systems. *Journal of Environmental Management*, 286,
812 112213. doi:<https://doi.org/10.1016/j.jenvman.2021.112213>
- 813 Ning, N. S. P., Petrie, R., Gawne, B., Nielsen, D. L., & Rees, G. N. (2015). Hypoxic blackwater events
814 suppress the emergence of zooplankton from wetland sediments. *Aquatic Sciences*, 77(2), 221-230.
815 doi:10.1007/s00027-014-0382-3
- 816 O'Connell, M., Baldwin, D. S., Robertson, A. I., & Rees, G. (2000). Release and bioavailability of
817 dissolved organic matter from floodplain litter: influence of origin and oxygen levels. *Freshwater*
818 *Biology*, 45(3), 333-342. doi:10.1111/j.1365-2427.2000.00627.x
- 819 Paerl, H. W. (2006). Assessing and managing nutrient-enhanced eutrophication in estuarine and
820 coastal waters: Interactive effects of human and climatic perturbations. *Ecological Engineering*,
821 26(1), 40-54. doi:<https://doi.org/10.1016/j.ecoleng.2005.09.006>
- 822 Pahor, S., & Newton, G. (2013). *Understanding blackwater events in rivers* (Vol. 22): Australian
823 Network for Plant Conservation Inc.
- 824 Pasco, T. E., Kaller, M. D., Harlan, R., Kelso, W. E., Rutherford, D. A., & Roberts, S. (2016). Predicting
825 Floodplain Hypoxia in the Atchafalaya River, Louisiana, USA, a Large, Regulated Southern Floodplain
826 River System. *River Research and Applications*, 32(5), 845-855. doi:<https://doi.org/10.1002/rra.2903>
- 827 Rabalais, N. N., Turner, R. E., & Wiseman Jr., W. J. (2002). Gulf of Mexico Hypoxia, A.K.A. "The Dead
828 Zone". *Annual Review of Ecology and Systematics*, 33(1), 235-263.
829 doi:10.1146/annurev.ecolsys.33.010802.150513

- 830 Rixen, T., Baum, A., Pohlmann, T., Balzer, W., Samiaji, J., & Jose, C. (2008). The Siak, a tropical black
831 water river in central Sumatra on the verge of anoxia. *Biogeochemistry*, *90*(2), 129-140.
832 doi:10.1007/s10533-008-9239-y
- 833 Saintilan, N., Kelleway, J. J., Mazumder, D., Kobayashi, T., & Wen, L. (2021). Incorporation of local
834 dissolved organic carbon into floodplain aquatic ecosystems. *Aquatic Ecology*, *55*(3), 779-790.
835 doi:10.1007/s10452-021-09860-7
- 836 Scavia, D., Bertani, I., Obenour, D. R., Turner, R. E., Forrest, D. R., & Katin, A. (2017). Ensemble
837 modeling informs hypoxia management in the northern Gulf of Mexico. *Proceedings of the National*
838 *Academy of Sciences*, *114*(33), 8823-8828. doi:doi:10.1073/pnas.1705293114
- 839 Southern Cross GeoScience. (2019). *Episodic estuarine hypoxia: resolving the geochemistry of coastal*
840 *floodplain blackwaters – Summary of project findings* (Southern Cross GeoScience Technical Report
841 No. 119). Retrieved from Southern Cross University, Lismore NSW 2480:
- 842 Steffe, A. S., Macbeth, W. G., & Murphy, J. J. (2007). Status of the recreational fisheries in two
843 Australian coastal estuaries following large fish-kill events. *Fisheries Research*, *85*(3), 258-269.
844 doi:<https://doi.org/10.1016/j.fishres.2007.02.003>
- 845 Tallis, H., Polasky, S., Hellmann, J., Springer, N. P., Biske, R., DeGeus, D., . . . Weaver, S. K. (2019). Five
846 financial incentives to revive the Gulf of Mexico dead zone and Mississippi basin soils. *Journal of*
847 *Environmental Management*, *233*, 30-38. doi:<https://doi.org/10.1016/j.jenvman.2018.11.140>
- 848 Townsend, S. A., Boland, K. T., & Wrigley, T. J. (1992). Factors contributing to a fish kill in the
849 Australian wet/dry tropics. *Water Research*, *26*(8), 1039-1044. doi:[https://doi.org/10.1016/0043-](https://doi.org/10.1016/0043-1354(92)90139-U)
850 [1354\(92\)90139-U](https://doi.org/10.1016/0043-1354(92)90139-U)
- 851 Townsend, S. A., & Edwards, C. A. (2003). A fish kill event, hypoxia and other limnological impacts
852 associated with early wet season flow into a lake on the Mary River floodplain, tropical northern
853 Australia. *Lakes & Reservoirs: Research & Management*, *8*(3-4), 169-176.
- 854 Tulau, M. J. (2011). Lands of the richest character: Agricultural drainage of backswamp wetlands on
855 the north coast of New South Wales, Australia: Development, conservation and policy change: An
856 environmental history.
- 857 Vaquer-Sunyer, R., & Duarte, C. M. (2008). Thresholds of hypoxia for marine biodiversity.
858 *Proceedings of the National Academy of Sciences*, *105*(40), 15452-15457.
859 doi:doi:10.1073/pnas.0803833105

- 860 Vithana, C. L., Sullivan, L. A., & Shepherd, T. (2019). Role of temperature on the development of
861 hypoxia in blackwater from grass. *Science of The Total Environment*, 667, 152-159.
862 doi:<https://doi.org/10.1016/j.scitotenv.2019.02.386>
- 863 Voß, K., Fernández, D., & Schäfer, R. B. (2015). Organic matter breakdown in streams in a region of
864 contrasting anthropogenic land use. *Science of The Total Environment*, 527-528, 179-184.
865 doi:10.1016/j.scitotenv.2015.04.071
- 866 Wallace, T. A., Ganf, G. G., & Brookes, J. D. (2008). A comparison of phosphorus and DOC leachates
867 from different types of leaf litter in an urban environment. *Freshwater Biology*, 53(9), 1902-1913.
868 doi:10.1111/j.1365-2427.2008.02006.x
- 869 Walsh, S., Copeland, C., & Westlake, M. (2004). Major fish kills in the northern rivers of NSW in 2001:
870 causes, impacts and responses. *NSW Department of Primary Industries. Fisheries final report*
871 *series(68)*, 55.
- 872 Watts, R. J., Kopf, R. K., McCasker, N., Howitt, J. A., Conallin, J., Wooden, I., & Baumgartner, L.
873 (2018). Adaptive Management of Environmental Flows: Using Irrigation Infrastructure to Deliver
874 Environmental Benefits During a Large Hypoxic Blackwater Event in the Southern Murray–Darling
875 Basin, Australia. *Environmental Management*, 61(3), 469-480. doi:10.1007/s00267-017-0941-1
- 876 Whitworth, K. L., & Baldwin, D. S. (2016). Improving our capacity to manage hypoxic blackwater
877 events in lowland rivers: the Blackwater risk assessment tool. *Ecological Modelling*, 320, 292-298.
- 878 Whitworth, K. L., Baldwin, D. S., & Kerr, J. L. (2012). Drought, floods and water quality: Drivers of a
879 severe hypoxic blackwater event in a major river system (the southern Murray–Darling Basin,
880 Australia). *Journal of Hydrology*, 450-451, 190-198.
881 doi:<https://doi.org/10.1016/j.jhydrol.2012.04.057>
- 882 Whitworth, K. L., Kerr, J. L., Mosley, L. M., Conallin, J., Hardwick, L., & Baldwin, D. S. (2013). Options
883 for managing hypoxic blackwater in river systems: case studies and framework. *Environmental*
884 *Management*, 52(4), 837-850.
- 885 Wong, V., Walsh, S., & Morris, S. (2018). Climate affects fish-kill events in subtropical estuaries of
886 eastern Australia. *Marine and Freshwater Research*, 69(11), 1641-1648.
- 887 Wong, V. N. L., Johnston, S. G., Burton, E. D., Bush, R. T., Sullivan, L. A., & Slavich, P. G. (2011).
888 Anthropogenic forcing of estuarine hypoxic events in sub-tropical catchments: Landscape drivers and
889 biogeochemical processes. *Science of The Total Environment*, 409(24), 5368-5375.

- 890 Wong, V. N. L., Johnston, S. G., Bush, R. T., Sullivan, L. A., Clay, C., Burton, E. D., & Slavich, P. G.
891 (2010). Spatial and temporal changes in estuarine water quality during a post-flood hypoxic event.
892 *Estuarine, Coastal and Shelf Science*, 87(1), 73-82. doi:<https://doi.org/10.1016/j.ecss.2009.12.015>
- 893 Xiong, S., Nilsson, C. (1997). Dynamics of leaf litter accumulation and its effects on riparian
894 vegetation: A review. . *The Botanical Review*, 69. doi:10.1071/mf17307
- 895 Zhang, M., Cheng, X., Geng, Q., Shi, Z., Luo, Y., & Xu, X. (2019). Leaf litter traits predominantly control
896 litter decomposition in streams worldwide. *Global Ecology and Biogeography*, 28(10), 1469-1486.