

1 **Influence of Solar Irradiation on Nitrous Acid Production in Western U.S. Wildfire Smoke**
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12 **Key Points:**

- 13 • Strong evidence for substantial daytime and nocturnal secondary production of nitrous
14 acid in biomass burning plumes.
- 15 • Daytime conversion of nitrogen dioxide to nitrous acid more than compensated for rapid
16 photolytic losses of HONO in Williams Flats fire.
- 17 • Ground surfaces likely provide the dominant surface area for the heterogenous secondary
18 production of HONO.
19

Abstract

21 Biomass burning is a primary emission source for a host of gas- and aerosol-phase
22 compounds, which can damage environmental and human health. During the FIREX-AQ
23 campaign in July and August of 2019, we measured reactive nitrogen species (NO_x , NO_2 ,
24 HONO, HNO_3 and p-NO_3^-), in wildfire plumes aboard NASA Langley's Mobile Aerosol
25 Characterization Laboratory (MACH-2). Nitrous acid (HONO) and nitric acid (HNO_3) mixing
26 ratios were measured at nominal 5-minute resolution using a dual mist chamber-ion
27 chromatograph from five separate areas of fire in the western US and are the primary focus of
28 this paper. Average HONO mixing ratios were significantly higher in young daytime smoke
29 compared to young nighttime smoke, while no statistical differences were observed between
30 young versus aged smoke during the day or night. In the largest fire sampled during the day, UV-
31 A irradiation was highly correlated ($R^2 = 0.91$) with HONO to nitrogen dioxide (NO_2) ratios
32 indicating that photo-enhanced heterogeneous NO_2 to HONO conversion, likely facilitated by
33 ground surfaces (e.g. soil, foliage, and dust), more than compensated for rapid photolytic loss of
34 HONO.

Plain Language Summary

36 Wildfires negatively impact environmental and human health. Reactive nitrogen pathways are
37 complex and research questions still remain regarding the formation of nitrous acid (HONO) in
38 wildfire plumes. Our observations show that HONO can be made in a smoke plume, even as that
39 smoke plume ages and mixes with surrounding air in both the daytime and nighttime. During the
40 day sunlight rapidly destroys HONO but our data suggests that HONO is being made faster than
41 it is lost in smoke close to the land surface. Moreover, the HONO is being produced by the
42 combination of nitrogen dioxide (NO_2) on surfaces such as soil, leaf foliage and dust.

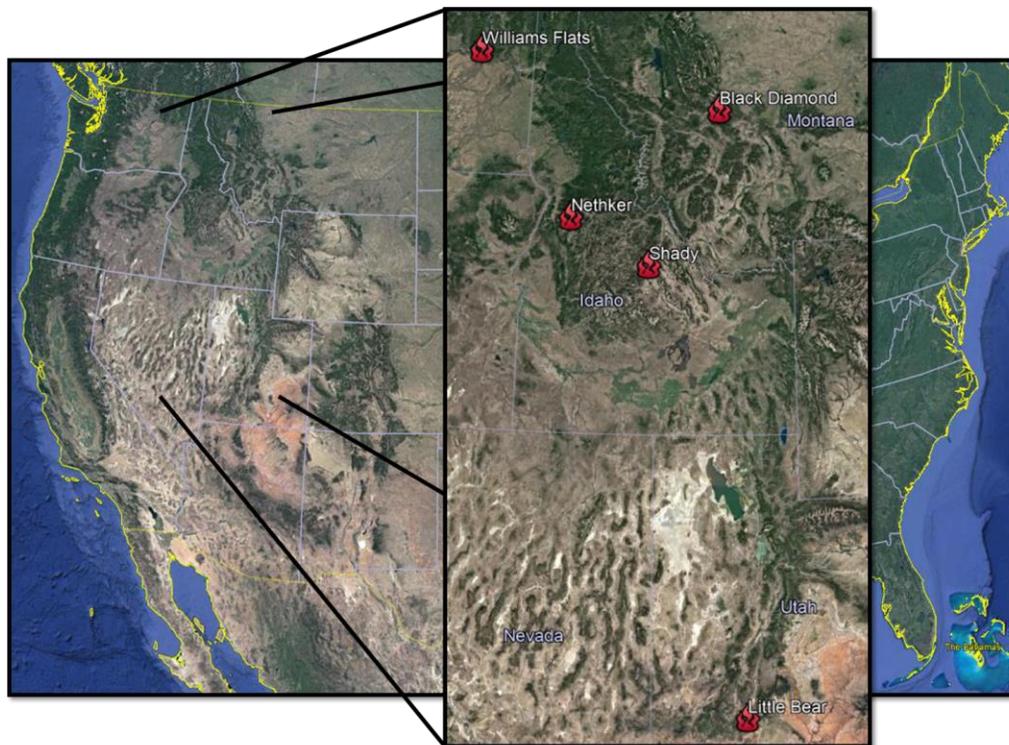
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54 **1. Introduction**

55 Biomass burning (BB) is a primary emission source for a multitude of gas- and aerosol-phase
56 compounds that impact environmental and human health (Grutzen & Andreae, 1990). Climate
57 change induced global temperature increases have resulted in rises in the total acreage burned in
58 the Western United States (Hoover & Hanson, 2019). A single degree centigrade of warming is
59 predicted to result in a two to four times increase in burned areas (Warneke et al., 2015). The
60 pressing need to better understand BB impacts on air quality and climate has motivated a number
61 of recent collaborative field campaigns including the Studies of Emissions, Atmospheric
62 Composition, Clouds and Climate Coupling by Regional Surveys (SEAC⁴RS); the fourth Fire
63 Lab at Missoula Experiment (FLAME-4); the Fire Influence on Regional to Global
64 Environments Experiment Fire Lab (FIREX-FIRELAB); Western wildfire Experiment for Cloud
65 chemistry, Aerosol absorption and Nitrogen (WE-CAN); and the Fire Influence on Regional to
66 Global Environments Experiment-Air Quality Campaign (FIREX-AQ) during which the data for
67 this paper were collected. FIREX-AQ addressed four main research topics: the dependence of
68 emissions on fuel sources, the difference in emissions between flaming and smoldering fires, the
69 mode of atmospheric transport of smoke plumes, and how plumes compositionally change as
70 they age (ESRL Chemical Sciences Division, 2019).

71 During FIREX-AQ collaborators from UNH, NASA Langley Aerosol Research Group
72 Experiment (LARGE), and Brown University deployed NASA Langley's Mobile Aerosol
73 Characterization Laboratory (MACH-2) during July and August of 2019 to measure the
74 compositional and microphysical properties of smoke in the vicinity of active forest fires.
75 Daytime and nighttime measurements were taken at the Shady, Black Diamond, Williams Flats,
76 Nethker and Little Bear fires which spanned four states (Idaho, Montana, Washington and Utah)
77 in the western U.S (Figure 1). Access to these fires varied depending on the availability of roads
78 as well as safety concerns, but in all cases, sampling was performed at the closest allowable
79 locations. MACH-2 is capable of stationary and mobile sampling, with stationary sampling
80 preferred as it limited issues with instrumentation caused by rough road conditions as well as
81 potential interference from sampling anthropogenic roadway emissions. However, the lofted
82 nature of most wildfire smoke plumes during the daytime required sampling while driving to
83 locations where smoke from the Williams Flats and Nethker fires was mixed down to the ground.
84 Most nighttime sampling was conducted while MACH-2 was parked in locations where it was

85 anticipated that smoke would settle into nearby valleys as fire intensity was lower than during
 86 late afternoon peak fire activity.

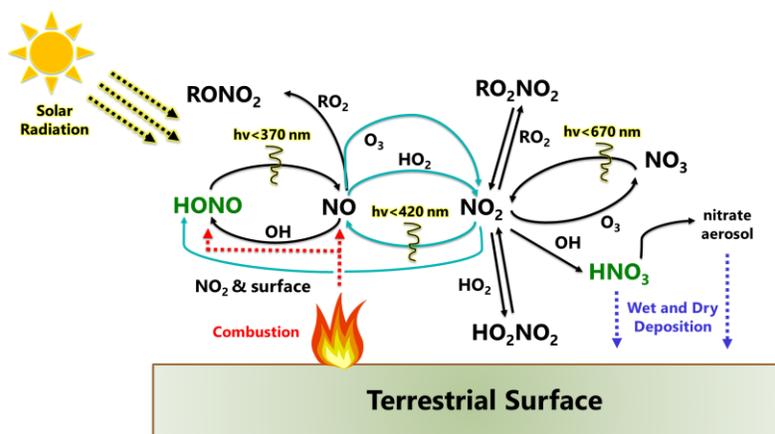


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 88 Figure 1. Locations of fires sampled by MACH-2 during FIREX-AQ.

89 This work examines the observed mixing ratios of nitrous acid (HONO), nitric acid
 90 (HNO₃), and particulate nitrate to establish the daytime and nighttime formation pathways and
 91 oxidative products of nitrogen oxides (NO_x). HONO and NO_x are of particular importance
 92 because their reactivity with other atmospheric species result in a net production of ozone (O₃) in
 93 the troposphere (Platt et al., 1980). Tropospheric O₃ is an oxidant and potent greenhouse gas that
 94 negatively impacts environmental and human health (Seinfeld & Pandis, 2016.). Enhanced O₃
 95 results in reduced photosynthesis and growth of vegetation, impairing the resiliency of
 96 ecosystems (Ainsworth et al., 2012). In humans, O₃ can cause muscle constrictions in airways
 97 leading to trapped air in the alveoli. Limited exposure thus makes breathing more difficult and
 98 results in inflammation and potential damage to airways. Prolonged exposure may result in more
 99 serious conditions such as lung disease, chronic obstructive pulmonary disease and aggravation
 100 of asthma (Nuvolone, Petri, & Voller, 2018).

101 Daytime reactions of HONO and HNO₃ are summarized in Figure 2. Hydroxyl radicals (OH)
 102 initially produced from the photolysis of HONO and O₃ (not shown) play a central role in the

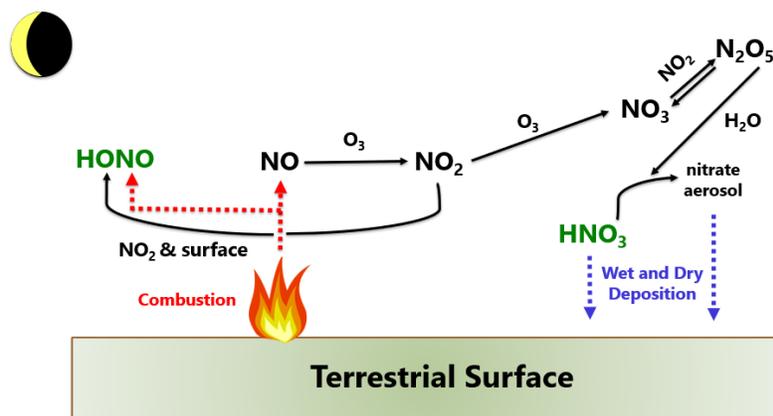
103 photocyclization of HONO and formation of HNO₃. Isotopic results from the 2016 FIREX-
 104 FIRELAB study show that at moderate combustion temperatures (~700 °C) the most preferable
 105 nitrogen reaction pathway post NO formation from BB is NO to NO₂ to HONO shown in light
 106 blue in Figure 2 (Chai et al., 2019). This reaction pathway is initiated via the reaction of NO with
 107 O₃ to form NO₂. NO₂ can then participate in a surface catalyzed heterogeneous reaction to form
 108 HONO which then is photolyzed back to NO. These surface catalyzed reactions, such as those
 109 fostered by quinone species in soil, occur regardless of the presense of solar radiation but are
 110 enhanced during solar flux periods (Scharko et al., 2017). While NO_x cyclization occurs, NO also
 111 reacts with OH forming HONO which rapidly photolyzes back to NO and OH (George et al.,
 112 2015). Peroxy radicals (HO₂ and RO₂) formed from the oxidation of volatile organic compounds
 113 form NO₂ from NO without consuming O₃ and thus lead to a net production of O₃ (George et al.,
 114 2015). HNO₃ is formed when NO₂ reacts with OH and serves as a major sink for NO_x as it is
 115 removed from the atmosphere by wet and dry deposition relatively rapidly (Adon et al., 2013;
 116 Ossohou et al., 2019).



117
 118 Figure 2. Summary of daytime pathways for reactive nitrogen in smoke plumes. Light blue arrows indicate the
 119 preferred pathway for HONO production under moderate combustion temperatures.
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121 Less is known about nighttime reactions involving HONO and HNO₃. At night, plumes settle
 122 down into the boundary layer and have major impacts on air quality close to the terrestrial
 123 surface. This provides a chance for ground-based mobile lab measurements to collect data
 124 inaccessible to airborne platforms due to nighttime flight restrictions and proximity of the plume
 125 to the ground. At night the lack of solar radiation ceases photolysis and therefore halts the NO_x
 126 cycling between NO and NO₂ (Figure 3). This generally leads to increasing concentrations of
 127 HONO until dawn when photolysis starts back up (Stutz et al., 2004). As the nighttime NO_x sink

128 occurs again via wet and dry deposition, these nitrate aerosols stem from the formation of NO_3
 129 radicals via reaction of NO_2 with O_3 , which can further react with NO_2 forming dinitrogen
 130 pentoxide (N_2O_5) (Brown & Stutz, 2012). Water on the surface of aerosol reacts with N_2O_5 to
 131 form gaseous HNO_3 as well as nitrate aerosol (Brown & Stutz, 2012).



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133 Figure 3. Summary of nocturnal pathways for reactive nitrogen in smoke plumes.

134 Here we present evidence for the secondary production of HONO in biomass burning plumes
 135 and explore the surface-based pathways responsible for that production. Results are reported for
 136 the Western U.S. wildfires sampled during FIREX-AQ (Figure 1) and a case study is presented
 137 for the most prominent of those fires (William Flats Fire). Observations from this work are
 138 crucial to understanding the evolution of wildfire-sourced reactive nitrogen, the implications for
 139 air quality in the human breathing zone and highlight the importance of ground-based research
 140 platforms.

141 2. Methodology

142 2.1 Mobile Aerosol Characterization (MACH-2) Laboratory and Sampling Logistics

143 The NASA mobile laboratory known as MACH-2 was equipped with a variety of aerosol and
 144 gas phase sensors, as well as a dual mist chamber-ion chromatograph (MC/IC) system and
 145 denuder collections for measurements of reactive nitrogen species. Further descriptions of the
 146 MACH-2 data set are included in section 3.1. MACH-2 was initially deployed to the Boise,
 147 Idaho area to sample surface level biomass burning plumes from target fires identified by
 148 FIREX-AQ campaign leadership. The majority of collections were conducted at the Williams
 149 Flats and Nethker fires, accounting for approximately 54% and 21% of MC/IC sampling,
 150 respectively. The Williams Flats fire was by far the largest, growing to a maximum of 44,446

151 acres (Table 1). Each fire presented distinct conditions resulting from available fuels, fire state
 152 (smoldering versus flaming) and sampling access. In this work the Williams Flats fire is used as
 153 a case study due to the duration and continuity of sampling. Comparisons between Williams
 154 Flats, Nethker and Little Bear fires are also made. Data for these fires are available in the
 155 Ecosystems Burned Analysis files in the FIREX-AQ data archive (see data acknowledgment
 156 below). The NASA DC-8 research aircraft, as part of FIREX-AQ, also sampled the Williams
 157 Flats fire which allowed for data set comparisons between the two platforms during overlap
 158 periods.

159 Table 1. MC/IC Sampling Dates, Fire Names, Approximate Locations and Maximum Burned Acreage.

Sampling Dates	Fire	Approximate Location	Max Burned Acreage
7/26 – 7/28	Shady	Challis, ID	6,286
8/2	Black Diamond	Lincoln, MT	36
8/3 – 8/7	Williams Flats	Fort Spokane, WA	44,446
8/9 – 8/16	Nethker	Burgdorf, ID	2,360
8/20 – 8/21	Little Bear	Bryce Canyon, UT	2,350

160 2.2 Nitrous and Nitric Acid Sampling Technique

161 HONO and HNO₃ concentrations were measured, in near real time, using a dual mist
 162 chamber-ion chromatograph (MC/IC) system resulting in approximately 1,200 total samples. The
 163 MC/IC sampling/analysis technique has been previously described in detail by Talbot et al.
 164 (1997). The MC/IC was shock mounted within the NASA mobile laboratory to reduce the
 165 potential for malfunctions during transits. The automated custom-built MC/IC is composed of
 166 Dionex analytical components and utilizes automated syringe pumps to inject samples and
 167 standards (Scheuer et al., 2003). The system is closed minimizing the potential for
 168 contamination. Injection volumes were fixed at 5 mL to allow for enough sensitivity to detect the
 169 targeted analytes. The use of electronic suppression via a Dionex Anion Self-Regenerating
 170 Suppressor minimized the background signal. Trifluoroacetate (TFA) added to the ultrapure
 171 sampling water filling the MC served as an internal tracer for sample solution volumes. The
 172 TFA concentration increased 10-20% due to evaporation caused by the sample air flow over a
 173 300 second sample period, allowing the final sample solution volume to be amended within ±
 174 3%. Between the 300 second samples there are 100 second gaps during which samples are
 175 injected into the IC and the MC is refilled before the next run. The IC was calibrated on a daily
 176 basis utilizing a series of eight known standards of aqueous NO₂⁻ and NO₃⁻. Eluents are

177 maintained under a pressurized helium environment. These solutions were replenished after two
178 calibration runs. This system was utilized during the initial phase of FIREX-AQ to analyze
179 smoke during controlled laboratory studies at the FIREX-FIRELAB in the fall of 2016 (Chai et
180 al., 2019) and more recently during the 2018 collaborative WE-CAN Campaign (Chai et al.,
181 2020).

182 The mist chambers collected gas soluble samples. The sample flow was pulled into parallel
183 streams through an externally mounted Fluoropore filter by use of a scroll pump at a controlled
184 nominal rate of 35 L/min. These filters were retained for the aerosol analysis described in section
185 2.3. Samples from one mist chamber were injected into one channel of the IC for HNO₃ and
186 HONO concentration measurements. Samples from the second mist chamber were transferred
187 via a syringe pump to an amber polyethylene sample bottle to be analyzed by Brown University
188 for isotopic analysis of HNO₃ if >10 nmol of nitrate was collected.

189 2.3 Filter Sampling and Extractions

190 A total of 37 Fluoropore filter samples were collected and extracted in the field to determine
191 the relative concentrations of aerosol associated ions (Cl⁻, NO₃⁻, SO₄²⁻, C₂O₄²⁻, Na⁺, NH₄⁺, K⁺,
192 Mg²⁺, Ca²⁺). Collection times per sample ranged from 0.5 – 15 hours depending on factors such
193 as the thickness of the plume, duration of the plume and MC/IC run time. Upon removal, filters
194 were stored in a freezer for preservation until extraction in batches of 16 filters. Due to the
195 hydrophobic nature of the filters, methanol (500 μL) was used to wet the filter allowing for the
196 contents to be dissolved in deionized water (20 mL). The solutions were then transferred to
197 sample bottles and preserved via the addition of chloroform (100 μL). Concentrations of the
198 ionic species were determined using a benchtop IC at the University of New Hampshire.

199 3. Ancillary Data Sets

200 3.1 MACH-2

201 All real time 1 Hz data collected on MACH-2 were averaged and merged to MC/IC run times
202 accounting for the 100 second sample breaks. Gas phase data includes mixing ratios for CO and
203 NO₂. Aerosol data includes PM_{2.5} mass loading and aerosol optical scattering coefficients at 450
204 nm. Furthermore, an AirMar sensor logged GPS coordinates. A full list of utilized merged
205 variables, units and the instruments used for measurements relevant to this analysis are included
206 in Table 2.

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Table 2. Variables, associated units and instruments used for measurements on MACH-2.

Variable	Units	Instrument
CO	ppmv	Los Gatos Research iCOS Spectrometer
NO ₂	ppbv	Los Gatos Research NO ₂ Spectrometer
PM _{2.5}	μg/m ³	Teledyne T640 PM Mass Monitor
AerOpt. Scat.	Mm ⁻¹	Air Photon Nephelometer (450 nm)
Lat. and Long.	Deg.	AirMar G2183 GPS Sensor

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3.2 Physical Distances

210 The Fire to Fuels group at NASA Langley provided the position of the maximum fire
 211 radiative potential (FRP) for the Williams Flats fire at the midpoint of each MC/IC sample,
 212 straight-line physical distances were determined between MACH-2 logged GPS coordinates and
 213 these locations of maximum FRP. FRP detections were combined from the Moderate Resolution
 214 Imaging Spectroradiometer (MODIS) and Visible Infrared Imaging Radiometer Suite (VIIRS)
 215 instruments. MODIS is a key instrument onboard NASA's Terra and Aqua satellites while
 216 VIIRS is housed onboard the Suomi National Polar-Orbiting Partnership spacecraft.

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3.3 UV-A Irradiance

218 In situ calibrated UV-A irradiance measured by an Ultraviolet Multifilter Rotating
 219 Shadowband Radiometer located in Pullman, WA approximately 65 miles from the Williams
 220 Flats fire was used to determine correlations between total UV-A irradiation and HONO/NO₂.
 221 Irradiance at 368 nm was used as a proxy for the full UV-A range (315 – 400 nm) which is
 222 known to drive heterogeneous reactions of NO₂ resulting in the production of HONO. These UV-
 223 A data is publicly available as part of the Colorado State University UV-B Monitoring and
 224 Research Program (Colorado State University, 2020)

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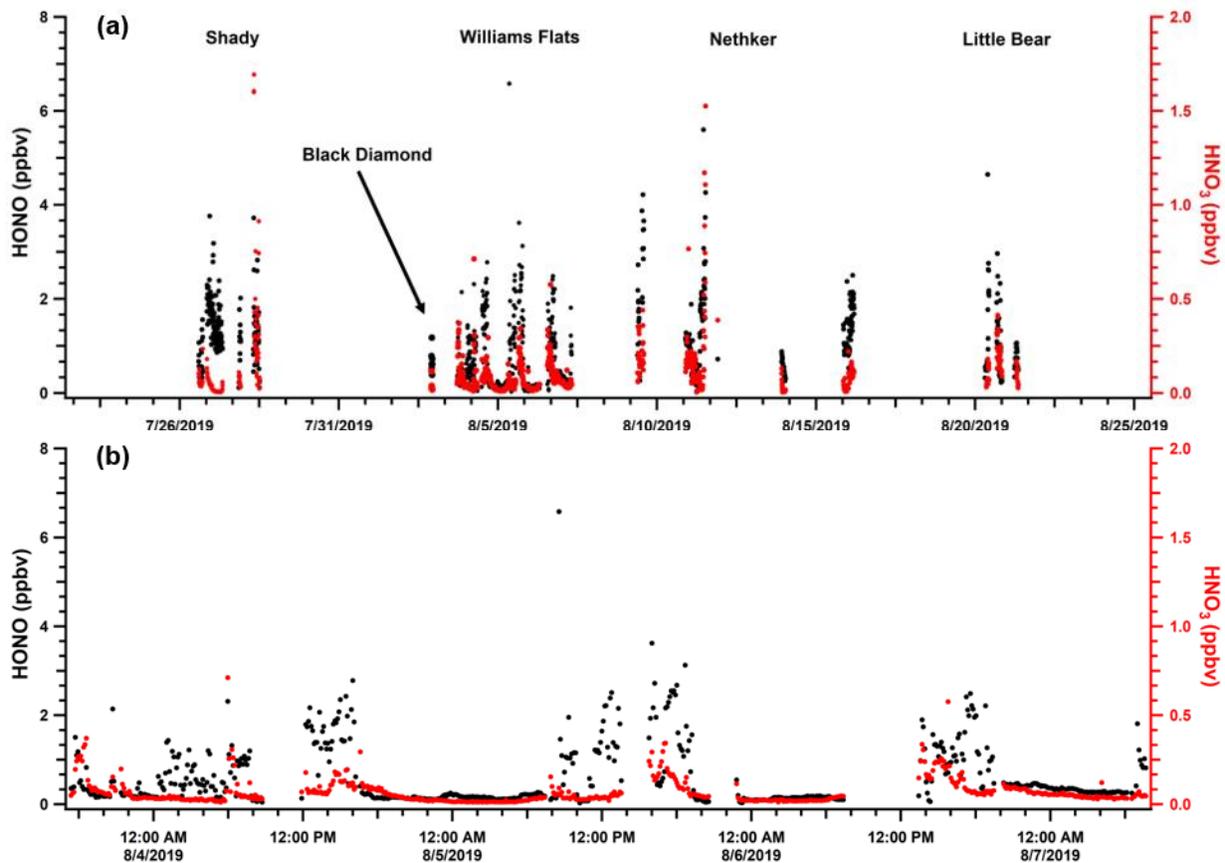
4. Results and Discussion

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4.1 HONO and HNO₃ Variations

227 Across the campaign HONO mixing ratios ranged from a minimum recorded concentration of
 228 0.03 ppbv to a maximum of 6.6 ppbv with an average value of 0.9 ppbv (Figure 4a). The most
 229 continuous and robust sampling was conducted at the Williams Flats fire (Figure 4b). HONO and
 230 HNO₃ exhibited corresponding diurnal variability (Figure 4b). Average mixing ratios were
 231 enhanced during the daytime compared to nocturnal levels at all fires where diurnal sampling
 232 was conducted. Both species display greater mixing ratio variability in the daytime as a result of

233 photolytic processes and the turbidity of the convective mixed layer. The lack of photolysis and
 234 convection produced much smaller and more constant mixing ratios at night. Daytime HONO
 235 mixing ratios were generally an order of magnitude greater than those measured for HNO₃. This
 236 matches the pathways presented in Figure 2, as HONO is directly emitted and has two expected
 237 formation pathways, while BB emissions of HNO₃ are minimal and its single production
 238 pathway in the atmosphere from NO₂ is outcompeted by HONO formation as the plume ages.
 239 Furthermore, HNO₃ undergoes wet and dry deposition and contributes to the formation of nitrate
 240 aerosol (p-NO₃⁻). Analysis of filter collected NO₃⁻ indicates that HNO₃ partitioning favored p-
 241 NO₃⁻ over the gas phase HNO₃ 60% of the time (Figure S1). The photolysis of p-NO₃⁻ presents
 242 another potential pathway for secondary HONO production (Ye et al., 2017). However, isotopic
 243 modeling conducted by collaborators at Brown University indicates that this pathway is a minor
 244 contributor to the overall HONO budget (Chai et al., 2020), thus we focus on other mechanisms
 245 to explain the observations.



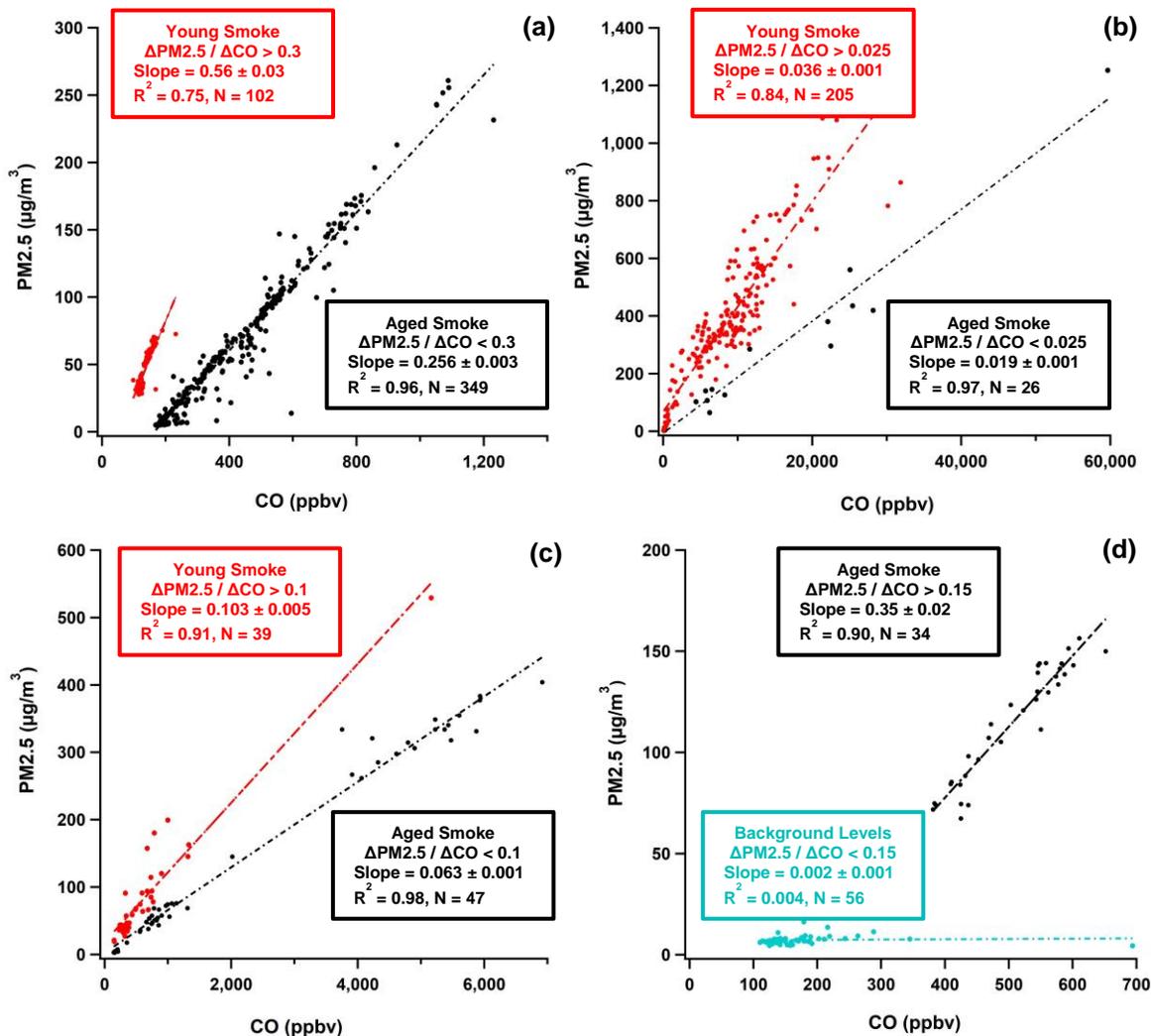
246 Figure 4. Timeseries of HONO and HNO₃ across the five fires sampled by the MACH-2 mobile laboratory
 247 platform during the 2019 FIREX-AQ campaign (a) and expanded plot for the Williams Flats fire (b).

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256 4.2 Young vs. Aged Smoke

257 A major FIREX-AQ goal was to quantify compositional differences between young and aged
258 smokes plumes. In these classifications, young plumes are more freshly emitted and aged plumes
259 correspond to a greater extent of chemical processing. A recently published paper by Selimovic
260 et al. (2020) suggests use of enhancement ratios between PM_{2.5} and CO as a valid method to bin
261 wildfire plume age. Applying that approach here, scatter plots of PM_{2.5} versus CO exhibit two
262 distinct highly linear populations for the Williams Flats, Nethker and Little Bear fires (Figure 5).
263 These differing fire conditions resulted in distinct ratios used to separate the populations. Of the
264 two, the population with the steeper slope (i.e., more PM_{2.5} (y-axis) loading per CO (x-axis)
265 levels) corresponds to young smoke while the other population is representative of aged smoke
266 (Figure 5a-c). The Shady fire proved to be a useful example for comparing background to smoke
267 impacts (Figure 5d). The plume from the Shady fire was very well defined and thus allowed for a
268 clean assessment of background levels which exhibit no correlation ($R^2 = 0.004$) between PM_{2.5}
269 and CO. Once the plume front moved into the valley where we were sampling, PM_{2.5} was
270 elevated and is tightly correlated with CO ($R^2 = 0.90$). This divergence in correlations highlights
271 that virtually all samples taken at the Williams Flats, Nethker and Little Bear fires were smoke
272 impacted. In each case $\Delta\text{PM}_{2.5}$ (enhancement from background PM_{2.5}) to ΔCO (enhancement
273 from background CO) ratios prove to be powerful metrics allowing the categorization and thus
274 quantification of differences present between younger and more processed plumes.

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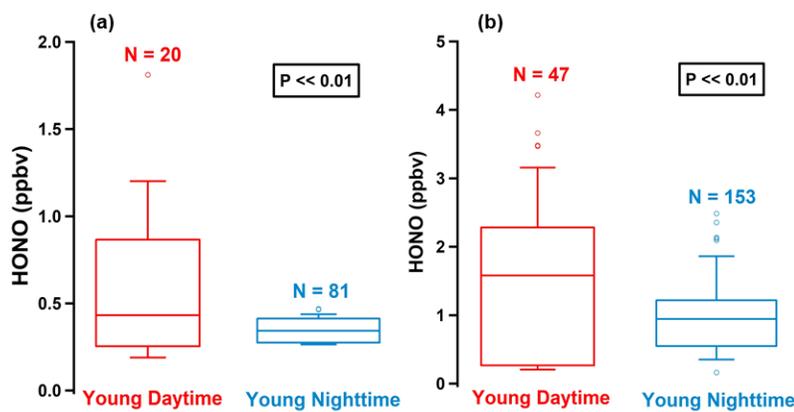
280 Figure 5. Young vs.aged $\Delta PM_{2.5}/\Delta CO$ binning for Williams Flats (a), Nethker (b), Little Bear (c) and Shady (d) fires. Williams
 281 Flats, Nethker and Little Bear all exhibit two strongly correlated populations (with the red dashed line indicative of “young”
 282 smoke and the black dashed line indicative of “aged” smoke). The Shady fire is included as an example of background levels.

283 To validate the $\Delta PM_{2.5}$ to ΔCO ratio as an indicator of processing, all values of the ratio
 284 measured in the Williams Flats smoke plume are plotted as a function of the physical distance
 285 between MACH-2 and the maximum FRP measured by MODIS and VIIRS (Figure S2). The red
 286 dashed line indicates the $\Delta PM_{2.5}$ to ΔCO ratio of 0.3 used to distinguish young from aged smoke
 287 for the Williams Flats data herein. All ratios categorized as being young were measured at a
 288 straight-line distance of 8 km while those categorized as aged exceeded 10 km spanning out
 289 towards 70 km at maximum. Note, the categorization between young and aged smoke is relative,
 290 hence the ratio applied for this distinction varies from one fire data set to the next (Figure 5).

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4.3 Implications of Secondary HONO Production

It is well established that the diurnal cycle impacts BB-sourced HONO concentrations. Surprisingly our results indicate that for Williams Flats and Nethker young daytime smoke contained significantly higher concentrations of HONO than young nighttime smoke (Figure 6a-b). Higher nighttime concentrations were expected due to a lack of nocturnal HONO loss pathways and the stability of the shallow nocturnal boundary layer. At Nethker, enhanced HONO in young daytime smoke can be attributed to differences in fire state between day and night. More flaming during the day results in proportionally greater primary HONO production. This primary production was accessible to MACH-2 at the Nethker fire, as sampling occurred where vegetation was actively flaming within 15 m of the mobile lab. Day/night differences for the Williams Flats smoke were less obvious. In the presence of solar irradiation, primary HONO has a lifetime ranging between 10 – 20 minutes (Spataro & Ianniello, 2014). At a straight-line distance of 8 km the average windspeed would need to exceed 24 km/hr for primary HONO to be measurable by MACH-2. Real-time weather data for the region indicates a maximum recorded windspeed of 24 km/hr during the sampling period. Therefore, it is unlikely that enhanced daytime primary HONO emission caused the difference at the Williams Flats fire. CO observations suggest that plume thickness was the dominant factor causing the statistical difference in HONO loading, as much higher CO was measured during the daytime (daytime average = 368 ppbv, nighttime average = 285 ppbv).



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Figure 6. The impact of diurnal cycle on average HONO concentrations in young smoke plumes for the Williams Flats (a) and Nethker (b) fires. Statistical P values are based off a 95% confidence interval and N values represent population sizes.

313 The lack of statistical differences between average HONO mixing ratios in young and aged
 314 smoke from a given fire is a strong indicator of substantial secondary HONO production under
 315 daytime and nocturnal conditions as shown by examples from the Williams Flats and Little Bear
 316 fires, given the short lifetime of HONO during both day (10-20 min) and night (1-2 hr) (Figure
 317 7a-c). This ground-based observation contrasts with previous findings from airborne platforms.

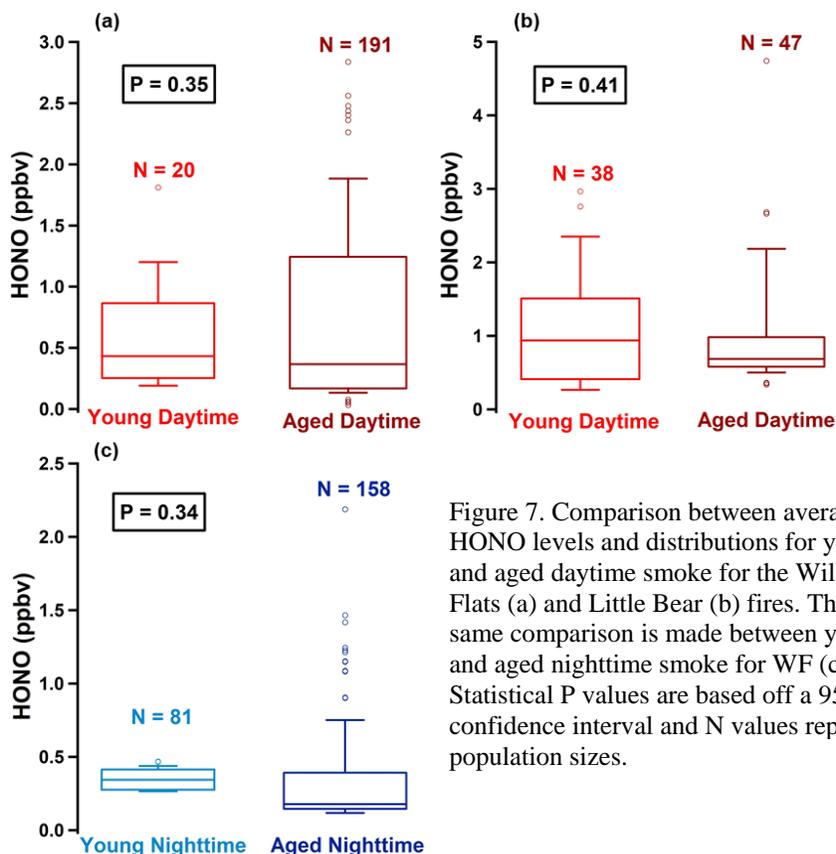


Figure 7. Comparison between average HONO levels and distributions for young and aged daytime smoke for the Williams Flats (a) and Little Bear (b) fires. The same comparison is made between young and aged nighttime smoke for WF (c). Statistical P values are based off a 95% confidence interval and N values represent population sizes.

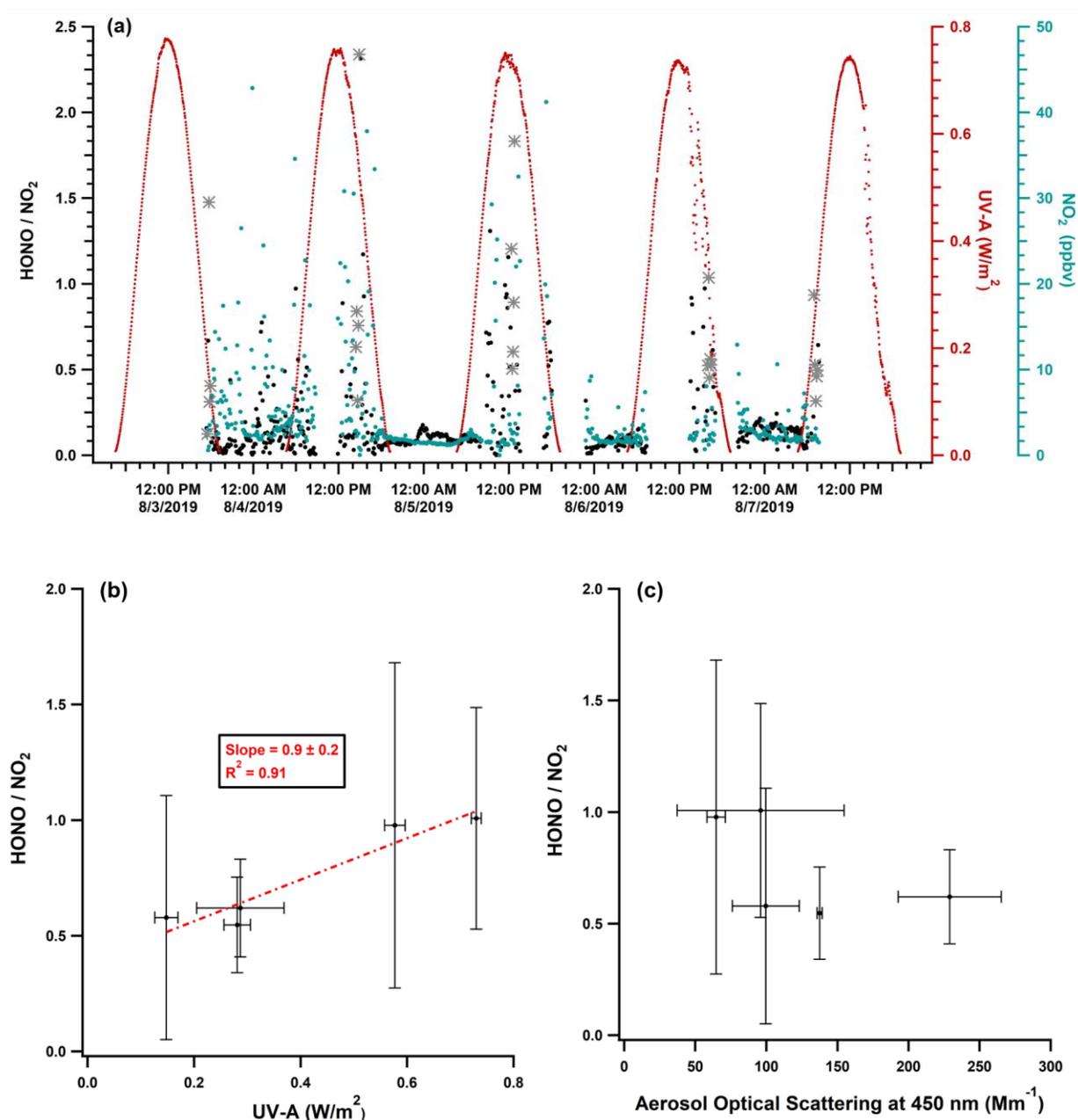
318 No substantial evidence was found for secondary HONO production in BB plumes sampled
 319 from the NOAA P3 aircraft during the Southeast Nexus Experiment (Neuman et al., 2016).
 320 Interestingly, results from NASA's DC-8, which took measurements of the Williams Flats fire at
 321 higher altitudes, also provide no clear evidence for substantial secondary production of HONO.
 322 However, there are several factors which could contribute to the differences in the budget and
 323 cycling of HONO between the ground-based MACH-2 mobile laboratory and the DC-8.

324 The atmospheric environment could be a substantial factor. In near-source plumes, which the
 325 DC-8 had better access to, the conditions are NO_x limited due to a large quantity of VOCs and
 326 SVOCs that store reactive nitrogen as temporary sinks. The secondary production of HONO is
 327 very limited in these NO_x deficient environments and could explain why DC-8 results do not

328 indicate secondary HONO production. Stark differences in measured NO₂ mixing ratios support
329 this hypothesis (Figure S3). Furthermore, heterogenous conversion on ground surfaces (soil,
330 foliage and dust) are also suggested as dominant pathways for the secondary production of
331 HONO (see section 4.4 below). These ground-based conversions would enhance HONO levels in
332 the boundary layer airmasses sampled by MACH-2. Overall, the divergent results indicate the
333 importance of ground-based studies and present an argument for mobile laboratories to
334 collaborate with aircraft during future campaigns.

335 4.4 Influence of UV-A Solar Irradiation and Aerosol Surface Area on Heterogenous NO₂ 336 to HONO Conversions

337 Heterogenous conversion of NO₂ to HONO has been represented in the literature by use of
338 HONO/NO₂ ratios (Li et al., 2012; Nie et al., 2015). The diurnal variation of these ratios exhibit
339 alignment with daily changes in UV-A solar flux across five days for the Williams Flats fire.
340 HONO/NO₂ declines in magnitude along the falling edge of UV-A on August 3rd and can be seen
341 increasing along the rising edge on the morning of August 7th with maxima aligning with peak
342 solar flux on August 5th (Figure 8a). This dependency was further probed by averaging ratios for
343 approximately 30-minute intervals across each day. The maximum 30-minute HONO/NO₂
344 average each day, which occurred at different times due to the MC/IC duty cycle, show a clear
345 dependency ($R^2 = 0.91$) on corresponding averaged UV-A flux (Figure 8b). This relationship
346 indicates that heterogenous conversions of NO₂ to HONO more than compensate for rapid
347 photolytic losses of HONO even when UV-A flux is at its highest intensity. A similar
348 observation was made in an urban setting by Wong et al. (2012) during the 2009 Study of
349 Houston Atmospheric Radical Precursor (SHARP) experiment. The resulting net HONO
350 production supports the lack of statistical differences in mean HONO mixing ratios between
351 young and aged daytime smoke for the William's Flats and Little Bear fires (Figure 7a-b).



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Figure 8. Time series HONO to NO₂ ratios (left), UV-A irradiation and NO₂ mixing ratios (right) (a), as well as the dependence of those HONO to NO₂ ratios on UV-A irradiation across five daytime collection periods at the Williams Flats fire (b). The same NO₂ to HONO conversion efficiencies are also included as a function of aerosol optical scattering at 450 nm (c). The gray stars in panel (a) indicate the HONO/NO₂ data points utilized for the averaging based analysis in panels (b) and (c). Error bars represent the standard deviations (1σ) over the averaged intervals.

361 A comprehensive review of HONO sources notes that aerosols can significantly contribute to
362 the secondary production of HONO if the aerosol loading is substantial, however ground sources
363 were found to overshadow aerosol impacts in urban and forest studies (Spataro & Ianniello,
364 2014). These conversions include those from redox chemistry facilitated by quinone species
365 present in humic acid (Han et al., 2016; Scharko et al., 2017), surface reactions promoted by
366 mineral dust (Ma et al., 2013; Ndour et al., 2008; Chai et al., 2020), as well as the photolysis of
367 adsorbed nitric acid on the surface of foliage (Zhou et al., 2011). Anthropogenic surfaces could
368 also be a contributor, but those surface site reactions have not been well constrained. In contrast
369 to urban and forest environments the aerosol loading and corresponding surface area in biomass
370 burning plumes is considerably higher. Thus, the contribution of aerosol to secondary HONO
371 production was explored using the same averaging method as was done for Figure 8b. No clear
372 relationship is present between HONO/NO₂ and aerosol optical scattering at 450 nm, which is
373 used here as a proxy for mass loading (Figure 8c). In fact, the highest HONO/NO₂ ratios
374 correspond with the lowest two scattering averages. This indicates that the terrestrial surface
375 provides the vast majority of available sites facilitating the daytime heterogeneous HONO
376 production. Under higher aerosol loading conditions there may be a dampening effect where
377 aerosol optical extinction limits the UV-A flux reaching the terrestrial surface, potentially
378 inhibiting the extent to which heterogeneous reactions occur there. It is possible that aerosol
379 surface sites also contribute to secondary HONO production although our data did not provide
380 evidence for this. More work is required to further assess the quantitative impact of aerosol
381 loading on ground-based secondary HONO formation.

382 **5. Conclusions**

383 Over the course of FIREX-AQ, MC/IC samples collected on board NASA's MACH-2
384 provided near-real time concentrations of HONO and HNO₃ for five distinct fires in the Western
385 U.S. This variety of fires offered a chance for high-level comparisons between the
386 aforementioned species. Relative plume age was determined by use of enhancement ratios
387 between PM_{2.5} and CO. Average HONO concentrations were significantly higher in young
388 daytime than young nighttime smoke while no statistical differences were observed upon
389 comparing young and aged smoke. The lack of statistical differences between young and aged
390 smoke under both daytime and nocturnal conditions is evidence that substantial secondary
391 HONO production is associated with ground-level smoke plumes.

392 The Williams Flats fire presented a unique opportunity for a case study given the fire's size,
393 the extent of sampling and overlap with the DC-8 platform. Across five days, heterogeneous
394 conversions of NO₂ to HONO are strongly correlated with UV-A flux indicating that these
395 conversions, likely occurring on the ground, more than compensate for rapid photolytic losses of
396 HONO.

397 Overall, this work presents strong evidence for secondary production of HONO occurring
398 through heterogeneous NO₂ surface site reactions. However, more studies utilizing ground-based
399 platforms are needed to verify whether the results from the Williams Flats fire case study are
400 universal to all wildfires. Furthermore, future collaborations between mobile laboratories such as
401 MACH-2 and aircraft will be essential for a comprehensive understanding of reactive nitrogen
402 evolution in smoke plumes and their implications for air quality.

403 **6. Data Availability Statement**

404 All data used for this analysis, except for UV-A solar irradiation, is publicly available at the
405 NASA Airborne Science Data for Atmospheric Composition database
406 (doi:10.5067/SUBORBITAL/FIREXAQ2019/DATA001). The UV-A data is publicly available
407 at the Colorado State UV-B Monitoring and Research Program.

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