

On the Thermal Signature of the Residual Foam in Breaking Waves

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Key Points:

- The thermal signature of cooling residual foam can be used to infer breaking wave dynamics.
- The onset of cooling coincides with the bubble plume subsidence and the end of foam generation and scales with the bubble plume decay time.
- The cooling time of the foam is not significantly affected by the presence of surfactants.

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Abstract

Quantifying energy dissipation due to wave breaking remains an essential but elusive goal for studying and modeling air-sea fluxes of heat, gas, and momentum. Previous observations have shown that lifetimes of bubble plumes and surface foam are directly related to the dissipated energy. Specifically, the foam decay time can be used to estimate the timescale of the subsurface bubble plume and the energy dissipated in the breaking process. A mitigating factor is that the foam decay time can be significantly affected by the surfactant concentration. Here we present an experimental investigation of a new technique that exploits the thermal signature of cooling foam to infer wave breaking dynamics. The experiments were conducted in a laboratory wave tank using artificial seawater with and without the addition of a surfactant. We show that the onset of cooling coincides with the bubble plume subsidence and the end of foam generation. The time from the start of the breaking process to the onset of cooling scales with the bubble plume decay time and the dissipated energy, and is not significantly affected by the presence of additional surfactants.

Plain Language Summary

Breaking waves cause mixing and are important for redistributing heat, transporting gases between the air and the water, and generating currents. Bubbles from breaking eventually rise and stay at the surface where they can be visually seen as foam. Scientists have found that the time it takes for the foam to disappear is related to the strength of the breaking waves. However, natural chemicals in the seawater can cause the bubbles to disappear more slowly, increasing the time they are seen at the surface. We present a new method to estimate when the bubbles stop rising based on the foam temperature. We generate breaking waves in a laboratory and use an infrared camera to measure the temperature of the foam and find that the foam cools when bubbles stop rising. We varied the strength of the breaking waves and measured the cooling time for the foam to show that larger, stronger breaking waves cause a longer time before the foam begins to cool. When we added chemicals to increase the time foam stays at the surface, we found that the cooling time remains about the same, even though the foam is still seen at the surface for a longer time.

1 Introduction

Wave breaking plays a critical role in air-sea interaction processes in both the open ocean and the surf zone. The energy transferred from the atmosphere to the ocean through wind-wave generation is ultimately dissipated by wave breaking. Therefore, quantifying the energy dissipation due to wave breaking is directly relevant to wave prediction models used for operational sea-state forecasting and the impact of waves on coastal regions. At high wind speeds, bubbles generated by large scale breaking waves are the primary mechanism for gas transfer and dominate the energy dissipation due to breaking (Lamarre & Melville, 1991). Bubbles generated by breaking waves also contribute to marine aerosol formation through spray droplets produced when foam bubbles burst at the surface (Veron, 2015; Erinin et al., 2019, and references therein). Foam generated by wave breaking has increased reflectivity of solar radiation that can affect the earth's albedo (Evans et al., 2010; Gordon & Jacobs, 1977) and the enhanced microwave emissivity of foam impacts space-borne radiometer measurements of wind speed. In short, wave breaking is an important mechanism for fluxes of momentum, gas, and heat across the air-water interface and for global ocean remote sensing applications. Here we focus on breaking waves that produce visible foam.

A wave begins to break when the forward face steepens and the crest becomes unstable. The morphology of an individual breaking wave that generates foam is generally categorized as either spilling or plunging. Plunging breakers occur when a wave crest forms an open curl and rapidly falls forward. Spilling breakers are characterized by a wave crest that spills forward and rolls down the face. For both types, subsurface bubble plumes are generated

61 by the impact of the overturning crest on the water surface. For plunging breakers, bubbles
 62 are also injected when the air pocket formed by the curling crest collapses, and by the jet
 63 impinging on the surface (Kiger & Duncan, 2012). As the actively breaking crest continues
 64 to propagate for the breaker lifetime, a bubbly turbulent wake is left behind. Bubbles rise to
 65 the surface and produce patches of residual surface foam as turbulence in the wake subsides.

66 The term whitecap has been used to describe both the foam generated by the actively
 67 breaking crest and the residual foam left behind in the wake. Monahan and Lu (1990)
 68 denoted the actively breaking period as stage-A and the period following as stage-B. Stage-
 69 A also has been referred to as the “acoustically active” period (Deane & Stokes, 2002),
 70 because the formation and fragmentation of bubbles generate underwater sound during this
 71 period. Stage-A includes the formation of the actively breaking crest and creation of the
 72 bubble plume and ends when air is no longer actively entrained. Stage-B includes the
 73 expansion and rise of the bubble plume as well as the formation and decay of the resulting
 74 residual foam and typically is of longer duration than stage-A (Kleiss & Melville, 2010;
 75 Monahan & Lu, 1990).

76 Whitecap coverage, W , is the percentage area of the sea surface covered by foam mea-
 77 sured from visible imagery. Techniques for measuring W have evolved from labor-intensive
 78 analysis of individual photographs (Monahan, 1969) to automated techniques using high-
 79 resolution digital imagery (Callaghan & White, 2009). Scanlon and Ward (2013) recently
 80 reported on a manual technique to separate active and maturing whitecaps, but automated
 81 processing of visible imagery remains a challenge. Most visible measurements of W include
 82 both stages of whitecap foam because of the difficulty of objectively and automatically dis-
 83 tinguishing between stage-A and stage-B foam. However, there is strong motivation to be
 84 able to separately measure whitecap coverage for stage-A, W_A , and for stage-B, W_B , in
 85 order to examine the different processes of interest that are associated with the different
 86 stages. For instance, W_A is the appropriate coverage to determine the breaking rate and for
 87 correlation with the energy dissipation (Kleiss & Melville, 2010) while W_B has been related
 88 to sea-salt aerosol production due to the preponderance of bursting bubbles as the foam
 89 dissipates (Callaghan, 2013; Monahan et al., 1986).

90 Since wave breaking is driven by wind stress, many authors have pursued a fundamental
 91 parameterization of W with wind speed. A compilation of historical data sets by Anguelova
 92 and Webster (2006) shows scatter of W versus wind speed of over three orders of magnitude,
 93 suggesting other factors need to be considered. Recent results by Callaghan et al. (2008,
 94 2012, 2013) indicate that potential contributors to the observed scatter of W with wind
 95 speed include environmental parameters such as wave field characteristics, breaker type,
 96 and surfactant effects as well as differences in image acquisition and analysis techniques.
 97 Variations in the decay time of oceanic whitecap foam led Callaghan et al. (2012) to speculate
 98 that the two primary mechanisms that cause scatter of whitecap coverage with wind speed
 99 are (i) the effect of surfactants on foam stability and (ii) differences between bubble plume
 100 characteristics caused by variation in breaker type.

101 In the laboratory, Callaghan et al. (2013) investigated these two mechanisms by compar-
 102 ing the visible decay times for foam, τ_{foam} , and for the bubble plume, τ_{plume} , for whitecaps
 103 generated by focused wave packets using clean and surfactant-contaminated (Triton X-100
 104 at $204 \mu\text{gr}\cdot\text{L}^{-1}$) seawater. They found that τ_{plume} was proportional to the increase in energy
 105 dissipated as the scale of breaking ranged from spilling to plunging. However, when surfac-
 106 tants were present, the scaling between τ_{foam} and τ_{plume} varied significantly. Surfactants
 107 act to stabilize the bubbles, causing them to persist at the surface after the bubble plume
 108 has decayed and the foam generation process has ceased. For clean conditions, the foam
 109 decay time can provide a direct estimate of the plume degassing time. In the presence of
 110 surfactants, their effect on increasing the foam decay time needs to be accounted for in order
 111 to infer the plume decay time (Callaghan et al., 2017).

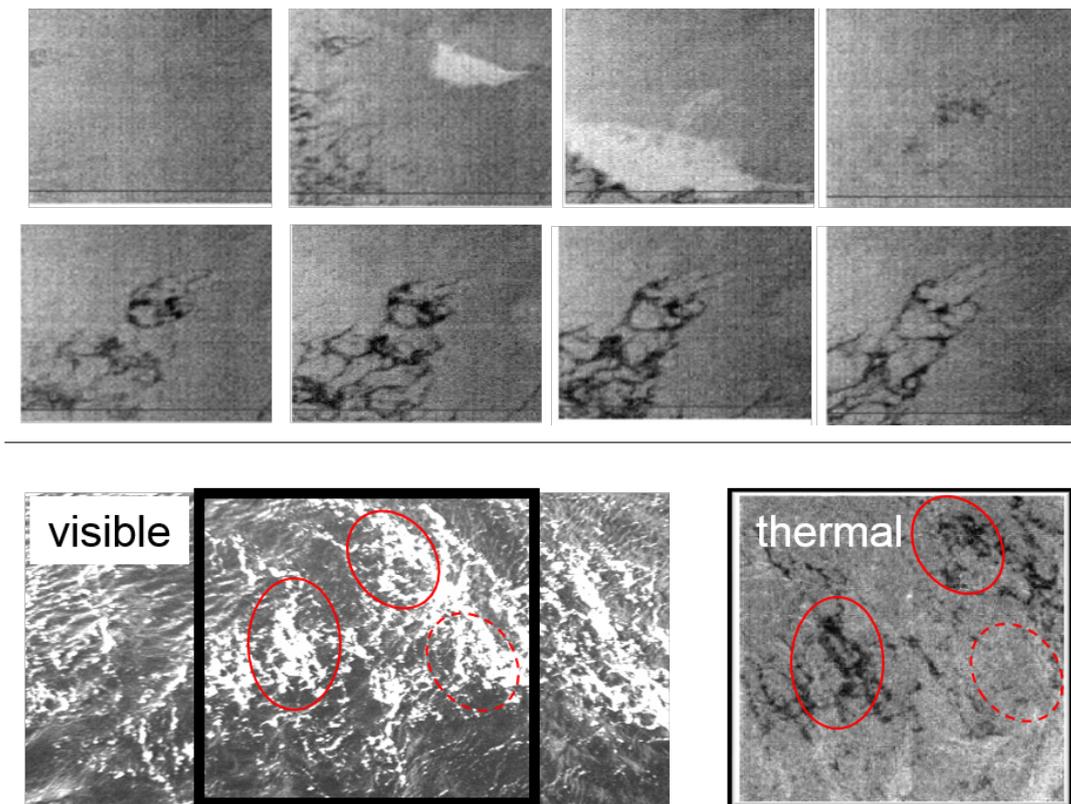


Figure 1. (Top) An examples of infrared image sequences showing the cooling of residual foam in the wake of a breaking wave in the open ocean. Time increases left to right then top to bottom. Lighter shades of gray are warm and darker shades are cold. Image size is approximately $5 \text{ m} \times 5 \text{ m}$. (Bottom) Simultaneous visible (left) and infrared (right) images of residual foam in the wake of a breaking wave in the open ocean. Solid ovals indicate locations where visible foam appears cool while foam does not appear cool in the dashed oval (Fogelberg, 2003).

112 The infrared image sequence in Figure 1 illustrates the rapid cooling of foam left behind
 113 after the passage of a breaking wave, similar to observations reported by Fogelberg (2003)
 114 and Marmorino and Smith (2005). Recent laboratory findings by Jessup et al. (2016) that
 115 the heat flux from foam is up to four times greater than foam-free water are consistent
 116 with the suggestion by Marmorino and Smith (2005) that the cooling is due to enhanced
 117 evaporation from bubbles. Jessup et al. (2016) also reported that the foam cooling begins
 118 after the foam-producing bubbles cease rising. The foam cooling phenomenon has been used
 119 recently to distinguish between the active and residual foam. In the open ocean, Potter et
 120 al. (2015) used infrared imagery to quantify the lifetime stages and characterize properties of
 121 the active and residual whitecaps. In the surf zone, Carini et al. (2015) used the difference in
 122 the thermal signature of active and residual foam to identify and extract the perpendicular
 123 crest length of the aerated breaking region.

124 A consistent observation in the field is that there is a momentary delay between when
 125 the foam appears in the visible to when it begins to cool. Additionally, not all residual
 126 foam from a given breaking event cools at the same time, as illustrated by the comparison
 127 of simultaneous visible and infrared images in Figure 1 (bottom). These characteristics of
 128 cooling foam and the finding that it cools more rapidly than foam-free water are consistent
 129 with the observation that the onset of cooling coincides with the cessation of foam generation

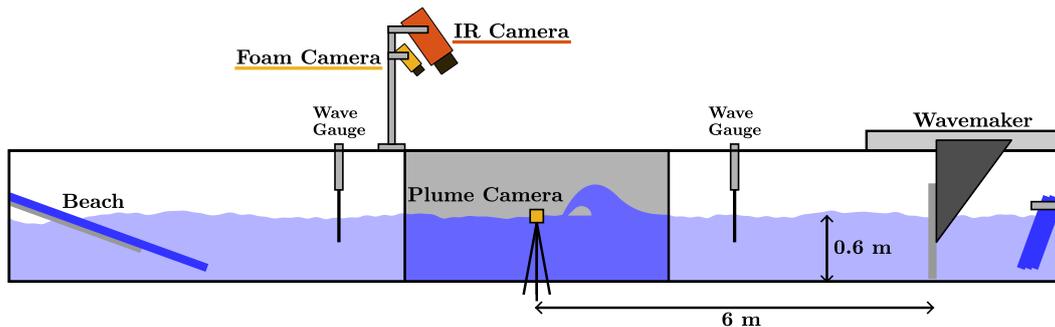


Figure 2. Schematic of the experimental setup. One side wall of the tank is made of glass for optical access. The other side wall is painted black in a 1.6 m long test section. The top of the tank is covered except for the test section. The waves are designed to break at a location approximately 6 m from the wave paddle. The tank width is 0.91 m.

130 by rising bubbles. We hypothesize that while the foam at the surface is being replenished
 131 from below, its temperature remains comparable to that of the water from which it is
 132 generated. The cooling begins only when the foam ceases being replenished by the rising
 133 bubbles, which implies that the foam replenishment rate compensates for or exceeds the
 134 rate of cooling. The start of the cooling is delayed for more energetic breaking waves that
 135 generate larger and deeper bubble plumes compared to less energetic breakers that generate
 136 smaller and shallower plumes. Thus the time from the start of breaking to when the foam
 137 begins to cool should scale with the bubble plume decay time.

138 Here we seek to exploit the thermal signature of the cooling of surface foam in the wake
 139 of a breaking wave to infer the subsurface plume dynamics. We present a new approach
 140 for estimating the time scale of the subsurface bubble plume based on the timescale of the
 141 cooling foam. Results from laboratory experiments show that the onset of cooling of the
 142 foam scales with the decay time of the bubble plume. Furthermore, the cooling time is not
 143 significantly affected by surfactants. We demonstrate that infrared imagery can provide the
 144 ability to infer the bubble plume decay time and thus a measure of wave energy dissipation.

145 2 Experimental Details

146 2.1 Setup

147 The experiments were performed in the Washington Air-Sea Interaction Research Fa-
 148 cility (WASIRF) wave flume at the University of Washington (Figure 2). The wave flume
 149 is 12 m long, 0.91 m wide, and 1.2 m tall, with one side wall made of glass that allows
 150 optical access. The top of the tank was covered with removable panels except for the test
 151 section that was left open for imaging. The facility includes a water circulation system that
 152 is equipped with an inline filter and electric heater. The flume was filled with salt water
 153 to a depth of 0.6 m using Instant Ocean and tap water. The salinity was set to 30 ppt
 154 and was frequently checked with a refractometer to ensure it remained constant during the
 155 experiments.

156 A programmable piston-type wavemaker at one end of the tank was used to generate
 157 breaking waves. The wavemaker consists of a flat rectangular paddle that is 0.9 m tall and
 158 spans the tank width and extends to the bottom of the tank. The motion of the wavemaker
 159 is controlled by an analog signal sent to the controller of the servo motor. Wave absorbing
 160 beaches were installed at both ends of the tank to diminish wave reflection from the end
 161 walls.

2.2 Wave Generation

Breaking waves are generated using the dispersive focusing wave packet technique used extensively in laboratory experiments (Rapp & Melville, 1990; Duncan et al., 1999; Drazen et al., 2008; Wang et al., 2018). In this technique, a packet is composed of many components and is designed such that all the components have the same phase at a prescribed “breaking” location. In the experiments presented here, the motion of the wavemaker can be described as

$$\eta_0 = \sum_{i=1}^N \frac{a_i}{a_i^{corr}} \cos(-k_i x_b - 2\pi f_i(t - t_b) - \phi_i^{corr}) \quad (1)$$

where η_0 is the wavemaker horizontal displacement, $N = 32$ is the total number of components, and for each component, a_i is the amplitude, a_i^{corr} is the amplitude correction factor found in the calibration process, k_i is the wavenumber, f_i is the frequency, and ϕ_i^{corr} is the phase correction. x_b and t_b are the theoretical breaking location and time, respectively. Equation (1) produces a periodic signal and needs to be windowed to provide the proper wavemaker motion. We used a window with hyperbolic tangent edges to taper the signal in a smooth fashion. The parameters that control the shape of the wave packet signal and the breaking location of the waves are the central frequency of the packet, $f_c = 0.88$ Hz, the frequency bandwidth, $\Delta f = 0.5f_c$, and the normalized breaking location, $x_b k_c = 33$, where k_c is the wavenumber corresponding to the central frequency. These parameters were chosen to generate breaking waves with similar shape but with considerable difference in foam generation, plume depth, and energy dissipation. The global slope of a wave packet is used as the control parameter for the scale of the breaking waves and is defined as:

$$S = \sum_{i=1}^N a_i k_i \quad (2)$$

Following Loewen (1991), all the components in the wave packet are chosen to have the same slope. Hence, increasing the global slope, S , is essentially equivalent to multiplying the signal by a constant factor without changing the overall shape.

2.3 Measurement Techniques

The bubble plume and the surface foam generated by breaking waves were measured using two identical visible cameras (Point Grey model Blackfly; 4 MP, 15 fps) to visualize the light scattered by the bubbles against the dark background in the test section, which was illuminated by two LED light sources. The bubble plume camera was located outside the glass side wall of the tank and was oriented normal to the wall. The horizontal center-line of the field of view coincided with the calm water surface. The field of view of this camera was approximately $1.2 \text{ m} \times 1.2 \text{ m}$. The foam camera was located approximately 2 m above the water surface and viewed the surface at an incidence angle of 30 degrees. The field of view of this camera extended to regions outside of the tank and these regions were masked in the analysis.

The surface temperature was measured using an infrared camera (DRS model UC640; 640×480 pixels, 30 fps, NEDT 25 mK) that was mounted adjacent to the foam camera. The foam and infrared cameras had overlapping fields of view and a transformation map between the two cameras was found in the calibration process. A circular metal target was attached to a float on the water surface and heated before being placed inside the tank so it would be visible to the infrared camera. The target was then moved in different parts of the field of view and imaged by the two cameras simultaneously. The center of the circular target was tracked through the image sequences and the transformation between the two cameras was found using a projective transformation between the pairs of target locations (Goshtasby, 1986). A second transformation was found and applied to all infrared and visible images to account for the oblique perspective of these cameras.

Table 1. Properties of the breaking waves used in the experiments where d_{\max} is the maximum depth of the bubble plume, ΔE is the total energy dissipation as estimated by upstream and downstream wave gauges, A_{plume}^{\max} is the maximum plume area, and A_{foam}^{\max} is the maximum whitecap area.

Slope	0.34	0.35	0.36	0.37
ΔE (J/m)	73.93	85.23	93.69	105.03
d_{\max} (cm)	15.05 ± 1.64	18.48 ± 1.71	24.79 ± 2.98	29.12 ± 3.21
A_{plume}^{\max} (m ²)	0.085 ± 0.0066	0.113 ± 0.0079	0.141 ± 0.0081	0.158 ± 0.0105
A_{foam}^{\max} (m ²)	0.599 ± 0.038	0.728 ± 0.045	0.808 ± 0.065	0.864 ± 0.058

The two visible cameras and the infrared camera were time-synchronized through the data collection software to record the breaking process simultaneously. Two capacitance-type wire wave gauges were mounted approximately 1.5 m upstream and downstream of the breaking location. The surface elevation data recorded at these locations were used to estimate the total energy dissipated by the breakers (Rapp & Melville, 1990).

2.4 Experimental Procedures and Conditions

For the experiments presented here, four breakers with slope values of $S = 0.34, 0.35, 0.36,$ and 0.37 were used. This range of slopes corresponds to plunging breakers that vary in intensity, amount of air entrained, and energy dissipation. This range of slopes corresponds to a range of 74-105 Joules per meter (along the crest of wave) in energy dissipation. Properties of the breakers are listed in Table 1.

The data collection process was automated so that many runs could be carried out unattended. The water in the tank was recirculated and filtered the night before each experiment day for about eight hours. In the morning of each experiment day, the water was heated to approximately 1 degree Celsius above the ambient air temperature. The air temperature varied slightly during the experiments due to the diurnal cycle. The temperature difference, $\Delta T = T_{\text{water}} - T_{\text{air}}$, was in the range of zero to 2 degrees Celsius for all the experimental runs presented here. The wavemaker and the data collection computers were set up to continuously generate breaking waves and collect data every ten minutes over the course of the day. This time between runs was found to be sufficient for the wave reflections to dissipate.

Two sets of experiments were carried out; in the first set, clean salt water was used (no additional surfactants), and in the second set, Triton X-100 was added to achieve a concentration of 200 $\mu\text{g/L}$. For each wave slope and for a condition with or without additional surfactants (eight total cases), between 50 to 60 runs were recorded and analyzed for a total of 462 individual breaking waves overall.

3 Results and Discussion

A sequence of visible bubble plume and foam images for a breaker propagating from left to right are shown in Figure 3 (see the supplementary material for movies corresponding to the image sequences). The images are separated by $\Delta t = 1/3$ s and the first image in the sequence is from $t = 1/3$ s. The time origin, $t = 0$, denotes the start of the breaking process and is found by manually inspecting the bubble plume camera images. For the example shown in Figure 3 (top), which is the largest wave slope used, a significant amount of air is entrained and left behind by the active breaker in (a) and (b). Two relatively

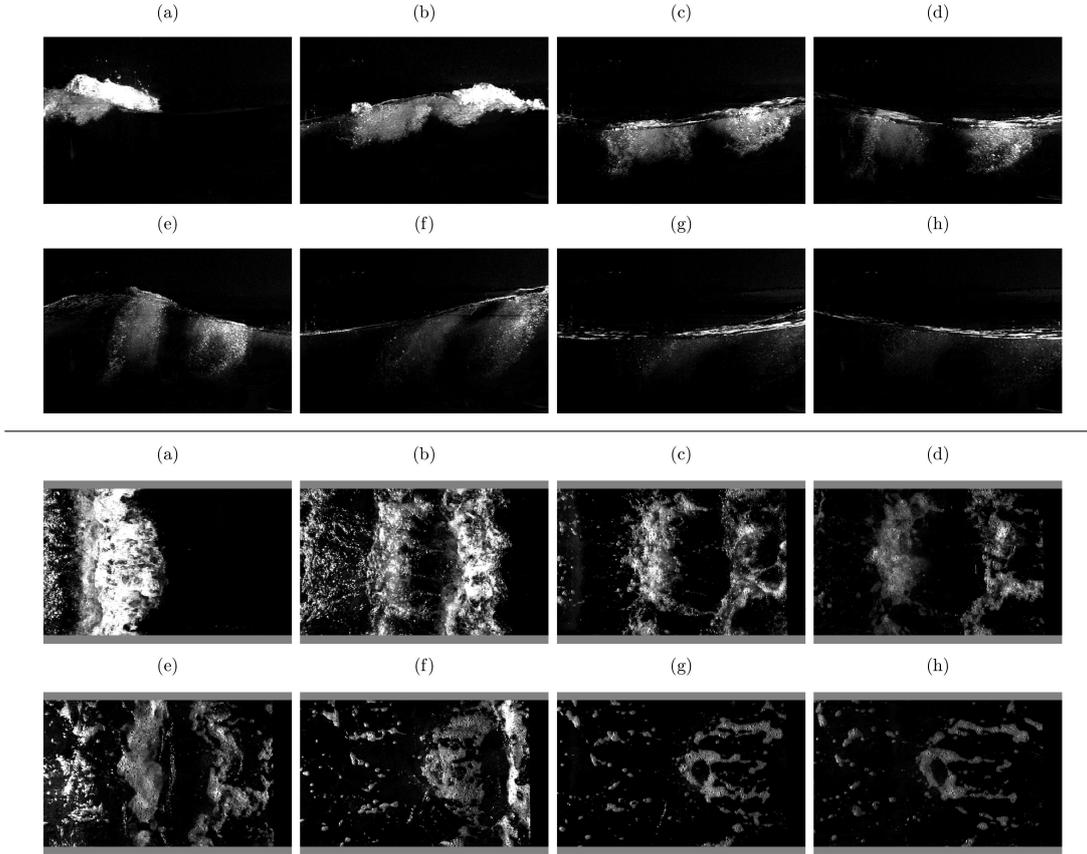


Figure 3. A sequence of images of a breaking wave with a slope of $S = 0.37$. The wave is propagating from left to right. The frames are separated in time by $1/3$ s. The wave packets are designed so that the breaking occurs at the edge of the field of view. (Top) bubble plume images taken from the camera that is looking through the glass wall of the tank. (Bottom) visible foam images taken by the camera that is looking down at the water surface. The foam images are shown in the coordinate system of the infrared camera images (see text). Each image is approximately 1.2 m long. This figure corresponds to Movie S2 in the supplementary information.

241 large and distinct bubble plumes occur for this breaker, as seen in images (c) and (d). The
 242 bubbles quickly rise to the surface in (e) and (f), and the residual surface foam left behind
 243 is apparent in (g) and (h).

244 The gray-scale visible images shown in Figure 3 were analyzed to obtain foam and plume
 245 area time series. First, the background was subtracted from all images in a sequence to
 246 enhance the signal and reduce the effect of non-uniformity in the lighting condition. Then,
 247 a manually determined intensity threshold was applied to segregate the bright foam and
 248 bubbles from the dark background, resulting in black and white (B/W) image masks. The
 249 same fixed threshold was used for all the runs since the lighting conditions were invariant.
 250 For the foam images, bright regions smaller than 200 pixels were removed from the images
 251 to reduce the speckle noise. Dark areas smaller than 200 pixels that are enclosed by bright
 252 foam were converted to white pixels. These regions are typically centers of large bubbles
 253 before they burst at the surface. The foam images were then transformed into the coordinate
 254 system of the infrared camera. A sequence of the resulting B/W bubble plume and foam
 255 masks is shown in Figure 4. This processing was done in the range of $t = -1$ s to $t = 10$ s
 256 for each run (166 frames per run).

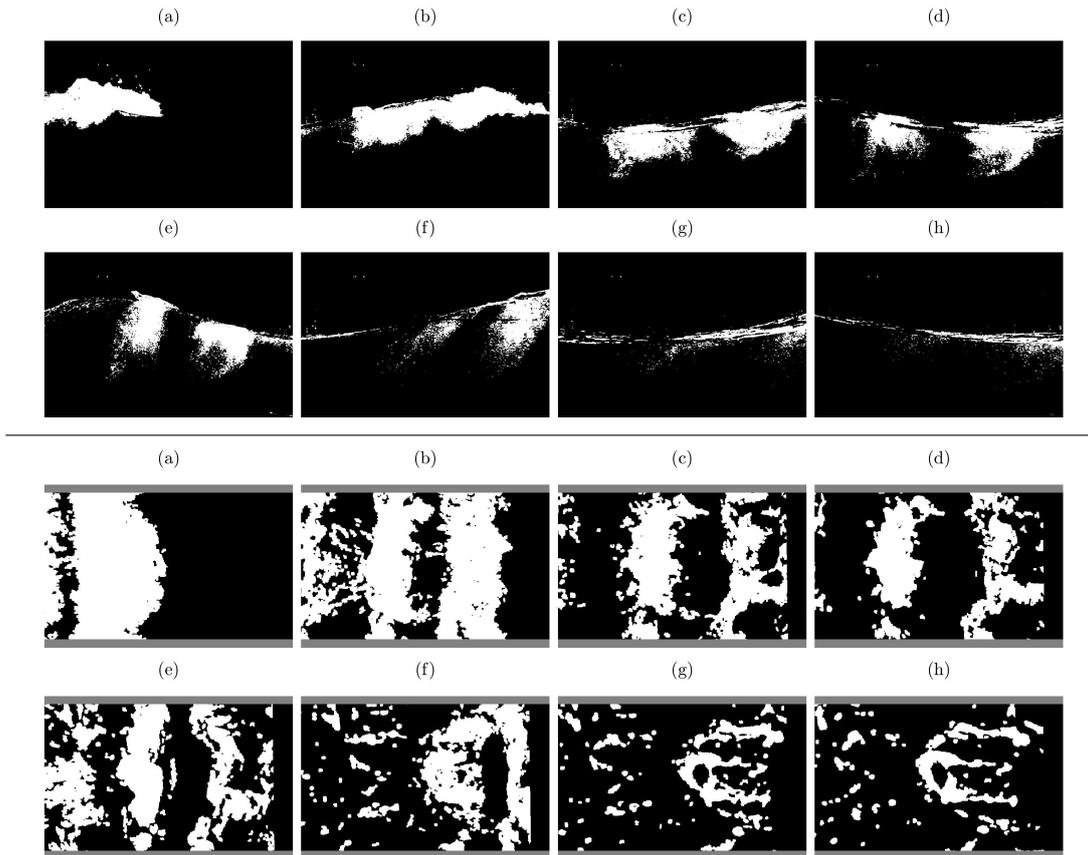


Figure 4. A sequence of thresholded B/W images of a breaking wave with a slope of $S = 0.37$ corresponding to the images in Figure 3.

257 The B/W masks were then used to calculate the foam and bubble plume areas. Figure
 258 5(a-d) shows the time series of the foam coverage for each wave packet. The foam coverage is
 259 defined as the fraction of the image area that is covered by the foam in each B/W mask. In
 260 each plot, the thick lines are the ensemble averages for each condition (slope and surfactant)
 261 and the shaded areas show one standard deviation of the samples. The oscillations in the
 262 foam coverage data and the existence of local peaks in the time series are due to the foam
 263 being advected in and out of the field of view of the camera by the orbital motion of the
 264 surface waves. The amount of foam generated by the breaking waves increases with the slope
 265 of the wave packet. Initially, there is little difference between clean water and surfactant-
 266 added cases. However, the longevity of the foam is increased for the cases with additional
 267 surfactants, as is apparent from the foam coverage values at later times. Furthermore, there
 268 is more variation in the amount of foam among individual runs with the same experimental
 269 condition for the cases with additional surfactants, especially at later times.

270 The bubble plume area is similarly extracted from the B/W bubble plume masks. The
 271 time series of the bubble plume area is shown in Figure 5(e-h). Similar to the foam coverage,
 272 the maximum bubble plume area increases with the wave packet slope but there is little
 273 difference between the clean water and surfactant-added cases, both in the amount and the
 274 persistence of the bubbles. The second peak in these plots (at $t \approx 2$ s) is caused by the
 275 rapid upward motion of the free surface that results in the stretching and dilation of the
 276 bubble plume (see Figure 4-e). The variations in the plume area values among different
 277 runs for $t > 4$ s is due to the residual surface foam appearing in the bubble-plume camera.

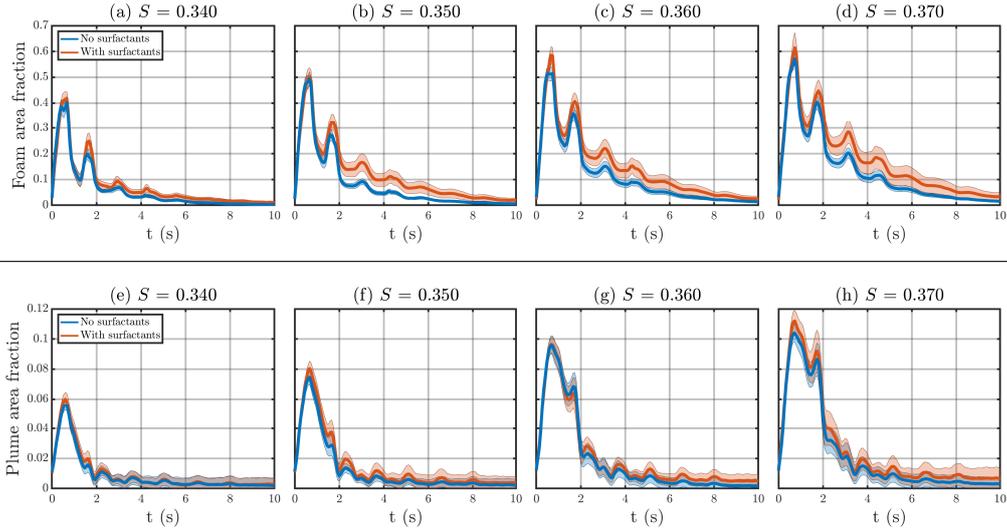


Figure 5. Time series of the foam area (top) and the bubble plume cross-sectional area (bottom). The area is normalized by the image size. The thick lines are the ensemble averages for each condition (slope and surfactant) and the shaded areas show one standard deviations of the samples.

278 This contamination of the plume area was reduced through processing and did not affect
 279 the correlation with the cooling time, presented below.

280 The visible foam timescale, τ_{foam} , and the bubble plume time scale, τ_{plume} , for each
 281 experimental run were calculated from their corresponding time series. An exponential
 282 function in the form of $A_{\text{plume}} = A_{\text{plume}}^{\text{max}} \exp(-t/\tau_{\text{decay}})$ was fit to the data between the time
 283 of maximum area and the time when the area drops below a threshold. The fitted curve was
 284 constrained to include the maximum area data point. The threshold used for the foam time
 285 series was five percent of the maximum area. For the plume area time series, the threshold
 286 varied for different wave packet slopes to reduce fitting error due to the influence of the
 287 residual foam at the surface which could be detected as bubble plume by the processing
 288 algorithm. The threshold values for these cases were between 0.2 and 0.4 of the maximum
 289 plume area. We define τ_{foam} and τ_{plume} as the sum of the growth and decay times for
 290 consistency with the definition of the parameter τ_{cool} , which is based on the time from the
 291 beginning of the breaking process. Thus, the time of maximum area is added to τ_{decay} to
 292 obtain the timescales for the foam and the plume. Figure 6 shows the visible foam timescale
 293 τ_{foam} versus the bubble plume time scale τ_{plume} . The small symbols correspond to the
 294 individual runs and the large symbols are the average timescales for each condition, with
 295 vertical and horizontal bars representing the standard deviation of the values. Consistent
 296 with the results of Callaghan et al. (2013), we find that τ_{foam} increases approximately
 297 linearly with τ_{plume} and there is an average increase of approximately 32 percent in τ_{foam}
 298 in the cases with added surfactants.

299 The sequence of surface temperature images in Figure 7 from an individual run cor-
 300 respond to the visible imagery in Figures 3 and 4. The blue lines are the boundaries of
 301 the regions covered by the visible foam and are derived from the masks in Figure 4. The
 302 temperature range is shown in the colorbar with dark corresponding to cold and bright to
 303 warm. For all the results that follow, the temperature is relative to a background refer-
 304 ence temperature for each individual run that was defined as the maximum value of the
 305 spatially-averaged temperature of the field in the time span of $0 < t < 10$ s.

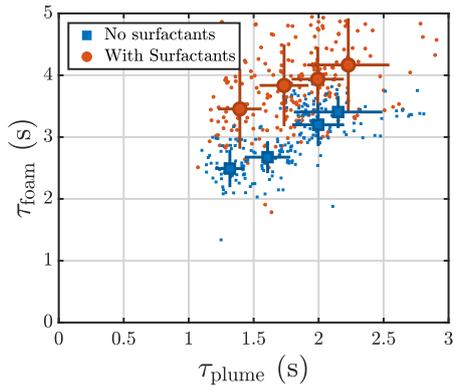


Figure 6. Visible foam timescale τ_{foam} versus bubble plume time scale τ_{plume} . Each small symbol represents one experimental run. The large symbols are the mean value for each wave packet slope S . The error bars show one standard deviations of the samples.

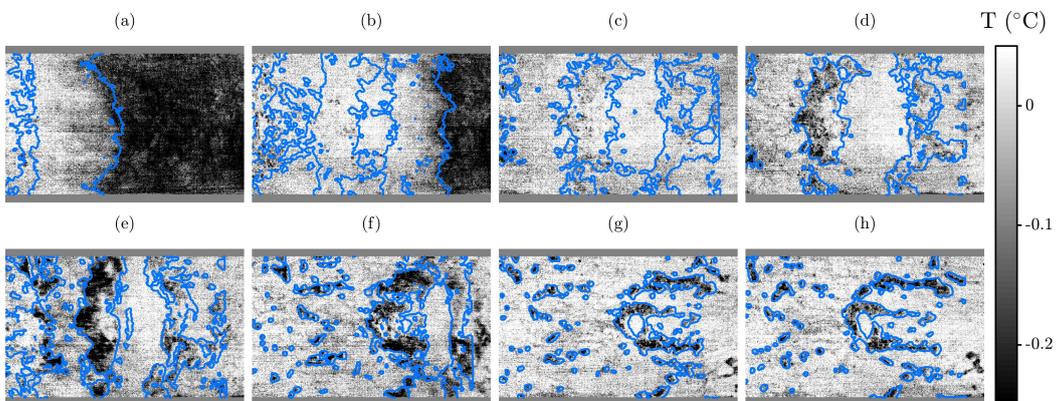


Figure 7. A sequence of infrared images corresponding to the images in Figure 3. The temperature range is shown in the colorbar with dark meaning cold and bright meaning warm. The blue outlines show the location of the foam extracted from the visible foam images (Figure 4). This figure corresponds to Movie S2 in the supplementary information.

306 As the crest begins breaking in Figure 7(a), the nearly uniform disruption of the cool
 307 skin layer produces a warm front over the entire width of the tank that advances with the
 308 crest. In frames (a)-(c), the temperature of the surface disrupted by the breaking crest is
 309 nearly uniform, regardless of whether it is foam-covered or foam-free. The foam begins to
 310 cool in (d) but the degree of cooling is not spatially uniform over the foam until the end
 311 of the sequence in (g)-(h). This spatial variation is apparent between the two main foam
 312 regions in (d) and (e). Eventually, the foam dissipates and after several minutes (not shown
 313 in the figure), the cool skin layer recovers and the surface temperature drops to its value
 314 before the disruption by the breaking wave.

315 The foam temperature was calculated for each run using the foam mask sequence extracted
 316 from the visible foam images (e.g. Figure 4). This mask was then applied to the
 317 corresponding frames of the infrared sequence to isolate the regions covered by the foam
 318 from the rest of the image. The mean temperature of the foam, T_{foam} , is plotted versus time
 319 in Figure 8 for each experimental condition. Immediately after the start of the breaking
 320 process, the temperature increases because of the disruption of the cool skin layer. The

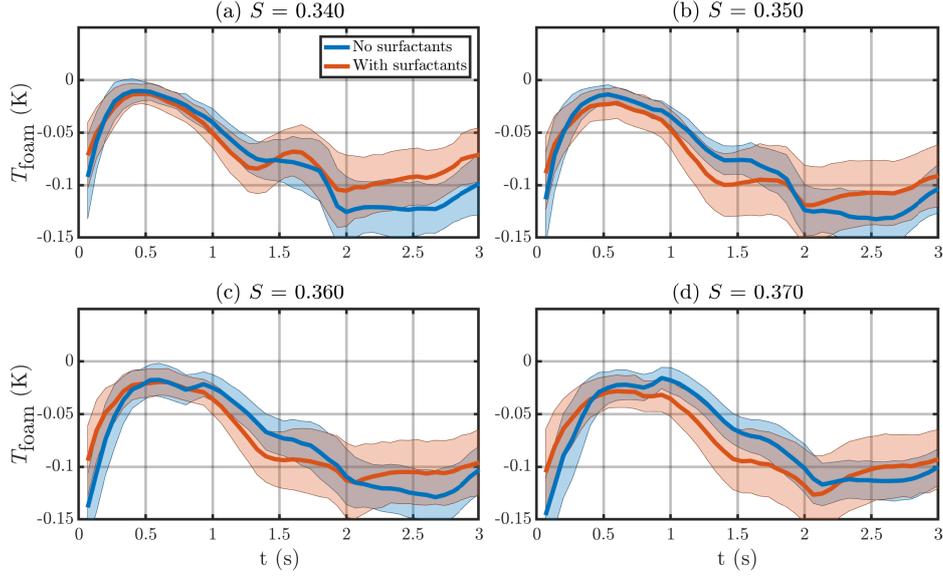


Figure 8. The mean temperature of the foam versus time, with and without added surfactants. The thick lines are ensemble averages and the shaded areas denote one standard deviations of the samples.

321 foam temperature plateaus for a short time and then starts to cool. These plots show that
 322 the duration of the plateau in foam temperature increases with the slope of the wave packet
 323 and the onset of the cooling of the foam is delayed for the larger breakers compared to the
 324 smaller ones. For the cases studied in these experiments, the time of the start of the cooling
 325 ranges approximately from $t \approx 0.5$ s to $t \approx 1.5$ s. Furthermore, the surfactant-free and
 326 surfactant-added cases follow each other closely in these plots.

327 We define the onset of cooling, τ_{cool} , as the time from the start of breaking to when
 328 the mean foam temperature T_{foam} falls below a fixed threshold equal to the noise level of
 329 the infrared camera of 0.025 K. Figure 9(a) shows the calculated value of τ_{cool} versus the
 330 decay time of the bubble plume, τ_{plume} . τ_{cool} varies approximately linearly with τ_{plume}
 331 and the difference between the surfactant-free and surfactant-added cases is within the
 332 experimental variation, indicated by the standard deviation of the ensemble. Figure 9(b)
 333 shows the individual data points from Figure 9(a) colored by the temperature difference
 334 between the water and the air, $\Delta T = T_{\text{water}} - T_{\text{air}}$, for each experimental run. This plot
 335 suggests that the onset of cooling is not strongly affected by the air-water heat flux, which
 336 is driven by the air-water temperature difference.

337 The method to generate these breaking events initially produces a single foam patch
 338 and corresponding bubble plume that then quickly separates into two distinct foam patch
 339 and bubble plume pairs, as shown in the time sequences from an individual run in Figure
 340 4. The distinct foam patches and associated bubble plumes are generated at different times
 341 and with different intensities and thus differ in their spatial and temporal evolution. For
 342 instance, the trailing bubble plume in Figure 4(c-e, bottom) is nearly dissipated in Figure
 343 4(f, bottom) when the leading plume is still robust, which is not necessarily reflected in
 344 the evolution of the foam patches in the corresponding panels in Figure 4. The occurrence
 345 of two separately evolving bubble plumes and corresponding foam patches suggests that
 346 separately tracking and measuring the evolution of τ_{cool} and τ_{plume} for one foam-plume pair
 347 may improve the correlation.

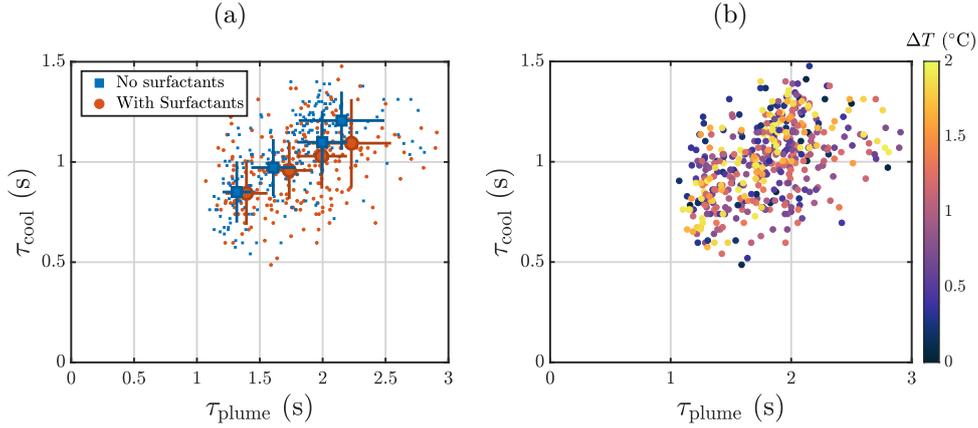


Figure 9. τ_{cool} versus τ_{plume} (a) The small symbols represent individual runs. The large symbols are the ensemble-averaged values at each condition and the error bars show one standard deviations. (b) Each circle represents one experimental run and is colored based on the water-air temperature difference, ΔT , shown in the colorbar.

348 The analysis region for individual foam patches was identified and tracked using ensemble-
 349 averaged intensity images for each slope and surfactant condition, as illustrated by the se-
 350 quence in Figure 10 (top) for $S = 0.37$ without additional surfactants. Each image is the
 351 result of averaging the same frame of foam masks relative to the start of breaking, among all
 352 the runs with the same conditions. Therefore, the intensity value at each pixel is equal to the
 353 fraction of runs in which the pixel was covered by foam or bubbles. The ensemble averaging
 354 of the foam images reveals the well defined two-dimensional structure of the foam. The
 355 single transverse strip of foam in Figure 10(a) quickly separates into two transverse strips in
 356 Figure 10(b) that continue to propagate down the tank as separate features. The rectangu-
 357 lar box overlaid on each image corresponds to the portion of the image used in analyzing the
 358 foam temperature in the individual run sequences of infrared imagery. The location and the
 359 width of the tracking box were defined manually by inspecting the ensemble-averaged foam
 360 images with the goal of containing most of the foam in a strip. Note that the tracking boxes
 361 were calculated based on the ensemble-averaged foam data and then applied to infrared
 362 images from individual runs with the same conditions. The corresponding bubble plume
 363 is similarly tracked using the ensemble-averaged bubble plume images (Figure 10, bottom)
 364 and then applied to individual runs. The location and width of the tracking window for the
 365 bubble plume were defined independently of the tracking window for the foam data. The
 366 reason was that the bubble plume is deformed greatly by the fluid motion at depth due to
 367 the waves and is not necessarily located directly beneath the foam strip.

368 The mean temperature of the foam in the tracking box is plotted versus time in Figure
 369 11(a-d). Similar to the results of Figure 8 for the whole field of view, the onset of cooling of
 370 the foam is delayed with the scale of the breaking strength, given by the wave packet
 371 slope, S . The difference between the surfactant-free and surfactant-added cases is once again
 372 relatively small. The bubble plume was similarly tracked as shown in Figure 10 (bottom).
 373 However, obtaining the time scale for individual bubble plumes proved to be problematic for
 374 the cases with the two larger slopes. The reason for this issue can be seen in the plume area
 375 time series shown in Figure 11(e-h). The individual bubble plume could only be tracked up
 376 to $t = 2$ s since the separate plumes merged together at that point. Fitting an exponential
 377 function to the plume area time series resulted in noisy data due to the lack of data at later
 378 stages of decay and the presence of a local peak in the time series caused by the stretching of
 379 the bubble plume. However, as can be seen in these plots, the time series for an individual

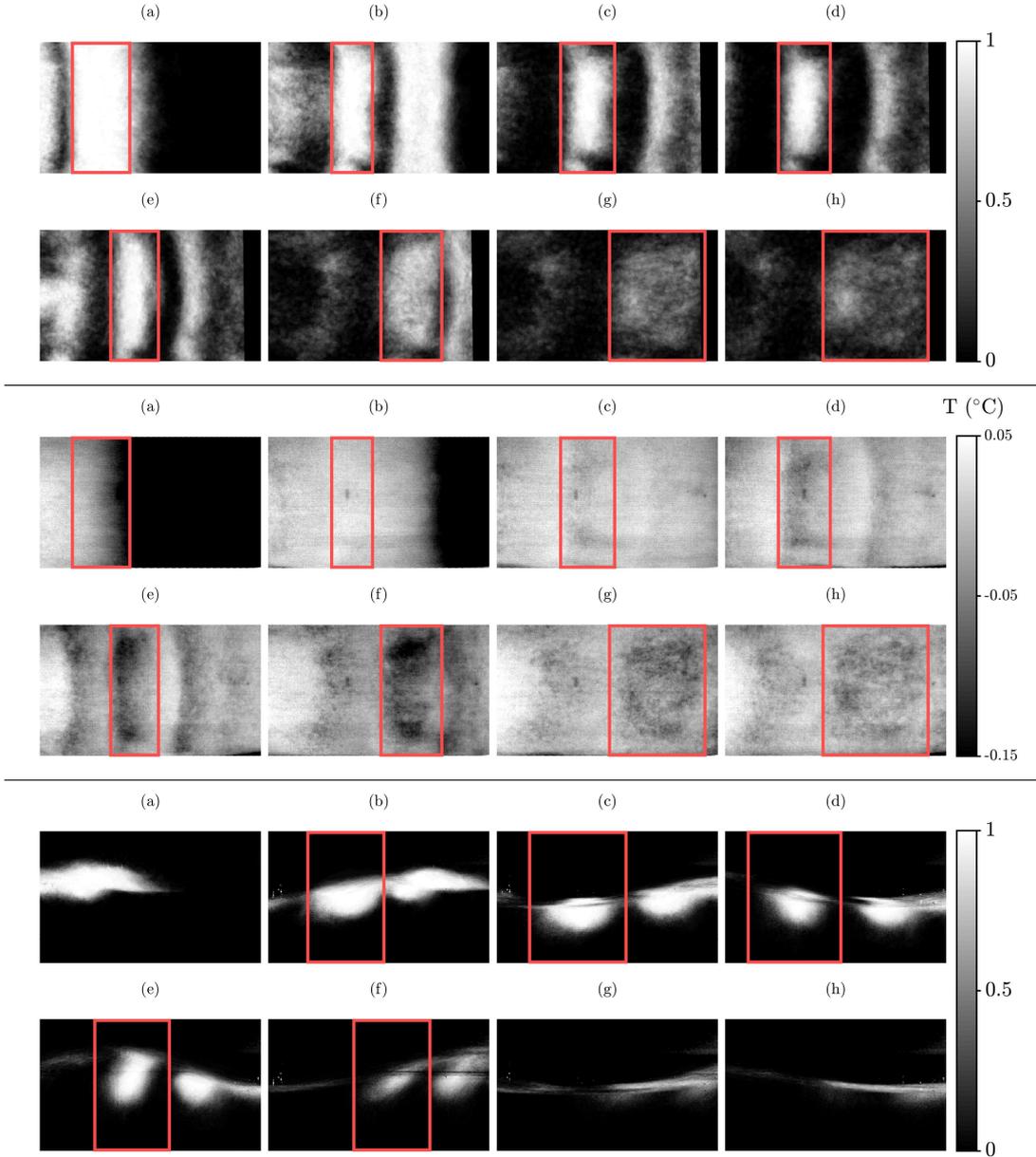


Figure 10. A sequence of ensemble-averaged intensity images of the foam (top), the surface temperature (middle), and the bubble plume (bottom) for the experimental condition with a slope of 0.37 without additional surfactants. The red box follows the most salient foam patch (and the corresponding bubble plume) of the breaker and shows the region that is used in analyzing the foam temperature and the plume area. This figure corresponds to Movie S4 in the supplementary information.

380 plume follows a trend similar to that for the whole image. Therefore, the maximum depth
 381 of the plume, d_{\max} is used instead of τ_{plume} to present the result of tracking an individual
 382 plume. Figure 12(a) and 12(b) show the onset of cooling, τ_{cool} , versus maximum depth
 383 of the plume for the whole field of view (i. e., both plumes) and for tracking the trailing
 384 plume, respectively. For both cases, τ_{cool} scales with d_{\max} and the difference between the
 385 two surfactant conditions is small (approximately 5 percent on average). Moreover, the

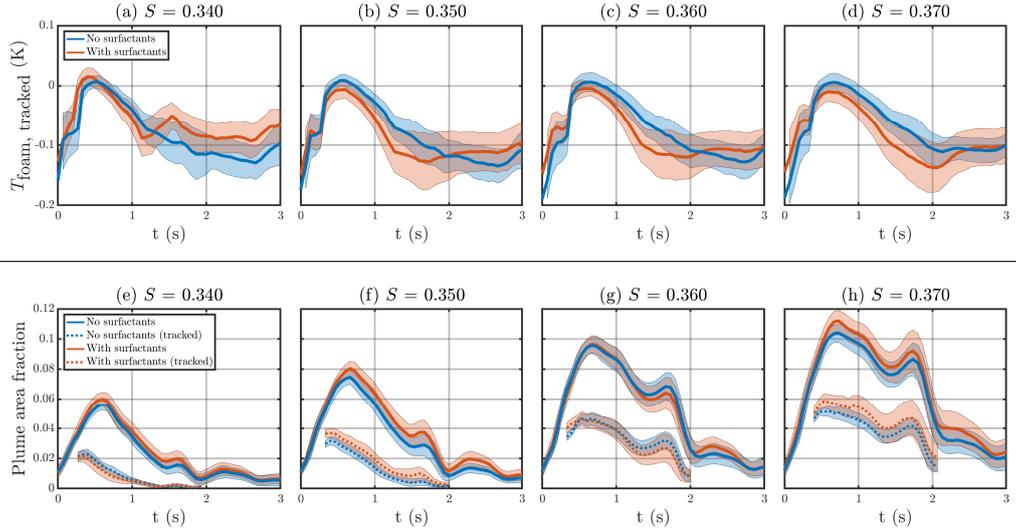


Figure 11. (Top) the mean temperature of the foam in the tracking box versus time, with and without additional surfactants. Thick lines are ensemble averages and the shaded areas are one standard deviations of samples. (Bottom) bubble plume area normalized by the image size. Solid lines are for the whole field of view and dotted lines are the data inside the tracking window.

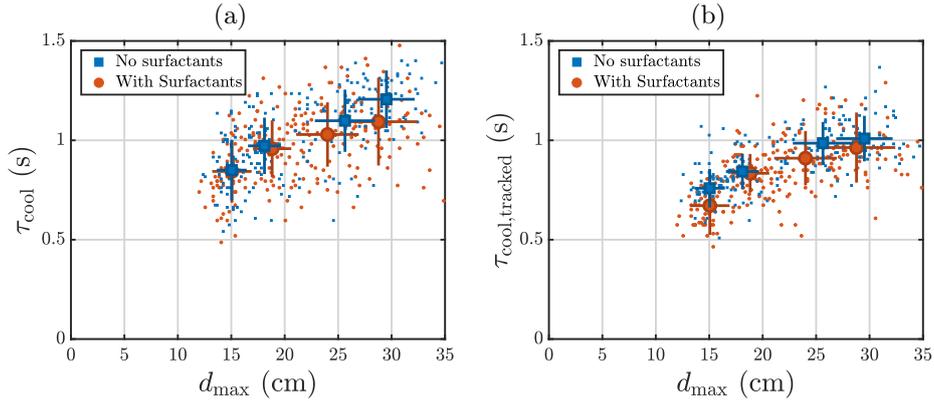


Figure 12. τ_{cool} versus the maximum depth of the plume, d_{max} (a) without tracking, and (b) with tracking. The small symbols represent individual runs. The large symbols are the ensemble-averaged values at each condition and the error bars show one standard deviations.

386 scatter of τ_{cool} data is less when an individual foam patch was tracked compared to the
 387 whole field of view (approximately 0.12 s compared to 0.16 s, respectively). The reduced
 388 scatter when tracking a single plume suggests that the spatial variation in τ_{cool} is related
 389 to the spatial variability of the bubble plume depth.

390 4 Conclusions

391 We presented an experimental investigation of the thermal signature of the residual
 392 foam left behind by breaking waves. The experiments were conducted in a saltwater wave
 393 tank and breaking waves were generated using the dispersive focusing wave packet technique.
 394 We used four different wave packets that had a similar shape but varied significantly in

395 breaking intensity, plume depth, and energy dissipation. For each packet, more than a
 396 hundred experiment runs were performed in salt water with and without added surfactants.
 397 The visible and thermal signatures of the surface foam produced by the breaking waves
 398 were measured. The foam area, bubble plume area, and foam temperature time series were
 399 calculated from the image sequence data for each experimental run. The visible foam and the
 400 bubble plume time scales, τ_{foam} and τ_{plume} , and the onset of cooling, τ_{cool} , were evaluated
 401 from their corresponding time series.

402 The time to the onset of cooling of the foam, τ_{cool} , was found to scale with the plume
 403 decay time, τ_{plume} , and the maximum plume depth. The cooling timescale was not signif-
 404 icantly affected by the environmental conditions of surfactant concentration and air-water
 405 temperature difference. Therefore, τ_{cool} can be used to infer sub-surface plume dynamics
 406 by quantifying the plume decay time and depth from above-surface observations.

407 Our results are consistent with the notion that surface foam cools faster than the sur-
 408 rounding clear water due to the enhanced cooling of the bubbles at the surface. Furthermore,
 409 they support the hypothesis that the surface foam begins to cool when the rising bubbles
 410 responsible for its generation are no longer present. Our results suggest that the observed
 411 spatial variability of τ_{cool} may provide information about the spatial variability of the bub-
 412 ble plume depth. Adequate investigation of this idea will require additional measurements
 413 with increased dynamic range of breaking intensity and significant spatial variability of the
 414 bubble plume depth.

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 421 <https://drive.google.com/open?id=1vvmYcuIyBRcGtfJQY4BddDx-HzLkAwUR> (we are in
 422 the process of placing the data in an appropriate repository).

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