

1 **New Findings from Explainable SYM-H Forecasting**
2 **using Gradient Boosting Machines**

3 **Daniel Iong¹, Yang Chen¹, Gabor Toth², Shasha Zou², Tuija Pulkkinen², Jiaen**
4 **Ren², Enrico Camporeale^{3,4}, Tamas Gombosi²**

5 ¹Dept. of Statistics, University of Michigan, Ann Arbor, MI, USA

6 ²Dept. of Climate and Space Sciences and Engineering, University of Michigan, Ann Arbor, MI, USA

7 ³CIRES, University of Colorado, Boulder, CO, USA

8 ⁴NOAA Space Weather Prediction Center, CO, USA

9 **Key Points:**

- 10 • We adapt gradient boosting machines (GBMs) for forecasting the SYM-H index
11 multiple hours ahead.
- 12 • We quantify feature contributions using Shapley additive explanation (SHAP) val-
13 ues to explain model predictions.
- 14 • Our proposed method has similiar accuracy to existing methods, while being more
15 interpretable.

Corresponding author: Daniel Iong, daniiong@umich.edu

Abstract

In this work, we develop gradient boosting machines (GBMs) for forecasting the SYM-H index multiple hours ahead using different combinations of solar wind and interplanetary magnetic field (IMF) parameters, derived parameters, and past SYM-H values. Using Shapley Additive Explanation (SHAP) values to quantify the contributions from each input to predictions of the SYM-H index from GBMs, we show that our predictions are consistent with physical understanding while also providing insight into the complex relationship between the solar wind and Earth’s ring current. In particular, we found that feature contributions vary depending on the storm phase. We also perform a direct comparison between GBMs and neural networks presented in prior publications for forecasting the SYM-H index by training, validating, and testing them on the same data. We find that the GBMs have a comparable root mean squared error as the best published black-box neural network schemes.

Plain Language Summary

Forecasting geomagnetic indices is crucial for mitigating potential effects of severe geomagnetic storms on critical infrastructures such as power grids. In this work, we adopt a machine learning method for SYM-H prediction hours ahead with various combinations of solar wind & interplanetary magnetic field parameters, past SYM-H values, and other derived parameters. The feature importance quantification that we derive provides important, new insight into the complex relationship between the solar wind and the Earth’s ring current.

1 Introduction

Geomagnetic storms are the largest geomagnetic disturbances, during which severe space weather threats can occur and disrupt our technological society. During geomagnetic storms, petajoules of energy enter the Earth’s magnetosphere from the solar wind, of which vast majority is stored in the ring current in the inner magnetosphere (Ganushkina et al., 2017). The ring current indices such as Dst and SYM-H provide essential information about the current strength and evolution as well as the energy budget, and thus are of crucial practical importance (Sugiura & Kamei, 1991). These ring current indices have been used in numerous space weather applications, such as in classification of storms, as inputs to empirical models of the magnetospheric magnetic topology (N. Tsyganenko, 1989; N. A. Tsyganenko, 1995, 2002a, 2002b), as features representing the geomagnetic activity level for machine learning forecasting the ionospheric total electron content (Liu et al., 2020), as parameters used for forecasting of the radiation belt energetic particle fluxes (Sakaguchi et al., 2015) and other magnetospheric quantities (Bortnik et al., 2018). Therefore, the ability to predict the ring current indices is crucial for space weather forecasts and end-users.

Several attempts have been made to use machine learning methods to forecast the SYM-H index. Cai et al. (2010) and Bhaskar and Vichare (2019) used a Nonlinear AutoRegressive with eXogeneous inputs (NARX) neural network to predict 5-minute averages of the SYM-H index one hour ahead using past SYM-H values, solar wind and IMF parameters as input. Cai et al. (2010) trained their neural networks with data from 67 geomagnetic storms from 1998 to 2006, while Bhaskar and Vichare (2019) used data from 25 additional geomagnetic storms from 2006 to 2013. With the goal of developing operationally feasible models, Siciliano et al. (2021) trained long short-term memory (LSTM) and convolutional (CNN) neural networks to predict the SYM-H index one hour ahead using only IMF parameters and past SYM-H values as input. Collado-Villaverde et al. (2021) took a similar approach to predict the SYM-H index several hours ahead, while also considering the effects of omitting past SYM-H values as input on predictive performance. Both Siciliano et al. (2021) and Collado-Villaverde et al. (2021) train and val-

66 idate their networks on 25 strong geomagnetic storms ($Dst < -100$ nT) from 1998 to 2017
67 and evaluate their performance using 17 strong test storms. To conduct a direct com-
68 parison of predictive performance, we use the same storms and features to train and test
69 our proposed model. For the rest of this article, we will use the terms features and (model)
70 inputs interchangeably. Comparison results are discussed in section 4.1.

71 Many machine learning approaches have been taken to forecast the Dst index and
72 other geomagnetic indices such as the Kp index. Attempts to apply machine learning
73 methods to forecast the Dst index date back to the works of Lundstedt and Wintoft (1994),
74 Gleisner et al. (1996), and Wu and Lundstedt (1997). These authors generally observed
75 that the initial and main phases were more accurately predicted than the recovery phase
76 when the Dst index is not used as an input due to the fact that the initial and main phases
77 are more strongly correlated with solar wind properties. Pallochia et al. (2006) advo-
78 cated for using only IMF parameters as inputs for operational forecasting of the Dst in-
79 dex because in situ solar wind plasma instruments tend to fail more often than space-
80 based magnetometers. This was also the motivation for using only IMF parameters and
81 past values to forecast the SYM-H index in Siciliano et al. (2021) and Collado-Villaverde
82 et al. (2021).

83 Although the majority of machine learning approaches to forecasting geomagnetic
84 indices use neural networks, other techniques have also been proposed: Chandorkar et
85 al. (2017) investigated the use of Gaussian Processes for forecasting the Dst index; Lu
86 et al. (2016) compared the use of support vector machines (SVM) with neural networks;
87 Boynton et al. (2011) employed the Nonlinear AutoRegressive Moving Average with eX-
88 ogeneous inputs (NARMAX) model to derive an analytic expression to forecast 1-hour-
89 ahead Dst as function of its past values and of the history of a solar wind-magnetosphere
90 coupling function. Xu et al. (2020) combined neural networks with SVM to construct
91 an ensemble model using bagging to predict the Dst index up to six hours ahead. We
92 also construct an ensemble model but use gradient boosting instead of bagging (see Bauer
93 and Kohavi (1999) for a detailed comparison between boosting and bagging). Another
94 difference is that we create an ensemble of many simple tree-based models as opposed
95 to a few complex models. A comprehensive review of machine learning models for geom-
96 agnetic indexes can be found in Camporeale (2019).

97 Despite the fact that data-driven machine learning methods have made a lot of progress
98 in many scientific fields and have become popular tools, the lack of interpretability has
99 been a major drawback. Even if machine learning methods have typically focused on pre-
100 dictive performance, there has been a recent surge in interest in making these methods
101 more interpretable (Molnar et al., 2020). The development of interpretable machine learn-
102 ing algorithms is of key importance especially in scientific fields such as space weather.
103 In spite of the fact that machine learning methods have repeatedly been shown to out-
104 perform operational models empirically, these methods have not been widely adopted
105 in an operational setting due to a lack of trust and skepticism from the space weather
106 community (Camporeale, 2019). Interpretability gives confidence to operational forecast-
107 ers that relevant physical processes are captured to some degree and encoded in a black-
108 box model, hence reassuring of its generalizability and robustness versus rare events, which
109 are the main focus of space weather forecasting. Gray-box approaches, which combine
110 physics-based models with black-box models, can also be used to make machine learn-
111 ing methods for space weather forecasting more reliable (Camporeale et al., 2020).

112 Explainability can be achieved by using either post-hoc explanation methods or
113 intrinsically interpretable models. Examples of intrinsically interpretable models include
114 linear regression, decision trees, and generalized additive models. Unfortunately, there
115 is often a tradeoff between intrinsic model interpretability and predictive performance
116 because interpretable models tend to make strong simplifying assumptions such as lin-
117 earity or additivity. Recent efforts have been made to close this gap, starting with ad-
118 ditive models that incorporate two-way feature interactions (Lou et al., 2013). Post-hoc

119 explanation methods, to some extent, can be used to explain the predictions made by
 120 more complex models, usually by constructing an approximate interpretable model af-
 121 ter training the original model. For an overview of interpretable machine learning meth-
 122 ods, see Molnar (2019). Several intrinsically interpretable models have previously been
 123 proposed for forecasting geomagnetic indices. Ayala Solares et al. (2016) proposed a Non-
 124 linear Autoregressive with Exogeneous Inputs (NARX) model to forecast the Kp index
 125 where the contribution of each model term to the output can be evaluated. Gu et al. (2019)
 126 proposed an interpretable NARX model the forecast the AE index that also includes un-
 127 certainty analysis.

128 In this work, we not only aim to obtain accurate predictions of the SYM-H index,
 129 but more importantly, to learn if the data-driven approach can reveal insights on the phys-
 130 ical mechanisms. In turn, these insights could then be used to inform future physics-based
 131 or grey-box models. We achieve this by using a post-hoc explanation method known as
 132 Shapley Additive Explanations (SHAP) to quantify the contributions from each input
 133 on the predictions made by gradient boosting machines (Lundberg & Lee, 2017). SHAP
 134 has been successfully used to explain predictions from tree-based models in other scien-
 135 tific fields such as medicine (Lundberg et al., 2018), solar power forecasting (Kuzlu et
 136 al., 2020; Mitrentsis & Lens, 2021), finance (Bluwstein et al., 2020; Mokhtari et al., 2019),
 137 and atmospheric science (Stirnberg et al., 2020). Section 3.2 continues this discussion
 138 on explainability and describes the SHAP method in detail.

139 The remainder of the paper is organized as follows. In Section 2, we introduce the
 140 data sources and our data processing procedures. In Section 3, we describe the gradi-
 141 ent boosting machine, hyperparameter tuning, and quantification of feature importance.
 142 In Section 4, we provide results of our predictions, compare them with those published
 143 in the existing literature, and most importantly, the new insights that we learn from the
 144 prediction model results. We conclude in Section 5 with a summary on key findings and
 145 some discussions on future work.

146 2 Data

147 The Disturbance Storm Time (Dst) index is computed as the H (magnetic north)
 148 component perturbation on equatorial magnetometers (Mayaud, 1980) on an hourly ba-
 149 sis, and is a characterization of a magnetic storm that has been used historically. The
 150 Dst index represents the longitudinally averaged part of the external geomagnetic field
 151 measured at the equator (Sugiura, 1964). As the index includes only the field variation,
 152 during geomagnetically quiet times, it hovers around zero. The typical definition of a
 153 geomagnetic storm is that the Dst index reaches values below -50 nT.

154 The SYM-H index is a high-time-resolution version of the original Dst index, and
 155 is given at 1-minute cadence (Iyemori, 1990; Wanliss & Showalter, 2006). The SYM-H
 156 index is compiled from 11 low- and mid-latitude magnetometer stations. Quiet time fields,
 157 including local time and seasonal quiet time Sq current effects, are removed, and the resid-
 158 uals are averaged together, divided by the cosine of the co-latitude of the station to yield
 159 the component parallel with the magnetic dipole. Geomagnetic storms can be classified
 160 based on the SYM-H values: moderate (-100 nT $<$ SYM-H $<$ -50 nT), intense (-250 nT
 161 $<$ SYM-H $<$ -100 nT), and superstorms (SYM-H $<$ -250 nT).

162 We extract the SYM-H index data from the OMNI dataset compiled at NSSDC
 163 (<https://spdf.gsfc.nasa.gov>) using the open-source Python library *swmcpy* (King,
 164 2005; Al Shidi, Qusai, 2020). We use the level-2 solar wind plasma and interplanetary
 165 magnetic field (IMF) parameters from the Advanced Composition Explorer (ACE) space-
 166 craft provided by the NASA Space Physics Data Facility ([https://cdaweb.gsfc.nasa](https://cdaweb.gsfc.nasa.gov/index.html/)
 167 [.gov/index.html/](https://cdaweb.gsfc.nasa.gov/index.html/)) as inputs in our models. The original dataset contains the IMF com-
 168 ponents from the ACE Magnetic Field Experiment (MAG) instrument (Smith et al., 1998)

169 at a 16-second cadence, as well as proton density, bulk speed, and ion temperature from
 170 the SWEFAM suite (McComas et al., 1998), at a 64-second cadence. In addition to so-
 171 lar wind plasma and IMF parameters, we also include derived quantities, in particular
 172 the solar wind dynamic pressure and electric field, as we expect them to be relevant in-
 173 put parameters for predicting geomagnetic storms (Newell et al., 2007).

174 Explanation methods, such as SHAP, allow us to confirm or disprove these expect-
 175 ations. To remove some of the high frequency variation inherent in high time resolu-
 176 tion data and to eliminate minor data gaps, we average the SYM-H index, solar wind
 177 and IMF parameters to a 5-min time resolution. This was also done by Collado-Villaverde
 178 et al. (2021); Siciliano et al. (2021).

179 For training and testing the GBMs discussed in section 3.1, we use 42 strong ge-
 180 omagnetic storms occurring between 1998 to 2018 which reached a minimum SYM-H in-
 181 dex value of less than -100 nT. Information about these storms are given in tables 1 and 2.
 182 We use 5-fold cross validation to optimize hyperparameters (see section 3.1) instead of
 183 using a separate set of storms for validation, which allows us to use more data for train-
 184 ing models. Descriptive statistics for the training and test storms are given in tables A1
 185 and A2.

Table 1. Storms used to train GBMs. These storms are identical to the ones used to train and
 validate models in Collado-Villaverde et al. (2021).

Storm #	Start date	End date	Min. SYM-H (nT)
1	1998-02-14	1998-02-22	-119
2	1998-08-02	1998-08-08	-168
3	1998-09-19	1998-09-29	-213
4	1999-02-16	1999-02-24	-127
5	1999-10-15	1999-10-25	-218
6	2000-07-09	2000-07-19	-335
7	2000-08-06	2000-08-16	-235
8	2000-09-15	2000-09-25	-196
9	2000-11-01	2000-11-15	-174
10	2001-03-14	2001-03-24	-165
11	2001-04-06	2001-04-16	-275
12	2001-10-17	2001-10-22	-210
13	2001-10-31	2001-11-10	-313
14	2002-05-17	2002-05-27	-113
15	2003-11-15	2003-11-25	-488
16	2004-07-20	2004-07-30	-208
17	2005-05-10	2005-05-20	-302
18	2006-04-09	2006-04-19	-110
19	1998-12-09	1998-12-19	-206
20	2012-03-01	2012-03-11	-149
21	1998-04-28	1998-08-05	-268
22	1999-09-19	1999-09-26	-160
23	2003-10-25	2003-11-03	-427
24	2015-06-18	2015-06-28	-207
25	2017-09-01	2017-09-11	-144

186 To predict SYM-H Δt hours ahead of time t , henceforth denoted as $y(t+\Delta t)$, we
 187 will consider different combinations of the features listed in table 3. We also consider lead
 188 times Δt of one and two hours. When the SYM-H index is included, the observations
 189 from the previous one hour are used as input. We set the history length for all other fea-

Table 2. Storms used to test GBMs. These storms are identical to the ones used to test models in Collado-Villaverde et al. (2021).

Storm #	Start time	End time	Min. SYM-H (nT)
26	1998-06-22	1998-06-30	-120
27	1998-11-02	1998-11-12	-179
28	1999-01-09	1999-01-18	-111
29	1999-04-13	1999-04-19	-122
30	2000-01-16	2000-01-26	-101
31	2000-04-02	2000-04-12	-315
32	2000-05-19	2000-05-28	-159
33	2001-03-26	2001-04-04	-434
34	2003-05-26	2003-06-06	-162
35	2003-07-08	2003-07-18	-125
36	2004-01-18	2004-01-27	-137
37	2004-11-04	2004-11-14	-393
38	2012-09-10	2012-10-05	-138
39	2013-05-28	2013-06-04	-134
40	2013-06-26	2013-07-04	-110
41	2015-03-11	2015-03-21	-233
42	2018-08-22	2018-09-03	-205

190 tures to be either two hours, if the SYM-H index is included, or 30 hours, if the SYM-
191 H index is excluded. The history length selections were motivated by Siciliano et al. (2021),
192 who examined the coefficient of determination R^2 that quantifies the amount of observed
193 variance that is explained by the predictions as a function of the history length, when
194 the SYM-H index was either included or excluded as an input. They found that R^2 started
195 to decrease when the history length was around 30 hours, if the SYM-H index was not
196 included as input. When the SYM-H index was included as input, the R^2 results for his-
197 tory lengths of 90 to 180 minutes were similar, while R^2 started to decrease for time in-
198 tervals longer than 180 minutes.

Table 3. Features used as input into our models.

Features	History length (in hours)
Past SYM-H index (nT)	1
IMF: B_x, B_y, B_z (nT)	2 or 30
Solar wind: V_x (km/s), ρ (amu/cm^3), T (K)	2 or 30
Derived quantities: ρV_x^2 (nPa), $E_s = \max(0, - V_x B_z)(\text{mV}/\text{m})$	2 or 30

199 The different sets of features used as inputs are listed in table 4. Using different
200 sets of features to train our models allows us to investigate how the inclusion of certain
201 features affects predictive performance and feature contributions. The choice to train our
202 models using only IMF parameters and past SYM-H (input set I_1 , table 4) was moti-
203 vated by the high percentage of missing observations for solar wind plasma parameters.
204 For IMF parameters and solar wind velocity, there is less than 2% of observations miss-
205 ing within our sample. However, this percentage is substantially higher (roughly 9%) for
206 solar wind density and temperature. Although our proposed model handles missing data
207 internally, we choose to impute missing observations using linear interpolation (see sec-
208 tion 3.4 in Chen and Guestrin (2016) for details).

209 Including solar wind plasma and derived parameters in input sets I_3 and I_4 allows
 210 us to investigate how these contribute to predictions. In particular, a sudden increase
 211 of dynamic pressure ρV_x^2 can compress the magnetosphere and cause a positive jump in
 212 SYM-H, which typically happens at the beginning of the geomagnetic storms (sudden
 213 storm commencement). Another physically important parameter is the y component of
 214 the interplanetary electric field $E_y = V_x B_z$ that characterizes the amount of north-south
 215 magnetic flux carried by the solar wind. Note that $V_x < 0$ in the geocentric-solar-magnetic
 216 (GSM) coordinate system used here. The rectified electric field $E_s = \max(0, E_y)$ is the
 217 same as E_y when the IMF has a southward component ($B_z < 0$), which facilitates the
 218 onset of dayside reconnection, and zero for northward IMF when dayside reconnection
 219 is limited to high latitudes beyond the polar cusps (Burton et al., 1975). Including E_s
 220 would allow us to compare and contrast its contribution to predictions using the Bur-
 221 ton equation (T. P. O’Brien & McPherron, 2000; T. P. O’Brien, 2002).

222 To examine how solar wind and IMF parameters influence predictions without knowl-
 223 edge of past SYM-H values, we train models with input sets I_2 and I_4 which exclude past
 224 SYM-H values (see Table 4).

Table 4. Various sets of features used as inputs to train our models.

Input set	Features included
I_1	IMF, past SYM-H
I_2	IMF
I_3	IMF/solar wind/derived quantities, past SYM-H
I_4	IMF/solar wind/derived quantities

225 3 Methods

226 3.1 Gradient Boosting Machines

227 Gradient boosting machines (GBMs), also known as gradient boosted trees, have
 228 had considerable success in prediction tasks across a wide range of domains (Natekin &
 229 Knoll, 2013). Shwartz-Ziv and Armon (2021) recently performed a rigorous study show-
 230 ing GBMs outperformed several neural network models in terms of accuracy in classi-
 231 fication and regression problems with tabular data. GBMs are consistently used in the
 232 winning solutions of various machine learning prediction competitions like Kaggle, show-
 233 ing its effectiveness in a wide range of problems (Chen & Guestrin, 2016). In the space
 234 sciences, GBMs and other ensemble methods have recently been used to predict ambi-
 235 ent solar wind flow (Bailey et al., 2021) and the Dst index (Xu et al., 2020).

236 In contrast to algorithms that construct one complex model, gradient boosting se-
 237 quentially constructs simple prediction models called base learners that improve upon
 238 previously constructed base learners and sums them together to obtain an ensemble model.
 239 This process is analogous to how gradient descent optimizes weights in a neural network.
 240 Seen as a form of “functional gradient descent”, gradient boosting minimizes an objec-
 241 tive function by iteratively adding a new base learner, usually a decision tree, that leads
 242 to the largest decrease in the loss function (Friedman, 2001). In the case of GBMs, the
 243 base learners are regression trees, which are a highly interpretable class of machine learn-
 244 ing models that mimic human decision-making but are often too simplistic for most pred-
 245 iction problems when used alone. Fortunately, ensembles of regression trees, like GBMs,
 246 are capable of producing highly accurate predictions while still taking advantage of the
 247 interpretability of regression trees. In addition to gradient boosting, bagging is another
 248 widely used ensemble method that constructs multiple base learners in parallel and ag-
 249 gregates them by averaging (Breiman, 1996).

The gradient boosting machines that we use to forecast SYM-H have the form

$$y(t + \Delta t) = \alpha + \sum_{m=1}^M T_m(I(t)) + \epsilon(t), \quad t = 1, \dots, N, \quad (1)$$

where $I(t)$ is a vector of inputs used at time t ; $\epsilon(t)$ is an error term at time t ; T_m 's are regression trees; M is the number of iterations (trees) in the training algorithm; N is the number of timepoints; and α is a constant intercept term. $I(t)$ depends on which input set from table 4 is used. For instance, if input set I_2 is used, $I(t) = (B_x(t), \dots, B_x(t-115), B_y(t), \dots, B_y(t-115), B_z(t), \dots, B_z(t-115))$, where, for example, $B_z(t-60)$ denotes the value of B_z 60 minutes prior. The regression trees can be written mathematically as

$$T(x) = w_{q(x)}, \quad (2)$$

250 where w are the leaf weights of the tree; and q represents the tree structure by mapping
 251 an input to its corresponding leaf node index. Figure 1 shows the tree structure of one
 252 of the trees in a GBM that we trained.

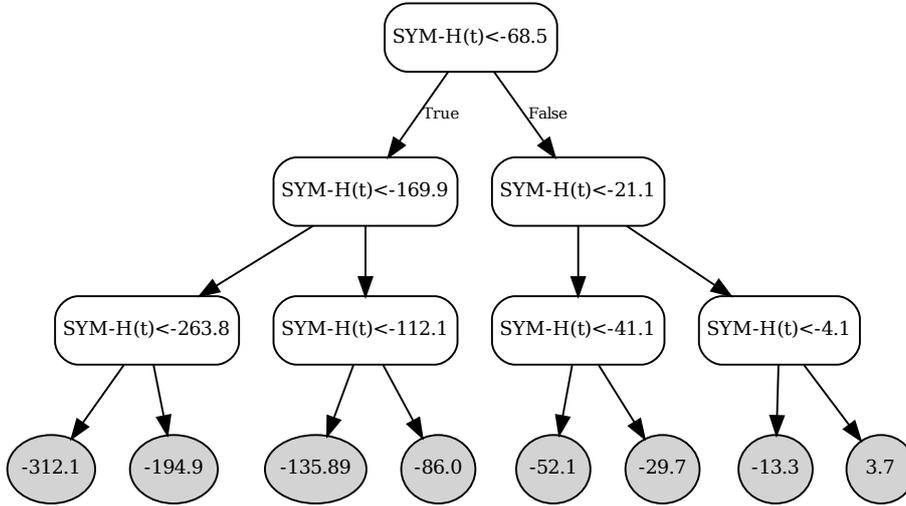


Figure 1. Structure of the first tree T_1 learned in a GBM trained with input set I_3 to predict the SYM-H index one hour ahead. The leaf nodes of the tree are shaded gray. The value in each leaf node is its corresponding leaf weight. Left splits correspond to the inequality in the previous node being true, and vice versa.

To train our GBMs, we use the open-source framework XGBoost that constructs the regression trees using gradient boosting and penalizes trees that are overly complex to avoid overfitting (Chen & Guestrin, 2016). More specifically, at each iteration m , we will construct a new regression tree T_m by minimizing the following objective function.

$$\mathcal{L}^{(m)}(T_m) = \sum_{t=1}^N \left\{ y(t + \Delta t) - [\hat{y}^{(m-1)}(t + \Delta t) + T_m(I(t))] \right\}^2 + \sum_{j=1}^m \Omega(T_j), \quad (3)$$

$$\text{where } \hat{y}^{(m-1)}(t + \Delta t) = \sum_{k=1}^{m-1} T_k(I(t)) \text{ and } \Omega(T_j) = \gamma K_j + \frac{1}{2} \lambda \sum_{k=1}^{K_j} w_{j,k}^2. \quad (4)$$

253 In eq. (4), K_j is the number of leaf nodes in T_j ; $w_{j,k}$'s are the leaf node weights in T_j ;
 254 and γ and λ are regularization hyperparameters. Ω is a regularization term that penal-
 255 izes the complexity of the regression trees by limiting the number of leaf nodes and shrink-
 256 ing the leaf weights. Increasing γ results in shallower trees while increasing λ leads to
 257 smaller leaf weights. An alternative method for controlling tree size is to explicitly set
 258 the maximum tree depth. Besides increasing λ , we can also reduce the influence of in-
 259 dividual trees by scaling their leaf weights by a learning rate. It is typically impossible
 260 to enumerate over all tree structures when constructing each regression tree. XGBoost
 261 takes a greedy approach that starts from a single leaf and iteratively adds branches to
 262 the tree that results in the largest loss reduction. This step involves finding the optimal
 263 feature and value to split the tree. Algorithms for splitting the tree are described in more
 264 detail in section 3 of Chen and Guestrin (2016).

265 To reduce the risk of overfitting, we control model complexity by optimizing sev-
 266 eral hyperparameters: learning rate, maximum tree depth, feature subsampling percent-
 267 age, minimum child weight, and number of boosting iterations (trees). We optimize these
 268 hyperparameters, except the number of iterations, using cross validation and a gradient-
 269 free optimization platform called Nevergrad (Rapin & Teytaud, 2018). To set the num-
 270 ber of iterations (trees), we monitor performance using cross validation at each iteration
 271 and terminate the algorithm when the performance stops improving. This technique is
 272 commonly referred to as early stopping in the machine learning literature (Zhang & Yu,
 273 2005). Cross validation is performed by first splitting the training storms in table 1 into
 274 5 sets. After that, we use each set for evaluation while training the model using the other
 275 4 sets. We repeat this procedure four times until all sets have been used for evaluation.
 276 Using cross validation, as opposed to a separate validation set, allows us to use more data
 277 when training the final model. The specific hyperparameter values we set are given in
 278 table 5.

Table 5. Hyperparameter values for training GBMs using the different input sets in table 4.

Input set	Hyperparameter	Value
I_1, I_2	Learning rate	0.072
	Max. tree depth	4
	Min. child weight	4
	Column subsampling %	0.78
	# of trees	84
I_3, I_4	Learning rate	0.147
	Max. tree depth	3
	Min. child weight	2
	Column subsampling %	0.894
	# of trees	291

279 GBMs have several advantages over competing machine learning methods. GBMs,
 280 and tree-based methods in general, are invariant to monotonic transformations of the
 281 features so it is better equipped to handle inputs on different scales. A practical conse-
 282 quence of this property is that the features don't have to be standardized before train-
 283 ing. GBMs are robust against issues arising from correlated features due to the greedy
 284 nature of gradient boosting and how regression trees are constructed. A downside of tree-
 285 based models for time series forecasting is that they produce predictions that are not smooth
 286 due to the tree structure of the model (Hastie et al., 2001). This can be seen in fig. 2,
 287 where the predictions from our GBM looks noisier than the ones from LSTM. Despite
 288 this property, GBMs are still able to produce highly accurate predictions. Another dis-
 289 advantage is that regression trees do not extrapolate well so they may exhibit sporadic

behavior when predicting with inputs that have values outside of the bounds of the inputs used for training. Fortunately, as seen in tables A1 and A2, the features in our test set are mostly within the bounds of the features in the training set.

GBMs can also suffer from over-specialization, wherein trees added in later iterations tend to only impact the predictions of a few instances (Korlakai Vinayak & Gilad-Bachrach, 2015). This may make the model highly sensitive to the contributions of the initially added trees. This issue is combated, to some extent, by selecting a small learning rate. To further alleviate this issue, we use a technique for employing dropouts in GBMs introduced by Korlakai Vinayak and Gilad-Bachrach (2015). Dropouts have been used successfully in neural networks, where a random subset of connections in the network is dropped during training (Srivastava et al., 2014). In the context of GBMs, at each training iteration, we replace $\hat{y}^{(m-1)}$ in eq. (3) with the sum of a random subset, instead of all, of the previously constructed trees and then normalize the newly constructed tree and dropped trees. Further details of this procedure can be found at (Korlakai Vinayak & Gilad-Bachrach, 2015).

3.2 Feature Importance

Methods for computing feature contribution, or feature importance, can be categorized as global versus local and model-specific versus model-agnostic. Global feature importance scores are used to explain a model’s overall behavior across the entire training dataset, while local feature importance scores tells you how individual features contributed to a single prediction. Model-specific feature importance is provided directly by the model, while model-agnostic methods, such as SHAP, typically construct an approximate interpretable model to explain predictions from the original model. For tree-based models, global feature importance can be calculated using information gain (Breiman et al., 1984), permutation (Breiman, 2001), or split count (Chen & Guestrin, 2016). In this paper, we will focus primarily on local feature importance as the contribution from each feature is likely to vary over time depending on the storm phase.

While there are several methods for computing local feature contribution in tree-based models (Molnar, 2019), we chose to use Shapley additive explanation (SHAP) because of its desirable theoretical properties (Lundberg & Lee, 2017). SHAP is based on Shapley values in cooperative game theory (Shapley, 1953), where they are used to fairly distribute payoffs in a game among a coalition of players with unequal contributions. In the case of SHAP, the payoff is the prediction and the players are the features. SHAP belongs to the class of additive feature attribution methods which assumes the following linear explanation model for an individual prediction.

$$g(\mathbf{z}) = \phi_0 + \sum_{i=1}^p \phi_i z_i, \quad (5)$$

where ϕ_0 is a reference value (e.g. mean); p is the number of input features; $\mathbf{z} = (z_1 \dots z_p)'$, where z_i is a binary variable indicating whether feature i is present; and ϕ_i is the contribution from feature i . SHAP yields the unique solution to eq. (5) that satisfies three desirable theoretical properties: local accuracy, missingness, consistency. The local accuracy property ensures that the sum of feature contributions for given inputs sum up to the prediction. The consistency property ensures that the SHAP value for a feature increases if the marginal contribution from that feature increases. Missingness is mainly a theoretical property that says a missing feature has zero contribution. The only alternative tree-specific local explanation method that we are aware of is Saabas (2014), which doesn’t have the consistency property. SHAP values describe a particular model’s decision-making process based on the data. Therefore, they can only be used to gain insight into the data-generating process when the model approximates the underlying process well enough. Furthermore, the effect that multicollinearity has on SHAP values depends on the particular model used (in our case, GBMs).

331 Although SHAP values can, in theory, be computed for any black box model, they
 332 are more computationally efficient for tree-based models like GBMs due to a model-specific
 333 algorithm for computing exact SHAP values known as TreeSHAP (Lundberg et al., 2019),
 334 which reduces the computational complexity from exponential to polynomial. For other
 335 complex models like neural networks, computing SHAP values would require refitting
 336 the model with many subsets of features, which is impractical if training is expensive and
 337 more than a few features are used. Unfortunately, a downside of using TreeSHAP is that
 338 non-contributing features can potentially have a non-zero contribution if they are cor-
 339 related with a contributing feature (Molnar, 2019).

340 4 Results

In this section, we will compare the predictive performance of GBMs with neural networks developed by Siciliano et al. (2021) and Collado-Villaverde et al. (2021), explain model predictions using the methods discussed in section 3.2, and discuss how predictions vary when the different set of features listed in table 4 are used as inputs. To evaluate the predictive accuracy of GBMs for forecasting the SYM-H index, we use the root mean squared error (RMSE) defined in eq. (6).

$$\text{RMSE}(y, \hat{y}) = \sqrt{\frac{1}{n} \sum_{i=1}^n (y_i - \hat{y}_i)^2} \quad (6)$$

341 The RMSE metric provides insight into how well predictions match observations on av-
 342 erage so a lower value is better.

To supplement the RMSE metric, we also use the forecast skill score (FSS) based on mean squared error (Murphy, 1988) using the Burton equation described in T. O'Brien and McPherron (2000) as a baseline defined as

$$\text{FSS}(y, \hat{y}, y_{\text{burton}}) = 1 - \frac{\text{MSE}(y, \hat{y})}{\text{MSE}(y, y_{\text{burton}})}, \quad (7)$$

343 where y_{burton} denotes the predictions from the Burton equation and $\text{MSE}(y, \hat{y}) = (1/n) \sum_{i=1}^n (y_i - \hat{y}_i)^2$. The Burton equation, which predicts the evolution of pressure-corrected Dst from the half-wave rectified solar wind motional electric field, is an appropriate baseline as it is derived from physical understanding and is thus also an interpretable method for predicting the SYM-H index. The metric in eq. (7) evaluates the performance of model predictions relative to the baseline predictions. If FSS is between 0 and 1 (inclusive), that means the considered model outperforms the baseline. However, if FSS is negative, that means the considered model performs worse than the baseline.

351 4.1 Comparison to existing methods

352 In this section, we compare the predictions obtained using our model with the neu-
 353 ral networks developed in Siciliano et al. (2021) (**LSTM1/CNN1**) and Collado-Villaverde
 354 et al. (2021) (**LSTM2**) on the 17 test storms in table 2 using the RMSE metric. Collado-
 355 Villaverde et al. (2021) considers 1-2 hours ahead prediction, whereas Siciliano et al. (2021)
 356 only considers 1-hour. On the other hand, Siciliano et al. (2021) trains models with and
 357 without the SYM-H index as an input, whereas Collado-Villaverde et al. (2021) only trains
 358 models with SYM-H. We train GBM models to predict 1-2 hours ahead with and with-
 359 out the SYM-H index as an input and compare them to the corresponding neural net-
 360 work models. All models were trained using data from the same storms in table 1. The
 361 RMSE values and forecast skill scores for each test storm and all considered models are
 362 shown in tables 6 to 9. Similar to Collado-Villaverde et al. (2021), we also compute the
 363 mean RMSE over all storms.

364 For each prediction scenario, we perform a paired t-test to determine if the mean
 365 difference in RMSEs across storms is statistically significant at a 5% significance level.
 366 A paired t-test can be used to compare two population means where you have two sam-
 367 ples with observations that can be paired with one another. It amounts to performing
 368 a one-sample t-test on the differences of the paired observations. In our case, we can match
 369 the RMSEs of different methods for the same storm together.

370 4.1.1 1-hour ahead predictions

371 Tables 6 and 7 show the RMSE values and forecast skill scores for 1 hr ahead pre-
 372 dictions with SYM-H included as an input using our GBM, LSTM1, LSTM2, and the
 373 simple persistence model. In this case, our GBM achieves the lowest mean and median
 374 RMSE among the considered models. Our GBM model has a 0.448 nT (5.7%) lower RMSE
 375 than LSTM2, a 1.138 nT (13.3%) lower RMSE than LSTM1, and a 1.942 nT (20.8%)
 376 lower RMSE than the persistence model. Furthermore, our GBM has the lowest RMSE
 377 and highest skill score for 14 out of 17 test storms (26-32, 35-38, 41, 42). Figure 2 shows
 378 the 1 hour ahead predictions from our GBM and LSTM2 during the main and recovery
 379 phases of the three strongest test storms with SYM-H < -300 nT (31, 33, 37) along with
 380 the corresponding prediction errors. The distribution of the prediction errors are roughly
 381 similar for these three test storms. For the March 2001 storm (second row; fig. 2), our
 382 GBM was able to accurately predict the minimum SYM-H of around -400 nT that was
 383 reached around 06:00 to 12:00 UT Mar 31 even though the timing is slightly off. A sim-
 384 ilar plot and analysis for the persistence model is given in appendix A1.

Table 6. RMSEs for 1-hour ahead prediction over the test storm set with our GBM model, LSTM1 (Siciliano et al., 2021) and LSTM2 (Collado-Villaverde et al., 2021) neural networks, Burton equation (T. O’Brien & McPherron, 2000) and simple persistence. Here, the GBM, LSTM1, and LSTM2 were trained with past SYM-H and IMF parameters as inputs. The lowest RMSE for each row is shown in **bold**.

Storm #	GBM	LSTM2	LSTM1	Burton	Persistence
26	5.863	6.630	6.700	6.839	7.631
27	7.729	8.913	8.900	7.954	9.623
28	4.281	5.858	5.400	5.697	5.814
29	5.833	6.683	7.200	6.511	7.174
30	4.927	5.200	5.600	4.614	4.810
31	8.277	8.584	10.700	8.838	10.429
32	6.841	7.259	8.300	9.487	10.528
33	14.492	13.340	16.300	16.630	21.167
34	10.190	10.034	11.300	10.888	10.913
35	7.154	7.693	8.500	7.918	8.011
36	8.512	9.525	8.700	9.082	9.708
37	14.548	15.184	17.500	15.713	19.698
38	3.886	4.080	4.200	4.572	4.842
39	5.901	6.431	5.600	6.663	7.597
40	4.976	4.673	5.500	5.371	5.057
41	7.558	7.882	9.000	8.358	9.984
42	5.030	5.669	5.900	5.549	6.036
Mean	7.412	7.860	8.550	8.276	9.354
Median	6.841	7.259	8.300	7.918	8.011
Min.	3.886	4.080	4.200	4.572	4.810
Max.	14.548	15.184	17.500	16.630	21.167
Std. error	0.763	0.713	0.901	0.840	1.131

Table 7. Forecast skill scores (using the Burton equation (T. O’Brien & McPherron, 2000) as the baseline) for 1-hour ahead prediction over the test storm set with our GBM model, LSTM1 (Siciliano et al., 2021) and LSTM2 (Collado-Villaverde et al., 2021) neural networks. Here, the GBM, LSTM1, and LSTM2 were trained with past SYM-H and IMF parameters as inputs. The highest skill score for each row is shown in **bold**.

Storm #	GBM	LSTM2	LSTM1
26	0.143	0.031	0.020
27	0.028	-0.120	-0.119
28	0.249	-0.028	0.052
29	0.104	-0.026	-0.106
30	-0.068	-0.127	-0.214
31	0.063	0.029	-0.211
32	0.279	0.235	0.125
33	0.129	0.198	0.020
34	0.064	0.078	-0.038
35	0.096	0.028	-0.074
36	0.063	-0.049	0.042
37	0.074	0.034	-0.114
38	0.150	0.108	0.081
39	0.114	0.035	0.160
40	0.074	0.130	-0.024
41	0.096	0.057	-0.077
42	0.094	-0.022	-0.063

385

4.1.2 2-hour ahead predictions

386

387

388

389

390

391

Tables 8 and 9 show the RMSE values and forecast skill scores for 2-hour ahead predictions from GBM and LSTM2 with past SYM-H included as an input. Our GBM model has a mean RMSE that is 3.585 nT (24.8%) lower than the mean RMSE for the simple persistence model. However, the mean RMSE for our GBM model is .328 nT (3.1%) greater than the one for LSTM2. Moreover, LSTM2 has a lower RMSE and higher skill score for 8 out of the 17 test storms (31-33, 36, 37, 39-41).

392

4.1.3 Predictions without past SYM-H

393

394

395

396

397

398

399

When we omit the SYM-H index as an input to predict 1-hour ahead, our GBM outperforms LSTM1 and has similar performance as CNN1. Table 10 shows the RMSE for 1-hour ahead predictions from GBM, LSTM1, and CNN1 and 2-hour ahead predictions from GBM. Our GBM model has a 3.5 nT (15.4%) lower mean RMSE than LSTM1 and a 1.6 nT (7.7%) lower mean RMSE than CNN1. Furthermore, the GBM model has the lowest RMSE for 11 out of 17 test storms. However, CNN1 achieves a lower RMSE for the 3 strongest test storms (33, 37, 40).

400

4.1.4 Statistical significance

401

402

403

404

405

Table 11 shows the p-values for the paired t-tests described in the second paragraph of section 4.1. From this table, we can see that the mean differences in RMSE across storms between GBM and competing methods for all prediction scenarios are statistically significant at a 5% significance level (p-value ≤ 0.05) except for 2 hr ahead prediction with LSTM2.

Table 8. RMSEs for 2-hour ahead prediction over the test storm set with our GBM model, the LSTM2 neural network (Collado-Villaverde et al., 2021), Burton equation (T. O’Brien & McPherron, 2000) and persistence. Here, the GBM and LSTM2 model were trained with past SYM-H and IMF parameters as inputs. The lowest RMSE for each row is shown in **bold**.

Storm #	GBM	LSTM2	Burton	Persistence
26	8.285	8.989	10.690	12.374
27	11.585	13.418	12.465	15.387
28	5.650	5.877	8.858	9.331
29	8.826	9.314	9.776	11.415
30	7.280	7.288	6.266	7.416
31	12.613	12.436	13.604	17.193
32	9.927	8.937	13.766	15.282
33	24.519	18.481	25.729	33.927
34	13.736	13.941	14.695	15.109
35	9.504	9.932	10.586	11.211
36	12.068	12.058	13.117	14.687
37	22.327	21.084	24.446	30.582
38	5.153	5.213	6.546	7.353
39	7.391	6.798	10.159	12.322
40	5.633	5.281	6.032	6.373
41	12.121	11.707	12.622	15.437
42	7.976	8.273	8.877	10.130
Mean	10.858	10.530	12.249	14.443
Median	9.504	9.314	10.690	12.374
Min.	5.153	5.213	6.032	6.373
Max.	24.519	21.0840	25.729	33.927
Std. error	1.310	1.077	1.338	1.808

406

4.2 Explaining predictions

407

408

409

In this section, we explain how the input features we use contributed to our model’s predictions using the methods discussed in section 3.2. To obtain the contributions from each feature in table 3, we sum up the contributions from the history of that feature.

410

411

412

413

414

415

416

417

418

419

420

421

422

423

424

425

426

Figure 3 shows the contributions to the 1-hour prediction from various features as a function of the SYM-H. Overall, the past SYM-H value dominates, which means that SYM-H varies smoothly at a 1-hour time scale. This also means that beating the persistence model is not easy. The second most important contribution comes from B_z , which is expected based on its importance in driving magnetic reconnection that allows energy entry into the magnetosphere. What is less expected is that the velocity V_x and the rectified electric field E_s are much less important for the storm peak values (SYM-H below -100 nT). In fact, the third most important feature is the dynamic pressure ρV_x^2 . One would expect the dynamic pressure to be most important during the sudden storm commencement that produces a positive jump in SYM-H. Interestingly, the contributions of ρV_x^2 and B_z are comparable even for predicting positive SYM-H, except for the most positive values. Overall, we find that past SYM-H and B_z are the most important features. Density, velocity, the derived dynamic pressure and rectified electric field are comparable. The rest of the features, such as B_x , B_y and temperature provide quite small contributions. Note that the rectified E_s is a less important contributor than B_z and the dynamic pressure, despite its physical significance of carrying the magnetic flux that induces dayside reconnection.

Table 9. Forecast skill scores (using the Burton equation (T. O’Brien & McPherron, 2000) as the baseline) for 2-hour ahead prediction over the test storm set with our GBM model and the LSTM2 neural network (Collado-Villaverde et al., 2021). Here, the GBM and LSTM2 model were trained with past SYM-H and IMF parameters as inputs. The highest skill score for each row is shown in **bold**.

Storm #	GBM	LSTM2
26	0.225	0.159
27	0.071	-0.076
28	0.362	0.337
29	0.097	0.047
30	-0.162	-0.163
31	0.073	0.086
32	0.279	0.351
33	0.047	0.282
34	0.065	0.051
35	0.102	0.062
36	0.080	0.081
37	0.087	0.138
38	0.213	0.204
39	0.272	0.331
40	0.066	0.125
41	0.040	0.072
42	0.101	0.068

427 Figure 4 shows the contribution of various features of the model that is not using
 428 past SYM-H. As expected, B_z becomes the most important feature. Now velocity and
 429 density are the next most important features, especially for moderate values of SYM-
 430 H, and the dynamic pressure by itself does not have enough information (unlike in the
 431 previous case that used past SYM-H). The rectified E_s is still a rather small contribu-
 432 tor compared to B_z . This can be explained by jointly examining the contributions of B_z
 433 and V_x : B_z becomes more and more dominant for larger negative SYM-H values. On
 434 the other hand, the contribution of V_x peaks at moderate storm with SYM-H above -100
 435 nT, and its contribution tapers off for the very strong storms. While the electric field
 436 E_s combines these two terms, one can see that their contributions are most effective in
 437 different severity of storms or different phases of the storm, suggesting that considering
 438 them as independent variables rather than as a single parameter provides more insight
 439 into the underlying physics. The strong contribution of density for small and positive
 440 SYM-H values speaks to the importance of density pulses that often are found at the lead-
 441 ing edges of solar wind structures impacting the Earth (Kilpua et al., 2017).

442 **4.2.1 November 2004 Storm**

443 We now look into how the prediction is obtained during the strongest test storm.
 444 Figure 5 shows the absolute and relative contributions of various features to the 1-hour
 445 and 2-hour ahead predictions of SYM-H during the November 2004 geomagnetic storm.
 446 The minimum SYM-H is close to -400 nT for this extreme event, so the RMSE of about
 447 30 nT for 1-hour and 39 nT for 2-hour forecast are quite accurate (top row). The abso-
 448 lute and relative contributions shown in the subsequent rows vary substantially during
 449 the storm. From 18:00 to 20:45 UT (following the Storm Sudden Commencement, SSC),
 450 the observed SYM-H is positive, and this is roughly captured by the model for 1-hour
 451 prediction, but is completely missed by the 2-hour forecast. This is not very surprising,

Table 10. RMSEs for 1- and 2-hour ahead predictions using only the IMF as input (I_2) with our GBM model and the LSTM1 and CNN1 models of Siciliano et al. (2021). For 1-hour ahead predictions, the lowest RMSE in each row is shown in **bold**.

Storm #	1-hour ahead			2-hour ahead
	GBM	LSTM1	CNN1	GBM
26	12.6	18.0	19.8	12.9
27	20.1	16.8	23.4	20.9
28	12.7	18.6	14.4	12.4
29	15.4	21.1	20.0	16.7
30	17.0	24.2	25.8	17.1
31	28.5	32.5	32.1	29.6
32	21.8	23.4	18.9	21.9
33	35.7	33.8	26.7	38.1
34	15.3	17.9	16.6	15.5
35	16.9	21.3	18.6	17.3
36	16.2	20.4	21.4	16.8
37	41.6	42.6	36.9	42.7
38	10.5	18.6	13.0	10.6
39	13.0	20.3	16.5	12.8
40	10.9	13.6	9.2	10.6
41	23.2	27.3	25.4	23.7
42	16.9	17.8	16.7	17.1
Mean	19.3	22.8	20.9	19.8
Median	16.9	20.8	19.9	17.1
Min.	10.5	13.6	9.2	10.6
Max.	41.6	42.6	36.9	42.7
Std. error	2.284	1.994	1.853	2.402

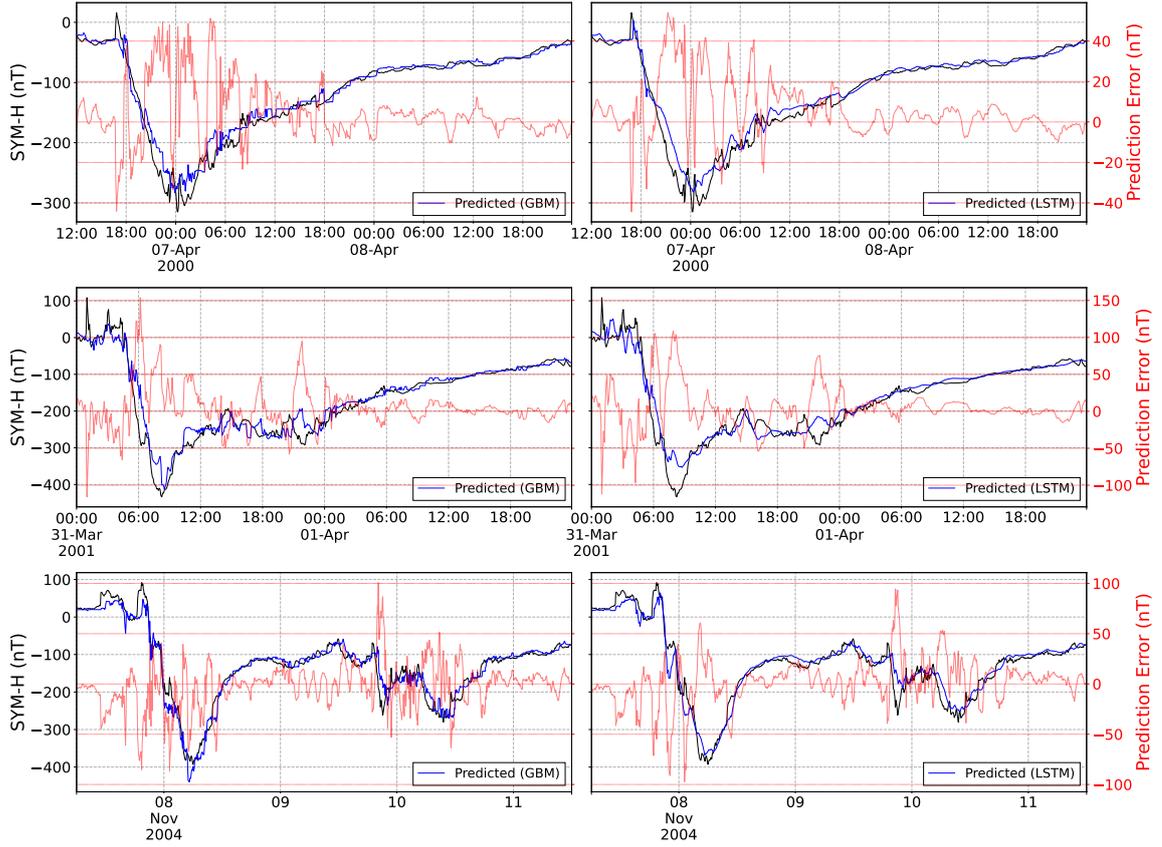


Figure 2. 1-hour ahead predictions for the 3 strongest geomagnetic storms in the test set during the main and recovery phases from our GBM (left column) and the LSTM2 developed by Collado-Villaverde et al. (2021) (right column). The observed SYM-H (black), the predicted SYM-H (blue) and the error (red) are shown for storms 31, 33, and 37 in the 3 rows, respectively.

452 since there is no information in the solar wind that would predict the sudden commence-
 453 ment prior to the arrival of the shock. The only reason the 1-hour prediction can get the
 454 SSC about half an hour rather than 1 hour late is the lead time provided by the time
 455 it takes the high speed solar wind to propagate from L1 to the Earth. The main con-
 456 tributors to the 1-hour prediction during this period are the density and dynamic pres-
 457 sure, and to some extent the IMF B_z . Based on our physical understanding, we would
 458 expect the dynamic pressure to be a more important predictor than the density, but that
 459 is clearly not the case, perhaps associated with the relatively constant value of the so-
 460 lar wind speed over that period.

461 During the main phase (22:00 Nov 7 to 06:00 Nov 8) of the storm, the SYM-H gradu-
 462 ally drops to its minimum value near -400 nT. Focusing on the two-hour prediction,
 463 the relative contribution of B_z peaks around 22:00 on November 7, and 01:00 and after
 464 04:00 UT. The first peak corresponds to the time when B_z decreases rapidly to nearly
 465 -50 nT value. The following period of very intense southward IMF shows initially low
 466 contribution from B_z , but then consistently high values with a peak at 04:00 close to the
 467 SYM-H minimum demarking the end of the storm main phase. The contribution from
 468 B_y , while generally low, has a broad peak between 20:00 and 00 UT on November 7. Dur-
 469 ing that period, B_y is first positive and then turns strongly negative. As the B_z is neg-
 470 ative during that time, the strong B_y component adds to the efficiency of the dayside

Table 11. P-values from paired t-tests for null hypothesis that the mean difference in RMSE across storms for GBM vs. competing methods is zero.

	1 hr ahead	2 hr ahead
LSTM2	0.008	0.419
LSTM1	0.000	
Persistence	0.000	0.000

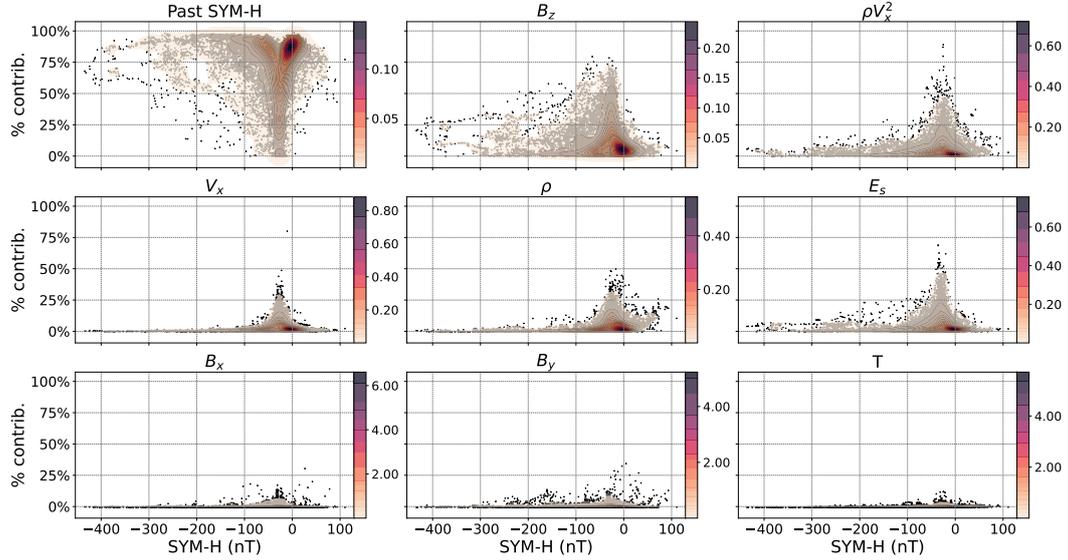


Figure 3. Scatter plot of percentage contributions (y-axis) against SYM-H (x-axis) for all the geomagnetic storms. The panels show the contributions of all considered features to the 1-hour ahead GBM prediction. Each prediction is represented as black dots. Kernel density estimates using a Gaussian kernel are shown in color with the corresponding color legend on the right of each scatter plot.

471 reconnection process, which may account for its independent role as a predictor. Finally,
 472 during the recovery phase the prior SYM-H dominates (SYM-H evolution dominated by
 473 internal ring current loss processes), with B_z playing a secondary role.

474 Figure 6 shows the contribution of features as a function of time when the prior
 475 SYM-H is not used. The RMSE values become 33 nT and 37 nT for the 1 and 2-hour pre-
 476 dictions, respectively. For the 1-hour prediction, RMSE slightly increases by about 3 nT,
 477 but for the 2-hour prediction, RMSE decreases by roughly 2 nT. This suggests that there
 478 is no additional information from the 2-hour old SYM-H compared to what the model
 479 can infer from a longer history of L1 observations, at least for this event. If this held in
 480 general, it would put a prediction window limit on using past SYM-H for data assim-
 481 ilation purposes. Another unexpected result is that the 1-hour prediction misses the posi-
 482 tive SYM-H period despite using the dynamic pressure. This is in contrast with the 1-
 483 hour prediction that includes past SYM-H, which produced a larger positive SYM-H, al-
 484 though still lower than observed.

485 The relative contributions (bottom row) show a rather complicated and interest-
 486 ing pattern. In the initial storm period 18:00 to 21:00 UT, when the observed SYM-H
 487 is positive, the main contributors are density and velocity. Once SYM-H goes negative,
 488 B_z gradually becomes the main contributing feature with E_s and, B_x (for 1-hour pre-

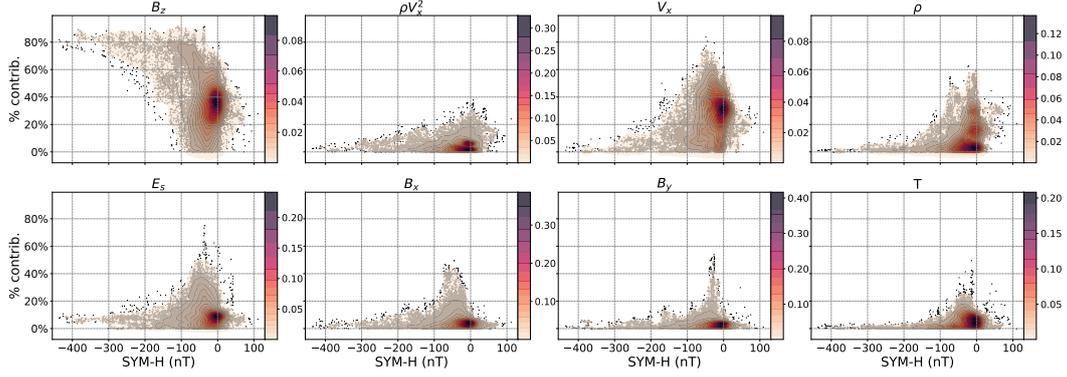


Figure 4. Scatter plot of percentage contributions (y-axis) against SYM-H (x-axis) from solar wind and IMF parameters for 1-hour ahead prediction from GBM using only solar wind and IMF parameters as input. Each prediction is represented as black dots. Kernel density estimates using a Gaussian kernel are shown in color with the corresponding color legend on the right of each scatter plot.

489 diction) and B_y (for 2-hour prediction) being the second and third most important. Once
 490 SYM-H drops below -100 nT, the contribution from B_z becomes dominant and this re-
 491 mains true during the whole recovery phase. The other features start to contribute more
 492 after 12:00 UT Nov 8 when B_z turns positive. Even with positive B_z , however, the main
 493 contributor remains B_z . This shows that the rectified E_s , which simply zeroes out the
 494 electric field for positive B_z , is throwing away potentially important information.

495 **4.2.2 January 2004 Storm**

496 Next, we study the storm of January 2004 that has a minimum SYM-H of about
 497 -140 nT, so it is an intense storm, but not as extreme as the November 2004 super storm.
 498 As shown in figure 7, this is a very complicated storm due to the highly variable B_z field
 499 in the CME sheath (00:00 UT to 11:00 UT Jan 22) preceding the magnetic cloud with
 500 consistently negative B_z . The model prediction has 14.22 nT and 19.96 nT RMSE for
 501 the 1- and 2-hour predictions, respectively, which is quite good for such a complicated
 502 event. In the ICME sheath, the main contributor is the previous SYM-H followed by the
 503 dynamic pressure.

504 The 1-hour ahead model predicts the jump of SYM-H from 0 to about $+30$ nT at
 505 2:00UT, which is about half an hour late compared to observations. This cannot be based
 506 on prior SYM-H that is observed 1 hour earlier, and it is clearly obtained from the dy-
 507 namic pressure as expected from physical understanding. The 2-hour prediction, how-
 508 ever, completely misses predicting positive SYM-H values (except for following the in-
 509 crease of the observed SYM-H with a 2-hour delay), similarly to the extreme event case.

510 Between 01:00 and 11:00 UT the main contributors are the prior SYM-H and the
 511 dynamic pressure, with B_z playing a minor role only. After 11:00 UT, however, B_z turns
 512 consistently negative and it becomes the main contributor of predicting the main phase
 513 of the storm 1 hour or 2 hours later for the two models, respectively. The 2-hour pre-
 514 diction also relies heavily on B_y between 10 and 12:00 UT. A possible explanation is that
 515 the strong magnetic field in the magnetic cloud rotates, so a strong signal in B_x or B_y
 516 may be a predictor for a strong, possibly negative, B_z value that has strong geomagnetic
 517 impact.

518 The model correctly predicts the minimum value of SYM-H, but it is late by an
 519 hour and two hours for the 1- and 2-hour predictions, respectively. This means that the
 520 prior SYM-H was the primary contributor to the prediction of the minimum SYM-H.
 521 We note that the last available B_z is negative, but has a small amplitude at this point
 522 (about -5 nT). Clearly the model is not capable of predicting the behavior of the storm
 523 very well during this time period for this particular event. The recovery phase is correctly
 524 captured with the prior SYM-H dominating, as expected. B_z becomes slightly more neg-
 525 ative from 19:00 to 23:00, and the importance of B_z and E_s becomes significant during
 526 this time correctly predicting the slow down of the recovery, although with considerable
 527 delay.

528 Figure 8 shows the model predictions for the January 2004 storm without relying
 529 on the prior SYM-H values. The RMSE is around 33 nT for both the 1-hour and 2-hour
 530 ahead forecast. The positive SYM-H values are missed by the model and in fact there
 531 is a considerable underprediction of SYM-H until 11:00 UT. The main phase of the storm
 532 corresponding the rapid decrease of SYM-H is quite well captured. It is slightly too early
 533 for the 1-hour prediction, and quite spot on for the 2-hour prediction. The minimum SYM-
 534 H is correctly predicted by both models with an hour delay, and it is actually somewhat
 535 better predicted by the 2-hour ahead model. The recovery phase is reasonably predicted,
 536 although the predicted recovery rate is somewhat slower than what is observed.

537 The main contributors to the prediction before 11:00 UT are velocity, the rectified
 538 electric field and density. During the main phase and the recovery, B_z becomes an im-
 539 portant contributor, but the velocity and E_s still play considerable roles. B_x becomes
 540 the most important contributor during the recovery phase. Figure 4 confirms that B_x
 541 and B_y become significant contributors when prior SYM-H is not used.

542 One of the surprises mentioned above was that B_z is a better predictor than E_s .
 543 However, these features are highly correlated so it is not clear if the GBM prefers B_z over
 544 E_s by chance only. To investigate this question, we have performed experiments to see
 545 whether B_z or E_y , or the rectified E_s is the best predictor out of the three for future SYM-
 546 H. To make E_y (or E_s) and B_z fully independent of each other, we have removed the
 547 V_x and ρV_x^2 features and used only one the three quantities (B_z , E_y , and rectified E_s)
 548 together with density and temperature while training the GBM. The RMSE values are
 549 shown in Table 12 including both cases with and without prior SYM-H.

Table 12. RMSE from models with only one of B_z , E_y , and E_s included as input calculated using all test storms. The RMSE from a model trained with B_z , E_s , and ρV_x^2 is shown in the last column as reference. For these experiments, density and temperature were also used as features.

	B_z	E_y	E_s	$B_z, E_s, \rho V_x^2$
Including SYM-H	7.35	8.00	8.26	7.26
Excluding SYM-H	20.84	21.12	21.45	18.39

550 Based on the RMSE values in the table, we conclude that B_z is the best predic-
 551 tor followed by E_y and the rectified E_s . It is also interesting to see that past SYM-H and
 552 B_z together are pretty much all that the model needs. The velocity V_x , for example, plays
 553 no significant role in contributing to the quality of the prediction as it only improves the
 554 RMSE from 7.35 to 7.26 nT. When past SYM-H is not used, the velocity plays a more
 555 important role by improving the RMSE from 20.84 to 18.39, but still much less impor-
 556 tant than B_z , E_y or E_s . A possible reason may be that V_x varies only about a factor of
 557 2 between about -350 km/s and -700 km/s even during storm events.

5 Discussion and conclusions

We apply an explainable machine learning method to quantify the contribution of prior SYM-H values, solar wind, IMF, and derived parameters to predictions of the SYM-H index 1 to 2 hours ahead. In particular, gradient boosting machines (GBM) are used and the explanation is based on the TreeSHAP method. We showed that gradient boosting machines have similar, if not better, performance compared to the less explainable but highly effective LSTM method for forecasting the SYM-H index.

From the quantified feature contributions, we were able to show that our proposed model makes predictions in a physically consistent manner, while also challenging some of the commonly assumed relationships among the interplanetary magnetic field, the solar wind and the formation of Earth's ring current. In particular, we found that past SYM-H and B_z are the most important features overall but feature contributions vary depending on the storm phase and the storm itself. During the storm sudden commencement, past SYM-H, density, velocity, and to some extent, dynamic pressure and electric field, became the main contributors to predictions. As SYM-H decreases during the main phase, past SYM-H and B_z played an increasingly larger role.

SHAP values revealed ways that our models made predictions during the two storms we investigated in detail: density and velocity had a larger independent contribution than dynamic pressure during the storm sudden commencement; B_y had a non-negligible contribution during the storm sudden commencement and main phase; and B_z was a better predictor than the rectified E_s . However, strong correlation among solar wind variables (Borovsky, 2018) may affect how SHAP values should be interpreted. A physically important feature may have a small contribution if a highly correlated feature is present and has a large contribution. For example, from figs. 3 and 4, we see that the contribution from V_x increases drastically when past SYM-H is omitted as an input, which is likely due to the correlation between SYM-H and V_x . Therefore, a low feature contribution should not simply be interpreted to mean the corresponding feature is not physically important without investigating how different features are correlated. Further efforts will be made to investigate the robustness of these findings and to perform a comparison of feature contributions for many different storms.

Along with gray-box approaches, this work takes the first steps in making machine learning methods more reliable and trustworthy for operational forecasting of geomagnetic activity. However, explanation methods like SHAP should be used with caution, especially in high-stakes decision making, as they do not always provide explanations that are faithful to the original model (Rudin, 2019). Thus, developing highly accurate but intrinsically interpretable models should be prioritized. In addition to interpretability, quantified uncertainty is also equally as important. Consequently, we will devote future efforts to developing interpretable methods for forecasting other types of geomagnetic indices and geomagnetic activity that also estimate predictive uncertainty.

Appendix A

A1 Graphical comparison with persistence model & Burton equation

Figure A1 shows the 1 hour ahead predictions from our GBM (with past SYM-H and IMF parameters as input) and the persistence model during the main and recovery phases of the three strongest test storms with SYM-H < -300 nT (31, 33, 37) along with the corresponding prediction errors. The difference in prediction error between our GBM and the persistence model is most notable during the main phases of the three storms considered. For example, during the main phase of storm 37, the persistence model has prediction errors reaching > 100 nT which means it severely overpredicts SYM-H during the main phase. Meanwhile, our GBM has prediction errors between around -100 to 40 nT, which means it tended to underpredict rather than overpredict SYM-H. Figure A2 shows the 1 hour ahead predictions from our GBM and the Burton equation during the same time periods. In these plots, the GBM seems to capture the timing of the storms slightly better than the Burton equation. However, they have similar predictive performance during these three storms as shown by their RMSEs in table 6.

A2 Descriptive statistics of solar wind & IMF parameters

Table A1. Descriptive statistics for the solar wind and IMF parameters in the 25 storms used for training listed in table 1. The minimum temperature (MK) is most likely a measurement error.

Parameter	Min.	25% Quantile	Median	75% Quantile	Max.
B_x (nT)	-43.700	-3.131	0.340	3.378	34.681
B_y (nT)	-51.968	-2.901	0.221	3.289	46.862
B_z (nT)	-77.258	-2.296	-0.092	2.179	38.717
V_x (km/s)	-1233.693	-539.489	-445.287	-384.021	-264.722
Density (amu/cm ²)	0.041	2.912	5.027	8.477	76.239
Temperature (MK)	0.0032	0.0385	0.0702	0.1262	1.0983

Table A2. Descriptive statistics for the solar wind and IMF parameters in the 25 test storms listed in table 2. The minimum temperature (MK) is most likely a measurement error.

Parameter	Min.	25% Quantile	Median	75% Quantile	Max.
B_x (nT)	-48.717	-2.868	0.221	3.444	33.827
B_y (nT)	-48.963	-2.816	-0.205	2.855	54.563
B_z (nT)	-48.585	-2.357	-0.084	1.933	53.002
V_x (km/s)	-887.784	-535.138	-424.304	-373.465	-251.481
Density (amu/cm ³)	0.295	2.760	4.424	7.643	113.982
Temperature (MK)	0.0052	0.037	0.0658	0.122	0.9909

Acknowledgments

We thank Austin Brenner, Qusai Al Shidi, and Professor Michael Liemohn from the Dept. of Climate and Space Sciences and Engineering (CLaSP) at the University of Michigan for helpful comments and fruitful discussions. This work was supported by NASA DRIVE Science Center grant 80NSSC20K0600, NASA MMS grant 80NSSC19K0564, NSF PRE-EVENTS grant 1663800 and NSF SWQU grant PHY-2027555. EC is partially funded

619 by NASA under grants 80NSSC20K1580 and 80NSSC20K1275. The ACE level 2 data
 620 used as inputs into our models in this study is available through NASA/GSFC's Space
 621 Physics Data Facility's (SPDF) Coordinated Data Analysis Web (CDAWeb) at [https://
 622 cdaweb.gsfc.nasa.gov/](https://cdaweb.gsfc.nasa.gov/). The SYM-H index data is available through SPDF's OMNI-
 623 Web at <https://omniweb.gsfc.nasa.gov/>. All relevant digital materials used in this
 624 manuscript will be permanently archived at the University of Michigan (UM) Library
 625 Deep Blue Data Repository (<https://deepblue.lib.umich.edu/data>), which is specif-
 626 ically designed for UM researchers to share their research data and to ensure its long-
 627 term viability. Processed data will be assigned Digital Object Identifiers (DOIs), which
 628 will serve as identifiers for the data, enabling them to be cited in publications.

629 References

- 630 Al Shidi, Qusai. (2020). *swmfpy*. Retrieved from <https://gitlab.umich.edu/swmf>
 631 [_software/swmfpy](https://gitlab.umich.edu/swmf/-/software/swmfpy) (version 2020.5)
- 632 Ayala Solares, J. R., Wei, H.-L., Boynton, R. J., Walker, S. N., & Billings, S. A.
 633 (2016). Modeling and prediction of global magnetic disturbance in near-Earth
 634 space: A case study for K_p index using NARX models: MODELING AND
 635 PREDICTION OF K_p INDEX. *Space Weather*, *14*(10), 899–916. doi:
 636 10.1002/2016SW001463
- 637 Bailey, R. L., Reiss, M. A., Arge, C. N., Möstl, C., Owens, M. J., Amerstorfer,
 638 U. V., ... Hinterreiter, J. (2021). Using gradient boosting regression to im-
 639 prove ambient solar wind model predictions. *arXiv:2006.12835 [astro-ph,*
 640 *physics:physics]*.
- 641 Bauer, E., & Kohavi, R. (1999). An Empirical Comparison of Voting Classifica-
 642 tion Algorithms: Bagging, Boosting, and Variants. *Machine Learning*, *36*(1/2),
 643 105–139. doi: 10.1023/A:1007515423169
- 644 Bhaskar, A., & Vichare, G. (2019). Forecasting of SYMH and ASYH indices for
 645 geomagnetic storms of solar cycle 24 including St. Patrick's day, 2015 storm
 646 using NARX neural network. *Journal of Space Weather and Space Climate*, *9*.
 647 doi: 10.1051/swsc/2019007
- 648 Bluwstein, K., Buckmann, M., Joseph, A., Kang, M., Kapadia, S., & Simsek,
 649 Ö. (2020). Credit Growth, the Yield Curve and Financial Crisis Predic-
 650 tion: Evidence from a Machine Learning Approach. *SSRN Journal*. doi:
 651 10.2139/ssrn.3520659
- 652 Borovsky, J. E. (2018). On the Origins of the Intercorrelations Between Solar Wind
 653 Variables. *Journal of Geophysical Research: Space Physics*, *123*(1), 20–29. doi:
 654 10.1002/2017JA024650
- 655 Bortnik, J., Chu, X., Ma, Q., Li, W., Zhang, X., Thorne, R. M., ... others (2018).
 656 Artificial neural networks for determining magnetospheric conditions. In *Ma-
 657 chine learning techniques for space weather* (pp. 279–300). Elsevier. doi: 10
 658 .1016/B978-0-12-811788-0.00011-1
- 659 Boynton, R., Balikhin, M., Billings, S., Sharma, A., & Amariutei, O. (2011). Data
 660 derived narmax dst model. In *Annales geophysicae* (Vol. 29, pp. 965–971).
- 661 Breiman, L. (1996). Bagging predictors. *Mach Learn*, *24*(2), 123–140. doi: 10.1007/
 662 BF00058655
- 663 Breiman, L. (2001). Random Forests. *Machine Learning*, *45*(1), 5–32. doi: 10.1023/
 664 A:1010933404324
- 665 Breiman, L., Friedman, J. H., Olshen, R. A., & Stone, C. J. (1984). *Classification
 666 and regression trees*. Wadsworth and Brooks.
- 667 Burton, R. K., McPherron, R. L., & Russell, C. T. (1975). An empirical relationship
 668 between interplanetary conditions and Dst. *Journal of Geophysical Research*,
 669 *80*, 4204. doi: 10.1029/JA080i031p04204
- 670 Cai, L., Ma, S. Y., & Zhou, Y. L. (2010). Prediction of SYM-H index during large
 671 storms by NARX neural network from IMF and solar wind data. *Annales Geo-*

- 672 *physicae*, 28(2), 381–393. doi: 10.5194/angeo-28-381-2010
- 673 Camporeale, E. (2019). The challenge of machine learning in space weather: Now-
 674 casting and forecasting. *Space Weather*, 17(8), 1166–1207. doi: 10.1029/
 675 2018SW002061
- 676 Camporeale, E., Cash, M. D., Singer, H. J., Balch, C. C., Huang, Z., & Toth, G.
 677 (2020). A Gray-Box Model for a Probabilistic Estimate of Regional Ground
 678 Magnetic Perturbations: Enhancing the NOAA Operational Geospace Model
 679 With Machine Learning. *J. Geophys. Res. Space Physics*, 125(11). doi:
 680 10.1029/2019JA027684
- 681 Chandorkar, M., Camporeale, E., & Wing, S. (2017). Probabilistic forecasting of the
 682 disturbance storm time index: An autoregressive gaussian process approach.
 683 *Space Weather*, 15(8), 1004–1019. doi: 10.1002/2017SW001627
- 684 Chen, T., & Guestrin, C. (2016). XGBoost: A Scalable Tree Boosting System. *Pro-
 685 ceedings of the 22nd ACM SIGKDD International Conference on Knowledge
 686 Discovery and Data Mining*, 785–794. doi: 10.1145/2939672.2939785
- 687 Collado-Villaverde, A., Muñoz, P., & Cid, C. (2021). Deep Neural Networks With
 688 Convolutional and LSTM Layers for SYM-H and ASY-H Forecasting. *Space
 689 Weather*. doi: 10.1029/2021SW002748
- 690 Friedman, J. H. (2001). Greedy function approximation: A gradient boosting ma-
 691 chine. *Ann. Statist.*, 29(5), 1189–1232. doi: 10.1214/aos/1013203451
- 692 Ganushkina, N., Jaynes, A., & Liemohn, M. (2017). Space Weather Effects Pro-
 693 duced by the Ring Current Particles. *Space Science Reviews*, 212(3-4), 1315–
 694 1344. doi: 10.1007/s11214-017-0412-2
- 695 Gleisner, H., Lundstedt, H., & Wintoft, P. (1996). Predicting geomagnetic storms
 696 from solar-wind data using time-delay neural networks. *Ann. Geophys.*, 14(7),
 697 679–686. doi: 10.1007/s00585-996-0679-1
- 698 Gu, Y., Wei, H.-L., Boynton, R. J., Walker, S. N., & Balikhin, M. A. (2019). System
 699 Identification and Data-Driven Forecasting of AE Index and Prediction Un-
 700 certainty Analysis Using a New Cloud-NARX Model. *Journal of Geophysical
 701 Research: Space Physics*, 124(1), 248–263. doi: 10.1029/2018JA025957
- 702 Hastie, T., Tibshirani, R., & Friedman, J. (2001). *The elements of statistical learn-
 703 ing*. New York, NY, USA: Springer New York Inc.
- 704 Iyemori, T. (1990). Storm-time magnetospheric currents inferred from mid-latitude
 705 geomagnetic field variations. *Journal of geomagnetism and geoelectricity*,
 706 42(11), 1249–1265.
- 707 Kilpua, E. K. J., Balogh, A., von Steiger, R., & Liu, Y. D. (2017). Geoeffective
 708 Properties of Solar Transients and Stream Interaction Regions. *Space Sci.
 709 Rev.*, 212, 1271-1314. doi: 10.1007/s11214-017-0411-3
- 710 King, J. H. (2005). Solar wind spatial scales in and comparisons of hourly Wind
 711 and ACE plasma and magnetic field data. *Journal of Geophysical Research*,
 712 110(A2), 2104. doi: 10.1029/2004JA010649
- 713 Korlakai Vinayak, R., & Gilad-Bachrach, R. (2015). DART: Dropouts meet multiple
 714 additive regression trees. In G. Lebanon & S. V. N. Vishwanathan (Eds.),
 715 *Proceedings of the eighteenth international conference on artificial intelligence
 716 and statistics* (Vol. 38, pp. 489–497). San Diego, California, USA: PMLR.
- 717 Kuzlu, M., Cali, U., Sharma, V., & Guler, O. (2020). Gaining Insight Into Solar
 718 Photovoltaic Power Generation Forecasting Utilizing Explainable Arti-
 719 ficial Intelligence Tools. *IEEE Access*, 8, 187814–187823. doi: 10.1109/
 720 ACCESS.2020.3031477
- 721 Liu, L., Zou, S., Yao, Y., & Wang, Z. (2020). Forecasting Global Ionospheric
 722 TEC Using Deep Learning Approach. *Space Weather*, 18(11). doi:
 723 10.1029/2020SW002501
- 724 Lou, Y., Caruana, R., Gehrke, J., & Hooker, G. (2013). Accurate intelligible models
 725 with pairwise interactions. In *Proceedings of the 19th ACM SIGKDD inter-
 726 national conference on Knowledge discovery and data mining* (pp. 623–631).

- 727 Chicago Illinois USA: ACM. doi: 10.1145/2487575.2487579
- 728 Lu, J., Peng, Y., Wang, M., Gu, S., & Zhao, M. (2016). Support vector machine
729 combined with distance correlation learning for dst forecasting during in-
730 tense geomagnetic storms. *Planetary and Space Science*, *120*, 48–55. doi:
731 10.1016/j.pss.2015.11.004
- 732 Lundberg, S. M., Erion, G. G., & Lee, S.-I. (2019). Consistent Individualized Fea-
733 ture Attribution for Tree Ensembles. *arXiv:1802.03888 [cs, stat]*.
- 734 Lundberg, S. M., & Lee, S.-I. (2017). A Unified Approach to Interpreting Model
735 Predictions. In *Advances in Neural Information Processing Systems* (Vol. 30).
736 Curran Associates, Inc.
- 737 Lundberg, S. M., Nair, B., Vavilala, M. S., Horibe, M., Eisses, M. J., Adams, T.,
738 ... Lee, S.-I. (2018). Explainable machine-learning predictions for the pre-
739 ventation of hypoxaemia during surgery. *Nat Biomed Eng*, *2*(10), 749–760. doi:
740 10.1038/s41551-018-0304-0
- 741 Lundstedt, H., & Wintoft, P. (1994). Prediction of geomagnetic storms from solar
742 wind data with the use of a neural network. *Ann. Geophys.*, *12*(1), 19–24. doi:
743 10.1007/s00585-994-0019-2
- 744 Mayaud, P. N. (1980). The dst index. In *Derivation, meaning, and use of geomag-
745 netic indices* (p. 115-129). American Geophysical Union (AGU). doi: 10.1002/
746 9781118663837.ch8
- 747 McComas, D. J., Bame, S. J., Barker, P., Feldman, W. C., Phillips, J. L., Riley,
748 P., & Griffee, J. W. (1998). Solar Wind Electron Proton Alpha Monitor
749 (SWEPAM) for the Advanced Composition Explorer. *Space Science Reviews*,
750 *86*, 563-612. doi: 10.1023/A:1005040232597
- 751 Mitrentsis, G., & Lens, H. (2021). An Interpretable Probabilistic Model for
752 Short-Term Solar Power Forecasting Using Natural Gradient Boosting.
753 *arXiv:2108.04058 [cs, stat]*.
- 754 Mokhtari, K. E., Higdon, B. P., & Başar, A. (2019). Interpreting financial time se-
755 ries with shap values. In *Proceedings of the 29th annual international confer-
756 ence on computer science and software engineering* (p. 166–172). USA: IBM
757 Corp.
- 758 Molnar, C. (2019). *Interpretable machine learning*. ([https://christophm.github
759 .io/interpretable-ml-book/](https://christophm.github.io/interpretable-ml-book/))
- 760 Molnar, C., Casalicchio, G., & Bischl, B. (2020). Interpretable Machine Learning –
761 A Brief History, State-of-the-Art and Challenges. *arXiv:2010.09337 [cs, stat]*.
- 762 Murphy, A. H. (1988). Skill scores based on the mean square error and their rela-
763 tionships to the correlation coefficient. *Monthly Weather Review*, *116*(12),
764 2417 - 2424. Retrieved from [https://journals.ametsoc.org/view/
765 journals/mwre/116/12/1520-0493_1988_116_2417_ssb0tm_2_0_co_2.xml](https://journals.ametsoc.org/view/journals/mwre/116/12/1520-0493_1988_116_2417_ssb0tm_2_0_co_2.xml)
766 doi: 10.1175/1520-0493(1988)116<2417:SSBOTM>2.0.CO;2
- 767 Natekin, A., & Knoll, A. (2013). Gradient boosting machines, a tutorial. *Front.
768 Neurorobot.*, *7*. doi: 10.3389/fnbot.2013.00021
- 769 Newell, P. T., Sotirelis, T., Liou, K., Meng, C. I., & Rich, F. J. (2007). A nearly
770 universal solar wind-magnetosphere coupling function inferred from 10 magne-
771 topheric state variables. *Journal of Geophysical Research*, *112*, 01206. doi:
772 10.1029/2006JA012015
- 773 O'Brien, T., & McPherron, R. L. (2000). Forecasting the ring current index dst in
774 real time. *Journal of Atmospheric and Solar-Terrestrial Physics*, *62*(14), 1295-
775 1299. Retrieved from [https://www.sciencedirect.com/science/article/
776 pii/S1364682600000729](https://www.sciencedirect.com/science/article/pii/S1364682600000729) (Space Weather Week) doi: [https://doi.org/10
777 .1016/S1364-6826\(00\)00072-9](https://doi.org/10.1016/S1364-6826(00)00072-9)
- 778 O'Brien, T. P. (2002). Seasonal and diurnal variation of Dst dynamics. *Journal of
779 Geophysical Research*, *107*(A11), 1341. doi: 10.1029/2002JA009435
- 780 O'Brien, T. P., & McPherron, R. L. (2000). An empirical phase space anal-
781 ysis of ring current dynamics: Solar wind control of injection and decay.

- 782 *Journal of Geophysical Research: Space Physics*, 105(A4), 7707–7719. doi:
783 10.1029/1998JA000437
- 784 Pallochia, G., Amata, E., Consolini, G., Marcucci, M. F., & Bertello, I. (2006). Ge-
785 omagnetic Dst index forecast based on IMF data only. *Ann. Geophys.*, 24(3),
786 989–999. doi: 10.5194/angeo-24-989-2006
- 787 Rapin, J., & Teytaud, O. (2018). *Nevergrad - A gradient-free optimization platform*.
788 <https://GitHub.com/FacebookResearch/Nevergrad>. GitHub.
- 789 Rudin, C. (2019). Stop explaining black box machine learning models for high stakes
790 decisions and use interpretable models instead. *Nature Machine Intelligence*,
791 1(5), 206–215. doi: 10.1038/s42256-019-0048-x
- 792 Saabas, A. (2014). *Interpreting random forests — Diving into data*. Retrieved from
793 <https://blog.datadive.net/interpreting-random-forests/>
- 794 Sakaguchi, K., Nagatsuma, T., Reeves, G. D., & Spence, H. E. (2015). Prediction
795 of MeV electron fluxes throughout the outer radiation belt using multivariate
796 autoregressive models. *Space Weather*, 13(12), 853–867. Retrieved 2021-08-05,
797 from <https://onlinelibrary.wiley.com/doi/10.1002/2015SW001254> doi:
798 10.1002/2015SW001254
- 799 Shapley, L. S. (1953). A Value for n-Person Games. In H. W. Kuhn & A. W. Tucker
800 (Eds.), *Contributions to the Theory of Games (AM-28), Volume II* (pp. 307–
801 318). Princeton University Press. doi: 10.1515/9781400881970-018
- 802 Schwartz-Ziv, R., & Armon, A. (2021). Tabular Data: Deep Learning is Not All You
803 Need. *arXiv:2106.03253 [cs]*.
- 804 Siciliano, F., Consolini, G., Tozzi, R., Gentili, M., Giannattasio, F., & De Michelis,
805 P. (2021). Forecasting SYM-H Index: A Comparison Between Long Short-
806 Term Memory and Convolutional Neural Networks. *Space Weather*, 19(2). doi:
807 10.1029/2020SW002589
- 808 Smith, C. W., L’Heureux, J., Ness, N. F., Acuña, M. H., Burlaga, L. F., & Scheifele,
809 J. (1998). The ACE Magnetic Fields Experiment. *Space Science Reviews*, 86,
810 613–632. doi: 10.1023/A:1005092216668
- 811 Srivastava, N., Hinton, G., Krizhevsky, A., Sutskever, I., & Salakhutdinov, R.
812 (2014). Dropout: A simple way to prevent neural networks from overfitting.
813 *Journal of Machine Learning Research*, 15(56), 1929–1958.
- 814 Stirnberg, R., Cermak, J., Kotthaus, S., Haeffelin, M., Andersen, H., Fuchs, J.,
815 ... Favez, O. (2020). *Meteorology-driven variability of air pollution (PM1)*
816 *revealed with explainable machine learning* (Preprint). Aerosols/Field Mea-
817 surements/Troposphere/Physics (physical properties and processes). doi:
818 10.5194/acp-2020-469
- 819 Sugiura, M. (1964). Oart 1. In *Hourly values of equatorial dst for the igy*. Pergamon
820 Press.
- 821 Sugiura, M., & Kamei, T. (1991). Equatorial dst index 1957–1986, iaga bull., 40. by
822 A. Berthelier and M. Menville (*Int. Serv. Geomagn. Indices Publ. Off., Saint*
823 *Maur, 1991*).
- 824 Tsyganenko, N. (1989). A magnetospheric magnetic field model with a warped tail
825 current sheet. *Planetary and Space Science*, 37(1), 5–20. doi: 10.1016/0032
826 -0633(89)90066-4
- 827 Tsyganenko, N. A. (1995). Modeling the Earth’s magnetospheric magnetic field
828 confined within a realistic magnetopause. *Journal of Geophysical Research*,
829 100(A4), 5599. doi: 10.1029/94JA03193
- 830 Tsyganenko, N. A. (2002a). A model of the near magnetosphere with a dawn-dusk
831 asymmetry 1. Mathematical structure: A NEW MAGNETOSPHERE MAG-
832 NETIC FIELD MODEL, 1. *Journal of Geophysical Research: Space Physics*,
833 107(A8), SMP 12–1–SMP 12–15. doi: 10.1029/2001JA000219
- 834 Tsyganenko, N. A. (2002b). A model of the near magnetosphere with a dawn-
835 dusk asymmetry 2. Parameterization and fitting to observations: A NEW
836 MAGNETOSPHERE MAGNETIC FIELD MODEL, 2. *Journal of Geophys-*

- 837 *ical Research: Space Physics*, 107(A8), SMP 10–1–SMP 10–17. Retrieved
838 2021-08-05, from <http://doi.wiley.com/10.1029/2001JA000220> doi:
839 10.1029/2001JA000220
- 840 Wanliss, J. A., & Showalter, K. M. (2006). High-resolution global storm index:
841 Dst versus sym-h. *Journal of Geophysical Research: Space Physics*, 111(A2).
842 Retrieved from [https://agupubs.onlinelibrary.wiley.com/doi/abs/
843 10.1029/2005JA011034](https://agupubs.onlinelibrary.wiley.com/doi/abs/10.1029/2005JA011034) doi: <https://doi.org/10.1029/2005JA011034>
- 844 Wu, J.-G., & Lundstedt, H. (1997). Geomagnetic storm predictions from solar wind
845 data with the use of dynamic neural networks. *J. Geophys. Res.*, 102(A7),
846 14255–14268. doi: 10.1029/97JA00975
- 847 Xu, S. B., Huang, S. Y., Yuan, Z. G., Deng, X. H., & Jiang, K. (2020). Prediction of
848 the Dst Index with Bagging Ensemble-learning Algorithm. *ApJS*, 248(1), 14.
849 doi: 10.3847/1538-4365/ab880e
- 850 Zhang, T., & Yu, B. (2005). Boosting with early stopping: Convergence and consis-
851 tency. *Ann. Statist.*, 33(4). doi: 10.1214/009053605000000255

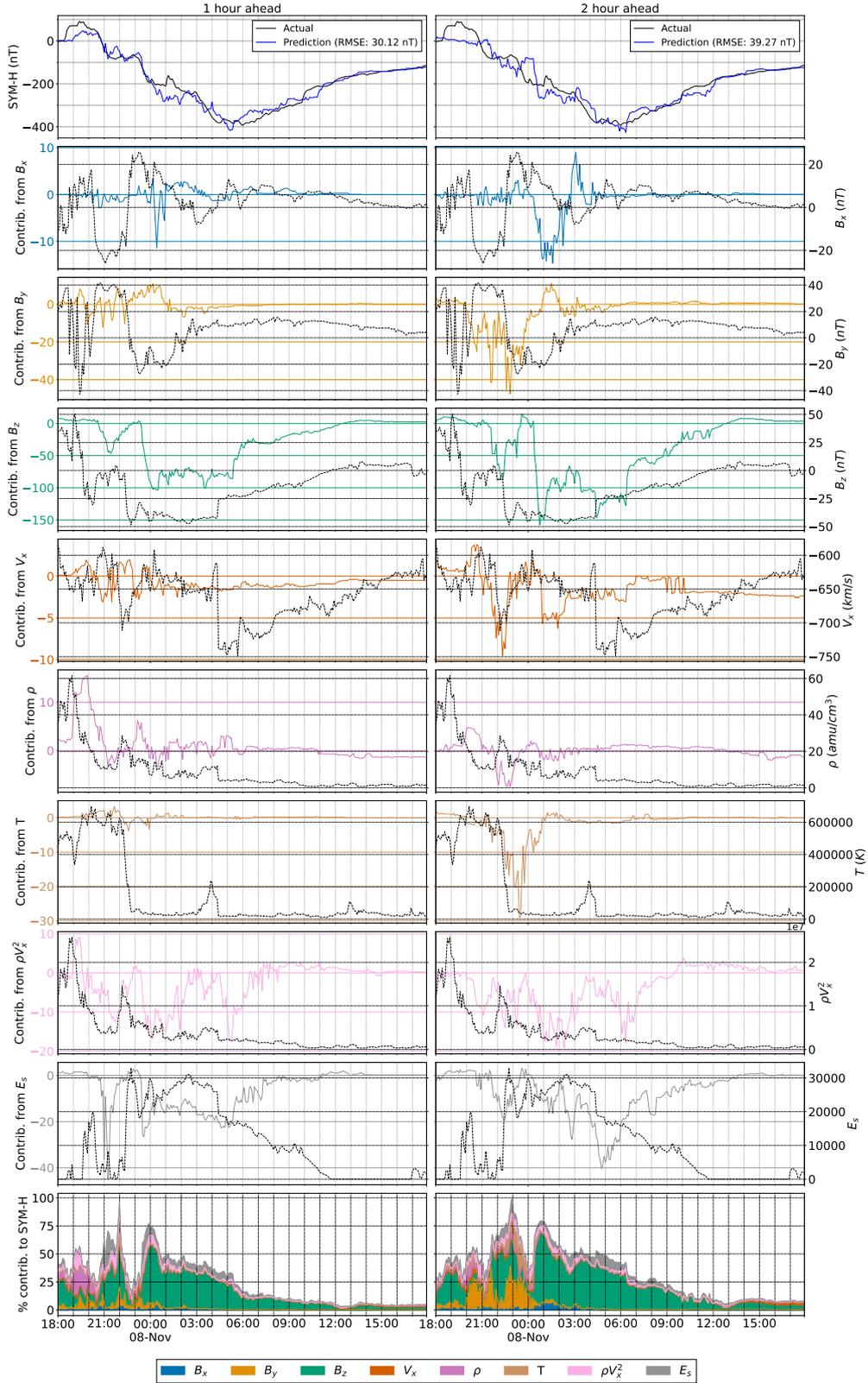


Figure 5. 1-hour (left) and 2-hour (right) ahead predictions for the Nov. 2004 storm using GBM trained on all considered features. The first row shows the observed (black) and predicted (blue) SYM-H values. Rows 2-9 show the contributions from each feature (left axis, colored) and its value (right axis, black). The percentage contributions are shown in the last row. The contribution from past SYM-H on predictions is omitted, but its percentage contribution is implicitly shown as the remaining white area in the last row.

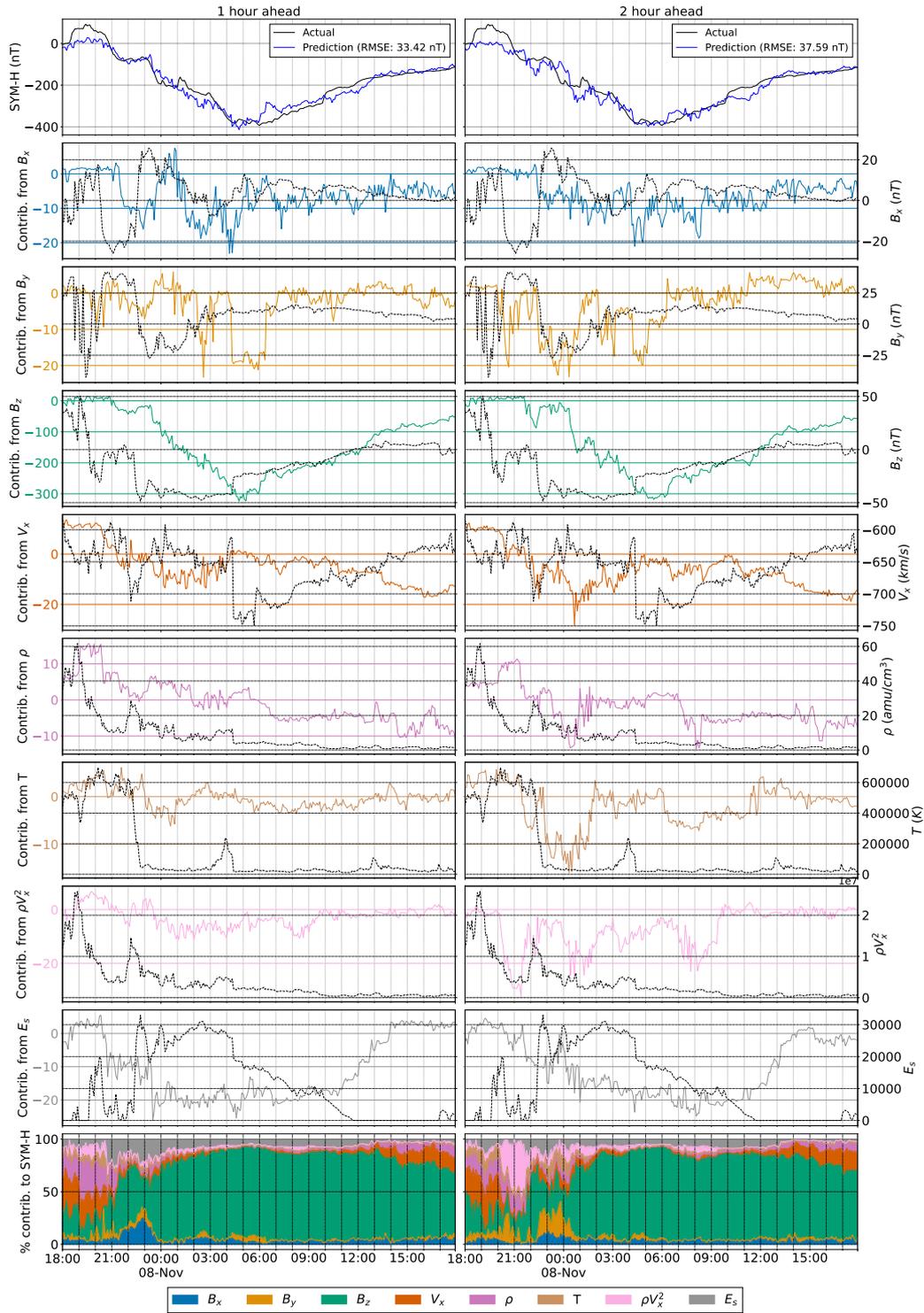


Figure 6. 1-hour (left) and 2-hour (right) ahead predictions for the Nov. 2004 storm using GBM trained on only solar wind and IMF parameters (first row), corresponding feature contributions and values (rows 2-9), and percentage contributions (last row).

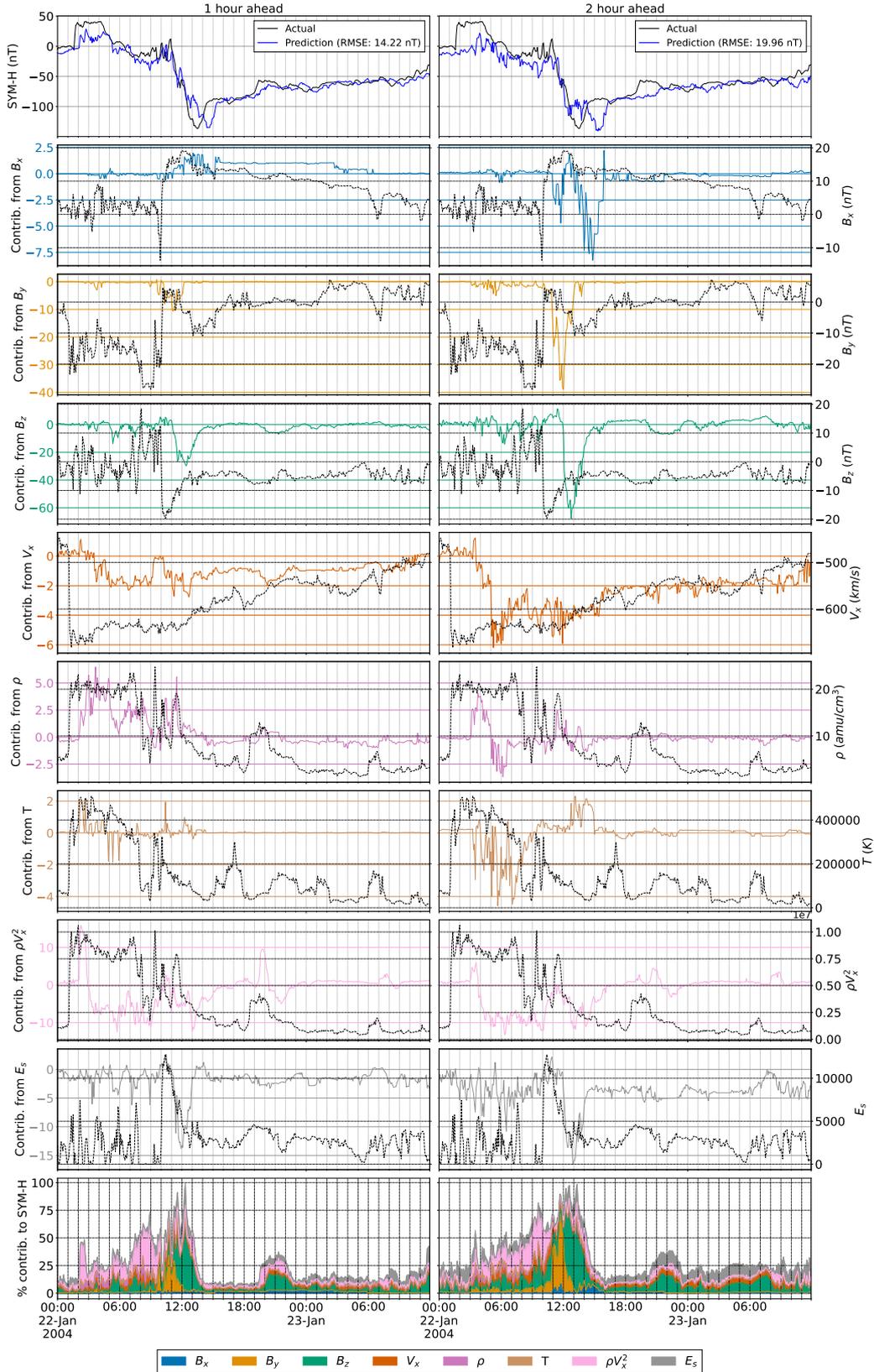


Figure 7. 1-hour (left) and 2-hour (right) ahead predictions for the Jan. 2004 storm using GBM trained on all considered features (first row), corresponding feature contributions and values (rows 2-9), and percentage contribution (last row). The contribution from past SYM-H on predictions is omitted but the percentage contribution is implicitly shown as the remaining white area in the last row.

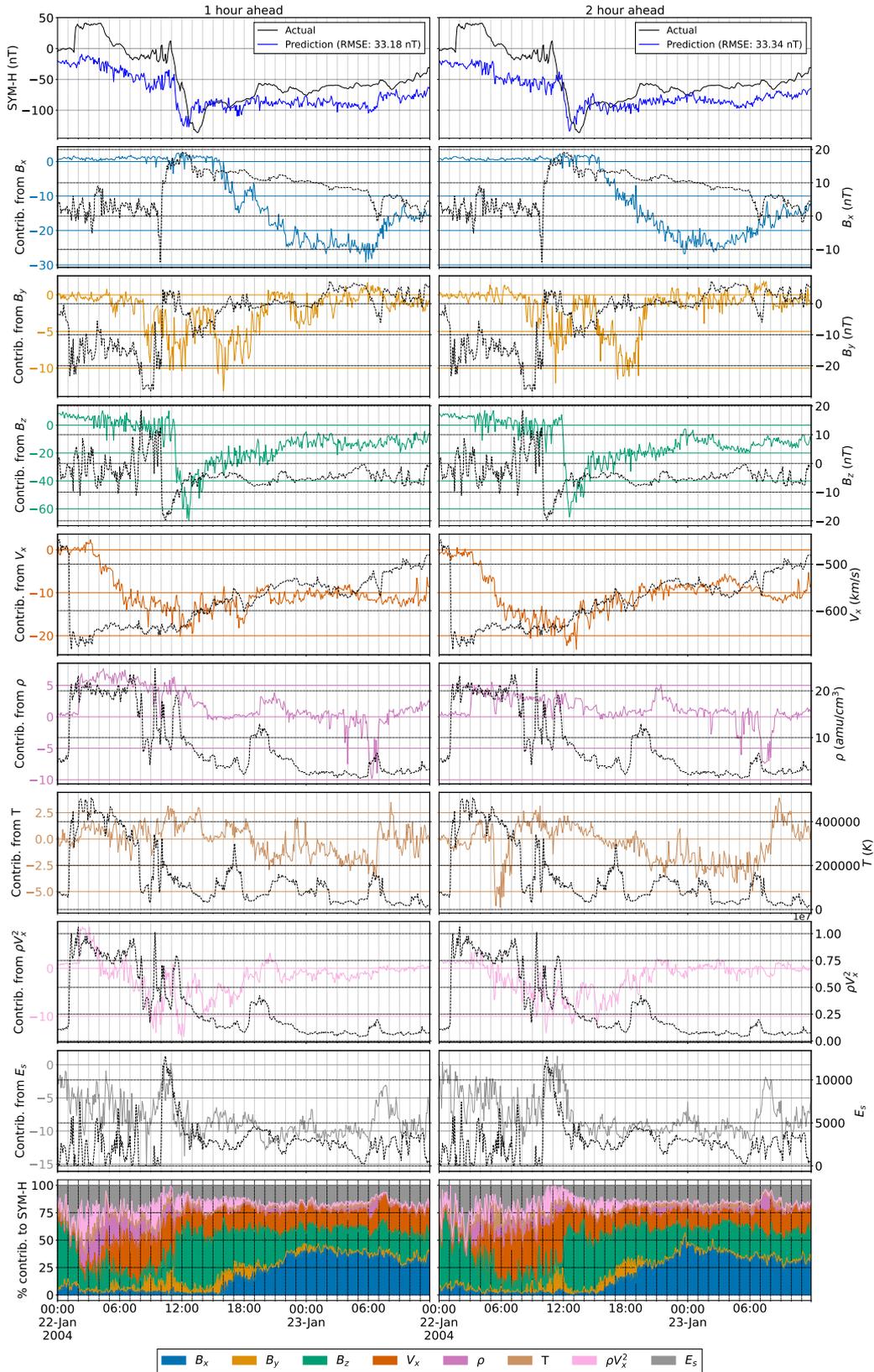


Figure 8. 1-hour (left) and 2-hour (right) ahead predictions for the Jan. 2004 storm using GBM trained on all considered features (first row), corresponding feature contributions (rows 2-9), and percentage contribution (last row).

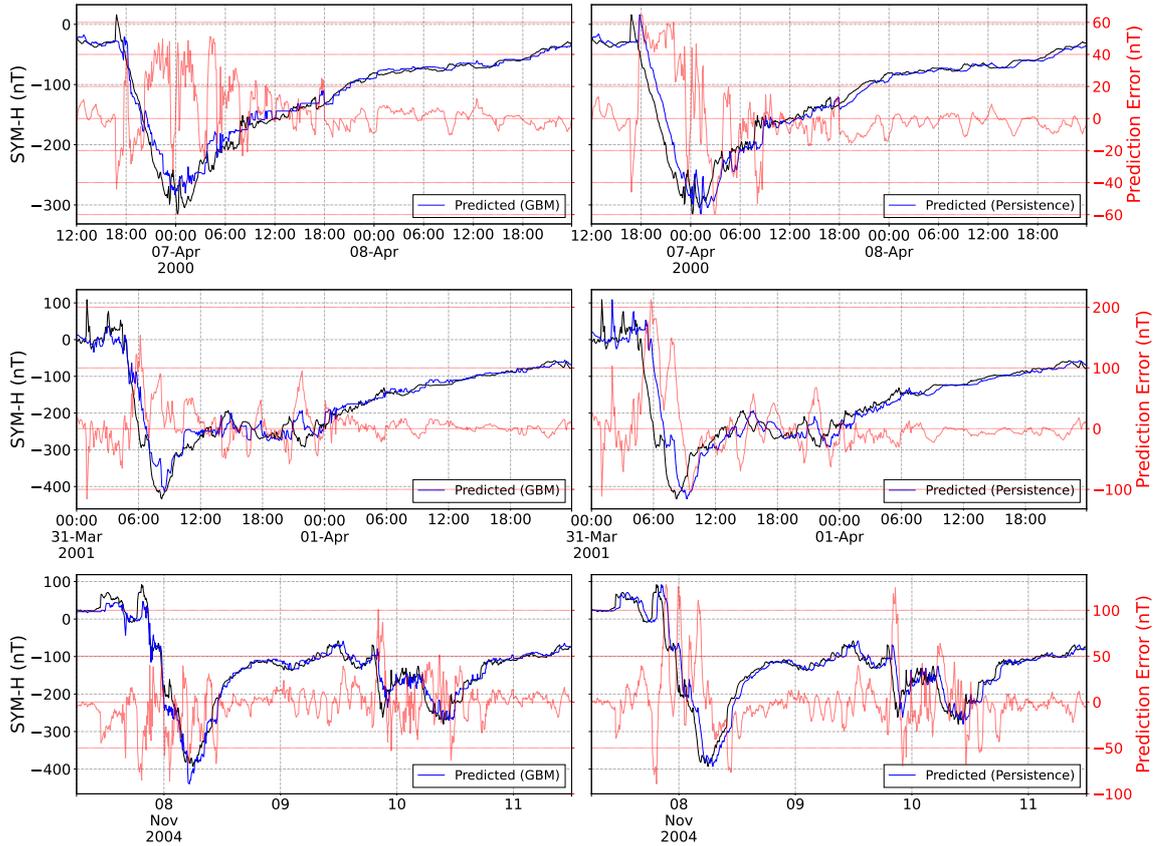


Figure A1. 1-hour ahead predictions for the 3 strongest geomagnetic storms in the test set during the main and recovery phases from our GBM with past SYM-H and IMF parameters as input (left column) and the persistence model (right column). The observed SYM-H (black), the predicted SYM-H (blue) and the error (red) are shown for storms 31, 33, and 37 in the 3 rows, respectively.

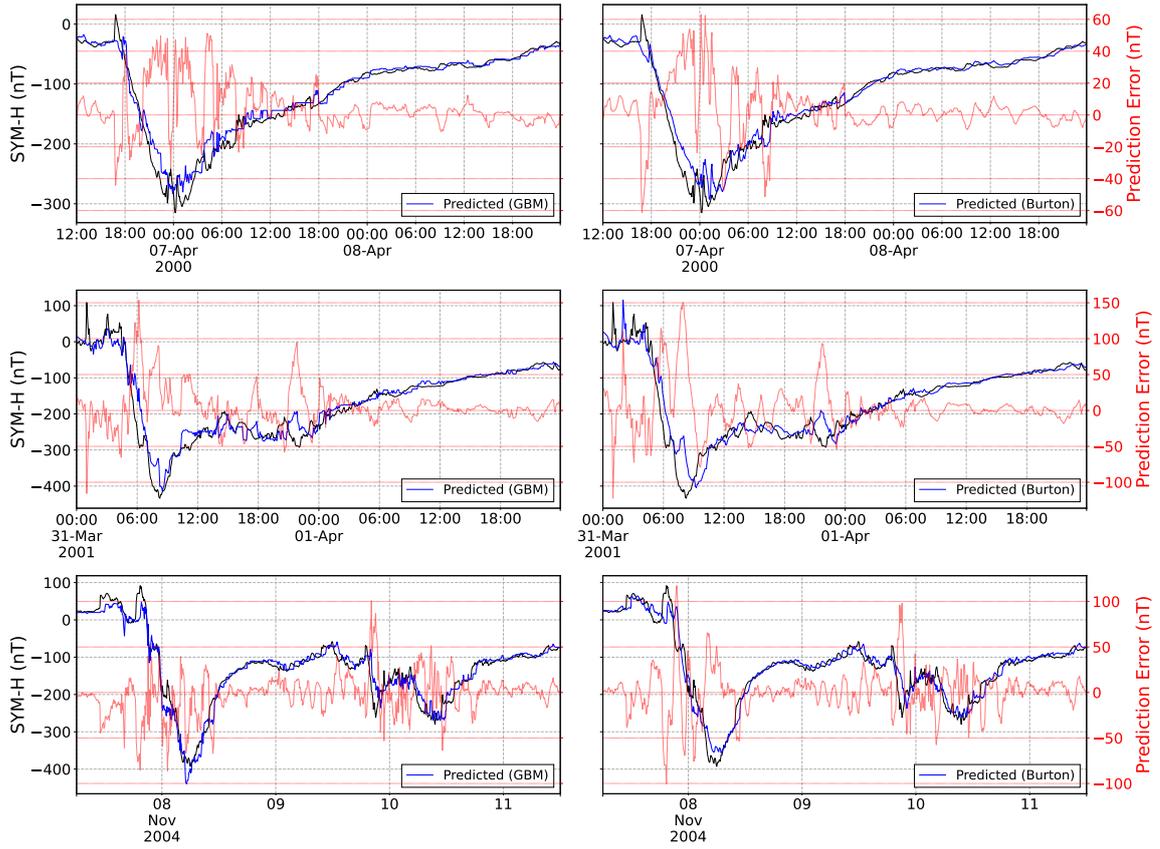


Figure A2. 1-hour ahead predictions for the 3 strongest geomagnetic storms in the test set during the main and recovery phases from our GBM with past SYM-H and IMF parameters (left column) and the Burton equation (right column). The observed SYM-H (black), the predicted SYM-H (blue) and the error (red) are shown for storms 31, 33, and 37 in the 3 rows, respectively.