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*Geophysical Research Letters*

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Supporting Information for

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**Months-long crustal deformation driven by aseismic slips**

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**and pore pressure transients triggered by local and regional**

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**earthquakes**

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19 **Contents of this file**

20           Figures S1 to S6

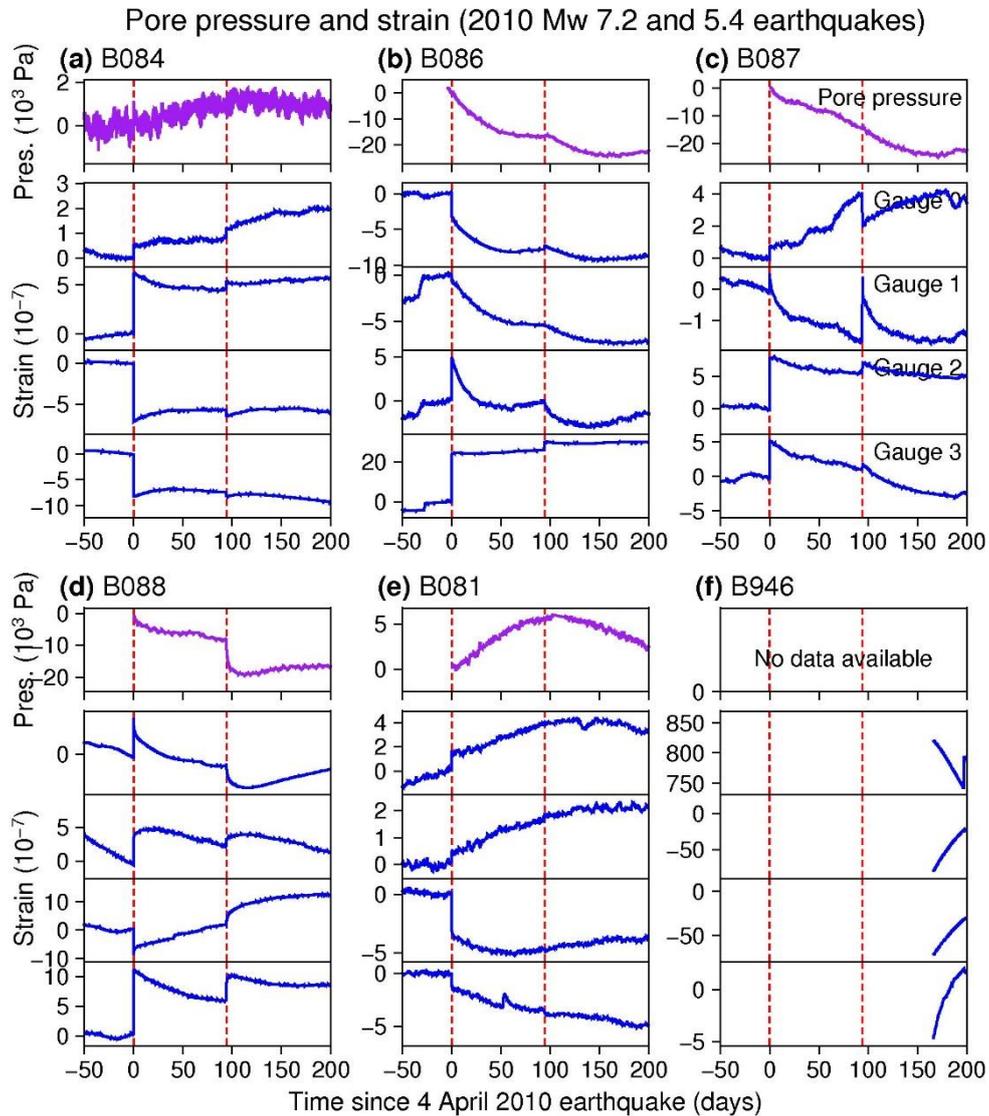
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22 **Introduction**

23 This supporting information (SI) provides 6 figures, including: **(1)** Pore pressure and  
24 strain before and after the 4 April 2010 El Mayor-Cucapah and 7 July 2010 Collins  
25 Valley earthquakes (Figure S1), the 11 March 2013 Borrego range earthquake (Figure  
26 S2), and the 10 June 2016 Borrego Springs earthquake (Figure S3), **(2)** cumulative static  
27 strains at strainmeters produced by the aftershocks of the 10 June 2016 earthquake  
28 (Figure S4), **(3)** L-curve analysis for the strain fitting (Figure S5), and **(4)** comparison  
29 between the postseismic strains observed at the strainmeters used for qualitative  
30 constraint and those produced by the best-fitting aseismic slip (Figure S6).

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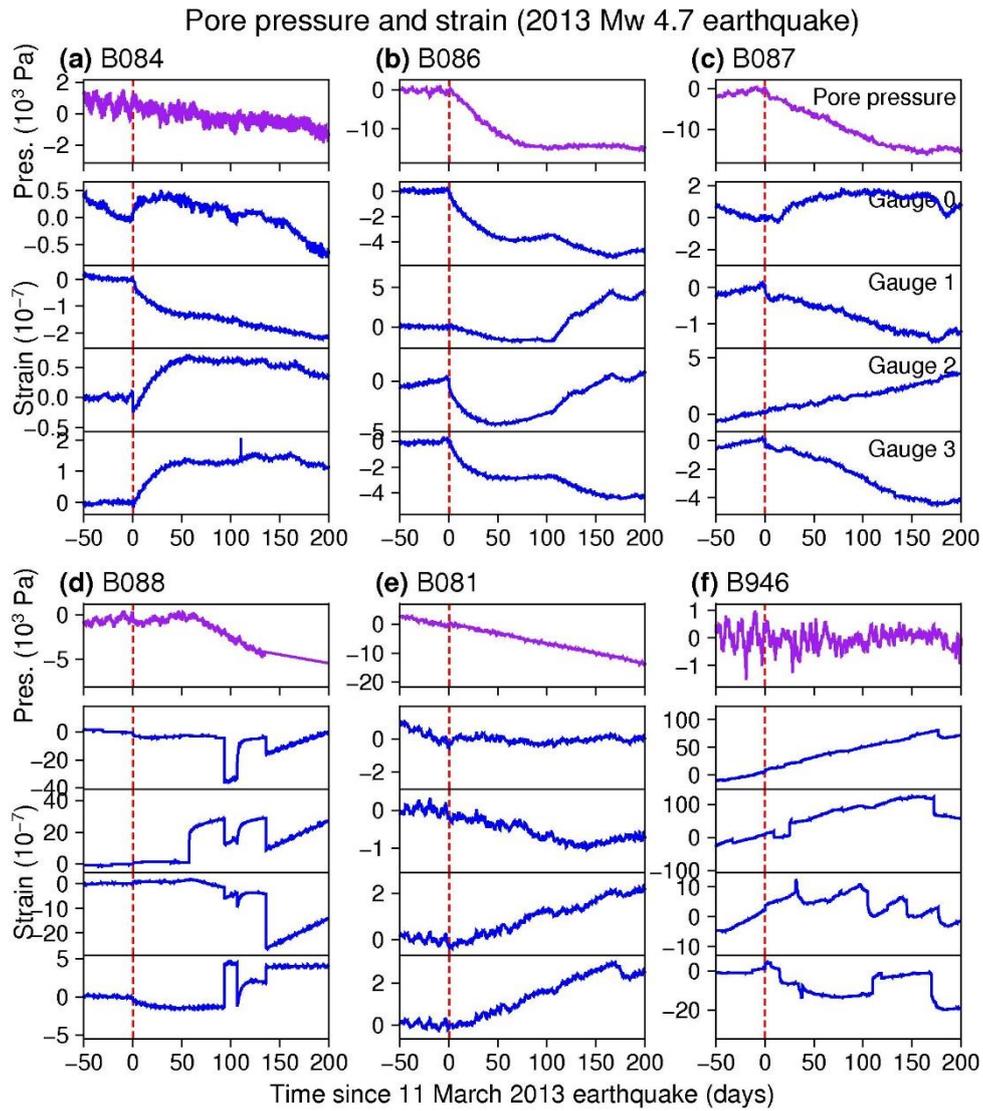


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35 **Figure S1. Pore pressure (purple curves) and strain (blue curves) before and after**  
 36 **the 4 April 2010 El Mayor-Cucapah (Mw 7.2) and 7 July 2010 Collins Valley (Mw**  
 37 **5.4) earthquakes observed at strainmeters B084, B086, B087, B088, B081 and B946.**

38 The vertical dashed lines in each panel mark the occurrence times of the two  
 39 earthquakes.

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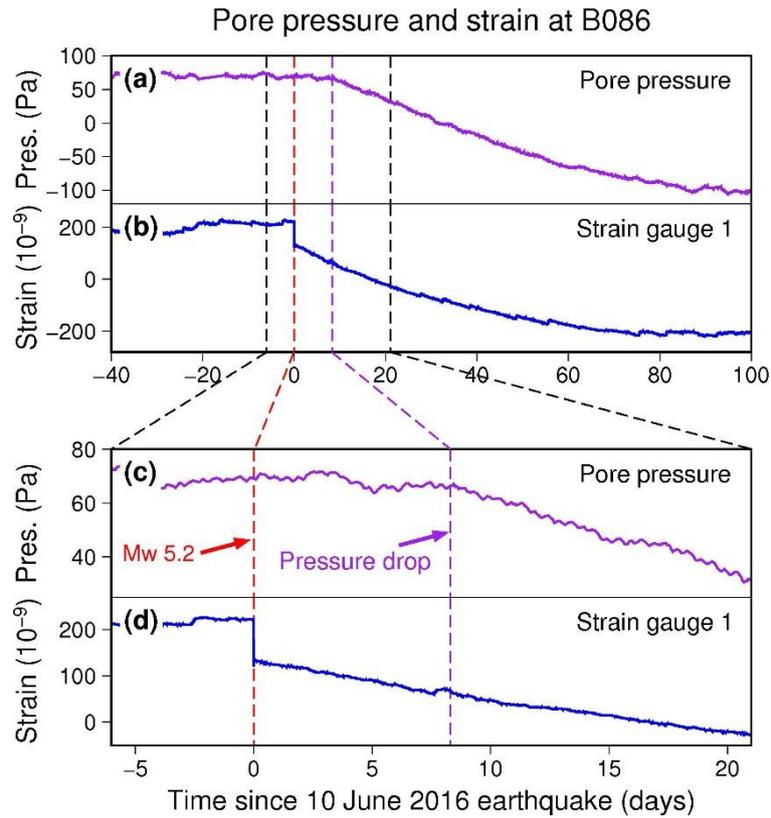


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42 **Figure S2. Same as Figure S1, except for the 11 March 2013 Borrego range (Mw**

43 **4.7) earthquake.**

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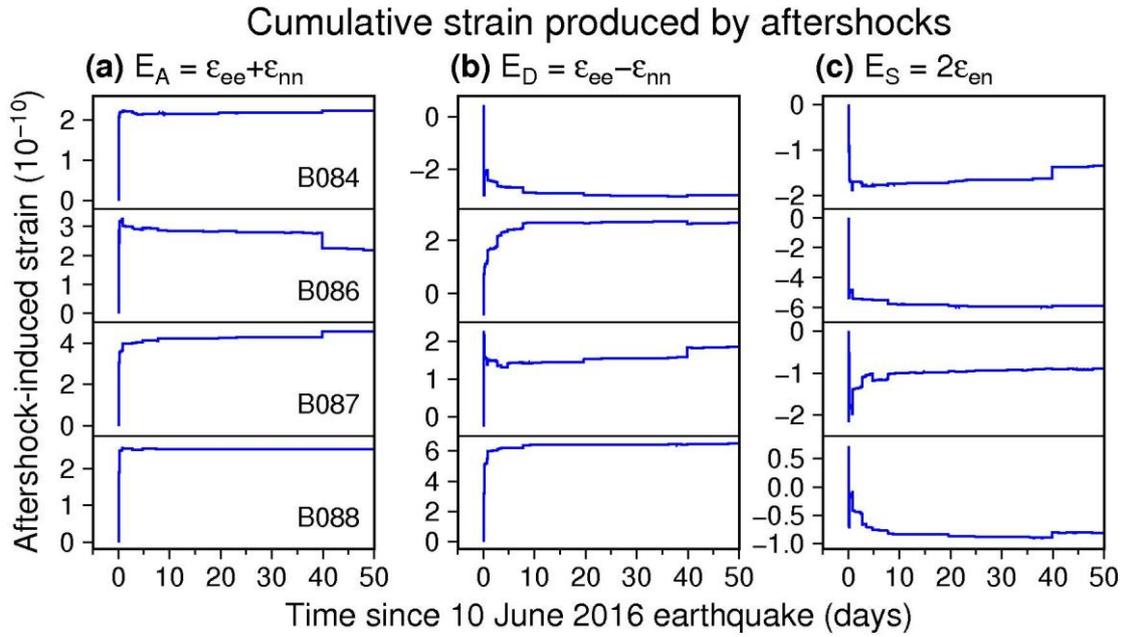
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46 **Figure S3. Pore pressure and strain (gauge 1) observed at B086 before and after**  
 47 **the 10 June 2016 Mw 5.2 earthquake.** Note that the strain changes immediately after  
 48 the earthquake (red dashed line), while the pore pressure remains at background level  
 49 for about 8 days after the earthquake before exhibiting a significant decrease (purple  
 50 dashed line).

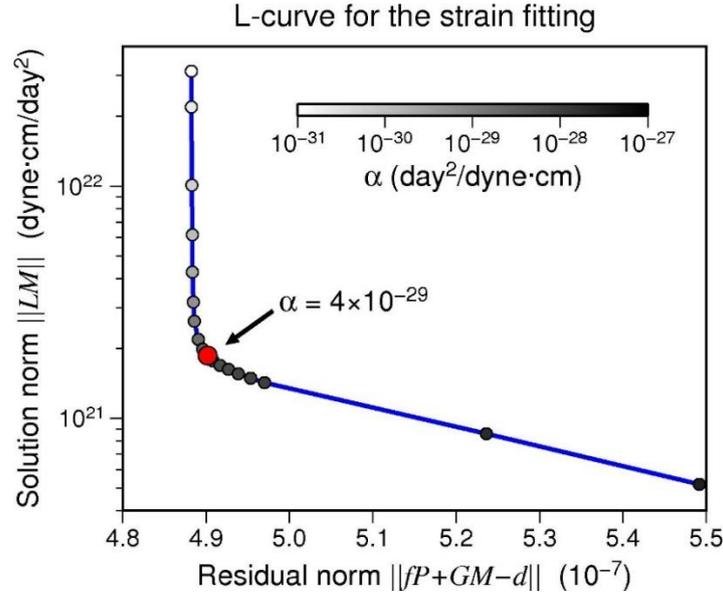
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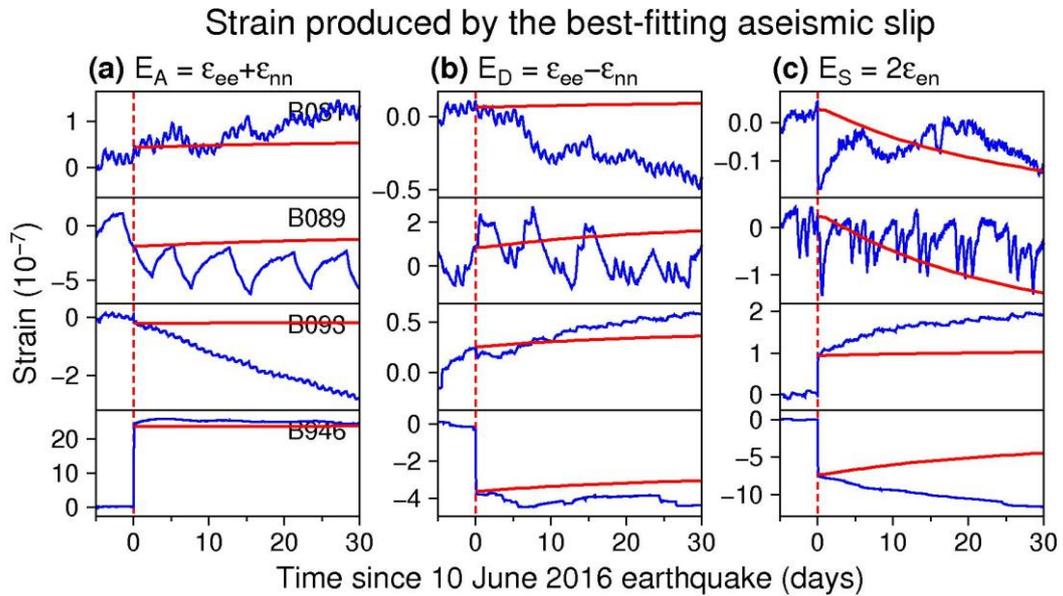
**Figure S4. Cumulative static strains at four strainmeters produced by the aftershocks of the 10 June 2016 Mw 5.2 earthquake.** The strains are calculated using Okada's method (Okada, 1985) in an elastic half-space Earth model, with the elastic moduli being  $\lambda = 37.2$  GPa and  $\mu = 36.8$  GPa (Laske et al., 2013). The hypocenters and focal mechanisms of the aftershocks are from the Southern California Earthquake Data Center (<https://service.scedc.caltech.edu/eq-catalogs/FMsearch.php>), with the focal mechanisms of the aftershocks with FMQ (focal mechanism quality) of C, D or F replaced by that of the mainshock (strike/dip/rake =  $306^\circ/68^\circ/179^\circ$ ).



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66 **Figure S5. L-curve analysis for the strain fitting based on Equation (2).** The corner  
 67 at  $\alpha = 4 \times 10^{-29}$  day<sup>2</sup>/dyne · cm is used as the smoothing parameter in the strain  
 68 fitting.

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71 **Figure S6. Comparison between the postseismic strains observed at the**  
 72 **strainmeters used for qualitative constraint based on Equation (1) (blue) and those**  
 73 **produced by the best-fitting aseismic slip (red).**

74 **References**

75 Laske, G., Masters., G., Ma, Z., & Pasyanos, M. (2013). *Update on CRUST1.0 - A 1-*

76 *degree global model of Earth's crust*. Paper presented at the EGU.

77 Okada, Y. (1985). Surface deformation due to shear and tensile faults in a half-space.

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