

Muon Imaging of Volcanic Conduit Explains Link between Eruption Frequency and Ground Deformation

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Key Points:

- An inverse correlation was observed between eruption frequency and ground surface deformation of Sakurajima volcano
- Beneath the crater, muon imaging visualized that the mass density increased during ground inflation and decreased during ground deflation
- Plugging of the conduit with dense, stiff magma during recharge caused inflation during quiescent periods

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Abstract

An inverse correlation was observed between eruption frequency and ground surface deformation of Sakurajima volcano (Japan) during November 2018 to April 2021. Over the same period, the mass density of magma in the upper conduit of the active crater was monitored via muography. Mass density increased significantly during inflation, when eruption frequency was low, and decreased during deflation, when eruption frequency was high. On the basis of the muography data, we find that periods of low eruption frequency are associated with the formation of a dense plug in the shallow conduit, which we infer caused inflation of the edifice by trapping pressurized magmatic gas. Conversely, periods of high eruption frequency are associated with the absence of a dense plug, which we infer allows gas to escape, leading to deflation. Muography thus reveals the in-conduit physical mechanism for the observed correlation, with implications for interpretation of deformation at other volcanoes.

Plain Language Summary

Monitoring changes in the ground-surface level at volcanoes can indicate changes in the likelihood of an eruption. In order to be confident of interpretations of monitoring data, we must understand the physical cause of the observed changes and their relation to ongoing volcanic phenomena. We measured changes in mass density in the magmatic plumbing system beneath Sakurajima volcano (Japan) using muons, which are cosmic particles that pass easily through rock. We detected muons that passed through Sakurajima, and constructed images of the density during dormant and active periods, and used this to explain the link between changes in the ground surface level and the number of eruptions. We found that the mass density increased during the dormant periods, when the ground surface of the volcano was high, and the mass density decreased during periods of frequent eruption, when the ground surface was low. Muon imaging helped to reveal the hidden volcanic processes: the volcano plugged during the dormant periods and recharge of fresh magma into the conduit resulted in uplift of the volcano's surface. During periods of eruption, gas pockets formed within the plug, driving explosions, and the release of pressure resulted in the downlift of the volcano's surface.

1 Introduction

Forecasting the location, onset, cessation and style of impending volcanic eruptions facilitates effective mitigation of the impact of associated hazards (Sparks, 2003; Poland & Anderson, 2020). Active volcanism is driven by the subsurface evolution and movement of magmatic materials, which may induce seismicity (Chouet, 1996; Brenguier et al., 2008; Dempsey et al., 2020), ground deformation (Biggs et al., 2014, 2016), gas emission (Fischer et al., 1994; Werner et al., 2013), and fumarolic activity (Francis et al., 1980). Monitoring of the signals induced by these phenomena is indirect and interpretation of the origin of the signals is challenging because a wide variety of factors influence the behaviour of magma and host rock in the run-up towards eruption (Woods & Koyaguchi, 1994; Melnik & Sparks, 2002; Caricchi et al., 2014). The complex structure of volcanic systems and the stochastic nature of the driving processes mean that interpretation of monitoring signals typically relies on correlation, rather than causation. This hinders the detection and interpretation of pre-eruptive phenomena (Sano et al., 2015) and may result in 'false positives': activity that is interpreted as a precursor, which does not foreshadow an eruption (Syahbana et al., 2019). The observed correlations between monitoring signals and eruptive activity would be more robust if they could be linked via the casual physical mechanism.

Near-real-time observations of ground surface deformations have revealed both subsidence of volcanic edifices during eruption (Massonnet et al., 1995) and inflation of volcanic edifices during quiescent periods preceding an eruption (Patané et al., 2003). Such

73 data sets have provided insights into how edifice deformation is linked to volcanism (Biggs
74 et al., 2014; Pinel et al., 2014). For instance, Biggs et al. (2014) present a probabilistic
75 analysis of interferometric synthetic aperture radar (InSAR) data collected at 198 vol-
76 canoes over an 18 year period, which showed that 94 % of the volcanoes that did not de-
77 form also did not erupt, whereas 46 % of volcanoes that deformed also erupted (tectonic
78 changes induced the remaining deformations). Modelling of magma flow and viscosity
79 changes and pressure changes occurring in the upper part of conduit have been used to
80 link short-term eruptive cycles to observed ground deformations, but the outcomes are
81 not always predictable (Albino et al., 2011). Combining ground deformation measure-
82 ments with other sensitive techniques can help us to understand the causal physical mech-
83 anism by which ground deformation and volcanic activity are linked, and lead towards
84 more robust and predictive interpretations of the monitoring signals. Geophysical mon-
85 itoring of combined magma and host rock mass density can reveal the underlying phys-
86 ical mechanism of volcanic activity by providing indirect information about the compo-
87 sition and spatio-temporal evolution of magma propagation to the surface (Poland & Car-
88 bone, 2016; Londoño & Kumagai, 2018). Here we use of muon imaging to determine changes
89 in density in the plumbing system of Sakurajima volcano (Kyushu, Japan) at high spa-
90 tial and temporal resolution, and relate the results to observations of both ground de-
91 formation and eruptive activity.

92 Eruptions at Sakurajima are dominantly vulcanian, comprising impulsive explo-
93 sions that typically last minutes-to-hours (Iguchi et al., 2008; Miwa & Toramaru, 2013;
94 Yokoo et al., 2013). Activity is cyclic, with typical inter-eruption interval of a few days (Gabellini
95 et al., 2022). Since 2006, eruptive activity at Sakurajima has alternated between the Mi-
96 namidake crater and the Showa crater (Japan Meteorological Agency, 2022). Periods of
97 relatively high eruption frequency at a particular crater alternate with periods of qui-
98 escence – both periods typically last for several months. The physical mechanism for the
99 hours-to-days cycles of vulcanian explosions has been extensively investigated through
100 analysis of visual observations, seismic and geodetic data, gas geochemistry data, and
101 characterization of eruption products (Iguchi et al., 2008; Miwa & Toramaru, 2013; Yokoo
102 et al., 2013; Gabellini et al., 2022). The prevailing physical model infers the presence of
103 a dense plug of viscous magma in the upper few tens of meters of the conduit, beneath
104 which a pressurized pocket of gas accumulates, which is a few hundred meters in ver-
105 tical extent; explosions result from failure of the plug (Iguchi et al., 2008; Miwa & Tor-
106 maru, 2013; Yokoo et al., 2013; Kazahaya et al., 2016; Gabellini et al., 2022). Similar
107 mechanisms have been invoked to explain vulcanian activity at other volcanoes, includ-
108 ing Semeru (Java, Indonesia), Suwanosejima (Kyushu, Japan), and Soufrière Hills (Montser-
109 rat, West Indies) among others (Watt et al., 2007; Iguchi et al., 2008; Burgisser et al.,
110 2011). The longer-term cycles of alternating periods of high and low eruption frequency
111 have received much less attention. Gabellini et al. (2022) study the morphological and
112 petrological characteristics of ash emitted during vulcanian activity at Sakurajima, and
113 conclude that the dense, viscous plug forms over a period of several months, character-
114 ized by quiescence or low eruptive frequency, and is progressively destroyed during pe-
115 riods of high eruption frequency. In this work, we use muography to investigate these
116 longer cycles.

117 2 Muon imaging of volcanic interior

118 Muography exploits naturally occurring cosmic-ray muons to reconstruct the av-
119 erage densities along the paths of muons across large-scale structures, producing muon
120 "radiographic" images (Tanaka et al., 2007). The constant flux and the high penetration
121 power of muons allow passive and remote imaging of the shallow density structures in
122 volcanoes at a spatial resolution of few meters (Nishiyama et al., 2014; Oláh et al., 2018;
123 Macedonio et al., 2022; Miyamoto et al., 2022). Muography has already been used to im-
124 age the spatio-temporal evolution of magmatic materials, – e.g., ascent and descent of

125 magma within a volcanic vent (Tanaka et al., 2014), magma degassing (Tanaka et al., 2009)
 126 and plug formation underneath deactivated craters (Oláh et al., 2019) – and to observe
 127 structural changes (Lo Presti et al., 2022; Tioukov et al., 2022) and hydrothermal activi-
 128 ties (Gibert et al., 2022) in volcanic systems. The early warning capabilities of muog-
 129 raphy have also been studied (Nomura et al., 2020; Leone et al., 2021; Oláh & Tanaka,
 130 2022a). We conducted muography of Sakurajima volcano over the period from Septem-
 131 ber 2018 to July 2021.

132 The elevation map of the observational site and schematic drawing of the exper-
 133 imental configuration are shown in Figure 1. We applied the Multi-Wire Proportional
 134 Chamber-based Muography Observation System (MMOS) (Oláh et al., 2018; Varga et
 135 al., 2020, 2022) of Sakurajima Muography Observatory for muographic monitoring of mass
 136 density changes underneath the active craters (Text S1 in the supporting information).
 137 The small black rectangle shows the location of the observatory (O) at latitude 31.557°N
 138 and longitude 130.650°E , at altitude 150 m above sea level. The MMOS was oriented
 139 towards the active craters at 30.25° from north (defined as $\tan(\theta_x) = 0$), as shown by
 140 the black arrow, and set to horizontal (defined as $\tan(\theta_y) = 0$). The MMOS collected
 141 muon tracks within ± 505 mrad in the horizontal direction and ± 353 mrad in the verti-
 142 cal direction. The points M, S, and R were selected on the northeast slope of the vol-
 143 cano to extract three cross-sections across Minamidake crater, Showa crater and Refer-
 144 ence region, respectively. Figure 1b shows the three selected cross-sections along the OM,
 145 OS and OR lines. The targeted region underneath the active crater is shown within the
 146 OP and OQ lines. We note that the vertical angle region of 0-150 mrad (see under the
 147 OP line) was also covered by the MMOS, but the excessive (> 2.5 km) rock thickness
 148 did not allow us to measure the density in this angular region beneath the crater. The
 149 MMOS measured the muon tracks continuously from January 2017 with only a few tech-
 150 nical stops of a few days each during installation of new MMOS modules or maintenance
 151 work (Oláh et al., 2019, 2021; Varga et al., 2020).

152 The data from the MMOS were processed to create muographic images that show
 153 the density structure through the crater region of Sakurajima volcano. Muographic im-
 154 ages use the natural coordinate system of the MMOS that is the tangents of the projec-
 155 tions of incoming muon directions with respect to the orientation of MMOS: $\tan(\theta_x)$ and
 156 $\tan(\theta_y)$ (Text S1 in the supporting information). Each muographic image was captured
 157 with a binning of 0.023 (that corresponds to a spatial resolution of approx. 60 m at the
 158 crater, which is located at a distance of 2.65 km from the MMOS) in both horizontal and
 159 vertical directions. Each image was determined for a period of five months. The muo-
 160 graphic image processing was based on comparison of modelled and measured muon fluxes.
 161 The measured trajectories of particles were reconstructed by 1+1-dimensional line fits
 162 onto coordinates on the tracking layers (Oláh et al., 2018), and the fluxes were calculated
 163 for each angular bins by taking into account the measurement time, dead time, and de-
 164 tector geometry (Oláh et al., 2019). Quality assurance of the data was performed by off-
 165 line analysis and low-quality data were removed from the analyzed data sets. The mod-
 166 elled fluxes were determined by integrating the differential muon spectra (Tang et al.,
 167 2006) from the threshold energies that were required for muons to penetrate through the
 168 volcanic edifice. The threshold energies were calculated (Oláh et al., 2021) by taking into
 169 account the stochastic energy loss processes of muons (Lipari & Stanev, 1991).

170 3 Results

171 Figures 2a-x show images selected from a period between November 2018 and March
 172 2021. The blue dashed line visualizes the shape of the crater along the blue dashed line
 173 of Figure 1a. Black rectangular outlines highlight a region underneath the active Minami-
 174 dake crater (M), a region underneath the dormant Showa crater (S) and a Reference re-
 175 gion (R) in which volcanism does not occur. The vertical range of these regions corre-
 176 sponds to the angular region between the OP and OQ lines in Figure 1a. The path-averaged

177 densities (ρ) ranged between 0.8 gcm^{-3} and 1.8 gcm^{-3} through the regions underlying
 178 Minamidake crater and Showa crater. Density values ranged from 0.8 gcm^{-3} to 1.25 gcm^{-3}
 179 through the Reference region. The lower densities through the Reference region likely
 180 arises because energetic ($> 1 \text{ GeV}$) muons scattered into the MMOS from the surface of
 181 the downward sloping parts of volcanic edifice (Ambrosino et al., 2015). The white-shaded
 182 regions without density values are due to the thickness that was not penetrated by muons
 183 during the data collection time. The muographic images show that the densities change
 184 over time through the region underneath the craters.

185 Vertical displacement of the volcanic edifice was determined over the same period
 186 as the muography measurements (e.g., red line in Figure 2y). Displacement was deter-
 187 mined relative to the ground level measured on 31 October 2018 at ten locations (red-
 188 coloured dots in Figure 1b) using the Phased Array type C-band Synthetic Aperture Radar
 189 images acquired by Sentinel-1 (The European Space Agency, 2022). Eruption frequency
 190 for each crater, over the same time period, was determined from the data base of Japan
 191 Meteorological Agency (2022). We found that ground level, averaged monthly, at the Mi-
 192 namidake crater (location 8 in Figure 1b and red-coloured line in Figure 2y) and erup-
 193 tion frequencies, binned monthly, correlated inversely with a Pearson's coefficient of -
 194 0.718 (Spearman's rank coefficient of -0.798). We found moderate and weak inverse cor-
 195 relations for the same quantities at the locations 1, 2, 3 and 9 with coefficients of -0.547
 196 (-0.595), -0.509 (-0.432), -0.514 (-0.421) and -0.564 (-0.685), respectively. No significant
 197 correlation was found at the remaining locations. Visual inspection of Figure 2 suggests
 198 that density beneath the active craters is high during periods of low eruption frequency
 199 and upward displacement of the ground surface, and vice versa.

200 The muographically measured density values were averaged for the three regions
 201 (M, S, R in Figure 2) to quantify their variations in relation to the eruption frequencies
 202 and ground deformation during periods of high eruption frequency and periods of qui-
 203 escence. Figures 3a-c show together: 1) five month average of densities with one stan-
 204 dard deviation relative to the densities measured for the first time sequence from 1 Septem-
 205 ber 2018 to 31 January 2019 ($\Delta\rho=\rho(t)-\rho(t_0)$, where $\rho(t_0)$ equals to 1.26 gcm^{-3} , 1.13 gcm^{-3}
 206 and 0.99 gcm^{-3} for the Minamidake crater, Showa crater and Reference region, respec-
 207 tively); 2) the monthly number of eruptions (black histogram); and 3) the vertical ground
 208 displacement (red-coloured lines) for each region. The relative density increased beneath
 209 the Minamidake crater (Figure 3a) throughout the periods of inflation and reduced erup-
 210 tion frequency (from March to September 2019 and from August 2020 to January 2021).
 211 The increase of average relative densities exceeded $0.2\text{-}0.35 \text{ gcm}^{-3}$ which is significantly
 212 above the systematic density error of 0.06 gcm^{-3} (see in Oláh and Tanaka (2022b) and
 213 Text S1 in the supporting information). The relative densities decreased beneath the Mi-
 214 namidake crater during periods in which the volcanic edifice deflated and the eruption
 215 frequency increased (from November 2019 to May 2020 and from November 2020 to May
 216 2021).

217 The densities beneath the dormant Showa crater (Figure 3b) slightly increased at
 218 the end of 2018 and remained below 0.1 gcm^{-3} during the observations periods. Across
 219 the Reference region (Figure 3c), the density changes were below 0.1 gcm^{-3} .

220 4 Discussion

221 We monitored the mass density changes through the upper conduit of Sakurajima
 222 volcano with muography during cyclic eruption episodes of Minamidake crater. We found
 223 that the trends in mass density were linked to trends in ground deformation and erup-
 224 tion frequency: the mass density increased during periods of inflation and low eruption
 225 frequency, and decreased during periods of deflation and high eruption frequency. These
 226 observed trends also correlate with other monitoring signals. During the periods from
 227 January to July 2019 and from June to November 2020 (roughly coincident with peri-

ods of high density, edifice inflation, and low eruption frequency), the time series of the vertical locations of seismic sources distributed at shallow depths underneath the Minamidake crater (middle right panel of Figure 3 in Japan Meteorological Agency (2021)) that also suggested the densification of this region. Furthermore, infrared thermal imaging revealed high-temperature regions in the Minamidake crater in August 2020 (Figure 3-1 in Japan Meteorological Agency (2020)), and glowing of Minamidake was observed in September and October 2020 (Figures 4-1, 4-2 in Japan Meteorological Agency (2020)). During periods of low density, edifice deflation, and high eruption frequency, increased trends were observed in the sulphur dioxide discharge mass rate by JMA (upper panel of Figure 8-2 in Japan Meteorological Agency (2021)).

The increase in mass density revealed by the muography data is consistent with the formation of a dense plug of magma in the shallow plumbing system (upper ~ 200 m) of Minamidake crater during periods of low eruption frequency, as proposed by Gabellini et al. (2022). Upward deformation of the volcanic edifice during this period is consistent with slow upward migration of the dense, stiff plug driven by pressure from below as the conduit refills with fresh magma. We note that five highly energetic explosions occurred in June 2020 (i.e., at the start of a period of inflation) from which volcanic ejecta reached the altitude of 1.5-3.7 km above the crater rim (Japan Meteorological Agency, 2020). These events may have been associated with the onset of magma intrusion into the gas pocket. The observations of high-temperature regions and glowing material in the crater indicates that portions of the dense, hot plug are extruded into the crater. Conversely, the decrease in mass density during periods of high eruption frequency is consistent with the presence of gas pockets in the conduit, as proposed in numerous studies (e.g., Iguchi et al. (2008), Miwa and Toramaru (2013), Yokoo et al. (2013), Kazahaya et al. (2016)). The associated deflation of the edifice may result from the cessation of recharge of magma, and the progressive outgassing of the magma within the conduit. This is consistent with the elevated sulphur dioxide discharge mass rate measured during these periods (Japan Meteorological Agency, 2020).

Concerning the Showa crater, the observed slight density increase is assumed to be a part of the plug formation process (Oláh et al., 2019) initiated at the end of 2017 by cessation of eruptive activity (Japan Meteorological Agency, 2022). Although the vertical uplifts increased from 2018 December to June 2020, no significant change was in the trend of density was observed during this period. The minor density changes ($< 0.1 \text{ gcm}^{-3}$) in the Reference region are consistent with its non-magmatic nature, and act as a control.

Lo Presti et al. (2022) measured the muon flux through the Northeast crater and the Voragine crater of Mount Etna at latitude 37.757°N and longitude 14.988°E . Data acquisition was conducted from July to November in 2018 (95 days) and between August and September in 2019 (51 days). The forward (volcano's direction) to backward (opposite to volcano's direction) muon flux ratios (C) were quantified for both periods. Figure 4 shows the ratio of the C measured in 2019 (C_{2019}) to the C measured in 2018 (C_{2018}) in the natural coordinate system of the muon tracker (same as the coordinate system of Figures 2a-x). The cross-sectional presentation of the crater floor in Etna is not shown in this figure since it was difficult for us to acquire the elevation data presenting the topographic shape right before the 2019 volcanic activity. This muographic image shows a region through the Voragine crater in which the C increased from 2018 to 2019. The increase in C represents the decrease in mass density. The eruption frequency of the Voragine crater increased also (Global Volcanism Program, 2022; National Institute of Geophysics and Volcanology, 2022): the crater did not erupt since December 2015. A new vent opened in August 2018 and degassing from the new vent initiated in January 2019. Another new vent opened in April 2019 from which two eruption sequences occurred in June and September 2019. Near surface relative density reduction observed at Voragine crater during the 2019 eruptions is in agreement with the relative density

281 reduction we observed during the eruptions of Sakurajima. A joint muographic obser-
 282 vation of the two volcanoes will may provide more insights into the shallow volcanic pro-
 283 cesses.

284 5 Conclusions and outlook

285 In this work we have demonstrated that muography can be a valuable tool for in-
 286 vestigating in-conduit processes at active volcanoes, particularly when combined with
 287 other monitoring data. The technique provides data for the evolution of mass density
 288 of magma in the shallow conduit at sufficiently high temporal and spatial resolution that
 289 changes in the eruption can be associated with changes in the state of the magma in the
 290 shallow conduit. This has great potential for elucidating the mechanisms through which
 291 changes in eruption frequency are mediated by physical processes in the shallow conduit
 292 which, in turn, will improve conceptual, physical, and numerical models of eruptive pro-
 293 cesses. Vulcanian activity similar to that at Sakurajima is common at volcanoes world-
 294 wide, and similar conceptual models for eruption activity have been proposed (e.g., Watt
 295 et al. (2007), Iguchi et al. (2008), Burgisser et al. (2011)). The model for long-period (multi-
 296 month) changes in eruptive activity that we develop here could therefore be applicable
 297 elsewhere.

298 This work also indicates that muography could be used to improve the intermediate-
 299 term assessment of the hazard levels at stratovolcanoes by allowing more meaningful in-
 300 terpretation of InSAR data. For instance, where edifice inflation is associated with an
 301 increase in magma density in the shallow conduit, it could indicate a plugging of the shal-
 302 low conduit and the onset of a period of reduced eruption frequency. Furthermore, muog-
 303 raphy has the potential for providing useful data even between the observation flights
 304 of space-borne InSAR for detecting rapid (a few days duration) changes in magma con-
 305 ditions underneath the active craters. Currently, the main limitation in muography is
 306 the relatively long measurement time; this can be reduced by enlarging the sensitive sur-
 307 face area of the available muographic observation systems, and we anticipate that tech-
 308 nological improvements in muography will further enhance its value to volcanology.

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 325 tws4j/?view_only=43ce27d1d94647eeaea9c3cdd117fe70](https://osf.io/tws4j/?view_only=43ce27d1d94647eeaea9c3cdd117fe70).

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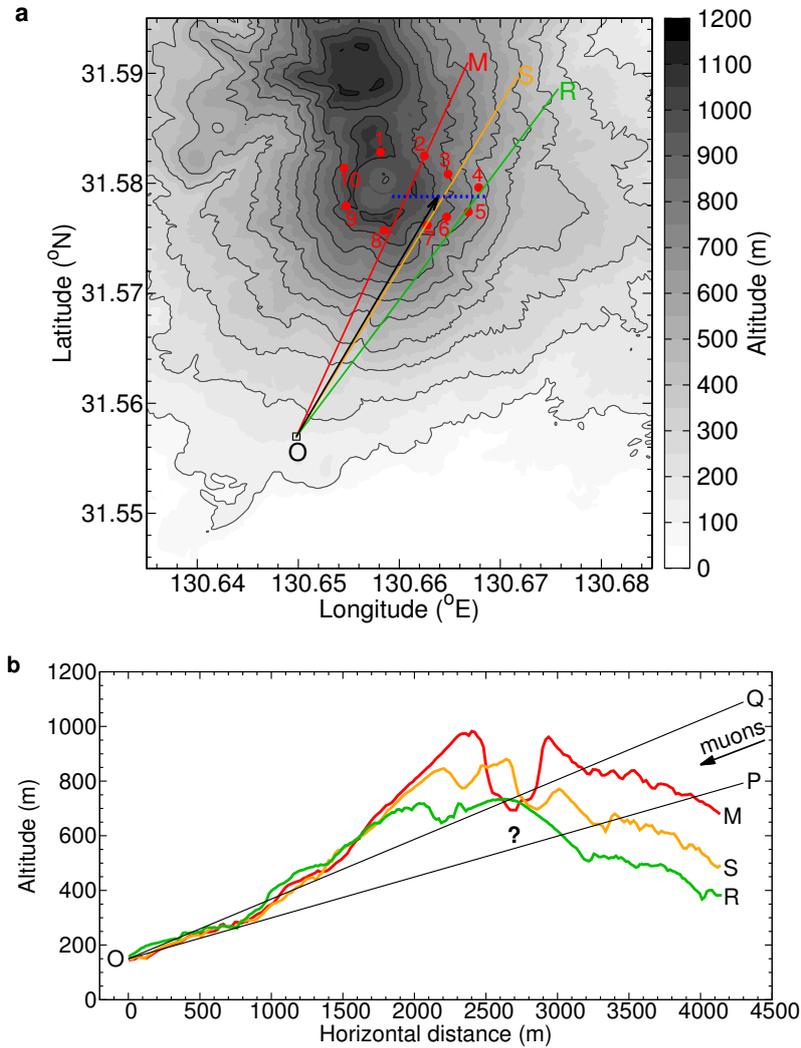


Figure 1. Map of measurement site and schematic of the experimental arrangement. **a**, The map of Sakurajima volcano is drawn based on the digital elevation data of Geospatial Information Authority of Japan (<http://www.gsi.go.jp/>). A small black rectangle shows the location of the muography observatory (O) at latitude 31.557 °N and longitude 130.650 °E, at 150 m altitude above sea level. The black arrow shows the azimuthal orientation of MMOS that was set to 30.25° from north (defined as $\tan(\theta_x) = 0$). The MMOS was oriented horizontally (defined as $\tan(\theta_y) = 0$). OM, OS, OR lines highlights three selected cross-sections across the Minamidake crater, Showa crater and a Reference region, respectively. The blue dashed line shows a selected cross-section across the crater region. Red-coloured dots and numbers refer respectively to the locations and identification numbers of selected sites where ground displacements were determined from the data collected by synthetic aperture radar. **b**, Three cross-sections of the measurement site are shown along the OM, OS and OR lines, respectively. The OP and OQ lines bound the vertical range of the studied region beneath the craters. Question mark shows the location of volcanic conduits within the studied angular region.

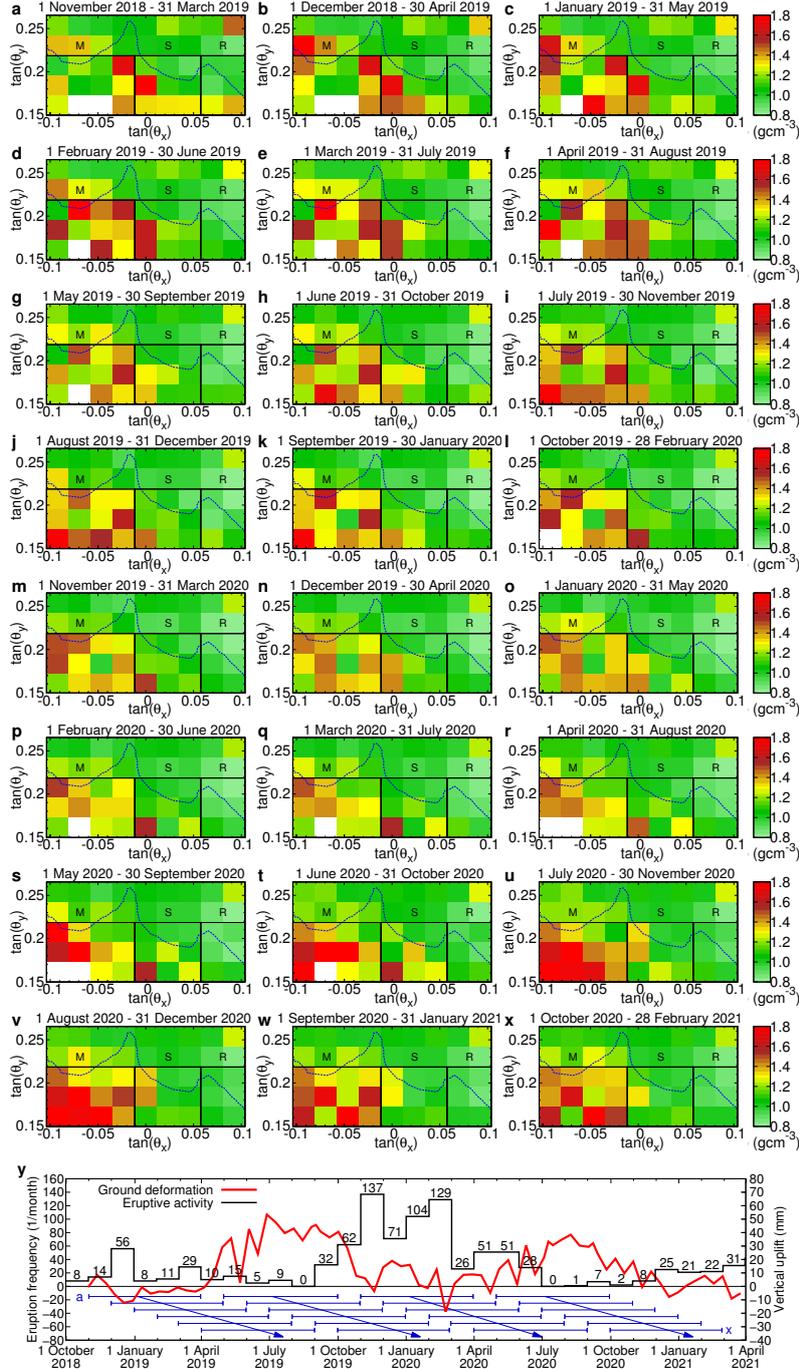


Figure 2. Time-sequential density images of Sakurajima volcano. **a-x**, The average density (ρ) values are plotted for the crater region as a function of horizontal and elevation directions for periods of 5 months from 1 November 2018 to 28 February 2021. The densities were calculated for angular bins with the size of $\Delta(\tan(\theta_x)) \times \Delta(\tan(\theta_y)) = 0.023 \times 0.023$ each. Blue-coloured dashed lines show the cross-section of craters along the blue-coloured dashed line of Figure 1a. Black rectangular outlines designate three angular regions beneath the Minamidake crater (M), the Showa crater (S) and the Reference region (R), respectively. **y**, Time-lines of ground deformation (red-coloured line) and eruptive frequency (black histogram) are shown for the Minamidake crater of Sakurajima volcano, Japan. Ground deformation data were recorded with InSAR (The European Space Agency, 2022). The eruptive frequency was determined using the data from (Japan Meteorological Agency, 2022). Blue arrows are drawn for comparing the **a-x** images with these time-lines.

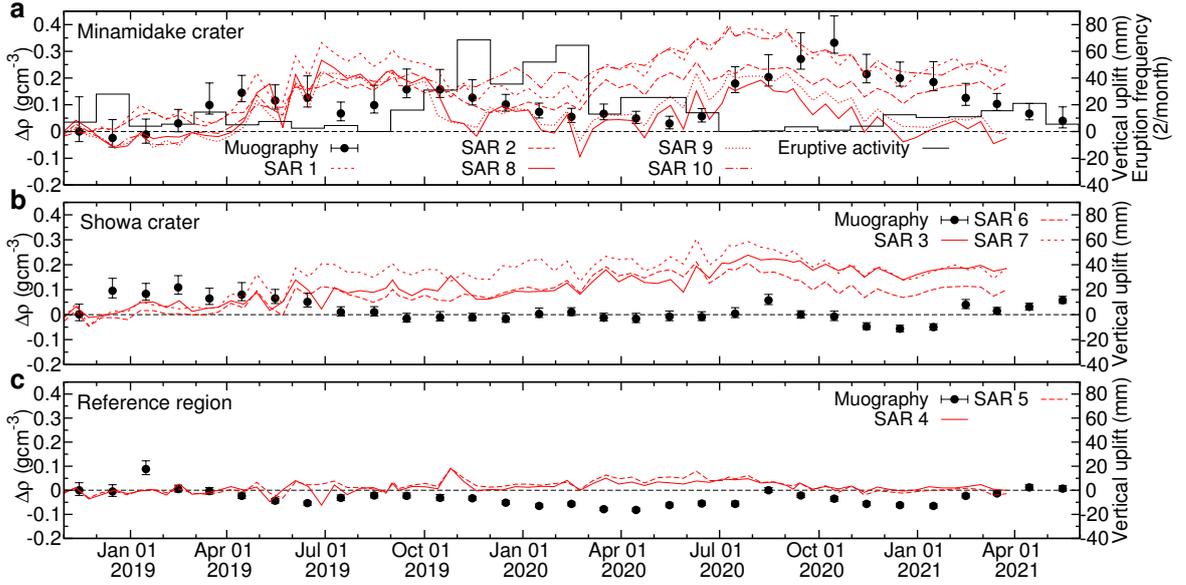


Figure 3. Time evolution of relative averaged densities and vertical uplifts through the three regions of Sakurajima volcano. The densities are shown with 1 standard deviation error bars from September 2018 to July 2021, relative to the averaged densities measured during the first time interval from 1 September 2018 to 31 January 2019 for **a**, the Minamidake crater, **b**, the Showa crater, and **c**, the Reference region, respectively. The points refer to the mids of time intervals. Vertical ground deformations measured at the ten selected locations (Figure 1a) are shown by the red-coloured lines. The eruption frequency of Minamidake crater is shown by the black histogram.

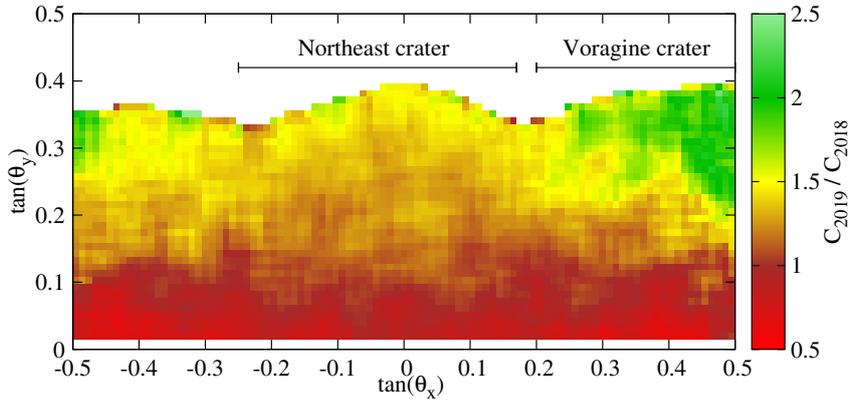


Figure 4. A muographic image of Mount Etna. The ratio of the forward to backward muon flux ratio measured for 2019 (C_{2019}) to the same quantity measured for 2018 (C_{2018}) is shown for each angular bin. The muographic observation was conducted at latitude 37.757°N and longitude 14.988°E , at a distance of 700 m in northwest from the Northeast Crater. Black arrows show the horizontal extensions of the Northeast crater and the Voragine crater, respectively. The current analysis was conducted based on the data provided by Lo Presti et al. (2022).