

Modeling Multiphase Flow Within and Around Deformable Porous Materials: A Darcy-Brinkman-Biot Approach

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Key Points:

- We developed a numerical solver to model multiphase flow through and around deformable porous media.
- The model replicates the Volume-of-Fluid approach and multiphase Biot Theory with a single equation.
- This open-source model can simulate complex phenomena including fracturing in ductile porous media.

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Abstract

We present a new computational fluid dynamics approach to simulating two-phase flow in hybrid systems containing solid-free regions and deformable porous matrices. Our approach is based on the derivation of a unique set of volume-averaged partial differential equations that asymptotically approach the Navier-Stokes Volume-of-Fluid equations in solid-free-regions and multiphase Biot Theory in porous regions. The resulting equations extend our recently developed Darcy-Brinkman-Biot framework to multiphase flow. Through careful consideration of interfacial dynamics (relative permeability and capillary effects) and extensive benchmarking, we show that the resulting model accurately captures the strong two-way coupling that is often exhibited between multiple fluids and deformable porous media. Thus, it can be used to represent flow-induced material deformation (swelling, compression) and failure (cracking, fracturing). The model's open-source numerical implementation, *hybridBiotInterFoam*, effectively marks the extension of computational fluid mechanics into modeling multiscale multiphase flow in deformable porous systems. The versatility of the solver is illustrated through applications related to material failure in poroelastic coastal barriers and surface deformation due to fluid injection in proplastic systems.

Plain Language Summary

Knowledge of how fluids flow through porous materials has significant implications for the design and operation of batteries, manufacturing plants, oil rigs, and biomedical devices. Even though scientists have been successful in creating computer models that capture fluid flow through *rigid* porous media, it has been very challenging to create models that can model flow through *deformable* porous media. In this paper, we describe a new model that can predict flow of immiscible fluids (say water and air, or oil and water) through and around deformable porous media. We derived this model by putting together separate conventional fluid-flow and solid-deformation models into a single simulation framework through a technique called volume averaging. The resulting model can capture complex multiscale, multiphysics phenomena such as hydraulic fracturing in the subsurface and its results on surface deformation and subsidence. Given the model's generality, successful verification, and open-source implementation, we are confident that this computational model can be used to study important phenomena in the fields of water and energy resources.

1 Introduction

Multiphase flow in deformable porous media is a ubiquitous phenomenon with important implications in many energy and environmental technologies including geologic CO₂ sequestration, soil bioremediation, water treatment, enhanced biochemical production, nuclear waste disposal, and battery technology (Bächer & Gekle, 2019; Bock et al., 2010; Cunningham et al., 2003; Räss et al., 2018; Towner, 1987). It also underlies iconic geophysical features at many scales, from coastal, riparian, and volcanic landforms to fractures in subsurface reservoirs, cracks in clay soils, and bubbles in soft sediments. An important and largely unresolved challenge in the areas outlined above is the difficulty of describing the inherently multiscale and multiphysics nature of situations where a mixture of several fluids interacts with a deformable porous material. For example, when modeling flow through biofilms or membranes it is imperative to understand how fluid flow behaves inside the microporous medium (in pores with length scales of $\sim 10^{-6}$ m) while simultaneously understanding how the deformation of this medium affects the overall flow field (often controlled by much larger flow paths with length scales on the order of $\sim 10^{-2}$ m) (Bottero et al., 2010). Similarly, the propagation of flow-driven fractures in porous materials and the propagation of waves in coastal barriers involve feedbacks between flow and mechanics in systems with characteristic pore widths that differ by three or more orders

of magnitude. In the present study, we develop a framework capable of representing multiphase flow and solid mechanics in systems with two characteristic pore length scales, as required to simulate many of the aforementioned phenomena.

The starting point for our study is based on the present ample understanding of multiphase flow dynamics within and around *static* porous materials, from viscous and capillary fingering (Ferer et al., 2004; Lenormand & Zarcone, 1989; Lenormand et al., 1988) to temperature and surface tension driven flows (Shih & Megaridis, 1996), all the way to turbulent multiphase flows (Colombo & Fairweather, 2015; Soulaine & Quintard, 2014). This knowledge, in conjunction with numerical techniques such as the Lattice Boltzmann Method, the Finite Volume Method, Homogenization Theory, and Averaging Theory, forms the basis of fast and accurate models that are routinely applied to help design and improve hydrocarbon production (Burrus et al., 1991; Mehmani & Tchelepi, 2019), CO₂ sequestration (Hassan & Jiang, 2012), and even nuclear reactors (Tentner et al., 2008). However, the study of multiphase flow across different scales remains limited as shown by the absence of well-established approaches to describe how bubbles or waves propagate into an unsaturated porous medium or how a multiphase fluid mixture is pushed out of a porous medium into open space. Understanding such processes would have a direct and immediate impact in the design of batteries, natural gas extraction from shales, biochemical gas production, fracturing systems, and coastal barriers.

A similar situation pertains with regard to the coupling between fluid flow and solid mechanics. Theoretical and numerical approaches based on Biot's Theory of poroelasticity (Biot, 1941), Terzaghi's effective stress principle (Terzaghi, 1943), and Mixture Theory (Siddique et al., 2017) have been successful at modeling systems with flow in deformable porous media including arteries, biofilms, boreholes, hydrocarbon reservoirs, seismic systems, membranes, soils, swelling clays, and fractures (Auton & MacMinn, 2017; Barry et al., 1997; Jha & Juanes, 2014; Lo et al., 2005, 2002; MacMinn et al., 2016; Mathias et al., 2017; Santillán et al., 2017). However, as mentioned above, we still have very little understanding of how flow-induced deformation of these solid materials affects the macroscopic flow around them (and thus their boundary conditions) or how fluid-fluid interfaces behave when pushed against a soft porous medium and vice-versa.

Three major approaches have been proposed to resolve the challenge posed by fluid flow in porous media containing both solid-free regions and microporous domains (hereafter referred to as multiscale systems). The most straightforward of these involves performing direct numerical simulations (DNS) throughout the entire multiscale domain, both within and outside the porous medium (Breugem & Boersma, 2005; Hahn et al., 2002; Krafczyk et al., 2015). Although rigorous, this technique is impractical in situations with a large difference in length scales between the largest and smallest pores, where it requires exceedingly fine grids and tremendous computational resources.

To save time and resources, other studies have relied on hybrid DNS-Darcy approaches, where fluid and solid mechanics within a porous medium are modeled as averaged quantities through Darcy's law, pore-network models, or Biot's theory of poroelasticity (Weishaupt et al., 2019; Ehrhardt, 2010). One such approach relies on the use of the Beavers-Joseph (BJ) boundary condition to couple fluid flow in solid-free domains (simulated using the Navier-Stokes Equations) and in microporous domains (simulated using Darcy's law) for *single* phase flow and *static* porous media (Beavers & Joseph, 1967; Fetzer et al., 2016). Recent studies have extended this BJ approach to allow multiphase flow in the solid-free domain (Baber et al., 2016) or to include the effects of poroelasticity within the porous medium (Lacis et al., 2017; Zampogna et al., 2019). However, to the best of our knowledge, no BJ based technique has yet been developed to couple solid mechanics with multiphase flow simultaneously within the solid-free and porous domains.

The Darcy-Brinkman (DB) approach presents a well-known alternative to the BJ interface matching technique. The crux of the DB approach is the use of a spatially de-

116 pendent penalization term within the Navier-Stokes fluid momentum equation. This term
 117 effectively creates an equation that approximates Navier-Stokes within solid-free domains
 118 and Darcy’s law within microporous domains. Although initially implemented as an em-
 119 pirical approach (Brinkman, 1947), this technique has since been formalized and rigor-
 120 ously derived from first principles through volume averaging theory (Soulaine et al., 2016;
 121 Whitaker, 2013). The resulting so-called “micro-continuum” approach has been exten-
 122 sively used to solve *single* phase flow through *static* multiscale porous media, such as flow
 123 in biofilms (Kapellos et al., 2007) and in rocks containing unresolved porosity (Guo et
 124 al., 2018; Kang et al., 2019; Singh, 2019). The approach has proved highly flexible as il-
 125 lustrated by its uses to represent embedded solid boundaries in low permeability media
 126 (Khadra et al., 2000) and the evolution of solid grain morphologies caused by mineral dis-
 127 solution (Soulaine et al., 2017, 2019).

128 Recently, a study by Carrillo and Bourg (2019) introduced a Darcy-Brinkman-Biot
 129 (DBB) approach capable of accurately representing *single* phase flow in multiscale *de-*
 130 *formable* media including elastic porous membranes and plastic swelling clays. Simultane-
 131 ously, studies by Soulaine et al. (2019) and Carrillo et al. (2020) extensively benchmarked
 132 and released an open source extension of the micro-continuum framework for *multiphase*
 133 flow in *static* multiscale porous media. This allowed accurate modeling of complex sys-
 134 tems such as multiphase flow in a fractured microporous medium, methane extraction
 135 from tight porous media, and wave absorption in coastal barriers. In the present paper,
 136 we build upon these previous studies to create the first model representing coupled fluid
 137 and solid mechanics during *multiphase* flow in *multiscale deformable* porous media: the
 138 multiphase Darcy-Brinkman-Biot model (Figure 1).

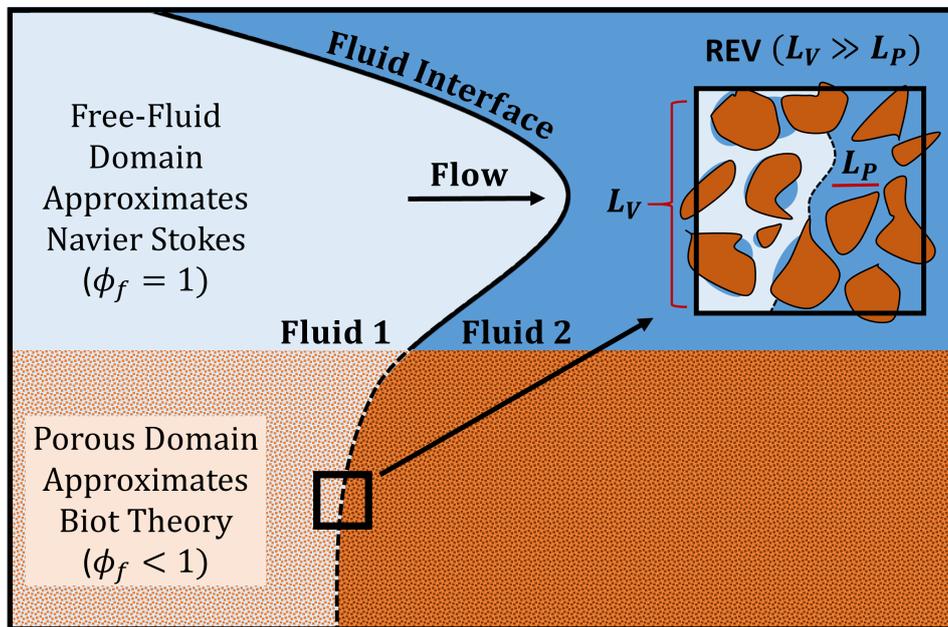


Figure 1. Conceptual representation of the multiphase Darcy-Brinkman-Biot model. The insert represents an exemplary Representative Elementary Volume (REV) within the microporous domain and ϕ_f is the porosity. The model considers wetting properties, interface mechanics, and irreducible saturations when averaging over the REV. Note that the stated relation between the averaging volume’s length scale L_V and the porous length scale L_P is required for the creation of a REV, and thus, for the application of this model.

139 This paper is organized as follows. Section 2 introduces the concept of volume aver-
 140 aging and describes the derivation of the governing equations for coupled fluid and solid
 141 mechanics. Section 3 explains the numerical implementation and algorithm development
 142 for the coupled mass and momentum equations and introduces the resulting open-source
 143 solver “*hybridBiotInterFoam*”. Section 4 presents five test cases that verify the implemen-
 144 tation of different coupling terms within the model, with an emphasis on fracturing me-
 145 chanics. Section 5 then presents two alternative applications that illustrate the versatility
 146 of the model, namely wave absorption in poroelastic coastal barriers and surface deforma-
 147 tion due to fluid injection in poroplastic geologic formations. Lastly, Section 6 concludes
 148 with a summary of the paper and a discussion on future work.

149 2 Model Derivation

150 2.1 Volume Averaging

151 In this section we introduce the concept of volume averaging. This technique forms
 152 the basis of the micro-continuum equations, as it allows the classical mass and momen-
 153 tum conservation equations to account for the coexistence of solid (s), wetting fluid (w),
 154 and non-wetting fluid (n) within a given control volume. It is well suited for use in con-
 155 junction with the Finite Volume Method (FVM) (Patankar, 1980), as the numerical grid
 156 elements used in the FVM provide an intuitive and straightforward numerical interpreta-
 157 tion of what we will define as the averaging volume (V). In keeping with standard volume
 158 averaging theory, we start by defining the volume averaging operator

$$\bar{\beta}_i = \frac{1}{V} \int_{V_i} \beta_i dV \quad (1)$$

159 where β_i is a function defined in each phase’s respective volume V_i ($i = w, n, s$). We also
 160 define the phase averaging operator

$$\bar{\beta}_i^i = \frac{1}{V_i} \int_{V_i} \beta_i dV \quad (2)$$

161 The volume and phase averaged variables associated with the fluids are intrinsically
 162 related by the porosity ($\phi_f = (V_w + V_n)/V$) and saturation fields ($\alpha_i = V_i/(V_w + V_n)$),
 163 such that $\bar{\beta}_i = \phi_f \alpha_i \bar{\beta}_i^i$ ($i = w, n$). For solid variables, the analogous relation involves only
 164 the solid fraction ϕ_s , such that $\bar{\beta}_s = \phi_s \bar{\beta}_s^s$. Note that $\phi_f + \phi_s = 1$ and $\alpha_w + \alpha_n = 1$;
 165 thus, knowledge of one of the ϕ_i or α_i variables implies knowledge of the other. Volume
 166 averaging then allows for the definition of several regions within a multiscale, multiphase
 167 system such as that represented in Figure 1:

$$\phi_f = \begin{cases} 1, & \text{in solid-free regions} \\]0; 1[, & \text{in porous regions} \end{cases} \quad (3)$$

$$\alpha_w = \begin{cases} 0, & \text{in regions saturated with non wetting fluid} \\]0; 1[, & \text{in unsaturated regions} \\ 1, & \text{in regions saturated with wetting fluid} \end{cases} \quad (4)$$

168 The application of an averaging transformation to fluid and solid conservation equa-
 169 tions will result in variables and equations that are weighted differently in each region.
 170 However, the averaging of differential equations is not straightforward, which is why we
 171 introduce the following spatial averaging theorems for volumes containing three distinct

172 phases (Howes & Whitaker, 1985; S. Whitaker, 1999)

$$\frac{\partial \overline{\beta_i}}{\partial t} = \frac{\partial \overline{\beta_i}}{\partial t} - \frac{1}{V} \int_{A_{i,j}} \beta_i \mathbf{v}_{i,j} \cdot \mathbf{n}_{i,j} dA - \frac{1}{V} \int_{A_{i,k}} \beta_i \mathbf{v}_{i,k} \cdot \mathbf{n}_{i,k} dA \quad (5)$$

$$\overline{\nabla \beta_i} = \nabla \overline{\beta_i} + \frac{1}{V} \int_{A_{i,j}} \beta_i \mathbf{n}_{i,j} dA + \frac{1}{V} \int_{A_{i,k}} \beta_i \mathbf{n}_{i,k} dA \quad (6)$$

$$\overline{\nabla \cdot \beta_i} = \nabla \cdot \overline{\beta_i} + \frac{1}{V} \int_{A_{i,j}} \beta_i \cdot \mathbf{n}_{i,j} dA + \frac{1}{V} \int_{A_{i,k}} \beta_i \cdot \mathbf{n}_{i,k} dA \quad (7)$$

173 where $A_{x,y}$ represents the interfacial area between phase x and y , $\mathbf{n}_{x,y}$ is a vector nor-
 174 mal to the interface and oriented toward phase y , and $\mathbf{v}_{x,y}$ is the velocity of the interface.
 175 These surface integrals are crucial components of the following derivations as they convert
 176 the boundary conditions at the fluid-fluid and fluid-solid interfaces into body forces within
 177 the averaged partial differential equations.

178 2.2 Derivation of the Fluid Mechanics Equations

179 We begin the derivation by stating the micro-continuum equations for two immis-
 180 sible incompressible fluids, which arise from applying the volume averaging operators to
 181 the classical Navier-Stokes mass and momentum conservation equations (Carrillo et al.,
 182 2020). This equation set can also be thought of as a modified and expanded version of the
 183 popular Volume-of-Fluid equations (Hirt & Nichols, 1981).

$$\frac{\partial \phi_f}{\partial t} + \nabla \cdot \mathbf{U}_f = 0 \quad (8)$$

$$\frac{\partial \phi_f \alpha_w}{\partial t} + \nabla \cdot (\alpha_w \mathbf{U}_f) + \nabla \cdot (\phi_f \alpha_w \alpha_n \mathbf{U}_r) = 0 \quad (9)$$

$$\begin{aligned} \frac{\partial \rho_f \mathbf{U}_f}{\partial t} + \nabla \cdot \left(\frac{\rho_f}{\phi_f} \mathbf{U}_f \mathbf{U}_f \right) = & -\phi_f \nabla p + \phi_f \rho_f \mathbf{g} + \nabla \cdot \overline{\mathbf{S}} \\ & + \mathbf{D}_{w,s} + \mathbf{D}_{n,s} + \mathbf{D}_{s,w} + \mathbf{D}_{s,n} + \mathbf{D}_{w,n} + \mathbf{D}_{n,w} \end{aligned} \quad (10)$$

184 where $\overline{\mathbf{S}} = \mu_f (\nabla \mathbf{U}_f + (\nabla \mathbf{U}_f)^T)$ is the averaged single-field viscous stress tensor, \mathbf{g} is
 185 gravity, μ_f is the arithmetic average of each fluid's viscosity $\mu_f = \alpha_w \mu_w + \alpha_n \mu_n$, and
 186 ρ_f is the arithmetic average of each fluid's density $\rho_f = \alpha_w \rho_w + \alpha_n \rho_n$. The single-
 187 field expressions for velocity \mathbf{U}_f , pressure p , and relative velocity \mathbf{U}_r are defined as the
 188 weighted averages of their respective phase averaged variables.

$$\mathbf{U}_f = \phi_f \left[\alpha_w \overline{\mathbf{U}}_w^w + \alpha_n \overline{\mathbf{U}}_n^n \right] \quad (11)$$

$$p = \alpha_w \overline{p}_w^w + \alpha_n \overline{p}_n^n \quad (12)$$

$$\mathbf{U}_r = \overline{\mathbf{U}}_w^w - \overline{\mathbf{U}}_n^n \quad (13)$$

189 We note that the single-phase velocity as defined above is equal to the sum of the
 190 Darcy filtration velocities: $\mathbf{U}_f = \bar{\mathbf{U}}_w + \bar{\mathbf{U}}_n$. Finally, the $\mathbf{D}_{i,k} = \frac{1}{V} \int_{A_{i,k}} \mathbf{n}_{i,k} \cdot (-\mathbf{I}p_i + \mathbf{S}_i) dA$
 191 values represent the momentum exchange from phase i to phase k ($i, k = w, n, s$). As de-
 192 scribed in Carrillo and Bourg (2019) and Carrillo et al. (2020), these terms can be recast
 193 into the following expression through asymptotic matching to the multiphase Darcy equa-
 194 tions:

$$\begin{aligned} \frac{\partial \rho_f \mathbf{U}_f}{\partial t} + \nabla \cdot \left(\frac{\rho_f}{\phi_f} \mathbf{U}_f \mathbf{U}_f \right) = & -\phi_f \nabla p + \phi_f \rho_f \mathbf{g} + \nabla \cdot \bar{\mathbf{S}} \\ & -\phi_f \mu k^{-1} (\mathbf{U}_f - \bar{\mathbf{U}}_s) + \phi_f \mathbf{F}_{c,1} + \phi_f \mathbf{F}_{c,2} \end{aligned} \quad (14)$$

195 where μk^{-1} is the drag coefficient (a function of the fluid viscosities and permeability k),
 196 $\bar{\mathbf{U}}_s$ is the averaged solid velocity, $\phi_f \mu k^{-1} (\mathbf{U}_f - \bar{\mathbf{U}}_s)$ is a solid-fluid momentum exchange
 197 term that accounts for a moving porous medium in an Eulerian frame of reference, and
 198 $\mathbf{F}_{c,i}$ represents the forces emanating from fluid-fluid and fluid-solid capillary interactions.
 199 As shown in Carrillo et al. (2020),

$$\mathbf{F}_{c,1} = \begin{cases} -\frac{\gamma}{\phi_f} \nabla \cdot (\mathbf{n}_{w,n}) \nabla \alpha_w & \text{in solid-free regions} \\ -p_c \nabla \alpha_w & \text{in porous regions} \end{cases} \quad (15)$$

$$\mathbf{F}_{c,2} = \begin{cases} 0 & \text{in solid-free regions} \\ M^{-1} (M_w \alpha_n - M_n \alpha_w) (\nabla p_c + (\rho_w - \rho_n) \mathbf{g}) & \text{in porous regions} \end{cases} \quad (16)$$

200 where, p_c is the average capillary pressure within a given averaging volume, γ is the
 201 fluid-fluid interfacial tension, M_i is the mobility of each fluid, and $M = M_w + M_n$ is the
 202 single-field mobility. Lastly, $\mathbf{n}_{w,n}$ is the unit normal direction of the fluid-fluid interface
 203 as calculated by the Continuum Surface Force (CSF) formulation (Brackbill et al., 1992).
 204 The equations presented above tend towards the standard Navier-Stokes Volume-of-Fluid
 205 approach in solid-free regions (where the drag term becomes negligible) and towards the
 206 multiphase Darcy equations in microporous regions (where the viscous stress tensor be-
 207 comes negligible under the scale-separation assumption) (Whitaker, 1986; Carrillo et al.,
 208 2020):

$$\text{Eqn. 14} \approx \begin{cases} (\mathbf{U}_f - \bar{\mathbf{U}}_s) = -\frac{k}{\mu} (\nabla p - \rho_f \mathbf{g} - \mathbf{F}_{c,1} - \mathbf{F}_{c,2}) & \text{in porous regions} \\ \left(\frac{\partial \rho_f \mathbf{U}_f}{\partial t} + \nabla \cdot (\rho_f \mathbf{U}_f \mathbf{U}_f) \right) = -\nabla p + \nabla \cdot \bar{\mathbf{S}} + \rho_f \mathbf{g} + \mathbf{F}_{c,1} & \text{in solid-free regions} \end{cases} \quad (17)$$

209 For clarity and conciseness μk^{-1} will be kept in its current form until the end of the
 210 derivation, at which point its full analytical form will be presented.

211 2.3 Derivation of the Solid Mechanics Equations

212 We proceed with the derivation of the micro-continuum solid mechanics equations
 213 by starting from the equations presented in Carrillo and Bourg (2019) for solid mass and
 214 momentum conservation in systems with a single solid phase.

$$\frac{\partial \phi_s}{\partial t} + \nabla \cdot (\phi_s \bar{\mathbf{U}}_s^s) = 0 \quad (18)$$

$$-\nabla \cdot \bar{\boldsymbol{\sigma}} = \phi_s \nabla \cdot \bar{\boldsymbol{\tau}}^s + \phi_s \rho_s \mathbf{g} + \mathbf{B}_{s,w} + \mathbf{B}_{s,n} \quad (19)$$

215 where $\bar{\boldsymbol{\sigma}}$ is the volume averaged solid elastic (or plastic) stress tensor and $\bar{\boldsymbol{\tau}}^s = \mathbf{P}_{conf} -$
 216 $\mathbf{I}p - \mathbf{I}p_{swell}$ is the Terzaghi stress tensor (a function of confining pressure \mathbf{P}_{conf} , fluid
 217 pressure p , and swelling or disjoining pressure p_{swell}). Here, the $\mathbf{B}_{s,i} = \frac{1}{V} \int_{A_{s,i}} (\boldsymbol{\tau} + \boldsymbol{\sigma}) \cdot \mathbf{n}_{s,i} dA$
 218 values represent the momentum exchange between the solid phase s and fluid phase i
 219 ($i = w, n$). Just as we did for the fluid equations, we will assume that the sum of the av-
 220 eraged stresses at the solid-fluid interface can be expressed as the sum of two independent
 221 terms: a drag force that captures shear-induced momentum exchange (\mathbf{B}_{drag}) and a cap-
 222 illary force originating from capillary pressure jumps across the integrated solid surfaces
 223 within the porous media (\mathbf{B}_{cap}).

$$\mathbf{B}_{drag} + \mathbf{B}_{cap} = \mathbf{B}_{s,w} + \mathbf{B}_{s,n} \quad (20)$$

224 We now seek closure of these two coupling terms. By conservation of momentum,
 225 we know that any drag-induced momentum lost by the fluid must be gained by the solid,
 226 thus (Carrillo & Bourg, 2019)

$$\mathbf{B}_{drag} = \phi_f \mu k^{-1} (\mathbf{U}_f - \bar{\mathbf{U}}_s) \quad (21)$$

227 Closure of the capillarity-induced interaction term \mathbf{B}_{cap} is obtained by combining
 228 the solid and fluid momentum equations within the porous medium at low Reynold num-
 229 bers and low permeability, which yields

$$-\nabla \cdot \bar{\boldsymbol{\sigma}} = \phi_s \nabla \cdot \bar{\boldsymbol{\tau}}^s - \phi_f \nabla p + (\phi_s \rho_s + \phi_f \rho_f) \mathbf{g} + \phi_f \mathbf{F}_{c,1} + \phi_f \mathbf{F}_{c,2} + \mathbf{B}_{cap} \quad (22)$$

230 In multiphase porous systems with incompressible grains and no swelling pressure
 231 (i.e. $\nabla \cdot \bar{\boldsymbol{\tau}}^s = -\nabla p$), Biot Theory states that $\nabla \cdot \bar{\boldsymbol{\sigma}} = \nabla p - \rho^* \mathbf{g} + p_c \nabla \alpha_w$, where $\rho^* =$
 232 $(\phi_s \rho_s + \phi_f \rho_f)$ and p_c is the capillary pressure (Jha & Juanes, 2014; Kim et al., 2013).
 233 This expression is satisfied by the previous equation in the absence of capillary forces,
 234 where $\mathbf{F}_{c,1}$, $\mathbf{F}_{c,2}$, \mathbf{B}_{cap} , and p_c equal zero (Carrillo & Bourg, 2019). In the presence of
 235 capillary forces, however, it imposes the following equality

$$\mathbf{B}_{cap} = -(\phi_f \mathbf{F}_{c,1} + \phi_f \mathbf{F}_{c,2} + p_c \nabla \alpha_w) \quad (23)$$

236 Given that $\mathbf{F}_{c,1} = -p_c \nabla \alpha_w$ in the porous domains (Carrillo et al., 2020), the previ-
 237 ous equation can be rearranged to obtain

$$\mathbf{B}_{cap} = \phi_s \mathbf{F}_{c,1} - \phi_f \mathbf{F}_{c,2} \quad (24)$$

238 Equation 24 gives closure to the last coupling parameter and marks the end of this
 239 derivation. The result is a solid conservation equation that tends towards Biot Theory in
 240 porous regions and towards an infinitely deformable solid with no momentum sources in
 241 solid-free regions.

2.4 Interfacial Conditions between Solid-Free Regions and Porous Regions

242 One of the most important features within the framework presented above is the ex-
 243 istence of an interface between solid-free and microporous domains. Although the creation
 244 of a rigorous un-averaged description of this interface is still an open question, we approx-
 245

246 imate a solution to it by guaranteeing its necessary components within our fluid and solid
247 averaged equations.

248 An accurate description of fluid behavior at the interface requires three compo-
249 nents: 1) mass conservation across the interface, 2) continuity of stresses across the inter-
250 face, and 3) an interfacial wettability condition. Components 1 and 2 are intrinsically ful-
251 filled by our solver due to its single-field formulation for velocity and pressure within the
252 fluid conservation equations (Eqns. 8 and 14). As shown in Neale and Nader (1974) and
253 Carrillo and Bourg (2019) these two components are necessary and sufficient to model
254 single-phase flow within a multiscale system. Furthermore, these conditions have also
255 been used for closure when modelling multiphase flow in moving porous media (Lacis
256 et al., 2017; Zampogna et al., 2019; Carrillo et al., 2020). The required wettability con-
257 dition at the porous interface (Component 3) is included in our model through the imple-
258 mentation of a penalized contact angle condition (Eqn. 33) following the steps outlined in
259 Horgue et al. (2014) and Carrillo et al. (2020).

260 The complementary solid conditions at the porous interface are very similar: 1)
261 solid mass conservation across the interface, 2) continuity of fluid-induced stresses across
262 the interface, and 3) a discontinuity of solid stresses at the interface. Just as before, the
263 first two conditions are intrinsically fulfilled through the use of a single set of mass and
264 momentum conservation equations across both domains and have also been used as clo-
265 sure conditions in previous studies (Lacis et al., 2017; Zampogna et al., 2019). The third
266 condition is enforced by the use of volume-averaged solid rheology models that tend to-
267 wards infinitely deformable materials in solid-free regions, as shown in Carrillo and Bourg
268 (2019). When volume-averaged, the behavior of the solid's stress tensor is domain depen-
269 dent (i.e. solid fraction dependent). Thus, in solid regions, the elasticity and viscosity of
270 the porous medium is determined by standard averaged rheological properties (the elastic
271 and viscoplastic moduli). Contrastingly, in solid-free regions, the solid fraction tends to
272 zero and, as such, said properties do as well. The result is a stress-free "ghost" solid that
273 does not apply resistance to the porous region, creating the required stress discontinuity at
274 the porous interface.

275 Although necessary, these conditions represent but an approximation to the complete
276 description of fluid and solid mechanics at the porous interface. However, to the best of
277 our knowledge, there does not exist an alternative set of boundary conditions that can or
278 have been used to model multiphase flow in multiscale porous media.

279 2.5 Model Summary

280 The final set of equations in our proposed multiphase DBB framework now fol-
281 lows. The combination of these solid and fluid conservation equations leads to a model
282 that tends towards multiphase Navier-Stokes in solid-free regions and towards Biot Theory
283 in porous regions, as described in Figure 1.

$$\frac{\partial \phi_f}{\partial t} + \nabla \cdot \mathbf{U}_f = 0 \quad (25)$$

$$\frac{\partial \phi_f \alpha_w}{\partial t} + \nabla \cdot (\alpha_w \mathbf{U}_f) + \nabla \cdot (\phi_f \alpha_w \alpha_n \mathbf{U}_r) = 0 \quad (26)$$

$$\begin{aligned} \frac{\partial \rho_f \mathbf{U}_f}{\partial t} + \nabla \cdot \left(\frac{\rho_f}{\phi_f} \mathbf{U}_f \mathbf{U}_f \right) = & -\phi_f \nabla p + \phi_f \rho_f \mathbf{g} + \nabla \cdot \bar{\mathbf{S}} \\ & -\phi_f \mu k^{-1} (\mathbf{U}_f - \bar{\mathbf{U}}_s) + \phi_f \mathbf{F}_{c,1} + \phi_f \mathbf{F}_{c,2} \end{aligned} \quad (27)$$

$$\frac{\partial \phi_s}{\partial t} + \nabla \cdot (\phi_s \bar{\mathbf{U}}_s) = 0 \quad (28)$$

$$-\nabla \cdot \bar{\boldsymbol{\sigma}} = \phi_s \nabla \cdot \bar{\boldsymbol{\tau}}^s + \phi_s \rho_s \mathbf{g} + \phi_f \mu k^{-1} (\mathbf{U}_f - \bar{\mathbf{U}}_s) - \phi_f \mathbf{F}_{c,1} + \phi_s \mathbf{F}_{c,2} \quad (29)$$

284 All that is left is stating the closed-form expressions of the multiscale parameters
 285 μk^{-1} , $\mathbf{F}_{c,i}$, and \mathbf{U}_r , which are defined differently in each region. A full derivation and
 286 discussion of these parameters can be found in Carrillo et al. (2020).

$$\mu k^{-1} = \begin{cases} 0 & \text{in solid-free regions} \\ k_0^{-1} \left(\frac{k_{r,w}}{\mu_w} + \frac{k_{r,n}}{\mu_n} \right)^{-1} & \text{in porous regions} \end{cases} \quad (30)$$

$$\mathbf{F}_{c,1} = \begin{cases} -\frac{\gamma}{\phi_f} \nabla \cdot (\mathbf{n}_{w,n}) \nabla \alpha_w & \text{in solid-free regions} \\ -p_c \nabla \alpha_w & \text{in porous regions} \end{cases} \quad (31)$$

$$\mathbf{F}_{c,2} = \begin{cases} 0 & \text{in solid-free regions} \\ M^{-1} (M_w \alpha_n - M_n \alpha_w) (\nabla p_c + (\rho_w - \rho_n) \mathbf{g}) & \text{in porous regions} \end{cases} \quad (32)$$

$$\mathbf{n}_{w,n} = \begin{cases} \frac{\nabla \alpha_w}{|\nabla \alpha_w|} & \text{in solid-free regions} \\ \cos(\theta) \mathbf{n}_{wall} + \sin(\theta) \mathbf{t}_{wall} & \text{at the interface between solid-free porous regions} \end{cases} \quad (33)$$

$$\mathbf{U}_r = \begin{cases} C_\alpha \max(|\mathbf{U}_f|) \frac{\nabla \alpha_w}{|\nabla \alpha_w|} & \text{in solid-free regions} \\ \phi^{-1} \begin{bmatrix} -(M_w \alpha_w^{-1} - M_n \alpha_n^{-1}) \nabla p + \\ (\rho_w M_w \alpha_w^{-1} - \rho_n M_n \alpha_n^{-1}) \mathbf{g} + \\ (M_w \alpha_n \alpha_w^{-1} + M_n \alpha_w \alpha_n^{-1}) \nabla p_c - \\ (M_w \alpha_w^{-1} - M_n \alpha_n^{-1}) p_c \nabla \alpha_w \end{bmatrix} & \text{in porous regions} \end{cases} \quad (34)$$

287 where C_α is an interface compression parameter (traditionally set to values between 1
 288 and 4 in the Volume-of-Fluid method), k_0 is the absolute permeability, $k_{r,i}$ and $M_i =$
 289 $k_0 k_{i,r} / \mu_i$ are the relative permeability and mobility of each fluid, and $M = M_w + M_n$.
 290 Lastly, θ is the imposed contact angle at the porous wall, and \mathbf{n}_{wall} and \mathbf{t}_{wall} are the nor-
 291 mal and tangential directions relative to said wall, respectively.

292 Finally, closure of the system of equations requires appropriate constitutive models
 293 describing the averaged behavior of the different phases within the porous regions. For the
 294 purpose of validating our multiphase DBB approach, in the present paper we use the fol-
 295 lowing well established constitutive models: absolute permeability is modeled as isotropic
 296 and porosity-dependent through the well-known Kozeny-Carman relation ($k_0 = k_0^0 \frac{\phi_f^3}{(1-\phi_f)^2}$);
 297 relative permeabilities and average capillary pressures within the porous domains are rep-
 298 resented using the Van Genuchten (van Genuchten, 1980) and Brooks-Corey (Brooks &
 299 Corey, 1964) models (Appendix A); plasticity is described through the Herschel-Bulkley
 300 model, where the solid viscously deforms only after local stresses become higher than the
 301 material yield stress (Appendix B1); the solid's yield stress and plastic viscosity are mod-
 302 eled as solid fraction-dependent based on the Quemada fractal model (Quemada, 1977;
 303 Spearman, 2017) (Appendix B2); finally, elastic solids are modeled as averaged linear-
 304 elastic materials, such that their averaged elastic coefficients scale linearly with respect to
 305 the solid fraction (Appendix B3). The last three choices imply that solid rheological prop-
 306 erties are modeled as isotropic and independent of saturation, a significant simplification

that is sufficient for the purpose of testing and validating the present framework. For the reader's convenience, a full implementation of this framework and its related models are included in the accompanying simulation "toolbox". If necessary, more complex constitutive models, such as the saturation-depended solid rheology models presented in Wan et al. (2014), Oldecop and Alonso (2003), Buscarnera and Einav (2012), and Di Donato et al. (2003) can be readily implemented into our code by virtue of its open-source implementation.

3 Numerical Implementation

3.1 Numerical Platform

The implementation of the multiphase DBB model was done in OpenFOAM®, a free, open-source, parallelizable, and widely used computational fluid mechanics platform. This C++ code uses the Finite Volume Method to discretize and solve partial differential equations in complex 3-D structured and unstructured grids. Its object-oriented structure and multitude of supporting libraries allows the user to easily customize each simulation's setup with different numerical discretization schemes, time-stepping procedures, matrix-solution algorithms, and supporting physical models. The implementation described below represents the natural extension of the multiphase micro-continuum toolkit "*hybridInterFoam*" (Carrillo et al., 2020) to systems with deformable solids. In particular, its solution algorithm stems directly from that used by "*hybridInterFoam*" and its precursor "*interFoam*".

3.2 Solution Algorithm

The solution of the governing equations is done in a sequential manner, starting with the fluid mechanics equations and following with the solid mechanics equations for every time step. Of particular importance is the handling and modification of the velocity-pressure coupling required for modeling incompressible fluids in conjunction with a moving solid matrix. For this step, we based our solution algorithm on the Pressure Implicit Splitting-Operator (PISO) (Issa, 1986). First, we explicitly solve the fluid saturation equation (Eqn. 9) for α_w^{t+1} through the Multidimensional Universal Limiter of Explicit Solution (MULES) algorithm (Márquez & Fich, 2013). This allows for stable numerical advection of the saturation field by the application of Flux Corrected Transport Theory (Rudman, 1997). Then, we update the boundary values of \mathbf{U}_f and \mathbf{U}_r in addition to the cell-centered values of the permeability k^{t+1} , density ρ_f^{t+1} , and viscosity μ_f^{t+1} based on the newly calculated saturation field α_w^{t+1} . The capillary forces $\mathbf{F}_{c,i}^{t+1}$ are also updated accordingly. After that, a preliminary value of the fluid velocity \mathbf{U}_f^* is calculated by implicitly solving the algebraically discretized form of the fluid momentum equation used in the Finite Volume Method.

$$a_p \mathbf{U}_f^* = \mathbf{H}(\mathbf{U}_f^*) + \rho_f^{t+1} \mathbf{g} + \mathbf{F}_{c,i}^{t+1} - \nabla p^t \quad (35)$$

where $\mathbf{H}(\mathbf{U}_f^*)$ contains inertial, convective, viscous, and drag source terms originating from neighboring cells and a_p represents these same terms but at the volume of interest. Note that the \mathbf{U}_f^* field does not follow mass conservation. To account for this, we use the fluid continuity equation (Eqn. 25) in conjunction with the previous equation (Eqn. 35) to update the velocity field \mathbf{U}_f^{**} and calculate a preliminary mass-conservative pressure field p^* . In other words, these fields must satisfy,

$$\mathbf{U}_f^{**} = \frac{1}{a_p} \left(\mathbf{H}(\mathbf{U}_f^*) + \rho_f^{t+1} \mathbf{g} + \mathbf{F}_{c,i}^{t+1} - \nabla p^* \right) \quad (36)$$

$$\nabla \cdot \mathbf{U}_f^{**} = - \frac{\partial \phi_f}{\partial t} \quad (37)$$

349 These equations can be recast into a single coupled equation which is then used to
 350 implicitly solve for pressure. This step can be done through several generalized matrix
 351 solvers that are standard in OpenFOAM®.

$$\nabla \cdot \left(\frac{1}{a_p} \left(\mathbf{H} \left(\mathbf{U}_f^* \right) + \rho_f^{t+1} \mathbf{g} + \mathbf{F}_{c,i}^{t+1} - \nabla p^* \right) \right) - \nabla \cdot \left(\frac{1}{a_p} \nabla p^* \right) = - \frac{\partial \phi_f}{\partial t} \quad (38)$$

352 After solving for pressure p^* , velocity can be re-calculated from Equation 36. This
 353 semi-implicit pressure-velocity correction step is repeated until the desired convergence is
 354 reached. It has been shown that at least two pressure-velocity correction loops are required
 355 to ensure mass conservation (Issa, 1986). At this point \mathbf{U}_f^{t+1} and p^{t+1} are set and used as
 356 input values for updating the drag and pressure source terms present in the solid mechan-
 357 ics momentum equation (Eqn. 29). Then, said equation is discretized in a similar way as
 358 the fluid momentum equation (Eqn. 27) and used to implicitly solve for \mathbf{U}_s^{t+1} . Finally, the
 359 updated solid velocity is used to “advect” the solid fraction field ϕ_s by solving the mass
 360 conservation equation (Eqn. 28). At this point the algorithm advances in time according
 361 to the imposed Courant-Friedrichs-Lewy (CFL) number. Further discussion regarding the
 362 discretization techniques and matrix-solution procedures can be found in Carrillo et al.
 363 (2020) and Jasak (1996).

364 3.3 Open-Source Implementation

365 The complete set of governing equations and solution algorithms, along with the
 366 necessary rheology, relative permeability, and capillary pressure models (Appendix A and
 367 B) were implemented into a single solver “*hybridBiotInterFoam*”. This solver, along with
 368 its representative tutorial cases, automated compilation and running procedures, and all
 369 the simulated cases presented in this paper were incorporated into an open-source CFD
 370 package of the same name. OpenFOAM® and our code are free to use under the GNU
 371 general public license and can be found at <https://openfoam.org/> and [https://](https://github.com/Franjcf)
 372 github.com/Franjcf (Carrillo & Bourg, 2020), respectively.

373 4 Model Validation

374 Most of the underlying components of the approach described above have been
 375 previously tested and verified. Carrillo and Bourg (2019) validated the momentum ex-
 376 change terms as an effective coupling mechanism between a *single* fluid phase and a *de-*
 377 *formable* plastic or elastic porous medium. The effects of confining and swelling pressures
 378 on porous media were also examined in said study. Then, Carrillo et al. (2020) extensively
 379 validated the extension of the Darcy-Brinkman equation into multiphase flow within and
 380 around *static* porous media by comparison with reference test cases in a wide range of
 381 flow, permeability, capillarity, and wettability conditions. Therefore, the only thing left to
 382 validate is the ability of the multiphase DBB model to accurately predict the behavior of
 383 multiscale systems that exhibit coupling effects between multiple fluids and a deformable
 384 porous matrix.

385 To that point, we begin with two validation cases relating to multiphase poroelas-
 386 ticity and the coupling between solid deformation and fluid pressure. Then, we proceed
 387 with two poroplastic cases that validate this framework for multiscale plastic systems. Fi-
 388 nally, we conclude with two additional cases that verify the implementation of the capil-
 389 lary force interaction terms. All of these can be found in the accompanying CFD simula-
 390 tion package.

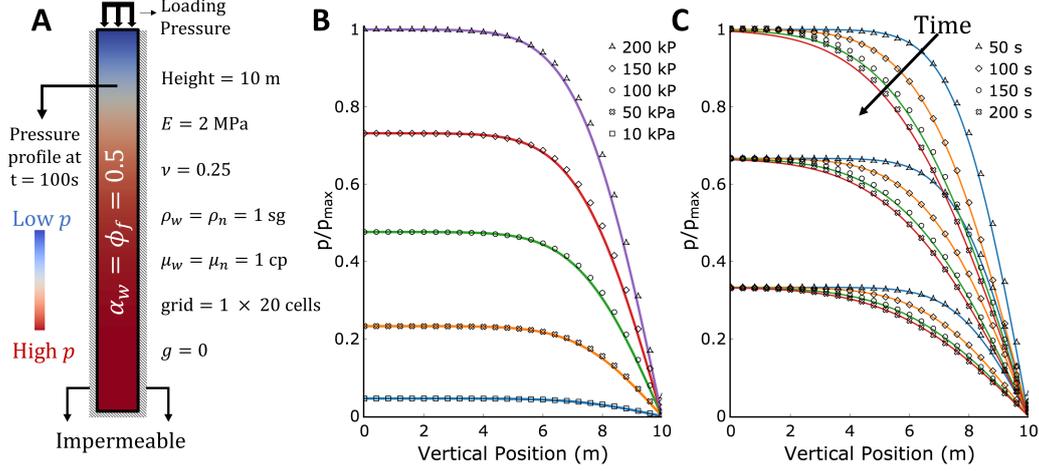


Figure 2. One-dimensional Terzaghi consolidation problem. (A) Simulation setup. (B) Analytical (solid lines) and numerical (symbols) pressure profiles at $t = 100$ s for different loading pressure values. (C) Time-dependent pressure profiles for different column porosity values (From top to bottom: $\phi_s = 0.75, 0.5, 0.25$).

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4.1 Terzaghi Consolidation Problem

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The Terzaghi uniaxial compaction test has been extensively used as a benchmark for the validation of numerical codes relating to poroelasticity (Terzaghi et al., 1996). Its main utility is to test the accuracy of the solid-fluid couplings that relate fluid pressure to solid deformation and *vice versa*. The problem consists of a constrained saturated elastic porous medium that is abruptly compressed from its upper boundary by a constant uniaxial load (Figure 2). This creates a sudden increase in pore pressure, which is then dissipated by flow through the upper boundary (all other boundaries have impermeable boundary conditions). In the case of a one-dimensional porous medium, the resulting temporal and spatial evolution in fluid pressure can be described by the following simplified analytical solution (Verruijt, 2013).

$$\frac{p}{p_{max}} = \operatorname{erf}\left(\frac{h-z}{2\sqrt{c_v t}}\right) \quad \text{for} \quad \frac{c_v t}{h^2} \ll 1 \quad (39)$$

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where $c_v = (k_0 E (\nu - 1)) / (\eta (2\nu^2 + \nu - 1))$ is the consolidation coefficient, k_0 is permeability, E is Young's modulus, ν is Poisson's ratio, η is the fluid's unit weight, h is the column height, and z is the vertical coordinate. Our equivalent numerical setup is shown in Figure 2. The values of the relevant parameters in our simulations are $h = 10$ m, $k_0 = 5 \times 10^{-11}$ m², $E = 2$ MPa, and $\nu = 0.25$. To show the accuracy of our model across different conditions, the loading pressure was varied from 10 to 200 kPa (Figure 2B) and the porosity from 0.25 to 0.75 (Figure 2C). Lastly, the column was partially saturated ($\alpha_w = 0.5$) with fluids with equal densities ($\rho_f = 1000$ kg/m³), viscosities ($\mu_f = 1$ cp), and negligible capillary effects. This last points allowed for testing the validity of the fluid-solid couplings irrespective of the simulated phases without violating any of the assumptions present in the analytical solution. Our numerical results show excellent agreement with Equation 39 for all tested conditions. Further verification of these terms for an oscillating linear elastic solid with pressure boundary conditions (as opposed to stress boundary conditions) can be found in the Supporting Information.

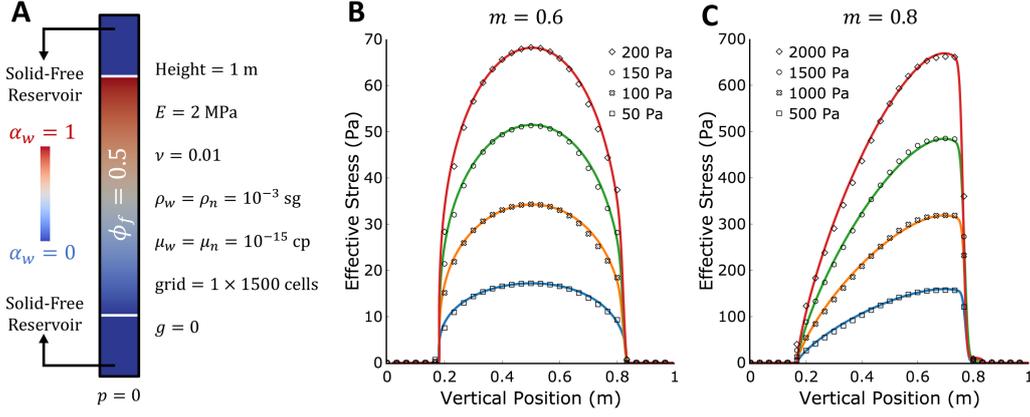


Figure 3. Capillary effects in a poroelastic column. (A) Simulation setup. (B & C) Analytical (solid lines) and numerical (symbols) effective stress profiles for different capillary pressure values ($p_{c,0} = 50$ to 2000 Pa) and Van Genuchten coefficients ($m = 0.6$ and 0.8).

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4.2 Capillary Pressure Effects in a Poroelastic Column

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Having verified the two-way coupling between solid deformation and fluid pressure, we now verify the implementation of the capillary pressure terms within the solid mechanics equation. To do so, we simulate a poroelastic column (1 m tall, 1500 Cells, $\phi_f = 0.5$) bounded by two non-wetting fluid reservoirs at its upper and lower boundaries. The column is initialized with a linear saturation profile spanning from $\alpha_w = 0$ to 1 (see Fig. 3). Fluid saturation is kept fixed by not solving Equation 26, and the mobilities of both fluids are set to very high values ($M_i = 1 \times 10^{10}$ m³/kg.s) to minimize drag-related effects. Under these conditions, the solid's effective stress is exclusively controlled by capillary effects and is described by the following analytical solution:

$$\text{Effective Stress} = \phi_s \times \alpha_w \times p_c \quad (40)$$

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We used the Van Genuchten capillary pressure model with $m = 0.6$ or 0.8 and $p_{c,0} = 50$ to 2000 Pa to calculate the solutions to said problem. The resulting agreement between the numerical and analytical solutions, shown in Fig. 3, confirms the accuracy of the fluid-solid capillary pressure coupling implemented in our model. Furthermore, the transitional behaviour of the effective stress at the macroscopic solid-fluid interface confirms the applicability of the interfacial condition described in Section 2.4: as expected, solid stresses are dictated by standard elasticity theory in the porous region and become negligible in solid-free regions.

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Given that the fluid-solid couplings in a poroelastic solid are now verified, we proceed to verify said terms for poroelastic materials.

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4.3 Fluid Invasion and Fracturing in a Hele-Shaw Cell

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The third verification case (and the first poroelastic case) consists in the qualitative replication of a set of fracturing experiments that examined the injection of aqueous glycerin into dry sand within a 30 by 30 by 2.5 cm Hele-Shaw cell (Huang et al., 2012a, 2012b). These experiments are inherently multiscale, in that the characteristic length scale of fractures in this system (\sim cm) is orders of magnitude larger than that of pores within the microporous matrix ($\sim \mu\text{m}$). They are also multiphysics, as they clearly exemplify the

443 drag-controlled transition from Darcy flow within the porous medium to Stokes flow in the
 444 open fractures and the coupling between the hydrodynamics of fluid flow and the mechan-
 445 ics of fracture propagation (Figure 4).

446 The experimental setup involved the injection of aqueous glycerin at various flow
 447 rates q between 5 and 50 ml/min while also varying the fluid's viscosity μ_{gly} between 5
 448 and 176 cp for different experiments. Our numerical simulations were parameterized us-
 449 ing measured values of the glycerin-air surface tension ($\gamma = 0.063 \text{ kg/s}^2$), the density of
 450 pure glycerin ($\rho_{gly} = 1250 \text{ kg/m}^3$), the density of air ($\rho_{air} = 1 \text{ kg/m}^3$), the viscosity
 451 of air ($\mu_{air} = 0.017 \text{ cp}$), and the average radius and density of sand grains ($100 \mu\text{m}$ and
 452 2650 kg/m^3 , respectively). To mimic the sand's experimental configuration and perme-
 453 ability, the simulated solid fraction field was set to a random initial normal distribution
 454 such that $\phi_s = 0.64 \pm 0.05$ and the permeability was modelled as a function of the solid
 455 fraction through the Kozeny-Carman relation with $k_0 = 6.7 \times 10^{-12} \text{ m}^2$. Relative per-
 456 meabilities were calculated through the Van Genuchten model with the Van Genuchten
 457 coefficient m set to 0.99 (see Appendix A), while capillary pressures were deemed neg-
 458 ligible (as $2\gamma r^{-1} \ll \mu k^{-1} U_f L$). Finally, the porous medium was modeled as a continu-
 459 ous Hershel-Bulkley-Quemada plastic (Appendix B) with yield stress $\tau_0 = 16.02 \text{ m}^2/\text{s}^2$
 460 (Quemada, 1977). Plasticity was used as the preferred mode of solid rheology due to its
 461 ability to account for the compressive and irreversible effects caused by fracturing within
 462 these experiments (Ahmed et al., 2007; van Dam et al., 2002).

463 Numerically speaking, the simulations were carried out in a 30 by 30 cm 2-D grid
 464 (500 by 500 cells) with constant velocity and zero-gradient pressure boundary conditions
 465 at the inlet, zero-gradient velocity and zero pressure boundary conditions at the boundary
 466 walls, and a solid velocity tangential slip condition at all boundaries (i.e. the solid can-
 467 not flow across the boundaries, but the fluids can). Lastly, to enable a closer comparison
 468 between our 2D simulation and the 3D experiment we added an additional drag term to
 469 the fluid momentum equation equal to $12\mu a^{-2} U_f$, which accounts for viscous dissipation
 470 through friction with the walls in a Hele-Shaw cell with aperture a (Ferrari et al., 2015).

471 As shown in Figure 4, a dramatic transition in the mode of fluid invasion is ob-
 472 served with increasing fluid injection velocity and viscosity. At low flow rates and low
 473 viscosity ($q = 5 \text{ ml/min}$, $\mu = 5 \text{ cp}$), there is no discernible solid deformation and the
 474 main mode of fluid flow is through uniform invasion of the porous medium (Figure 4A).
 475 At intermediate flow rates and low viscosity ($q = 25 \text{ ml/min}$ to 30 ml/min , $\mu = 5 \text{ cp}$),
 476 we still observe a uniform invasion front, but small fractures begin to appear (Figure 4B,
 477 C). At high viscosity ($\mu = 176 \text{ cp}$), we see clear fracturing patterns preceded by a non-
 478 uniform fluid invasion front (Figure 4H, I).

479 Figure 4 shows that our simulation predictions are qualitatively consistent with the
 480 experiments presented in Huang et al. (2012a) with regard to both the stability of the cap-
 481 illary displacement front and the observed fracturing transition behavior. As suggested
 482 above, accurate prediction of this transition requires not only proper handling of fluid-
 483 fluid interactions (surface tension and relative permeability effects), but also accurate de-
 484 scriptions of their relationship with solid mechanics (drag) and the proper implementation
 485 of a solid rheological model that can replicate irreversible and unstable fracturing pro-
 486 cesses. We note that in our simulations, fracture initialization and propagation are pre-
 487 dicted based on continuum-scale equations for the rheology and mechanics of the bulk
 488 microporous solid, with no specific treatment of grain-scale mechanics. Grid-level insta-
 489 bilities are brought about by the normally distributed porosity and permeability fields, as
 490 shown in Appendix C. The microstructural differences between the experiments and our
 491 simulations (most clear in Figure 4C, F, and H, K) likely arise at least in part from the
 492 fact that the solid is modelled as a continuum rather than a granular material.

493 This section demonstrates that the multiphase DBB model can be used to replicate
 494 and predict the main mode of fluid flow and solid deformation within fracturing systems.

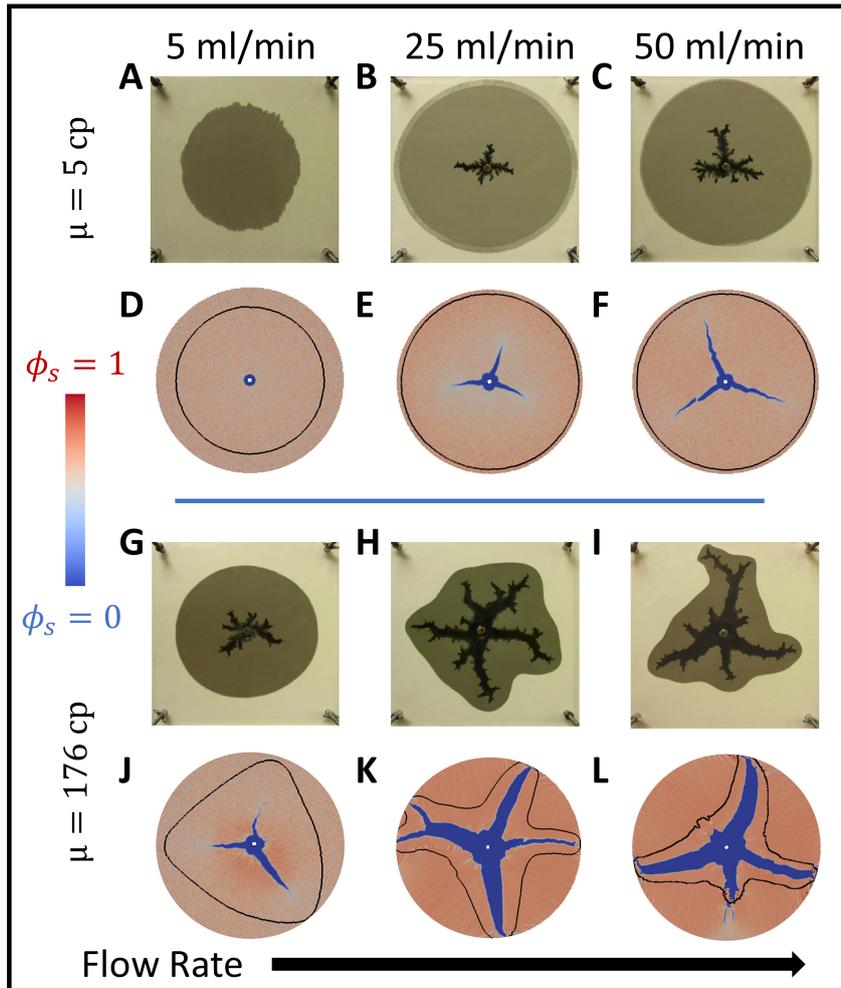


Figure 4. Comparison of experimental (A, B, C, G, H, I) and simulated (D, E, F, J, K, L) fracturing in a Hele-Shaw cell. The color bar represents the solid fraction within the simulations (where red implies a pure solid and blue pure fluids) and the black lines represent the advancing glycerin saturation front. The experiments shown here are part of the results presented in Huang et al. (2012a).

495 A comprehensive study of the controlling parameters for multiphase fracturing in the presence of both viscous and capillary stresses will be the focus of an adjacent study.
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497 **4.4 Modeling Fracturing Wellbore Pressure**

498 Having shown that our model can qualitatively predict fracturing behavior, we now
499 aim to determine whether it can do so in a quantitative matter. As depicted in Figure 5,
500 fluid-induced fracturing of low-permeability rocks proceeds through the following well-
501 established series of stages: First, fluid pressure increases linearly as fracturing fluid is
502 injected into the wellbore. Second, as wellbore pressure increases and approaches the
503 leak-off pressure, a small amount of pressure is propagated by fluid leakage into the rock.
504 Third, fluid pressure continues to increase until it reaches the breakdown pressure, at
505 which point it is high enough to fracture the rock. Fourth, a fracture is initiated and prop-
506 agates; the wellbore pressure slowly decreases. Fifth, injection stops, fracture propagation

507 stops, and wellbore pressure rapidly dissipates (Abass et al., 2007; Ahmed et al., 2007;
 508 Huang et al., 2012a; Papanastasiou, 2000; Santillán et al., 2017).

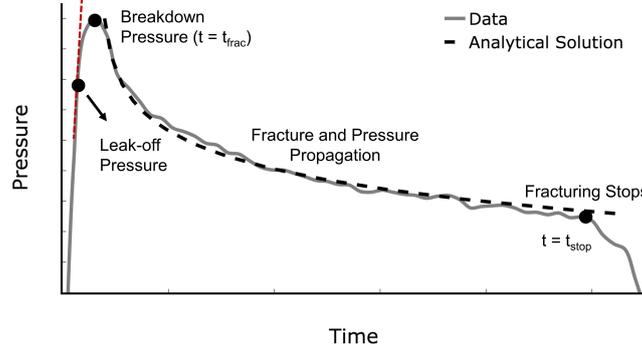


Figure 5. Conceptual representation of wellbore pressure evolution during fluid-induced fracturing of low permeability rocks. In this section, we are interested in modeling the behavior between t_{fract} and t_{stop} .

509 In this section we aim to numerically replicate the time-dependent fracturing well-
 510 bore pressure during fracture propagation (i.e., the fourth stage outlined above) as de-
 511 scribed by an analytical solution presented in Barros-Galvis et al. (2017).

$$p_{well} = p_0 - \frac{\mu q}{4\pi k_0 h} \left[\ln \left(\frac{t k_0 \tau_0}{\phi_f \mu r_{well}^2} \right) + 0.81 \right] \quad (41)$$

512 where t is the time elapsed since fracture initialization, q is the fluid injection rate, p_{well}
 513 is the wellbore pressure, p_0 is the minimum pressure required for starting a fracture (a
 514 function of the solid's yield stress τ_0), h is the formation thickness, and r_{well} is the well-
 515 bore radius. The remaining variables follow the same definitions described earlier.

516 The general numerical setup is almost identical to the one presented in the previ-
 517 ous section. The key difference is that we now inject aqueous glycerin into a strongly-non
 518 wetting (and thus almost impermeable) porous material. This is done to ensure an accu-
 519 rate replication of the analytical solution and its related assumptions, where fracturing is
 520 the main mode of fluid flow and there is virtually no fluid invasion into the porous ma-
 521 trix. The exact simulation parameters are $q = 46$ to 110 ml/min, $\tau_0 = 0.2$ or 2 m²/s²,
 522 $k_0 = 6.7 \times 10^{-11}$ or 6.7×10^{-12} m², $\mu_{gly} = 5$ cp, and $m = 0.05$. Note that low values of m
 523 indicate that the porous formation is strongly non-wetting to the injected fluid. All other
 524 parameters are as in the previous section.

525 Lastly, as hinted at before, a notable characteristic of our model is that different
 526 normally-distributed solid fraction field initializations give different fracturing results (Ap-
 527 pendix C). For this reason, we performed four simulations for each parameter set. In Fig-
 528 ure 6, we present the average predicted wellbore pressure evolution with errors bar repre-
 529 senting the 95% confidence interval.

530 Figure 6 shows that our model can accurately and reliably predict the pressure and
 531 deformation behavior of a variety of fracturing systems, as all curves exhibit excellent
 532 agreement with their respective analytical solution. Note that the length of each curve re-
 533 lates inversely to the injection speed. This is because fractures at higher injection rates
 534 consistently reach the system's boundary faster than their counterparts, at which point
 535 there is a sharp decrease in pressure and the analytical solution no longer applies. There-
 536 fore, each curve's cutoff point represents the time at which the fracture effectively be-

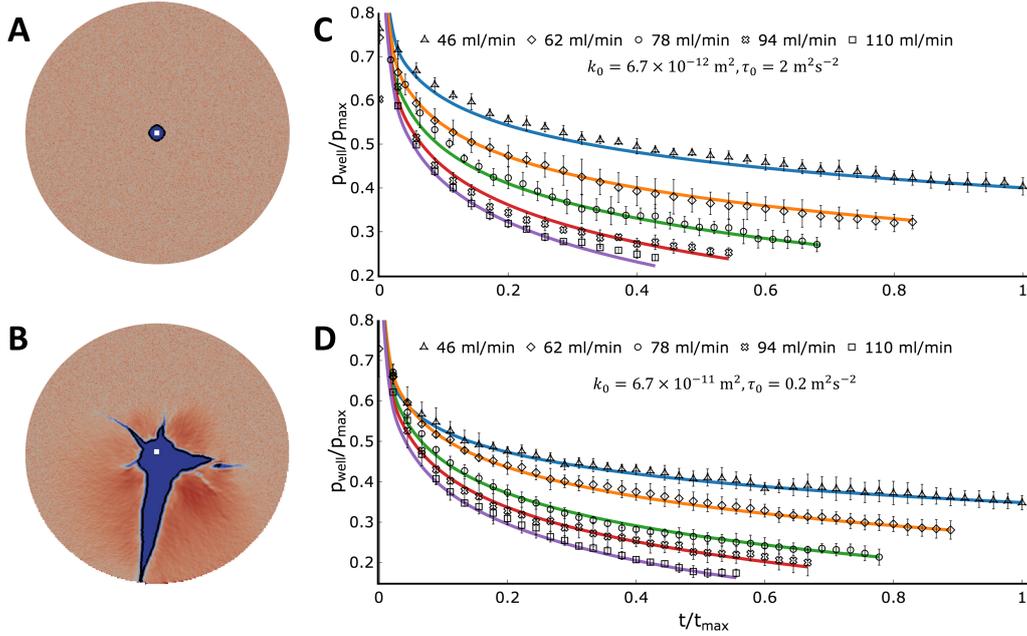


Figure 6. Wellbore pressure as a function of injection rate and time. (A) The initial simulation setup showing the initial wellbore radius $r_{well} = 1.3$ cm, as well as the normally distributed solid fraction field. (B) The fractured system, where the thin black line represents the position of the advancing glycerin saturation front. C and D show the wellbore pressure as a function of time for different flow rates and different combinations of solid yield stress and permeability. Solid curves represent analytical solutions, while symbols represent simulation predictions. The color scheme in A and B is the same as in Figure 4, and p_{max} is the maximum analytically-predicted pressure in each simulation.

537 comes an open channel between the wellbore and the outer boundary, normalized to the
 538 average value of that time for the slowest-moving fracture (i.e. $t = t_{max}$).

539 The successful replication of the analytical pressure profiles in this section verifies
 540 the model components pertaining to the pressure-velocity-deformation coupling and the
 541 two-way momentum transfer between the fluid and solid phases (drag). Therefore, the only
 542 model component left to verify is the implementation of the capillary force terms during
 543 fracturing of a plastic solid.

544 4.5 Capillary Effects on Fracturing Wellbore Pressure

545 Our fifth verification systematically varies the capillary entry pressure within non-
 546 wetting fracturing systems to quantify its effects on wellbore pressure. For this, we con-
 547 sider two different complementary cases: one where capillary forces are comparable to
 548 their viscous counterparts, and another where they are significantly larger than them. All
 549 parameters are the same as in the previous experiments (Section 4.4) unless otherwise
 550 specified.

551 The first set of experiments expands the previous analysis (Section 4.4) into strongly
 552 non-wetting systems with the addition of a constant capillary pressure jump at the frac-
 553 ture interface imposed by a flat capillary pressure curve ($p_c = p_{c,0} = 1$ to 2 kPa, $\tau_0 =$
 554 $2 \text{ m}^2/\text{s}^2$, $k_0 = 6.7 \times 10^{-12} \text{ m}^2$, $m = 0.05$, and $q = 78 \text{ ml/min}$). In this case, all the
 555 assumptions present in the fracturing analytical solution (Eqn. 41) are satisfied. However,

556 said solution still does not account for capillarity. For constant flow in non-wetting sys-
 557 tems, the addition of a constant capillary entry pressure jump at the fluid-solid interface
 558 would increase the calculated propagation pressure in Eqn. 41 by said value such that
 559 $p_{well}^{new} = p_{well} + p_c$. This effect is exemplified in Figure 7A, where we present the up-
 560 dated analytical results in conjunction with our equivalent numerical results, demonstrating
 561 excellent agreement between them. Note that the predicted linear relationship between
 562 wellbore pressure and capillary entry pressure is not explicitly imposed in the numer-
 563 ical model. On the contrary, it arises naturally from the balance of viscous, capillary, and
 564 structural forces in Eqns. 25-29.

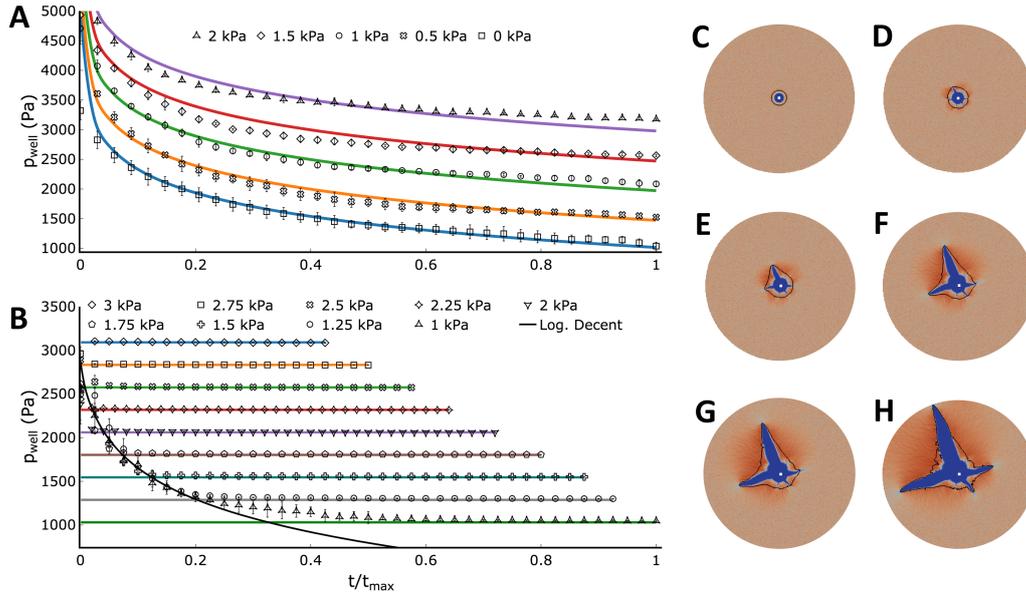


Figure 7. Effect of capillary entry pressure on fracturing wellbore pressure. (A-B) Wellbore pressure as a function of time and entry pressure for low and high permeability systems, respectively. In B, curves at increasingly high pressures were cut off for illustrative purposes and the solid line represents a fitted reference logarithmic pressure descent curve. (C-H) Time evolution of fractured system with a 1 kPa capillary entry pressure and high permeability. (C) Initial fluid invasion ($t/t_{max} < 0$): at early times the wellbore pressure rises rapidly and becomes larger than the entry capillary pressure. The fluid invades the porous formation symmetrically. (D) Fracture initiation ($t/t_{max} = 0$): The wellbore pressure continues to rise until it is larger than the breakdown pressure, at which point small fractures start to form. Fluid invasion continues. (E-F) Fracture propagation ($t/t_{max} > 0 \mid p_{well} > p_{c,0}$): the wellbore pressure drops as fractures propagate. Fluid invasion continues asymmetrically around said fractures. (G) Fluid invasion stops ($t/t_{max} > 0 \mid p_{well} \sim p_{c,0}$): As the wellbore pressure keeps dropping, the entry capillary pressure condition at the porous interface ensures that that wellbore pressure never goes below $p_{c,0}$, at which point fluid invasion stops. (H) Fracture reaches the simulation boundary ($t/t_{max} = 1$). The color convention in Figures C-H is the same as in Figure 4.

565 The second set of experiments modifies the previous experiments by making the
 566 porous medium significantly more permeable, while still maintaining a constant capillary
 567 pressure jump at the fracture interface ($p_c = p_{c,0} = 1$ to 3 kPa, $\tau_0 = 0.2$ m²/s², $k_0 =$
 568 6.7×10^{-11} m², $m = 0.99$, and $q = 78$ ml/min). This results in a set of cases where the
 569 wellbore pressure is increasingly controlled by the capillary pressure drop rather than by
 570 the viscous pressure drop across the fracture and porous formation.

Figure 7 demonstrates precisely this effect. Our simulations show that the wellbore pressure always decays towards the capillary entry pressure once viscous effects are dissipated by fracture growth, i.e., we observe a transition between viscous- and capillary-dominated regimes. At low values of $p_{c,0}$ (< 2500 Pa) the entry pressure is not high enough to prevent fluid flow into the surrounding porous matrix during fracturing (Figure 7B-H). The resulting pressure drop cannot be modeled by the previously presented analytical solution (as it violates the no leak-off assumption), but still follows a logarithm-type curve that is characteristic of flow in fracturing systems. With increasing fracture propagation, the viscous pressure drop decreases until the wellbore pressure equals the entry pressure, which is, by definition, the minimum pressure drop required for fluid flow in highly permeable non-wetting systems. Finally, we note that in cases where capillary entry pressure is high relative to the pressure required to fracture the solid (i.e., at $(p_{c,0} > 2.250$ Pa in the conditions simulated in Fig. 7b), fracturing begins before the wellbore pressure can exceed $p_{c,0}$. This prevents essentially all flow into the porous formation, and the wellbore pressure is immediately stabilized at $\sim p_{c,0}$. For all cases, fractures continue to propagate until they reach the system boundary, at which point the pressure drops rapidly as noted in Section 4.4.

In this section we reduced the inherent complexity of the model’s capillary force terms $F_{c,i}$ (Eqns. 31-32) into a simple set of intuitive verifications. The quantitative agreement between these two analytical cases and their corresponding numerical simulations validate the implementation of the impact of capillary pressure effects on the mechanics of a ductile porous solid within our model.

5 Illustrative Applications

Having verified and tested the model, we now proceed with two illustrations that demonstrate how *hybridBiotInterFoam* enables the simulation of relatively complex coupled multiphase multiscale systems. The following cases serve as illustrative examples of our model’s features and capabilities as well as tutorial cases within the accompanying toolbox.

5.1 Elastic Failure in Coastal Barriers

Coastal barriers are ubiquitous features in coastal infrastructure development. When designed appropriately, these structures can be very effective in regulating water levels and protecting against inclement weather (Morton, 2002). However, accurate prediction of the coupled fluid-solid mechanics of these structures (which can lead to barrier failure) is inherently challenging as it requires modeling large-scale features (waves) while also considering small-scale viscous and capillary interactions within the barrier.

The following case represents the continuation of the three-dimensional coastal barrier illustration presented in Carrillo et al. (2020) with the addition of linear-elastic poromechanics. As such, the simulation was created by initializing a heterogeneous porosity field (with $k_0 = 2 \times 10^{-8}$ m² and $\phi_f = 0.5$) in the shape of a barrier within a 8.3 by 2.7 by 0.25 m rectangular grid (1600 by 540 by 50 cells). The relevant solid mechanics parameters were $E = 5$ MPa, $\nu = 0.45$, and $\rho_s = 2350$ kg/m³. Relative permeabilities and capillary pressures were evaluated through the Van Genuchten model with $m = 0.8$ and $p_{c,0} = 1$ kPa. Before the start of the simulation, the water level was set to partially cover the barrier and then allowed to equilibrate. A single wave was then initialized at $t = 0$. This results in a simulation that exhibits a clear wave absorption cycle that gradually dissipates in time, as seen in Figure 8. Detailed discussion on the fluid mechanics of this problem can be found in Carrillo et al. (2020).

Here, however, we are interested in evaluating the barrier’s propensity to failure. We do this by applying the Von Mises yield criterion, which is commonly used to pre-

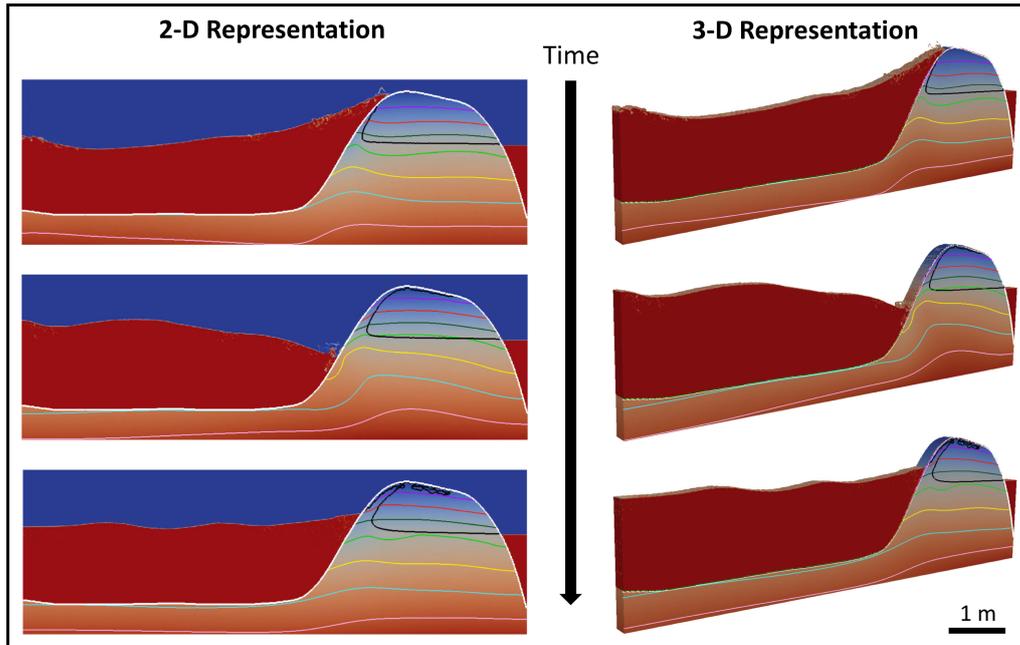


Figure 8. Waves crashing against a poroelastic coastal barrier. Here, the thin black line represents the water-air interface ($\alpha_w = 0.5$) and red-blue colors outside the coastal barrier represent water and air, respectively. Colored contours within the barrier are the calculated Von Mises stresses and are shown in 5 kPa increments in the general downwards direction. Note that the largest stresses are seen during the initial wave crash and increase towards the base of the barrier due to gravitational effects.

620 dict material failure in elastic systems. It states that if the second invariant of the solid's
 621 deviatoric stress (the Von Mises stress) is greater than a critical value (the yield strength)
 622 the material will begin to deform non-elastically (Von Mises, 1913). Although we do not
 623 specify said critical value within our simulations, we can map the time-evolution of Von
 624 Mises stresses within the coastal barrier as a result of a wave absorption cycle (Figure 8).
 625 Our results illustrate the potential utility of our simulation framework in predicting the loca-
 626 tion and time-of-formation of stress induced defects within coastal barrier as a function
 627 of wave characteristics, permeability, and barrier geometry.

628 5.2 Flow-Induced Surface Deformation

629 Surface deformation due to subsurface fluid flow is a common geological phenomenon
 630 occurring in strongly coupled systems and has clear implications in studies related to in-
 631 duced seismicity (Shapiro & Dinske, 2009), CO₂ injection in the subsurface (Morris et
 632 al., 2011), land subsidence (Booker & Carter, 1986), and the formation of dykes and vol-
 633 canoes (Abdelmalak et al., 2012; Mathieu et al., 2008). In order to properly model these
 634 systems, it is necessary to be able to capture the time-evolution of surface uplift, cracks,
 635 and hydraulic fractures, as well as the effects that these features have on the overall flow
 636 field. Here, we use the terms hydraulic fracture vs. crack to refer to solid failure at vs.
 637 away from the injected fluid, respectively.

638 This illustrative case was inspired by the experiments reported by Abdelmalak et
 639 al. (2012), where the authors injected a highly viscous fluid into a dry silica powder in
 640 a Hele-Shaw cell in order to study the impact of hydraulic fractures on surface deforma-
 641 tion, e.g., during the creation of volcanic structures. The system also bears some anal-

642 ogy to situations involving the injection of fluids into subsurface reservoirs, e.g., dur-
 643 ing geologic CO₂ sequestration (Rutqvist, 2012). The base case of our simulations consists
 644 of an impermeable rectangular container (50 by 30 cm, 500 by 300 cells) that is
 645 open to the atmosphere, is partially filled with a dry porous medium ($\phi_s = 0.6 \pm 0.05$,
 646 $\rho_s = 2650 \text{ kg/m}^3$, $k_0 = 5 \times 10^{-11} \text{ m}^2$), and has an injection well at its lower boundary
 647 that injects water at $q = 6.5 \text{ ml/s}$ (Figure 9). To account for irreversible solid deformation,
 648 the porous medium is modeled as a plastic with yield stress $\tau_0 = 0.22 \text{ m}^2/\text{s}^2$. The solid
 649 is represented as impermeable to the invading fluid through the use of the Van Genuchten
 650 model with $m = 0.05$ and $p_c = 0$. Then, using this base case as a standard, we individ-
 651 ually varied each of the main parameters (q , k_0 , τ_0 , m , ϕ_s , μ_{water}) over several simu-
 652 lations in order to model the resulting solid deformation processes: fracturing, cracking,
 653 surface uplift, and subsidence (Figure 9).

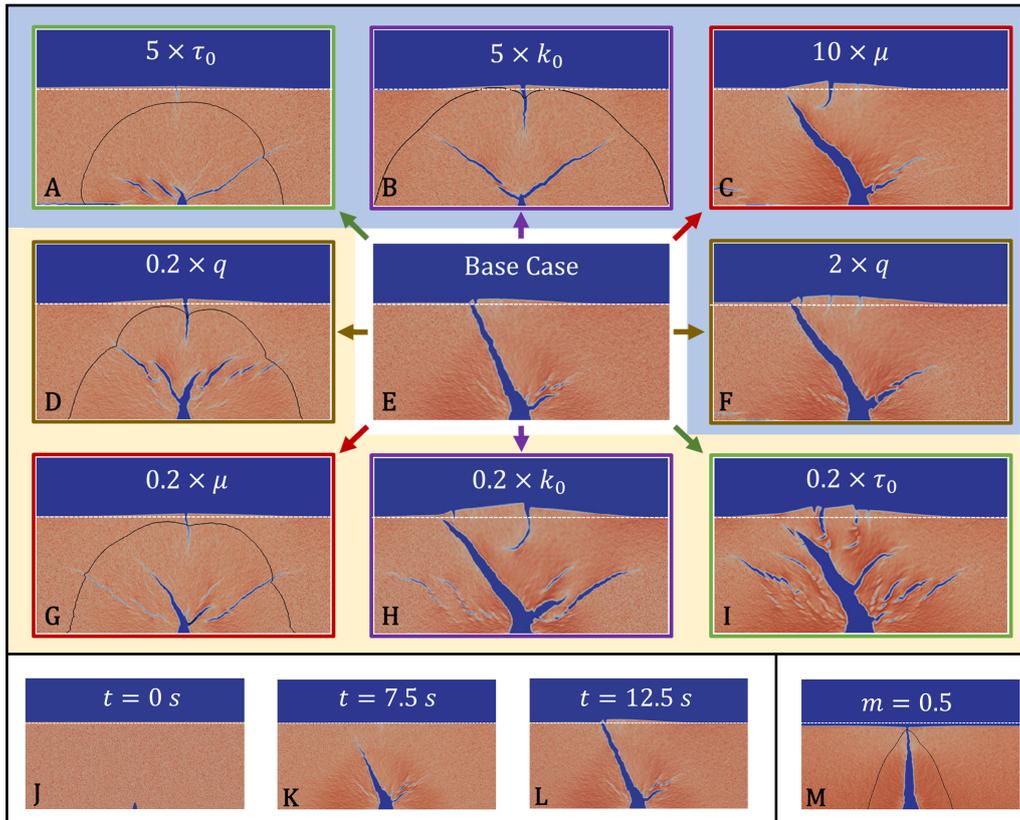


Figure 9. Study of hydraulic fracturing and cracking on surface deformation. (A-I) Representative cases showing the effects of changing permeability k_0 (purple), solid yield stress τ_0 (green), injection rate q (brown), and injected fluid viscosity μ (red) on surface deformation. The blue and yellow subsections contain the results of increasing or decreasing the controlling parameters, respectively. (J-L) Time evolution of the fracturing base case. (M) Surface subsidence example. The difference between the base case (E) and all other simulations is shown in each case's legend. Dotted white lines represent the surface height of the initial solid fraction configuration. Note that the color scheme in all simulations is the same as in Figure 4.

654 The resulting cases demonstrate that cracking (solid failure away from the injected
 655 fluid) is strictly dependent on the number and orientation of existing hydraulic fractures,
 656 as it only occurs when there is more than one fracture branching off from the main injec-
 657 tion point (Figure 9B, C, D, H, and I). This is likely because in cases presenting a single

vertical fracture solid displacement is almost exclusively perpendicular to the fracturing direction, leading to virtually no surface deformation or cracking (Figure 9A, E and M). Contrastingly, the creation of inclined fractures exerts vertical forces on the solid, resulting in surface uplift and crack formation. The above diagram strongly suggests that deformation is controlled by the balance between viscous and structural forces: larger fractures occur within softer solids with higher momentum transfer, and smaller fractures occur in tougher solids with lower momentum transfer. As stated above, a comprehensive examination of the parameters that control solid fracturing will be the focus of an adjacent paper.

In addition to the surface uplift presented above, subsurface subsidence is observed in the simulated system in conditions where the porous solid is rendered permeable to the invading fluid (i.e., $m \gg 0.05$). This phenomenon is not primarily controlled by momentum transfer, but rather by a gravitational effect whereby the displacement of air by water within the porous medium around the advancing hydraulic fracture renders the solid-fluid mixture heavier. Once it is heavy enough to overcome the plastic yield stress, the solid sinks and compresses around the fluid source (Figure 9M).

With these last two illustrative examples, we have shown that our modeling framework is flexible and readily applicable to a large variety of cases within elastic and plastic systems. We invite the interested reader to tune, adapt, and expand the present illustrative simulations, which are included in the accompanying CFD toolbox.

6 Conclusions

We derived, implemented, benchmarked, and applied a novel CFD package for simulation of multiscale multiphase flow within and around deformable porous media. This micro-continuum modeling framework is based on elementary physics and was rigorously derived through the method of volume averaging and asymptotic matching to the multiphase Volume of Fluid equations in solid-free regions and multiphase Biot Theory in porous regions. The result is a single set of partial differential equations that is valid in every simulated grid cell, regardless of content, which obviates the need to define different meshes, domains, or complex boundary conditions within the simulation. The solver's numeric and algorithmic development were also discussed and implemented into *hybridBiotInterFoam*, an open-source package accessible to any interested party.

Throughout this paper and its of predecessors (Carrillo & Bourg, 2019; Carrillo et al., 2020), we show that the Multiphase DBB model can be readily used to model a large variety of systems, from single-phase flow in static porous media, to elastic systems under compression, to viscosity- or capillarity-dominated fracturing systems, all the way up to multiscale wave propagation in poroelastic coastal barriers.

We note, however, that the solver presented here cannot be liberally applied to any porous system, as it comes with the following inherent limitations. First, closure of the system of equations requires appropriate constitutive and parametric relations that describe fluid pressure, permeability, capillarity, and rheology within volume averaged porous regions. Therefore, the assumptions present in each of these models should be carefully considered. Second, volume averaging imposes important length scale restrictions in order to fulfill the scale separation hypothesis, where the pore sizes within the averaging volume must be substantially smaller than the chosen REV, and the REV must be substantially smaller than the macroscopic length scale. Third, as implemented here, the multiphase DBB framework only represents continuum-level elastic or plastic solid mechanics. As such, it cannot be used to model phenomena originating from sub-REV heterogeneities such as fluidization or granular mechanics (Meng et al., 2020), except insofar as they are captured in an averaged manner at the REV scale. Fourth, the use of the CSF as a representation of capillary forces within solid-free regions enforces mass conservation, but it creates a diffuse fluid-fluid interface that may generate spurious and parasitic currents.

708 Finally, although the modeling framework developed here opens up significant new
 709 possibilities in the simulation of coupled fluid-solid mechanics, it also creates a need
 710 for the development of constitutive relations describing the coupling between multiphase
 711 flow and poromechanics. Of particular importance is the formulation of saturation and
 712 deformation-dependent solid rheological models (both plastic and elastic), as well as the
 713 rigorous derivation of the interfacial condition between solid-free and deformable porous
 714 regions. In this paper we proposed a suitable approximation for said boundary condition
 715 based on our single-field formulation, the implementation of a wettability boundary condi-
 716 tion, and the previous work done by Neale and Nader (1974) and Zampogna et al. (2019).
 717 However, the accuracy and validity of such an approximation is still an open question, one
 718 that is at the frontier of our modeling and characterization capabilities (Qin et al., 2020).
 719 The derivation and implementation of said boundary condition, along with the addition of
 720 erosion and chemical reactions into this modeling framework, will be the focus of subse-
 721 quent papers.

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 727 this manuscript is archived at <https://doi.org/10.5281/zenodo.4013969> (Carrillo &
 728 Bourg, 2020) and can also be found at <https://github.com/Franjcf>.

729 A Relative Permeability and Capillary Pressure Models

730 A1 Relative Permeability Models

731 The two relative permeability models used in this paper and implemented in the ac-
 732 companying code depend on defining an effective saturation in order to account for the
 733 presence of irreducible saturations within a porous medium

$$\alpha_{w,eff} = \frac{\alpha_w - \alpha_{w,irr}}{1 - \alpha_{w,irr} - \alpha_{w,irr}}$$

734 here, $\alpha_{w,eff}$ is the wetting fluid's effective saturation, which is the wetting fluid's satura-
 735 tion normalized by each fluid's irreducible saturation $\alpha_{i,irr}$. The Brooks and Corey (1964)
 736 model relates each phase's relative permeability to saturation through the following expres-
 737 sions

$$k_{r,n} = (1 - \alpha_{w,eff})^m$$

$$k_{r,w} = (\alpha_{w,eff})^m$$

738 where m is a non-dimensional coefficient that controls how sensitive the relative perme-
 739 ability is with respect to saturation. The van Genuchten (1980) model calculates relative
 740 permeabilities in the following way

$$k_{r,n} = (1 - \alpha_{w,eff})^{\frac{1}{2}} \left((1 - \alpha_{w,eff})^{\frac{1}{m}} \right)^{2m}$$

$$k_{r,w} = (\alpha_{w,eff})^{\frac{1}{2}} \left(1 - \left(1 - (\alpha_{w,eff})^{\frac{1}{m}} \right)^m \right)^2$$

In this case, m controls how wetting (or non-wetting) the porous medium is to a given wetting (or non-wetting) fluid. High values of m indicate high relative permeabilities for the non-wetting fluid, while low values of m indicate very low relative permeabilities for the same fluid.

A2 Capillary Pressure Models

The implemented capillary pressure models also depend on an effective wetting-fluid saturation $\alpha_{w,pc}$,

$$\alpha_{w,pc} = \frac{\alpha_w - \alpha_{pc,irr}}{\alpha_{pc,max} - \alpha_{pc,irr}}$$

here, $\alpha_{pc,max}$ is the maximum saturation of the wetting fluid and $\alpha_{pc,irr}$ is its irreducible saturation. The Brooks and Corey (1964) model uses the following expression to calculate the capillary pressures within a porous medium

$$p_c = p_{c,0}(\alpha_{w,pc})^{-\beta}$$

where $p_{c,0}$ is the entry capillary pressure, and β is a parameter depending on the pore size distribution. Conversely, the van Genuchten (1980) model calculates the capillary pressure with the following relation

$$p_c = p_{c,0} \left((\alpha_{w,pc})^{-\frac{1}{m}} - 1 \right)^{1-m}$$

B Solid Rheology Models

B1 Hershel-Bulkley Plasticity

A Bingham plastic is a material that deforms only once it is under a sufficiently high stress. After this yield stress is reached, it will deform viscously and irreversibly. The Herschel-Bulkley rheological model combines the properties of a Bingham plastic with a power-law viscosity model, such that said plastic can be shear thinning or shear thickening during deformation. In OpenFOAM® this model is implemented as follows:

$$\boldsymbol{\sigma} = \mu_s^{eff} \left(\nabla \mathbf{U}_s + (\nabla \mathbf{U}_s)^T - \frac{2}{3} \nabla \cdot (\mathbf{U}_s \mathbf{I}) \right)$$

where μ_s^{eff} is the effective solid plastic viscosity, which is then modeled through a power law expression:

$$\mu_s^{eff} = \min \left(\mu_s^0, \frac{\tau}{\eta} + \mu_s \eta^{n-1} \right)$$

where μ_s^0 is the limiting viscosity (set to a large value), τ is the yield stress, μ_s is the viscosity of the solid once the yield stress is overcome, n is the flow index ($n = 1$ for constant viscosity), and η is the shear rate.

B2 Quemada Rheology Model

The Quemada rheology model (Quemada, 1977; Spearman, 2017) is a simple model that accounts for the fact that the average yield stress and effective viscosity of a plastic

769 are functions of the solid fraction. These two quantities are large at high solid fractions
 770 and small at low solid fractions, as described by the following relations

$$\tau = \tau_0 \left(\frac{(\phi_s / \phi_s^{max})}{(1 - \phi_s / \phi_s^{max})} \right)^D$$

$$\mu_s = \frac{\mu_0}{\left(1 - \frac{\phi_s}{\phi_s^{max}}\right)^2}$$

771 here, ϕ_s^{max} is the maximum solid fraction possible (perfect incompressible packing), τ_0 is
 772 the yield stress at $\phi_s = \phi_s^{max}/2$, μ_0 is the viscosity of the fluid where the solid would be
 773 suspended at low solid fractions (high fluid fractions), and D is a scaling parameter based
 774 on the solid's fractal dimension.

775 B3 Linear Elasticity

776 A linear elastic solid assumes that a solid exhibits very small reversible deformations
 777 under stress. Linear elasticity is described by the following relation:

$$\sigma = \mu_s \nabla \mathbf{u}_s + \mu_s (\nabla \mathbf{u}_s)^T + \lambda_s \text{tr}(\nabla \mathbf{u}_s) \mathbf{I}$$

778 where \mathbf{u}_s is the solid displacement vector (not to be confused with solid velocity U_s), and
 779 μ_s and λ_s are the Lamé coefficients (Jasak & Weller, 2000).

780 C Fracturing Instabilities

781 The following figures demonstrate how different fracturing patterns can result from
 782 different solid fraction initializations. Here we set up two sets of four identical experi-
 783 ments. In the first set, the only difference between cases is the value of the standard devia-
 784 tion of their respective normally-distributed solid fraction field (all centered at $\phi_s = 0.64$).
 785 These experiments follow the same simulation setup used for the fracturing case shown in
 786 Figure 4K.

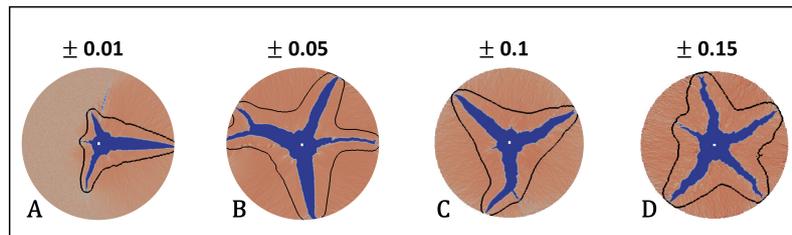


Figure A. Effects of the solid fraction field's standard deviation on fracturing.

787 In the second set of experiments we simulated the base case presented in Figure 9
 788 with different solid fraction profiles picked from the same normal distribution $\phi_s = 0.6 \pm$
 789 0.05 .

790 Figures A and B clearly show that the created fractures are dependent on the initial
 791 solid fraction distribution.

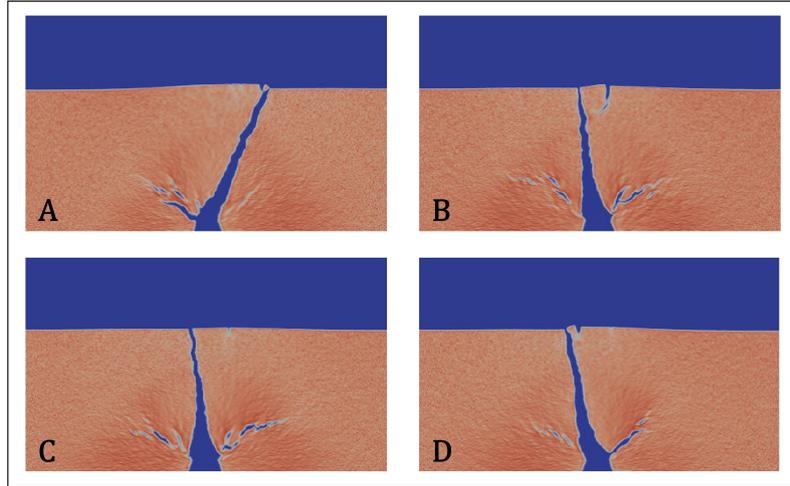


Figure B. Effects of different solid fraction field initializations on fracturing.

792

List of Symbols

793

α_n Saturation of the non-wetting phase

794

α_w Saturation of the wetting phase

795

σ Elastic (or plastic) solid stress tensor in the grid-based domain (Pa)

796

τ Terzaghi stress tensor in the grid-based domain (Pa)

797

$B_{i,k}$ Drag force exerted by phase k on phase i (Pa/m)

798

$D_{i,k}$ Drag force exerted by phase k on phase i (Pa/m)

799

$F_{c,i}$ Surface tension force in the grid-based domain (Pa.m⁻¹)

800

g Gravity vector (m.s⁻²)

801

$n_{i,j}$ Normal vector to the i - j interface in the continuous physical space

802

n_{wall} Normal vector to the porous surface

803

S Single-field fluid viscous stress tensor in the grid-based domain (Pa)

804

t_{wall} Tangent vector to the porous surface

805

U_f Single-field fluid velocity in the grid-based domain (m/s)

806

U_r Relative velocity in the grid-based domain (m/s)

807

U_s Solid velocity in the grid-based domain (m/s)

808

$v_{i,j}$ Velocity of the i - j interface in the continuous physical space (m/s)

809

γ Interfacial tension (Pa.m)

810

μ_f Single-field viscosity (Pa.s)

811

μ_i Viscosity of phase i (Pa.s)

812

ν Poisson's ratio

813

\overline{U}_i^i Phase-averaged velocity of phase i in the grid-based domain (m/s)

814

\overline{U}_i Superficial velocity of phase i in the grid-based domain (m/s)

815

ϕ_f Porosity field

816

ϕ_s Solid fraction field

817

ρ_f Single-field fluid density (kg/m³)

818

ρ_i Density of phase i (kg/m³)

819

τ_0 Plastic yield stress (Pa)

820

θ Surface contact angle

821

$A_{i,j}$ Interfacial area between phase i and j (m²)

822

C_α Parameter for the compression velocity model

823	E	Young's modulus (Pa)
824	I	Identity matrix
825	k	Apparent permeability (m^2)
826	k_0	Absolute permeability (m^2)
827	$k_{r,i}$	Relative permeability with respect to phase i
828	M	Total mobility ($\text{kg}^{-1}\text{m}^3\text{s}^{-1}$)
829	m	Van Genuchten coefficient
830	M_i	Mobility of phase i ($\text{kg}^{-1}\text{m}^3\text{s}^{-1}$)
831	p	Single-field fluid pressure in the grid-based domain (Pa)
832	p_c	Capillary pressure (Pa)
833	$p_{c,0}$	Entry capillary pressure (Pa)
834	V	Volume of the averaging-volume (m^3)
835	V_i	Volume of phase i in the averaging-volume (m^3)

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