Chemical characterization of organic compounds involved in iodine-initiated new particle formation from coastal macro-algal emission

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Abstract

Iodine-initiated new particle formation (I-NPF) has long been recognized in coastal hotspot regions. However, no prior work has studied the exact chemical composition of organic compounds and their role in the coastal I-NPF. Here we present an important complementary study to the ongoing laboratory and field researches of iodine nucleation in coastal atmosphere. Oxidation and NPF experiments with vapor emissions from real-world coastal macroalgae were simulated in a bag reactor. On the basis of comprehensive mass spectrometry measurements, we reported for the first time a variety of volatile precursors and their oxidation products in gas and particle phases in such a highly complex system. Organic compounds overwhelmingly dominated over iodine in the new particle growth initiated by iodine species. The identity and transformation mechanisms of organic compounds were identified in this study to provide a more complete story of coastal NPF from low-tide macroalgal emission.

1. Introduction

Coastal new particle formation (NPF) may be driven by daytime low-tide emission of iodine species from macroalgae fully or partially exposed to the air. The phenomenon was reported in hotspot locations of west Europe, Australia and polar regions (Allan et al., 2015; Baccarini et al., 2020; Beck et al., 2021; Heard et al., 2006; McFiggans et al., 2010; O'Dowd et al., 2002; Sipilä et al., 2016; Whitehead et al., 2009). In the southeast coastline of China, we reported intense iodine-initiated NPF based on particle number size distribution and iodine measurements (Yu et al.,

32   2019).

33   To simulate iodine-initiated NPF (I-NPF) in controlled laboratory conditions, I_2 or CH_2I_2 vapor
34   was usually photolyzed in the presence of ozone to provide nucleation precursors (Burkholder et al.,
35   2004; He et al., 2021; Huang et al., 2022; Jimenez et al., 2003; Martín et al., 2020; Monahan et al.,
36   2012; O'Dowd et al., 2004; Saunders and Plane, 2005). Ashu-Ayem et al. (2012); McFiggans et al.
37   (2004); Monahan et al. (2012); Sellegri et al. (2005) and Sellegri et al. (2016) also investigated the
38   NPF from the vapors emitted by real-world macroalgal specimens or seawater in laboratory chamber
39   or apparatus. However, the focus of all above studies are emission rate, oxidation mechanisms or
40   nucleation pathways of iodine species. For example, positive correlations between particle
41   concentrations and I_2 or CH_2I_2 mixing ratios were usually observed (Burkholder et al., 2004; Jimenez
42   et al., 2003; Monahan et al., 2012; Sellegri et al., 2005). Kinetic studies in flow tube or CERN
43   CLOUD chamber proposed the clustering of iodine oxides (I_xO_y) or iodine oxoacids (HIO_3, HIO_2) as
44   nucleation mechanisms on the basis of photoionization TOF-MS (Martín et al., 2020), Api-TOF and
45   nitrate-Chemical Ionization Mass Spectrometer (CIMS) measurements (He et al., 2021). A recent
46   chamber study showed heterogeneous reaction between iodine oxide nanoparticle, meso-erythritol
47   (or glyoxal) and dimethylamine accelerated nanoparticle growth (Huang et al., 2022).

48   Until now, no prior work has investigated the exact chemical identity of organic compounds
49   (other than iodinated methane) and their role in I-NPF. The role of biogenic terpenes and
50   anthropogenic aromatics in continental NPF has been recognized for a long time (Donahue et al.,
51   2013). Their ozonolysis or photochemistry products have been investigated in depth by using
52   Electrospray Ionization Mass Spectrometry (ESI-MS) and more recently, CIMS (Ehn et al., 2014;
53   Faxon et al., 2018; Kundu et al., 2012; Kundu et al., 2017; Nguyen et al., 2010; Riva et al., 2017;
54   Wang et al., 2020; Yan et al., 2020). It is very likely that certain volatile organic compounds (VOCs)
55   emitted mutually with iodine or iodinated methane from coastal biota or biologically active sea
56   surface may also be involved in coastal I-NPF process and promote the growth of iodine particles.

57   To test this hypothesis, we conducted oxidation and NPF experiments with vapor emissions from
58   real-world coastal macroalgae in a bag reactor. A suite of mass spectrometric methods including
59   Inductively Coupled Plasma-MS (ICP-MS), Gas Chromatography-MS (GC-MS), iodide-CIMS and
60   ESI-orbitrap MS were applied to measure vapor precursors, gaseous products and particulate
61   products during the NPF process. Mass concentrations of total organic carbon (TOC) and total iodine
62   (TI) of new particles were compared to evaluate the relative importance of organics and iodine in
63   new particle growth. The identity and transformation mechanisms of organic compounds were
64   identified to provide a more complete story of coastal NPF from low-tide macroalgal emission. Our

study is thus complementary to prior laboratory and field studies of I-NPF, but has an emphasis on organics.

2. Experiments

2.1 Experimental apparatus and sample collection

Similar to Potential Aerosol Mass (PAM) Oxidation Flow Reactor, a bag reactor was designed to provide an oxidizing environment for simulating atmospheric oxidation processes of algae-emitted VOCs. The bag reactor was made from 75 μm-thick fluorinated ethylene propylene (FEP) Teflon (1.2 m×1.5 m, flat dimension). The volume of the bag at full inflation was determined experimentally to be about 200 L. The bag was suspended vertically (Figure 1) and kept in the dark or directly exposed to room light of fluorescent lamp. Because the purpose of this study is to qualitatively measure the oxidation products of algae-emitted VOCs, wall loss, production rate and other kinetic factors in the bag reactor were not evaluated. Fresh macroalgae (*Undaria pinnatifida*) was collected from intertidal zone at Xiangshan gulf of east China coast and stored at -10 °C until the experiments. 2 kg macroalgae was put in a 20 L Pyrex glass bottle that was filled with ~1 L natural seawater. The specimens was partially exposed to the air to simulate tidal emersion of macroalgae. A flow of particle-free ultra high purity (UHP) air blew algae-emitted VOCs out of the bottle and merged with a diluting air flow before entering the bag reactor.

Two types of experiments were conducted. In the three ozonolysis experiments, ozone (O_3) was generated by flowing an UHP air flow through a 5 Watts 185 nm UV lamp. The O_3 flow was fed just before the bag reactor was fully inflated. Final O_3 concentration in the bag reactor was measured to be ~200 ppbv using an ozone analyzer (Model 49i, Thermo-Fisher Scientific Inc.). In an additional OH-enhanced experiment, the O_3/VOC mixture flow was directed through a 254 nm UV light before entering the bag reactor. OH radicals were produced via the reaction $O_3+h\nu\rightarrow O_2+O(^1D)$ and $O(^1D)+H_2O\rightarrow 2OH$.

Before each experiment, the bag was purged for several hours to reduce background particle concentrations to below 1 cm^{-3}. The bag reactor was first operated in a static mode to monitor the time evolution of gaseous products and particle size, and then in a dynamic mode to collect enough particles for offline chemical analysis. In the static mode, the bag was first filled to full inflation with the VOCs/O_3 flows. The flows were then shut down; a Scanning Mobility Particle Sizer (SMPS, model 3936, TSI Inc., Shoreview, MN, USA) and an Aerodyne iodide-CIMS pulled two flows of 0.3 liters per minute (lpm) and 1.8 lpm out of the bag, respectively. The SMPS measured the particle

96 number size distribution from 14 to 600 nm.

97 In the dynamic mode, the VOCs/O_3 flow of 3 lpm was fed to the bag continuously, while the
98 SMPS and a vacuum pump (GAST Group Ltd.) pulled sample flows of 0.3 and 2.7 lpm, respectively,
99 out of the bag reactor. This resulted in an overall residential time of 67 minutes for the O_3/VOC
100 mixture in the fully inflated bag. The particles in the 2.7 lpm sample flow were collected onto a
101 Zefluor® PTFE membrane filter mounted in a filter inlet for gases and aerosols (FIGAERO) for
102 iodide-CIMS analysis, or alternatively, onto 47 mm diameter double quartz fiber filter pack mounted
103 in a filter holder for ESI-orbitrap MS, ICP-MS and TOC analysis. The front filter of the double filter
104 pack collected the particles, while the back filter placed downstream of the front filter was supposed
105 to adsorb the same amount of volatile species as the front filter.

2.2 Chemical analysis

107 Before the ozonolysis experiments, the algae-emitted VOCs in the bag reactor was collected by a
108 6-liter pre-evacuated stainless-steel canisters (Entech Instruments, Inc., Simi Valley, CA, USA) and
109 was analyzed using a quadrupole GC-MS system (model TH-300B, Wuhan Tianhong Instruments Co.
110 Ltd., Wuhan, China). The algae-emitted VOCs, as well as their gaseous and particulate products,
111 were also measured by the FIGAERO-iodide-CIMS. Iodide-adduct chemical ionization is well suited
112 for measuring oxygenated or acidic compounds with minimal fragmentation. More details of the
113 GC-MS and FIGAERO-iodide-CIMS measurements can be found in Supporting Material. The
114 theory and design of the two instruments were described by Wang et al. (2014) and Lopez-Hilfiker et
115 al. (2014).

116 The particles collected on quartz fiber filters were sent for offline quantification of TOC and TI,
117 as well as non-target analysis of organic compounds using ESI-orbitrap MS. The front and back
118 filters were treated, separately, following the procedure as below: the filter was ultrasonicated twice
119 with 10-mL water and acetone nitrile solvent mixture (v:v=1:1). The extract was filtered by a 0.2 μm
120 PTFE syringe filter and evaporated in a rotary evaporator to 0.5 mL. After being centrifuged for 30
121 min at 12000 rpm, the supernatant was collected for TI analysis by Agilent 1100 HPLC-7900
122 ICP-MS (Agilent Technologies, Santa Clara, CA, USA) and TOC analysis by a TOC analyzer
123 (Model TOC-5000A, Shimadzu, Japan). TI or TOC in the particles was obtained by subtracting the
124 amount on the back filter from that on the front filter. Nontarget analysis of organic compounds in
125 the supernatant was conducted using a Q Exactive hybrid Quadrupole-Orbitrap mass spectrometer
126 (Thermo Scientific, Bremen, Germany). The supernatant was directly infused by a syringe pump and
127 ionized in negative ESI source. All the ions in the m/z range from 50 to 500 Th were scanned with a

128 mass resolution of 70000. The chemically sound CHO molecular formulas were computed with a
129 mass tolerance of ±2 ppm for these ions. Only the compounds that existed solely in the front filter or
130 with ion intensity in the front filter higher than that in the back filter by a factor of 3 were regarded
131 as the organic compounds in the particle phase.

3. Results and discussion

3.1 Relative mass contribution of organic carbon and iodine to new particles

134 Typical banana-shape particle size spectrum observed in the static mode of an ozonolysis
135 experiment is shown in Figure 2a. In the presence of room light, new particles larger than 14 nm
136 were observed only 58 minutes after the injection of ozone flow. This relative long time is due to the
137 build-up of O_3 concentration and subsequent accumulation of oxidation products. No particles were
138 formed in the absence of room light or ozone. In the dynamic mode experiments, O_3 in the bag
139 reactor was kept at its maximum concentration 200 ppbv. With a prolonged residential time of 67
140 min, the particles grew to 102±23 nm, which was measured by the SMPS at the outlet of the bag
141 reactor. The TOC and TI measurements show that organic compounds contributed more particle
142 mass than iodine with TOC/(I+TOC) ratio of 96.1±2.9% (Table 1).

143 In the OH-enhanced experiment (dynamic mode), more particulate products were generated with
144 enhanced oxidation capacity: TI in the particles increased by a factor of 10.8; TOC increased by a
145 factor of 2.7; particle number concentration increased by a factor of 7.4. On the other hand, particle
146 size decreased to 73 nm and TOC/(TI+TOC) ratio decreased to 92.9% (Table 1). These differences
147 indicate that more iodine nuclei were produced with enhanced OH concentration, probably via
148 OIO+OH$\rightarrow$ $HOIO_2$ (Plane et al., 2006). Competitive uptake of condensing organic vapors onto these
149 iodine nuclei limited the growth of individual new particles. Nevertheless, organic compounds
150 overwhelmingly dominated over iodine in the mass contribution to new particle growth.

151 The significant organic contribution observed in the laboratory condition is generally consistent
152 with TOC/(I+TOC) ratio of 98.2% in 10-56 nm new particles collected during a coastal I-NPF event in
153 China (Yu et al., 2019), although TOC and TI during the field event are two orders of magnitude
154 lower than those in the bag reactor (Table 1). Mean diameter of new particles was observed to be
155 only 16 nm during the field event. But those small new particles are expected to grow into CCN
156 active sizes, given longer residence time and uptake of more condensing vapors in the atmosphere.

3.2 Macroalgal emission

It is of particular interest to know what VOCs are emitted from coastal macroalgae. They are potential precursors of iodine particle nucleation and growth. The canister sampling followed by GC-MS analysis showed that the top 9 non-CHO compounds with highest TIC peak areas (Table 2) are C_5 alkanes, C_{10} alpha-pinene and halogenated C_1, C_3 and C_5 alkanes. The top 10 CHO compounds are C_2-C_6 alcohols and carbonyls with saturated or unsaturated carbon chain.

Iodide-CIMS is more sensitive to more oxygenated or acidic compounds and thus complementary to the GC-MS measurement. The 76 organic precursors detected by iodide-CIMS before ozone addition were characterized by $C_{1,2,3,6}$ and O_{2-3} formulas (Figure 3a). The top 7 compounds with highest ion intensities were CH_2O_2, $C_2H_4O_2$, $C_3H_6O_3$, $C_6H_{10}O_3$, $C_2H_6O_2$, $C_4H_8O_2$ and $C_6H_{12}O_3$, which accounted for 82.5% of total ion intensity. They are C_1-C_6 mono-carboxylic acids, hydroxyl carboxylic acids or oxo-carboxylic acids with 2 to 3 oxygen atoms (Table 2). Their carbon atom numbers are in general consistent with the VOCs detected by GC-MS.

Relatively high signals of HNO_3 were observed as NO_3^- and HNO_3I^- before the addition of ozone to the bag reactor. Because HNO_3 and HNO_2 were also observed as deprotonated ions or I^- clusters in the particle phase during the NPF, HNO_3 is also an important precursor of particle formation.

3.3 Gaseous and particulate products

3.3.1 Inorganic molecules and radicals

Being different from nitrate-CIMS, our iodide-CIMS did not detect nucleating clusters of iodine oxides or oxyacids after the addition of ozone. Instead, dozens of new inorganic molecules or radicals were observed as clusters with I^-, NO_3^- or deprotonated ions in the gas or particle phase (Figure 4). We grouped these species by elemental composition and investigated their role in the NPF by observing how their gaseous ion intensities evolved during the NPF event in the bag reactor (Figure 2b-2f).

1. Cl, I, Cl_2 and ClI in the gas phase: the intensities of I and Cl increased ca. 10 minutes before 14 nm particles appeared and decreased as the particles grew up. Based on prior work of Burkholder et al. (2004); Jimenez et al. (2003); O'Dowd et al. (2004), we suggested the photolysis of CH_2Cl_2, CHBrCl, CH_3I and C_3H_7I was the source of halogen atoms (e.g., $CH_3I+h\nu \rightarrow CH_3+I$). There was a time lag of 20-25 minutes between the appearances of Cl and I and those of ClI and Cl_2, which were probably from the recombination of Cl and I atoms.

187   2. IO_2, IO, ClIO, INO_2 and $ClNO_2$ in the gas phase: these species showed a similar time series to I

188   and Cl atoms. IO, IO_2 and ClIO could be from the reactions between I, ClI and O_3. INO_2 is usually

189   thought to form upon the reaction $I+NO_2+M \rightarrow INO_2+M$ (Saiz-Lopez et al., 2012). $ClNO_2$ was likely

190   to form upon similar reaction between Cl and NO_2 in the bag reactor.

191   3. HIO_3 and INO_3: the two species seem to be the end products of above intermediates, because

192   their intensities kept on increasing during new particle growth. INO_3, which is iodine nitrate $IONO_2$,

193   was detected in both gas and particle phases. $IONO_2$ probably formed upon the recombination of IO

194   and NO_2 ($IO+NO_2+M \rightarrow IONO_2+M$) (Saiz-Lopez et al., 2012). HIO_3 was likely to form via

195    $OIO+OH \rightarrow HOIO_2$ or $I + H_2O + O_3 \rightarrow HOIO_2 + OH$ (Martín et al., 2020; Plane et al., 2006). HIO_3

196   was not detected in particle phase by iodide-CIMS, which is contrary to the observation by

197   HPLC-ICP-MS that total iodine was mostly dominated IO_3^- peak. The signals of IO^-, IO_2^- and

198    $HIONO_3^-$ in the particle phase are therefore most likely to result from thermal decomposition of

199   HIO_3 to HIO and HIO_2 in the FIGAERO thermal desorption process.

200   4. CH_3SO_3H, S_2^-, S_3^-, SO_3^-: We observed methane sulfonic acid (CH_3SO_3H, MSA) in both gas and

201   particle phases. Gaseous MSA increased in the beginning, but decreased after new particles appeared

202   (Figure 2f). Apparently, our measurement suggested MSA contributed to the growth of new particles,

203   but it is unknown if it also participated in nucleation. We suggested S_2^-, S_3^-, SO_3^- ions observed in the

204   particle phase were thermal decomposition products of MSA.

205   **3.3.2 Gaseous organic products**

206   After ozone addition, a gradual transformation from C_1-C_3 precursors to C_5-C_8 gaseous products

207   was observed during the NPF process (Figure 2h). In the meanwhile, the oxygen atom number of the

208   compounds increased from 2-3 to 4-7 (Figure 2g). The formation of compounds with more carbon

209   atoms than the parent VOCs is unlikely in the gas phase, except bimolecular reactions of stabilized

210   Criegee intermediates (SCIs) that typically form upon alkene ozonolysis. Similar to isoprene

211   ozonolysis (Inomata et al., 2014; Riva et al., 2017), we propose the SCI addition mechanism can

212   also explain the transformation observed in our case: (1) C_4 SCIs formed upon the ozonolysis of

213   CHO precursors with C=C double bonds (e.g., those observed by GC-MS in Table 2). (2) the

214   insertion of C_4 SCIs into carboxylic acid precursors (e.g., those observed by CIMS in Table 2)

215   produced oligomeric hydroperoxides. An example was shown in Scheme I for the reactions of most

216   abundant ethyl vinyl carbinol ($C_5H_{10}O$), ozone and formic acid (CH_2O_2), but the same mechanism is

217   also applicable for ethyl vinyl ketone (C_5H_8O) and other abundant C_2-C_5 carboxylic acids and

218   hydroxyl carboxylic acids. As a result, a series of gaseous oligomeric hydroperoxides $C_5H_{10}O_5$,

219 $C_6H_{10}O_5$, $C_6H_{12}O_5$, $C_7H_{12}O_6$, $C_7H_{14}O_6$, $C_8H_{14}O_5$, $C_8H_{16}O_6$, $C_8H_{16}O_5$ and $C_9H_{16}O_6$ were observed with

220 high intensity by iodide-CIMS.

221   Scheme I
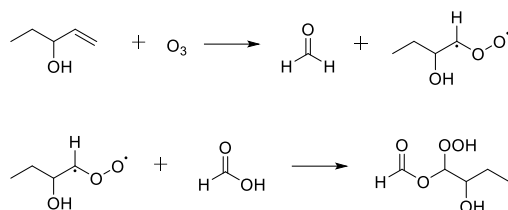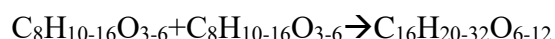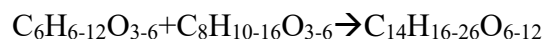
3.3.3 Particulate organic products

223   In the end of a typical ozonolysis experiment (dynamic mode), 100 and 364 new formulas were

224 observed in the gas and particle phases, respectively, including 73 semi-VOCs appeared in both gas

225 and particle phases. Those semi-VOCs accounted for 81 and 20% of total ion intensities of gaseous

226 and particulate products, respectively. Being different from unimodal atom number distributions of

227 gaseous products (C_{max}= 7 and O_{max}=5, Figure 3b), particulate products were characterized by

228 distinct bimodal or trimodal distribution of carbon number (C_{max}=8, 14 and 16, Figure 3c) and

229 oxygen number (O_{max}=4 and 8), implying possible dimer formation via accretion reactions in the

230 particle phase.

231   ESI-Orbitrap MS differs from FIGAERO-iodide-CIMS in extraction method (ultrasonic solvent

232 extraction from quartz fiber filter *vs.* thermal desorption from PTFE membrane filter), ionization

233 source (electrospray ionization *vs.* iodide-adduct chemical ionization) and MS resolving power

234 (70000 *vs.* 4500). The result showed that ESI-orbitrap MS and FIGAERO-iodide-CIMS detected 336

235 and 364 organic formulas, respectively, in the particle phase. 167 organic formulas were commonly

236 observed by both methods, which accounted for 87% and 54% of total ion intensity of organic

237 formulas by the two methods, respectively (Figure S1). As shown in Figure 3c and 3d,

238 FIGAERO-iodide-CIMS had better sensitivity toward the organic compounds with more oxygen

239 atoms (e.g., O≥8) and carbon atoms (e.g., C≥10). As a result, bimodal carbon and oxygen atom

240 number distributions were observed by FIGAERO-iodide-CIMS, but not ESI-orbitrap MS.

241   The measurement by ESI-orbitrap MS provided more insights about the formation mechanism of

242 particulate products. We compared the 336 formulas detected by ESI-orbitrap MS in our study with

243 the 414 formulas of isoprene ozonolysis SOA products (Nguyen et al., 2010) and 922 formulas of

244 alpha-pinene ozonolysis SOA products (Putman et al., 2012) measured by the ESI-orbitrap MS. It

245 was found that 72% of the formulas in this study can also be found in isoprene SOA, but only 39%

246 can be found in alpha-pinene SOA. This seems to imply that some similar alkene ozonolysis

247 reactions occurred in our system and isoprene ozonolysis.

248 For such a highly complex system full of various algae-emitted precursors, it is impossible to
249 simply propose a reaction mechanism to explain the formation of all particulate products, nor to list
250 all reactions occurring in the bag reactor. On the basis of particle-phase oligomer chemistry (Seinfeld
251 and Pandis, 2016), especially the well-understood isoprene ozonolysis SOA chemistry (Inomata et al.,
252 2014; Nguyen et al., 2010; Riva et al., 2017), we suggest a variety of accretion reactions without
253 uniform oligomerization pattern (e.g., esterification, aldol condensation, hemiacetal reactions,
254 peroxyhemiacetal formation and SCI reactions, etc.) transformed $O_{max}=4$ and $C_{max}=8$ multifunctional
255 monomers (like alcohols, carbonyls, hydroperoxides, carboxylic acids) to $O_{max}=8$ and $C_{max}=14$ or 16
256 dimers. Scheme II illustrated addition type self- and cross-oligomerization between C_6 and C_8
257 monomers produces C_{14} and C_{16} dimers. All the formulas in Scheme II are among the most abundant
258 ones observed in the particle phase by the iodide-CIMS.

259 $$C_6H_{6-12}O_{3-6}+C_8H_{10-16}O_{3-6}\rightarrow C_{14}H_{16-26}O_{6-12}$$

260 $$C_8H_{10-16}O_{3-6}+C_8H_{10-16}O_{3-6}\rightarrow C_{16}H_{20-32}O_{6-12}$$ Scheme II

261 **4. Conclusions**

262 Using a suite of mass spectrometers, we reported, for the first time, the chemical compositions of
263 volatile precursors emitted by real-world coastal macroalgae and their gaseous and particulate
264 oxidation products. In the presence of room light and ozone, the photolysis of halogenated $C_{1,3,5}$
265 alkanes ends up as HIO_3 and INO_3. It was most likely HIO_3 initiated NPF and provided nuclei for the
266 further condensation of other products like INO_3, MSA and CHO compounds. Gas-phase SCI
267 reactions and particle-phase accretion reactions transformed C_1-C_6 and O_2-O_3 precursors gradually to
268 particulate products with $C_{max}=8$, 14 and 16 and $O_{max}=4$ and 8. As a result, organic carbon were
269 found to overwhelmingly dominated over iodine in the mass contribution to the new particle growth.
270 Although our instruments did not allow the detection of nucleating clusters of iodine oxides or
271 oxyacids, our study provided important complementary information to the ongoing laboratory and
272 field researches of coastal I-NPF.

273 **Data Availability Statement**

274 All data related to figures and tables in this study are archived and made available through
275 Zenodo data repository https://doi.org/10.5281/zenodo.6965859.

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
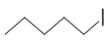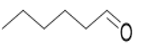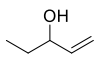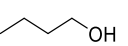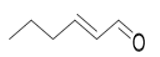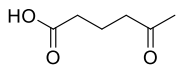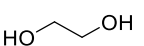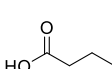Kundu, S., R. Fisseha, A. L. Putman, T. A. Rahn, and L. R. Mazzoleni (2017), Molecular formula composition of β-caryophyllene ozonolysis SOA formed in humid and dry conditions, *Atmospheric Environment*, *154*, 70-81, doi:https://doi.org/10.1016/j.atmosenv.2016.12.031.

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Table 1. Particle number concentration (N), mean diameter (D_p), total organic carbon (TOC) and total iodine (TI) of new particles with a residential time of 67 min in the bag reactor in the ozonolysis experiments and OH-enhanced experiment (dynamic mode). Those of 10-56 nm new particles collected by a nano Micro-Orifice Uniform Deposit Impactor (nano-MOUDI, MSP, Inc.) during an I-NPF event at a coastal site of Ningbo, China (Yu *et al.*, 2019) were also listed.

	TOC (μg m^{-3})	TI (μg m^{-3})	TOC/(TI+TOC)	N (cm^{-3})	D_p (nm)
ozonolysis experiments	45.6±9.7	0.88±0.34	96.1±2.9%	(5.58±2.04)×10^4	102±23
OH-enhanced experiment	125.3	9.5	92.9%	4.16×10^5	73
I-NPF event at a coastal site of China	0.7	0.0135	98.2%	6.00×10^5	16

Table 2. Major volatile organic compounds emitted by macroalgae as potential NPF precursors, sorted by TIC peak area measured by GC/MS or MS peak intensity measured by iodide-CIMS

|	Formula	Structure	Peak area/MS peak intensity
1	C_5H_{12}		1.90×10^6
2	C_5H_{10}		1.59×10^6
3	CH_3I		1.37×10^6
4	C_3H_7I		7.60×10^5
5	$CHBr_3$		4.71×10^5
6	$C_5H_{11}I$		3.75×10^5
7	$CHBr_2Cl$		2.71×10^5
8	CH_2Cl_2		2.55×10^5
9	$C_{10}H_{16}$		2.26×10^5
1	C_2H_6O		1.70×10^7
2	C_3H_6O		1.38×10^7
3	$C_4H_6O_2$		1.30×10^7
4	$C_6H_{12}O$		1.03×10^7
5	$C_5H_{10}O$		1.00×10^7
6	$C_4H_{10}O$		5.16×10^7
8	C_2H_4O		3.46×10^7
9	$C_6H_{10}O$		2.88×10^7
7	C_5H_8O		1.45×10^7
10	C_4H_8O		1.37×10^7
1	CH_2O_2		1.58×10^6
2	$C_2H_4O_2$		9.52×10^5
3	$C_3H_6O_3$		9.21×10^5
4	$C_6H_{10}O_3$		4.44×10^5
5	$C_2H_6O_2$		2.88×10^5
6	$C_4H_8O_2$		1.17×10^5

7    $C_6H_{12}O_3$
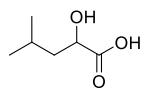
1.12×10^5
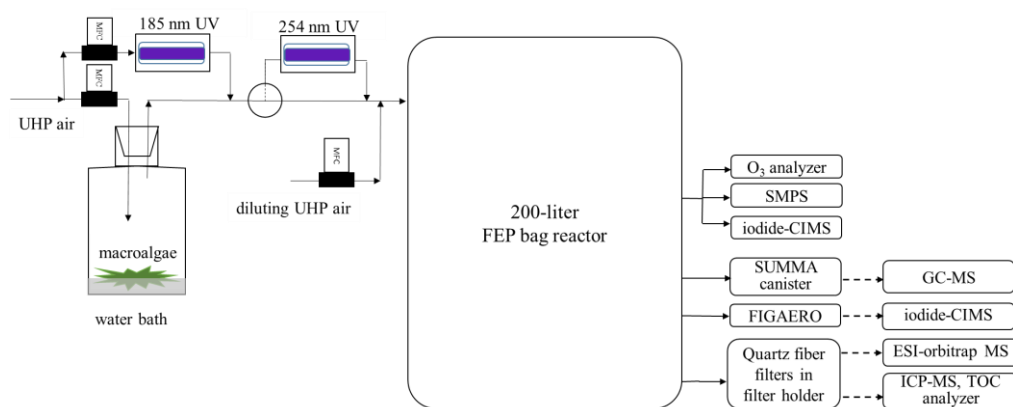
Figure 1. Schematic of experimental setup. Solid line: air flows. Dashed lines: sent for offline chemical analysis.
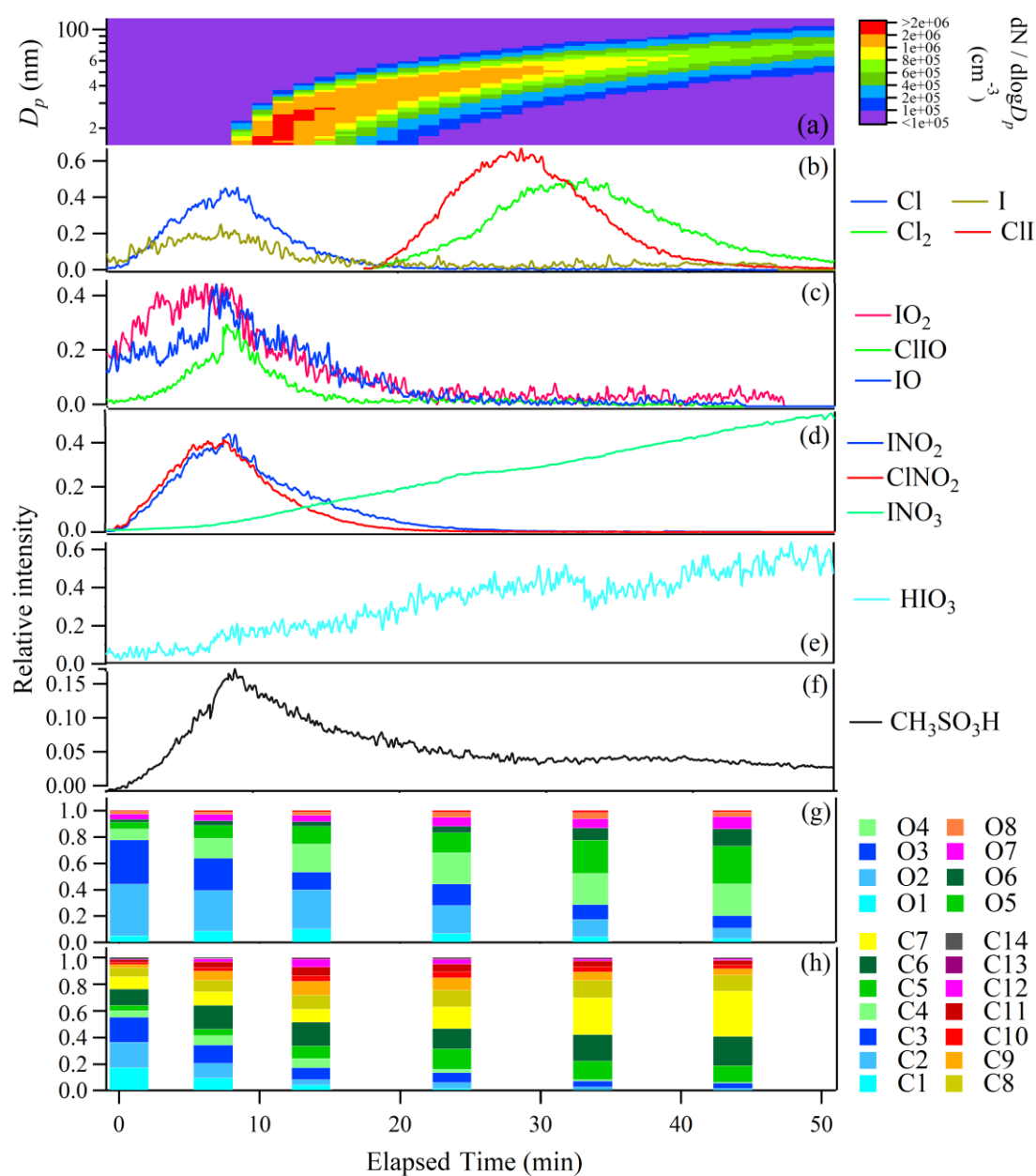
Figure 2. Time evolution of particle number size distribution (a) and relative intensities of gaseous molecules and radicals (b-f); the fractions of organic compounds grouped by O and C atom numbers in the selected time points (g-h) in a typical ozonolysis experiment (static mode). Time zero was chosen as the start time when HIO_3 was observed.
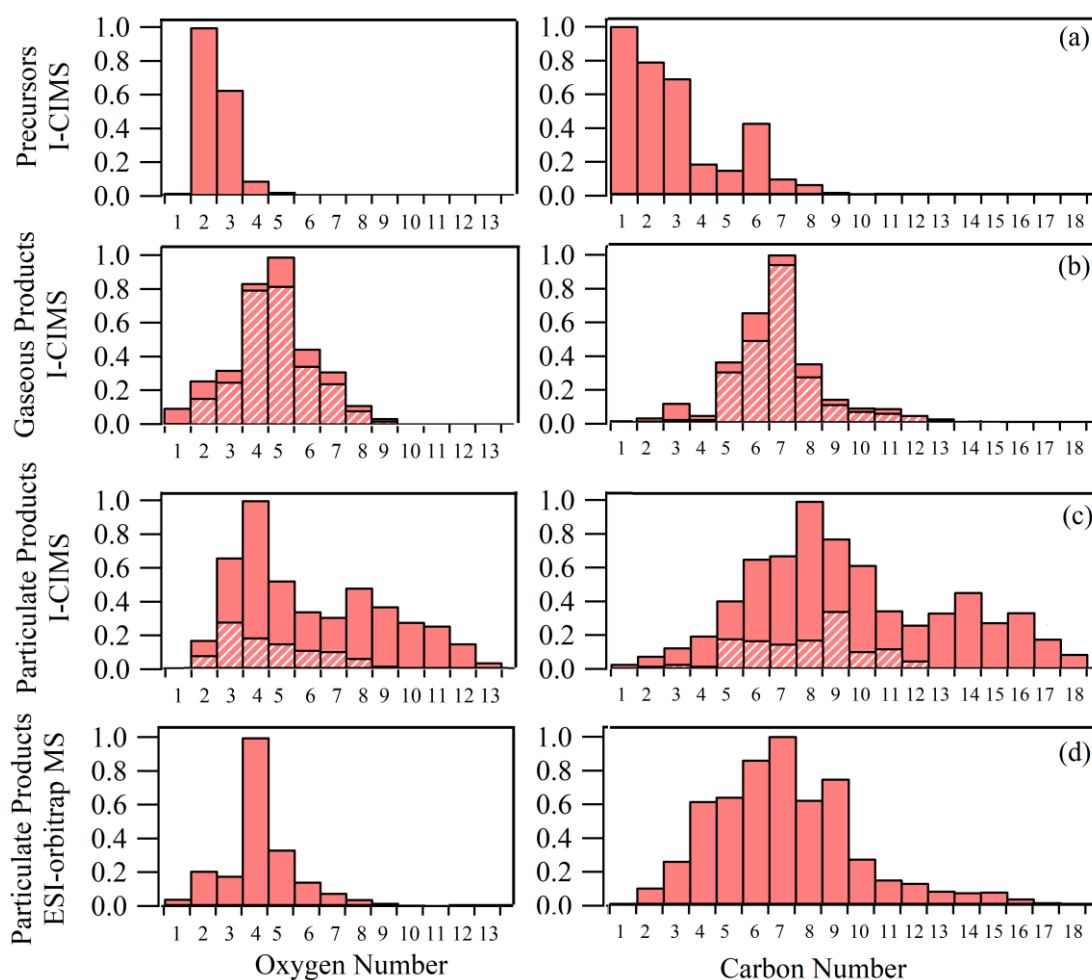
Figure 3. Oxygen and carbon atom number distributions of potential VOC precursors (a), gaseous products (b) and particulate products measured by iodide-CIMS (c), as well as the particulate products measured by ESI-orbitrap MS (d) in a typical ozonolysis experiment (dynamic mode). Hatched bars indicate the fractions of organic formulas observed in both gas and particle phases by iodide-CIMS.
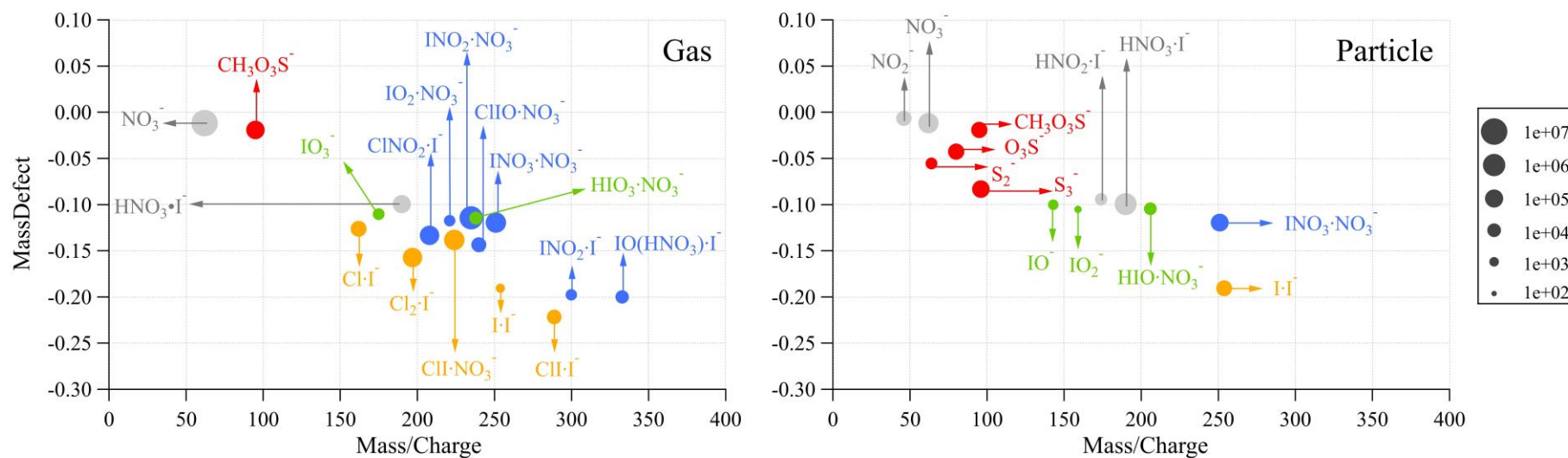
Figure 4. Integrated ion intensities of inorganic molecules and radicals in the gas phase (static mode) and particle phase (dynamic mode) measured by iodide-CIMS in a typical ozonolysis experiment. The ions were coded in color according to their elemental composition