

# Organic compounds from macroalgal emission dominate new particle growth initiated by iodine species in coastal atmosphere

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## Abstract

Iodine-initiated new particle formation (I-NPF) has long been recognized in coastal hotspot regions. However, no prior work has studied the exact chemical composition of organic compounds and their role in the coastal I-NPF. Here we present an important complementary study to the ongoing laboratory and field researches of iodine nucleation in coastal atmosphere. Oxidation and NPF experiments with vapor emissions from real-world coastal macroalgae were simulated in a bag reactor. On the basis of comprehensive mass spectrometry measurements, we reported for the first time a series of volatile precursors and their oxidation products in gas and particle phases in such a highly complex system. Organic compounds overwhelmingly dominated over iodine in the new particle growth initiated by iodine species. This study provided a more complete story of coastal NPF from low-tide macroalgal emission.

## 1. Introduction

Coastal new particle formation (NPF) may be driven by daytime low-tide emission of iodine species from macroalgae fully or partially exposed to the air. The phenomenon was reported in hotspot locations of west Europe, Australia and polar regions [Allan *et al.*, 2015; Baccharini *et al.*, 2020; Beck *et al.*, 2021; Heard *et al.*, 2006; McFiggans *et al.*, 2010; O'Dowd *et al.*, 2002; Sipilä *et al.*, 2016; Whitehead *et al.*, 2009]. In the southeast coastline of China, we reported intense iodine-initiated NPF based on particle number size distribution and iodine measurements [Yu *et al.*, 2019].

To simulate iodine-initiated NPF (I-NPF) in controlled laboratory conditions, I<sub>2</sub> or CH<sub>2</sub>I<sub>2</sub> vapor

32 was usually photolyzed in the presence of ozone to provide nucleation precursor [Burkholder et al.,  
33 2004; He et al., 2021; Huang et al., 2022; Jimenez et al., 2003; Martín et al., 2020; Monahan et al.,  
34 2012; O'Dowd et al., 2004; Saunders and Plane, 2005]. Ashu-Ayem et al. [2012]; McFiggans et al.  
35 [2004]; Monahan et al. [2012]; Sellegri et al. [2005] and Sellegri et al. [2016] also investigated the  
36 NPF from the vapors emitted by real-world macroalgal specimens or seawater in laboratory chamber  
37 or apparatus. However, the focus of all above studies are emission rate, oxidation mechanisms or  
38 nucleation pathways of iodine species. For example, positive correlations between particle  
39 concentrations and I<sub>2</sub> or CH<sub>2</sub>I<sub>2</sub> mixing ratios were usually observed [Burkholder et al., 2004; Jimenez  
40 et al., 2003; Monahan et al., 2012; Sellegri et al., 2005]. Kinetic studies in flow tube or CERN  
41 CLOUD chamber proposed the clustering of iodine oxides (I<sub>x</sub>O<sub>y</sub>) or iodine oxoacids (HIO<sub>3</sub>, HIO<sub>2</sub>) as  
42 nucleation mechanisms on the basis of photoionization TOF-MS [Martín et al., 2020], Api-TOF and  
43 nitrate-Chemical Ionization Mass Spectrometer (CIMS) measurements [He et al., 2021]. A recent  
44 chamber study showed heterogeneous reaction between iodine oxide nanoparticle, meso-erythritol  
45 (or glyoxal) and dimethylamine accelerated nanoparticle growth [Huang et al., 2022].

46 Until now, no prior work has investigated the exact chemical identity of organic compounds  
47 (other than iodinated methane) and their role in I-NPF. The role of biogenic terpenes and  
48 anthropogenic aromatics in continental NPF has been recognized for a long time [Donahue et al.,  
49 2013]. Their ozonolysis or photochemistry products have been investigated in depth by using  
50 Electrospray Ionization Mass Spectrometry (ESI-MS) and more recently CIMS [Ehn et al., 2014;  
51 Faxon et al., 2018; Kundu et al., 2012; Kundu et al., 2017; Nguyen et al., 2010; Riva et al., 2017;  
52 Wang et al., 2020; Yan et al., 2020]. It is very likely that certain volatile organic compounds (VOCs)  
53 emitted mutually with iodine or iodinated methane from coastal biota or biologically active sea  
54 surface may also be involved in coastal I-NPF process and promote the growth of iodine particles.

55 To test this hypothesis, we conducted oxidation and NPF experiments with vapor emissions from  
56 real-world coastal macroalgae in a bag reactor. A suite of mass spectrometric methods including  
57 Inductively Coupled Plasma-MS (ICP-MS), Gas Chromatography-MS (GC-MS), ESI-orbitrap MS  
58 and iodide-CIMS were applied to measure vapor precursors, gaseous products and particulate  
59 products during the NPF process. Mass concentrations of total organic carbon (TOC) and total iodine  
60 (TI) of new particles were compared to evaluate the relative importance of organics and iodine in  
61 new particle growth. The identity and transformation mechanisms of organic compounds were  
62 identified to provide a more complete story of coastal NPF from low-tide macroalgal emission. Our  
63 study is thus complementary to prior laboratory and field studies of I-NPF, but has an emphasis on  
64 organics.

## 65 2. Experiments

### 66 2.1 Experimental apparatus and sample collection

67 Similar to Potential Aerosol Mass (PAM) Oxidation Flow Reactor, a bag reactor was designed to  
68 provide an oxidizing environment for simulating atmospheric oxidation processes of algae-emitted  
69 VOCs. The bag reactor was made from 75  $\mu\text{m}$ -thick fluorinated ethylene propylene (FEP) Teflon  
70 (1.2 m $\times$ 1.5 m, flat dimension). The volume of the bag at full inflation was determined  
71 experimentally to be about 200 L. The bag was suspended vertically (Figure S1) and kept in the dark  
72 or directly exposed to room light of fluorescent lamp. Because the purpose of this study is to  
73 qualitatively measure the oxidation products of algae-emitted VOCs, wall loss, production rate and  
74 other kinetic factors in the bag reactor were not evaluated. Fresh macroalgae (*Undaria pinnatifida*)  
75 was collected from intertidal zone at Xiangshan gulf of east China coast and stored at -10  $^{\circ}\text{C}$  until the  
76 experiments. 2 kg macroalgae was put in a 20 L Pyrex glass bottle that was filled with  $\sim$ 1 L natural  
77 seawater. The specimens was partially exposed to the air to simulate tidal emersion of macroalgae. A  
78 flow of particle-free ultra high purity (UHP) air blew the algae-emitted VOCs out of the bottle and  
79 merged with a diluting air flow before entering the bag reactor.

80 Two types of experiments were conducted. In the three ozonolysis experiments, ozone ( $\text{O}_3$ ) was  
81 generated by flowing an UHP air flow through a 5 Watts 185 nm UV lamp. The  $\text{O}_3$  flow was fed just  
82 before the bag reactor was fully inflated. Final  $\text{O}_3$  concentration in the bag reactor was measured to  
83 be  $\sim$ 200 ppbv using an ozone analyzer (Model 49i, Thermo-Fisher Scientific Inc.). In an additional  
84 OH-enhanced experiment, the  $\text{O}_3/\text{VOC}$  mixture flow was directed through a 254 nm UV light before  
85 entering the bag reactor. OH radicals were produced via the reaction  $\text{O}_3+h\nu\rightarrow\text{O}_2+\text{O}(^1\text{D})$  and  
86  $\text{O}(^1\text{D})+\text{H}_2\text{O}\rightarrow 2\text{OH}$ .

87 Before each experiment, the bag was purged for several hours to reduce background particle  
88 concentrations to below  $1\text{ cm}^{-3}$ . The bag reactor was first operated in a static mode to monitor the  
89 time evolution of gaseous products and particle size, and then in a dynamic mode to collect enough  
90 particles for offline chemical analysis. In the static mode, the bag was first filled to full inflation with  
91 the VOCs/ $\text{O}_3$  flows. The flows were then shut down; a Scanning Mobility Particle Sizer (SMPS,  
92 model 3936, TSI Inc., Shoreview, MN, USA) and an Aerodyne iodide-CIMS pulled two flows of 0.3  
93 liters per minute (lpm) and 1.8 lpm out of the bag, respectively. The SMPS measured the particle  
94 number size distribution from 14 to 600 nm.

95 In the dynamic mode, the VOCs/ $\text{O}_3$  flow of 3 lpm was fed to the bag continuously, while the

96 SMPS and a vacuum pump (GAST Group Ltd.) pulled sample flows of 0.3 and 2.7 lpm, respectively,  
97 out of the bag reactor. This resulted in an overall residential time of 67 minutes for the O<sub>3</sub>/VOC  
98 mixture in the fully inflated bag. The particles in the 2.7 lpm sample flow were collected onto a  
99 Zefluor® PTFE membrane filter mounted in a filter inlet for gases and aerosols (FIGAERO) for  
100 iodide-CIMS analysis, or alternatively, onto 47 mm diameter double quartz fiber filter pack mounted  
101 in a filter holder for ESI-orbitrap MS, ICP-MS and TOC analysis. The front filter of the double filter  
102 pack collected the particles, while the back filter placed downstream of the front filter was supposed  
103 to adsorb the same amount of volatile species as the front filter.

## 104 2.2 Chemical analysis

105 Before the ozonolysis experiments, the algae-emitted VOCs in the bag reactor was collected by a  
106 6-liter pre-evacuated stainless-steel canisters (Entech Instruments, Inc., Simi Valley, CA, USA) and  
107 was analyzed using a quadrupole GC-MS system (model TH-300B, Wuhan Tianhong Instruments Co.  
108 Ltd., Wuhan, China). The algae-emitted VOCs, as well as their gaseous and particulate products,  
109 were also measured by the FIGAERO-iodide-CIMS. Iodide-adduct chemical ionization is well suited  
110 for measuring oxygenated or acidic compounds with minimal fragmentation. More details of the  
111 GC-MS and FIGAERO-iodide-CIMS measurements can be found in Supporting Material. The  
112 theory and design of the two instruments were described by *Wang et al. [2014]* and *Lopez-Hilfiker et*  
113 *al. [2014]*.

114 The particles collected on quartz fiber filters were sent for offline quantification of TOC and TI,  
115 as well as non-target analysis of organic compounds using ESI-orbitrap MS. The front and back  
116 filters were treated, separately, following the procedure as below: the filter was ultrasonicated twice  
117 with 10-mL water and acetone nitrile solvent mixture (v:v=1:1). The extract was filtered by a 0.2 µm  
118 PTFE syringe filter and evaporated in a rotary evaporator to 0.5 mL. After being centrifuged for 30  
119 min at 12000 rpm, the supernatant was collected for TI analysis by Agilent 1100 HPLC-7900  
120 ICP-MS (Agilent Technologies, Santa Clara, CA, USA) and TOC analysis by a TOC analyzer  
121 (Model TOC-5000A, Shimadzu, Japan). TI or TOC in the particles was obtained by subtracting the  
122 amount on the back filter from that on the front filter. Nontarget analysis of organic compounds in  
123 the supernatant was conducted using a Q Exactive hybrid Quadrupole-Orbitrap mass spectrometer  
124 (Thermo Scientific, Bremen, Germany). The supernatant was directly infused by a syringe pump and  
125 ionized in negative ESI source. All the ions in the m/z range from 50 to 500 Th were scanned with a  
126 mass resolution of 70000. The chemically sound CHO molecular formulas were computed with a  
127 mass tolerance of ±2 ppm for these ions. Only the compounds that existed solely in the front filter or

128 with ion intensity in the front filter higher than that in the back filter by a factor of 3 were regarded  
129 as the organic compounds in the particle phase.

### 130 **3. Results and discussion**

#### 131 **3.1 Relative mass contribution of organic carbon and iodine to new particles**

132 Typical banana-shape particle size spectrum observed in the static mode of an ozonolysis  
133 experiment is shown in Figure 1a. In the presence of room light, new particles larger than 14 nm  
134 were observed only 58 minutes after the injection of ozone flow. This relative long time is due to the  
135 build-up of O<sub>3</sub> concentration and subsequent accumulation of oxidation products. No particles were  
136 formed in the absence of room light or ozone. In the dynamic mode experiments, O<sub>3</sub> in the bag  
137 reactor was kept at its maximum concentration 200 ppbv. With a prolonged residential time of 67  
138 min, the particles grew to 102±23 nm, which was measured by the SMPS at the outlet of the bag  
139 reactor. The TOC and TI measurements show that organic compounds contributed more particle  
140 mass than iodine with TOC/(I+TOC) ratio of 96.1±2.9% (Table 1).

141 In the OH-enhanced experiment (dynamic mode), more particulate products were generated with  
142 enhanced oxidation capacity: TI in the particles increased by a factor of 10.8; TOC increased by a  
143 factor of 2.7; particle number concentration increased by a factor of 7.4. On the other hand, particle  
144 size decreased to 73 nm and TOC/(TI+TOC) ratio decreased to 92.9% (Table 1). These differences  
145 indicate that more iodine nuclei were produced with enhanced OH concentration, probably via  
146  $\text{OIO} + \text{OH} \rightarrow \text{HOIO}_2$  [Plane et al., 2006]. Competitive uptake of condensing organic vapors onto these  
147 iodine nuclei limited the growth of individual new particles. Nevertheless, organic compounds  
148 overwhelmingly dominated over iodine in the mass contribution to new particle growth.

149 The significant organic contribution observed in the laboratory condition is generally consistent  
150 with TOC/(I+TOC) ratio of 98.2% in 10-56 nm new particles collected during a coastal I-NPF event in  
151 China [Yu et al., 2019], although TOC and TI during the field event are two orders of magnitude  
152 lower than those in the bag reactor (Table 1). Mean diameter of new particles was observed to be  
153 only 16 nm during the field event. But those small new particles are expected to grow into CCN  
154 active sizes, given longer residence time and uptake of more condensing vapors in the atmosphere.

#### 155 **3.2 Macroalgal emission**

156 It is of particular interest to know what VOCs are emitted from coastal macroalgae. They are  
157 potential precursors of iodine particle nucleation and growth. The canister sampling followed by

158 GC-MS analysis showed that the top 9 non-CHO compounds with highest TIC peak areas (Table 2)  
159 are C<sub>5</sub> alkanes, C<sub>10</sub> alpha-pinene and halogenated C<sub>1</sub>, C<sub>3</sub> and C<sub>5</sub> alkanes. The top 10 CHO  
160 compounds are C<sub>2</sub>-C<sub>6</sub> alcohols and carbonyls with saturated or unsaturated carbon chain.

161 Iodide-CIMS is more sensitive to more oxygenated or acidic compounds and thus complementary  
162 to the GC-MS measurement. The 76 organic precursors detected by iodide-CIMS before ozone  
163 addition were characterized by C<sub>1,2,3,6</sub> and O<sub>2-3</sub> formulas (Figure 2a). The top 7 compounds with  
164 highest ion intensities were CH<sub>2</sub>O<sub>2</sub>, C<sub>2</sub>H<sub>4</sub>O<sub>2</sub>, C<sub>3</sub>H<sub>6</sub>O<sub>3</sub>, C<sub>6</sub>H<sub>10</sub>O<sub>3</sub>, C<sub>2</sub>H<sub>6</sub>O<sub>2</sub>, C<sub>4</sub>H<sub>8</sub>O<sub>2</sub> and C<sub>6</sub>H<sub>12</sub>O<sub>3</sub>,  
165 which accounted for 82.5% of total ion intensity. They are C<sub>1</sub>-C<sub>6</sub> mono-carboxylic acids, hydroxyl  
166 carboxylic acids or oxo-carboxylic acids with 2 to 3 oxygen atoms (Table 2). Their carbon atom  
167 numbers are in general consistent with the VOCs detected by GC-MS.

168 Relatively high signals of HNO<sub>3</sub> were observed as NO<sub>3</sub><sup>-</sup> and HNO<sub>3</sub>I<sup>-</sup> before the addition of ozone  
169 to the bag reactor. Because HNO<sub>3</sub> and HNO<sub>2</sub> were also observed as deprotonated ions or I<sup>-</sup> clusters in  
170 the particle phase during the NPF, HNO<sub>3</sub> is also an important precursor of particle formation.

### 171 3.3 Gaseous and particulate products

#### 172 3.3.1 Inorganic molecules and radicals

173 Being different from nitrate-CIMS, our iodide-CIMS did not detect nucleating clusters of iodine  
174 oxides or oxyacids after the addition of ozone. Instead, dozens of new inorganic molecules or  
175 radicals were observed as clusters with I<sup>-</sup>, NO<sub>3</sub><sup>-</sup> or deprotonated ions in the gas or particle phase  
176 (Figure S2). We grouped these species by elemental composition and investigated their role in the  
177 NPF by observing how their gaseous ion intensities evolved during the NPF event in the bag reactor  
178 (Figure 1b-1e).

179 1. Cl, I, Cl<sub>2</sub> and ClI in the gas phase: the intensities of I and Cl increased ca. 10 minutes before 14  
180 nm particles appeared and decreased as the particles grew up. Based on prior work of *Burkholder et al.*  
181 *[2004]*; *Jimenez et al. [2003]*; *O'Dowd et al. [2004]*, we suggested the photolysis of CH<sub>2</sub>Cl<sub>2</sub>,  
182 CHBrCl, CH<sub>3</sub>I and C<sub>3</sub>H<sub>7</sub>I was the source of halogen atoms (e.g., CH<sub>3</sub>I+hν→CH<sub>3</sub>+I). There was a  
183 time lag of 20-25 minutes between the appearances of Cl and I and those of ClI and Cl<sub>2</sub>, which were  
184 probably from the recombination of Cl and I atoms.

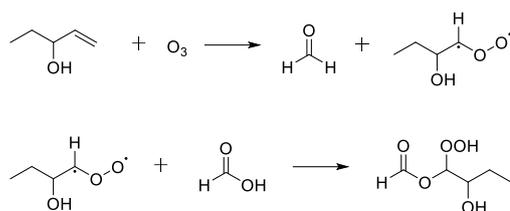
185 2. IO<sub>2</sub>, IO, ClIO, INO<sub>2</sub> and ClNO<sub>2</sub> in the gas phase: these species showed a similar time series to I  
186 and Cl atoms. IO, IO<sub>2</sub> and ClIO could be from the reactions between I, ClI and O<sub>3</sub>. INO<sub>2</sub> is usually  
187 thought to form upon the reaction I+NO<sub>2</sub>+M →INO<sub>2</sub>+M [*Saiz-Lopez et al., 2012*]. ClNO<sub>2</sub> was likely  
188 to form upon similar reaction between Cl and NO<sub>2</sub> in the bag reactor.

189 3. HIO<sub>3</sub> and INO<sub>3</sub>: the two species seem to be the end products of above intermediates, because  
190 their intensities kept on increasing during new particle growth. INO<sub>3</sub>, which is iodine nitrate IONO<sub>2</sub>,  
191 was detected in both gas and particle phases. IONO<sub>2</sub> probably formed upon the recombination of IO  
192 and NO<sub>2</sub> (IO+NO<sub>2</sub>+M→IONO<sub>2</sub>+M) [Saiz-Lopez *et al.*, 2012]. HIO<sub>3</sub> was likely to form via  
193 OIO+OH→HOIO<sub>2</sub> or I + H<sub>2</sub>O + O<sub>3</sub>→HOIO<sub>2</sub> + OH [Martín *et al.*, 2020; Plane *et al.*, 2006]. HIO<sub>3</sub>  
194 was not detected in particle phase by iodide-CIMS, which is contrary to the observation by  
195 HPLC-ICP-MS that total iodine was mostly dominated IO<sub>3</sub><sup>-</sup> peak. The signals of IO<sup>-</sup>, IO<sub>2</sub><sup>-</sup> and  
196 HIONO<sub>3</sub><sup>-</sup> in the particle phase are therefore most likely to result from thermal decomposition of  
197 HIO<sub>3</sub> to HIO and HIO<sub>2</sub> in the FIGAERO thermal desorption process.

198 4. CH<sub>3</sub>SO<sub>3</sub>H, S<sub>2</sub><sup>-</sup>, S<sub>3</sub><sup>-</sup>, SO<sub>3</sub><sup>-</sup>: We observed methane sulfonic acid (CH<sub>3</sub>SO<sub>3</sub>H, MSA) in both gas and  
199 particle phases. Gaseous MSA increased in the beginning, but decreased after new particles appeared  
200 (Figure 1f). Apparently, our measurement suggested MSA contributed to the growth of new particles,  
201 but it is unknown if it also participated in nucleation. We suggested S<sub>2</sub><sup>-</sup>, S<sub>3</sub><sup>-</sup>, SO<sub>3</sub><sup>-</sup> ions observed in the  
202 particle phase were thermal decomposition products of MSA.

### 203 3.3.2 Gaseous organic products

204 After ozone addition, a gradual transformation from C<sub>1</sub>-C<sub>3</sub> precursors to C<sub>5</sub>-C<sub>8</sub> gaseous products  
205 was observed during the NPF process (Figure 1h). In the meanwhile, the oxygen atom number of the  
206 compounds increased from 2-3 to 4-7 (Figure 1g). The formation of compounds with more carbon  
207 atoms than the parent VOCs is unlikely in the gas phase, except bimolecular reactions of stabilized  
208 Criegee intermediates (SCIs) that typically form upon alkene ozonolysis. Similar to isoprene  
209 ozonolysis [Inomata *et al.*, 2014; Riva *et al.*, 2017], we propose the SCI addition mechanism can  
210 also explain the transformation observed in our case: (1) C<sub>4</sub> SCIs formed upon the ozonolysis of  
211 CHO precursors with C=C double bonds (e.g., those observed by GC-MS in Table 2). (2) the  
212 insertion of C<sub>4</sub> SCIs into carboxylic acid precursors (e.g., those observed by CIMS in Table 2)  
213 produced oligomeric hydroperoxides. An example was shown in Scheme I for the reactions of most  
214 abundant ethyl vinyl carbinol (C<sub>5</sub>H<sub>10</sub>O), ozone and formic acid (CH<sub>2</sub>O<sub>2</sub>), but the same mechanism is  
215 also applicable for ethyl vinyl ketone (C<sub>5</sub>H<sub>8</sub>O) and other abundant C<sub>2</sub>-C<sub>5</sub> carboxylic acids and  
216 hydroxyl carboxylic acids. As a result, a series of gaseous oligomeric hydroperoxides C<sub>5</sub>H<sub>10</sub>O<sub>5</sub>,  
217 C<sub>6</sub>H<sub>10</sub>O<sub>5</sub>, C<sub>6</sub>H<sub>12</sub>O<sub>5</sub>, C<sub>7</sub>H<sub>12</sub>O<sub>6</sub>, C<sub>7</sub>H<sub>14</sub>O<sub>6</sub>, C<sub>8</sub>H<sub>14</sub>O<sub>5</sub>, C<sub>8</sub>H<sub>16</sub>O<sub>6</sub>, C<sub>8</sub>H<sub>16</sub>O<sub>5</sub> and C<sub>9</sub>H<sub>16</sub>O<sub>6</sub> were observed with  
218 high intensity by iodide-CIMS.



Scheme I

219

### 220 3.3.3 Particulate organic products

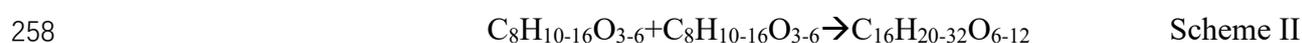
221 In the end of a typical ozonolysis experiment (dynamic mode), 100 and 364 new formulas were  
 222 observed in the gas and particle phases, respectively, including 73 semi-VOCs appeared in both gas  
 223 and particle phases. Those semi-VOCs accounted for 81 and 20% of total ion intensities of gaseous  
 224 and particulate products, respectively. Being different from unimodal atom number distributions of  
 225 gaseous products ( $C_{\max}=7$  and  $O_{\max}=5$ , Figure 2b), particulate products were characterized by  
 226 distinct bimodal or trimodal distribution of carbon number ( $C_{\max}=8, 14$  and  $16$ , Figure 2c) and  
 227 oxygen number ( $O_{\max}=4$  and  $8$ ), implying possible dimer formation via accretion reactions in the  
 228 particle phase.

229 ESI-Orbitrap MS differs from FIGAERO-iodide-CIMS in extraction method (ultrasonic solvent  
 230 extraction from quartz fiber filter *vs.* thermal desorption from PTFE membrane filter), ionization  
 231 source (electrospray ionization *vs.* iodide-adduct chemical ionization) and MS resolving power  
 232 (70000 *vs.* 4500). The result showed that ESI-orbitrap MS and FIGAERO-iodide-CIMS detected 336  
 233 and 364 organic formulas, respectively, in the particle phase. 167 organic formulas were commonly  
 234 observed by both methods, which accounted for 87% and 54% of total ion intensity of organic  
 235 formulas by the two methods, respectively (Figure S3). As shown in Figure 2c and 2d,  
 236 FIGAERO-iodide-CIMS had better sensitivity toward the organic compounds with more oxygen  
 237 atoms (e.g.,  $O \geq 8$ ) and carbon atoms (e.g.,  $C \geq 10$ ). As a result, bimodal carbon and oxygen atom  
 238 number distributions were observed by FIGAERO-iodide-CIMS, but not ESI-orbitrap MS.

239 The measurement by ESI-orbitrap MS provided more insights about the formation mechanism of  
 240 particulate products. We compared the 336 formulas detected by ESI-orbitrap MS in our study with  
 241 the 414 formulas of isoprene ozonolysis SOA products [Nguyen *et al.*, 2010] and 922 formulas of  
 242 alpha-pinene ozonolysis SOA products [Putman *et al.*, 2012] measured by the ESI-orbitrap MS. It  
 243 was found that 72% of the formulas in this study can also be found in isoprene SOA, but only 39%  
 244 can be found in alpha-pinene SOA. This seems to imply that some similar alkene ozonolysis  
 245 reactions occurred in our system and isoprene ozonolysis.

246 For such a highly complex system full of various algae-emitted precursors, it is impossible to

247 simply propose a reaction mechanism to explain the formation of all particulate products, nor to list  
248 all reactions occurring in the bag reactor. On the basis of particle-phase oligomer chemistry [*Seinfeld*  
249 *and Pandis, 2016*], especially the well-understood isoprene ozonolysis SOA chemistry [*Inomata et*  
250 *al., 2014; Nguyen et al., 2010; Riva et al., 2017*], we suggest a variety of accretion reactions without  
251 uniform oligomerization pattern (e.g., esterification, aldol condensation, hemiacetal reactions,  
252 peroxyhemiacetal formation and SCI reactions, etc.) transformed  $O_{\max}=4$  and  $C_{\max}=8$  multifunctional  
253 monomers (like alcohols, carbonyls, hydroperoxides, carboxylic acids) to  $O_{\max}=8$  and  $C_{\max}=14$  or 16  
254 dimers. Scheme II illustrated addition type self- and cross-oligomerization between  $C_6$  and  $C_8$   
255 monomers produces  $C_{14}$  and  $C_{16}$  dimers. All the formulas in Scheme II are among the most abundant  
256 ones observed in the particle phase by the iodide-CIMS.



#### 259 **4. Conclusions**

260 Using a suite of mass spectrometers, we reported, for the first time, the chemical compositions of  
261 volatile precursors emitted by real-world coastal macroalgae and their gaseous and particulate  
262 oxidation products. In the presence of room light and ozone, the photolysis of halogenated  $C_{1,3,5}$   
263 alkanes ends up as  $HIO_3$  and  $INO_3$ . It was most likely  $HIO_3$  initiated NPF and provided nuclei for the  
264 further condensation of other products like  $INO_3$ , MSA and CHO compounds. Gas-phase SCI  
265 reactions and particle-phase accretion reactions transformed  $C_1$ - $C_6$  and  $O_2$ - $O_3$  precursors gradually to  
266 particulate products with  $C_{\max}=8, 14$  and 16 and  $O_{\max}=4$  and 8. As a result, organic carbon were  
267 found to overwhelmingly dominated over iodine in the mass contribution to the new particle growth.  
268 Although our instruments did not allow the detection of nucleating clusters of iodine oxides or  
269 oxyacids, our study provided important complementary information to the ongoing laboratory and  
270 field researches of coastal I-NPF.

#### 271 **Data Availability Statement**

272 All data related to figures and tables in this study are archived and made available through  
273 Zenodo data repository <https://doi.org/10.5281/zenodo.6965859>.

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276 42175131) and Start-up research funding from China University of Geosciences.

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279 during ACCACIA, *Atmos. Chem. Phys.*, *15*(10), 5599-5609, doi:10.5194/acp-15-5599-2015.
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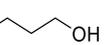
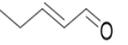
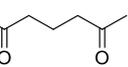
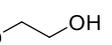
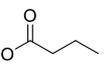
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Table 1. Particle number concentration ( $N$ ), mean diameter ( $D_p$ ), total organic carbon (TOC) and total iodine (TI) of new particles with a residential time of 67 min in the bag reactor in the ozonolysis experiments and OH-enhanced experiment (dynamic mode). Those of 10-56 nm new particles collected by a nano Micro-Orifice Uniform Deposit Impactor (nano-MOUDI, MSP, Inc.) during an I-NPF event at a coastal site of Ningbo, China (Yu *et al.*, 2019) were also listed.

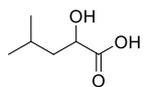
	TOC ( $\mu\text{g m}^{-3}$ )	TI ( $\mu\text{g m}^{-3}$ )	TOC/(TI+TOC)	$N$ ( $\text{cm}^{-3}$ )	$D_p$ (nm)
ozonolysis experiments	45.6±9.7	0.88±0.34	96.1±2.9%	(5.58±2.04)×10 <sup>4</sup>	102±23
OH-enhanced experiment	125.3	9.5	92.9%	4.16×10 <sup>5</sup>	73
I-NPF event at a coastal site of China	0.7	0.0135	98.2%	6.00×10 <sup>5</sup>	16

Table 2. Major volatile organic compounds emitted by macroalgae as potential NPF precursors, sorted by TIC peak area measured by GC/MS or MS peak intensity measured by iodide-CIMS

	Formula	Structure	Peak area/MS peak intensity
1	C <sub>5</sub> H <sub>12</sub>		1.90×10 <sup>6</sup>
2	C <sub>5</sub> H <sub>10</sub>		1.59×10 <sup>6</sup>
3	CH <sub>3</sub> I		1.37×10 <sup>6</sup>
4	C <sub>3</sub> H <sub>7</sub> I		7.60×10 <sup>5</sup>
5	CHBr <sub>3</sub>		4.71×10 <sup>5</sup>
6	C <sub>5</sub> H <sub>11</sub> I		3.75×10 <sup>5</sup>
7	CHBr <sub>2</sub> Cl		2.71×10 <sup>5</sup>
8	CH <sub>2</sub> Cl <sub>2</sub>		2.55×10 <sup>5</sup>
9	C <sub>10</sub> H <sub>16</sub>		2.26×10 <sup>5</sup>
1	C <sub>2</sub> H <sub>6</sub> O		1.70×10 <sup>7</sup>
2	C <sub>3</sub> H <sub>6</sub> O		1.38×10 <sup>7</sup>
3	C <sub>4</sub> H <sub>6</sub> O <sub>2</sub>		1.30×10 <sup>7</sup>
4	C <sub>6</sub> H <sub>12</sub> O		1.03×10 <sup>7</sup>
5	C <sub>5</sub> H <sub>10</sub> O		1.00×10 <sup>7</sup>
6	C <sub>4</sub> H <sub>10</sub> O		5.16×10 <sup>7</sup>
8	C <sub>2</sub> H <sub>4</sub> O		3.46×10 <sup>7</sup>
9	C <sub>6</sub> H <sub>10</sub> O		2.88×10 <sup>7</sup>
7	C <sub>5</sub> H <sub>8</sub> O		1.45×10 <sup>7</sup>
10	C <sub>4</sub> H <sub>8</sub> O		1.37×10 <sup>7</sup>
1	CH <sub>2</sub> O <sub>2</sub>		1.58×10 <sup>6</sup>
2	C <sub>2</sub> H <sub>4</sub> O <sub>2</sub>		9.52×10 <sup>5</sup>
3	C <sub>3</sub> H <sub>6</sub> O <sub>3</sub>		9.21×10 <sup>5</sup>
4	C <sub>6</sub> H <sub>10</sub> O <sub>3</sub>		4.44×10 <sup>5</sup>
5	C <sub>2</sub> H <sub>6</sub> O <sub>2</sub>		2.88×10 <sup>5</sup>
6	C <sub>4</sub> H <sub>8</sub> O <sub>2</sub>		1.17×10 <sup>5</sup>

7

$C_6H_{12}O_3$



$1.12 \times 10^5$

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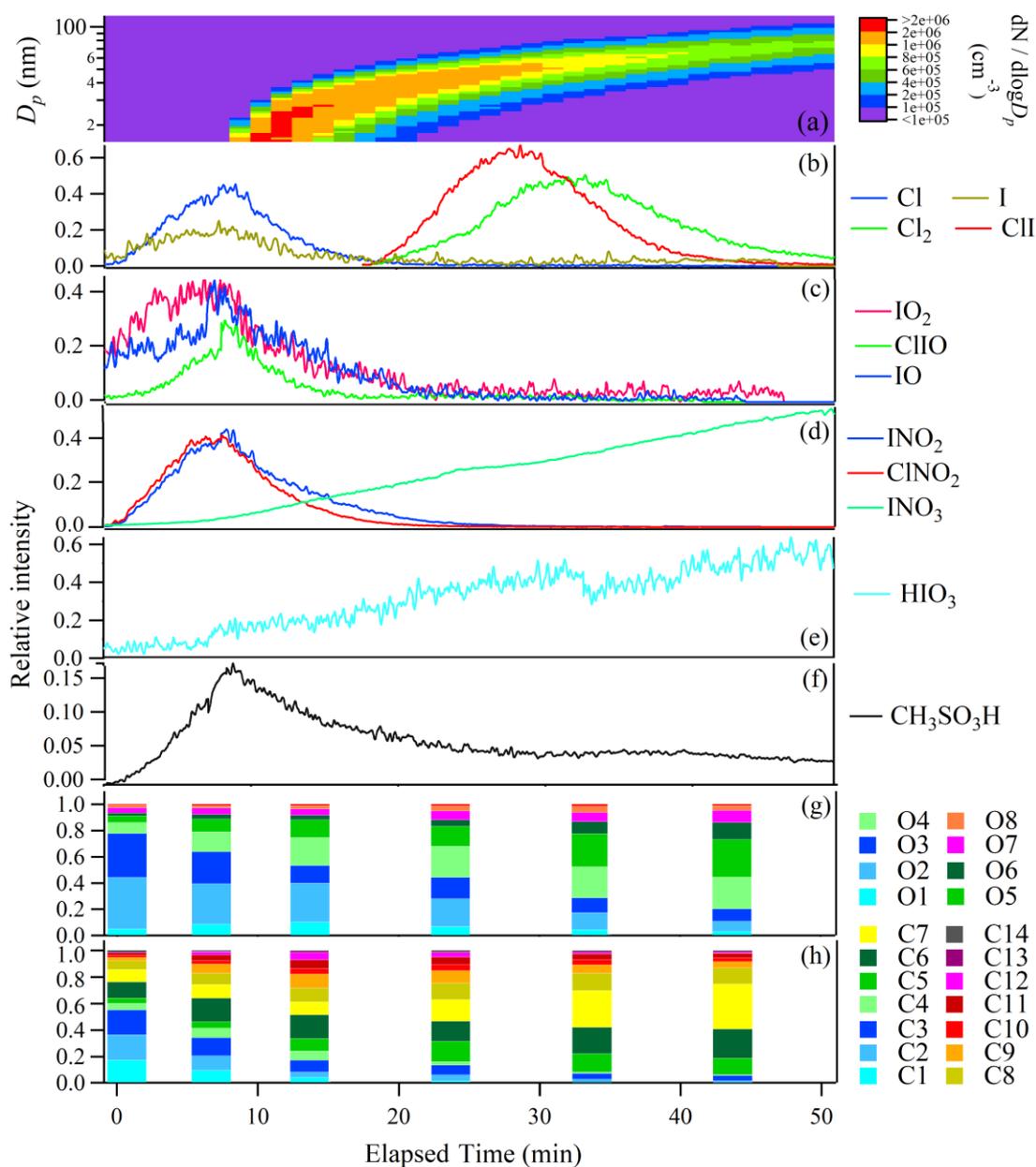


Figure 1. Time evolution of particle number size distribution (a) and relative intensities of gaseous molecules and radicals (b-f); the fractions of organic compounds grouped by O and C atom numbers in the selected time points (g-h) in a typical ozonolysis experiment (static mode). Time zero was chosen as the start time when HIO<sub>3</sub> was observed.

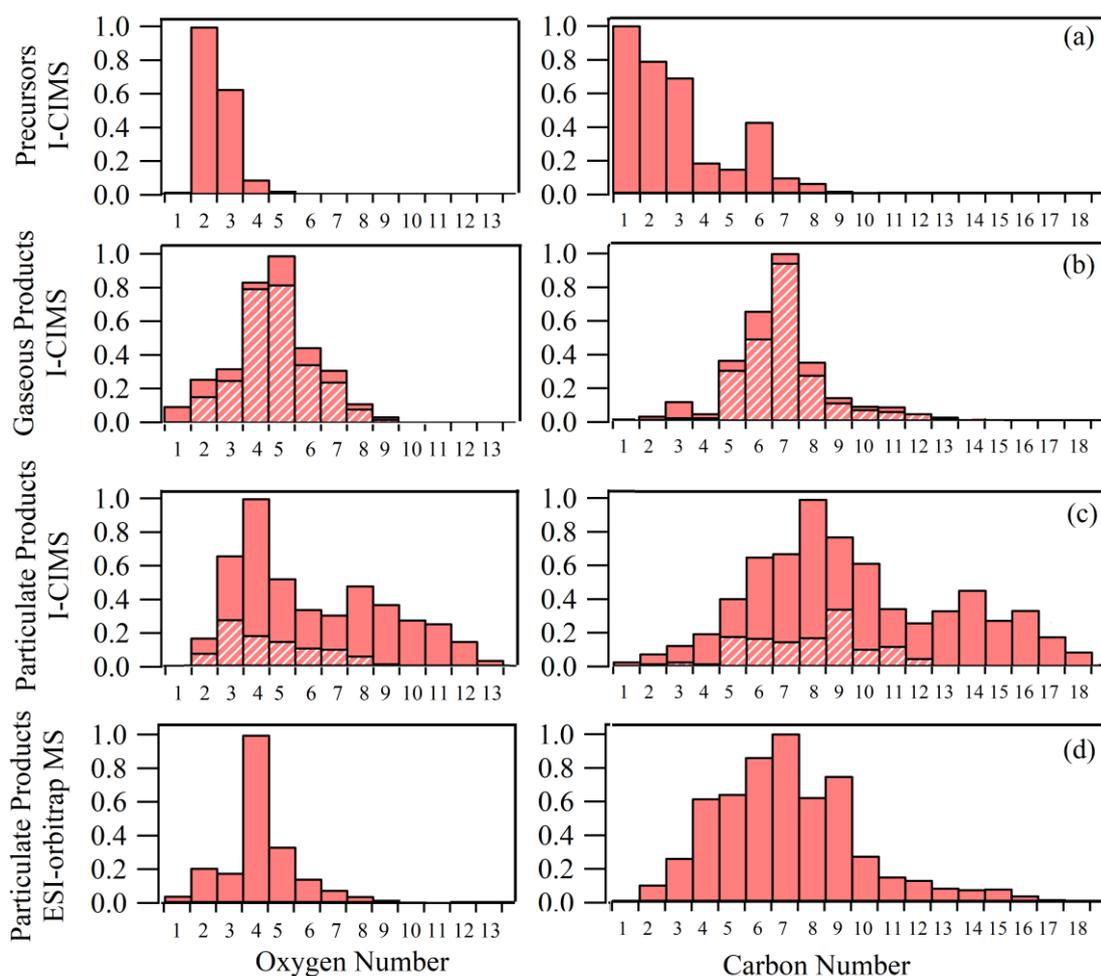


Figure 2. Oxygen and carbon atom number distributions of potential VOC precursors (a), gaseous products (b) and particulate products (c) measured by iodide-CIMS, as well as the particulate products measured by ESI-orbitrap MS (d) in a typical ozonolysis experiment (dynamic mode). Hatched bars indicate the fractions of organic formulas observed in both gas and particle phases by iodide-CIMS.