

Stratospheric Circulation Changes Associated with the Hunga Tonga-Hunga Ha'apai Eruption

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Key Points:

- Extreme perturbations in the stratospheric winds and temperatures are linked to a volcanic eruption.
- Hunga Tonga-Hunga Ha'apai eruption effects are unique in the stratospheric record.
- Data assimilation can track temperature and wind perturbations, even when some physics is missing.

Abstract

The 15 January 2022 eruption of the Hunga Tonga-Hunga Ha'apai underwater volcano injected a record amount of water directly into the stratosphere. This study attempts to quantify this impact on the temperature, as well as the subsequent changes to the stratospheric circulation, during the months following the eruption based on reanalysis fields. The extreme nature of the temperature, wind, and circulation changes are tracked through comparisons of the first six months of 2022 with the past 42 years. Examination of the data assimilation process shows that at 20 hPa the thermal observations are forcing temperatures to cool significantly, compensating for the absence of the excess stratospheric moisture in the model used for the reanalysis, resulting in unusually low temperatures. In response to this cooling the atmosphere adjusts by creating strong westerly winds above the temperature anomaly and large changes to the downward and poleward mean meridional circulation.

Plain Language Summary

The 15 January 2022 eruption of the Hunga Tonga-Hunga Ha'apai underwater volcano injected a record amount of water directly into the stratosphere. This study attempts to quantify this impact on the temperature, as well as the subsequent changes to the stratospheric circulation, during the months following the eruption. The extreme nature of the stratospheric temperature, wind, and circulation changes are tracked through comparisons of the first six months of 2022 with the past 42 years. Details of the stratospheric perturbations in latitude and pressure are presented for June 2022, where anomalously low temperatures are found at near 20 km altitude from 60°S to 30°S. In response to this cooling the atmosphere adjusts by creating strong westerly winds above the temperature anomaly and large changes to the downward and poleward mean meridional circulation.

1 Introduction

The 15 January 2022 eruption of the Hunga Tonga-Hunga Ha'apai underwater volcano injected an unprecedented amount of water directly into the stratosphere (Millán et al., 2022; Xu et al., 2022; Carr et al., 2022). While the initial injection plume at 20°S reached to the upper stratosphere (Carr et al., 2022), Millán et al. (2022), showed that after three months this excess water vapor settled near 20 hPa altitude in a latitude band from 30°S to 5°N. With no major thermodynamic or photochemical sinks, this excess moisture is expected to remain in the stratosphere for two to three years. Water vapor is radiatively active in the infrared, contributing to the total radiative cooling in the stratosphere, which is dominated by the effects of carbon dioxide and ozone (e.g., Gille and Lyjak, 1986). These large perturbations in water vapor are expected to increase the amount of radiation lost to space, locally cooling the stratosphere. This study attempts to quantify this impact on the temperature, as well as the subsequent changes to the stratospheric circulation, during the first six months after the eruption.

The MERRA-2 (Modern-Era Retrospective Analysis for Research and Applications, Version 2) reanalysis (Gelaro et al., 2017) provides the circulation fields (temperatures and winds) for this study. While MERRA-2 assimilates a number of in-situ and space-based observations that constrain tropospheric moisture, stratospheric water vapor observations are not assimilated. Stratospheric moisture is closely constrained to monthly climatologies imposed by a

relaxation constraint in the GEOS model, so that the stratospheric water vapor in MERRA-2 does not respond to the volcanic perturbation. This precludes the use of MERRA-2 to directly infer the thermal impacts of the water vapor increase on the circulation. However, numerous nadir- and limb-sounding microwave observations are assimilated in the stratosphere (McCarty et al., 2016) and these data will constrain the analyzed temperatures, which can be used to indirectly infer the thermal and dynamical response of the stratosphere in MERRA-2. Additionally, a more recent data assimilation system, M2-SCREAM (described below), does contain additional water vapor and is used to compare with the MERRA-2 analysis and radiative temperature tendencies.

2 Assimilation Products

The 43-year (1980-2022) climate record from MERRA-2 is used to assess the anomalies of 2022. Two sets of monthly averaged files are used: The assimilation files for winds and temperatures (GMAO, 2015a), and the temperature tendency files to obtain the analysis temperature tendencies (GMAO, 2015b). The residual zonal mean circulations for each month are calculated from the monthly averaged assimilation files as, in addition to winds and temperatures, these files also contain the heat and momentum fluxes needed for evaluation of the residual circulation (see Andrews et al., 1987, page 128). For this study we examine the stream function of the residual circulation as well as the residual mean meridional and vertical winds.

The MERRA-2 Stratospheric Composition Reanalysis of Aura Microwave Limb Sounder (M2-SCREAM: Wargan et al., 2022) is a new stratosphere-focused reanalysis product developed at NASA's Global Modeling and Assimilation Office. M2-SCREAM assimilates stratospheric profiles of ozone, water vapor, hydrogen chloride, nitric acid, and nitrous oxide from version 4.2 retrievals of the Microwave Limb Sounder (MLS: Waters et al., 2006; Livesey et al., 2020) observations, the same as those used in Millán et al. (2022), alongside total ozone observations from the Ozone Monitoring Instrument (Levelt et al., 2006, 2018). Temperature, winds, surface pressure, and tropospheric water vapor in M2-SCREAM are constrained by the MERRA-2 assimilated fields. M2-SCREAM covers the MLS period from September 2004 to May 2022. Details of the M2-SCREAM system are described in Wargan et al., (2020; 2022). Because this reanalysis assimilates MLS data, unlike MERRA-2, it contains a representation of the water vapor enhancement from the Hunga Tonga eruption in agreement with MLS. The total increase of stratospheric water vapor mass calculated from M2-SCREAM is 10% as in Millán et al. (2022). As shown below, the radiative transfer model in M2-SCREAM responds to the moisture enhancement by producing long wave cooling in better agreement with the observations.

3 Results

The assimilation technique in MERRA-2 imposes an additional forcing, the analysis tendency, to force the GEOS model towards the observations. This additional forcing can often be regarded as a compensation for inadequacies in the representation of processes in the model, and can manifest as either random error (with a complex spatial structure) or as a mean bias, with a coherent spatio-temporal structure. Because the GEOS model uses climatological water vapor fields in the stratosphere of MERRA-2, any anomalous radiative forcing caused by the volcanic eruption will not be captured in the radiative tendencies, and will instead be captured by the bias

100 in the analysis tendency term. The global analysis temperature tendencies at 20hPa, shown for
 101 each month of MERRA-2 (Fig. 1a), reveals the anomalous values seen in 2022. Cooling by the
 102 increments is below average in January 2022, one of the strongest cooling years in February, and
 103 lies well below the average in March through June. In May 2022 the cooling is over three
 104 standard deviations below the mean. These anomalies at 20 hPa coincide with the peak moisture

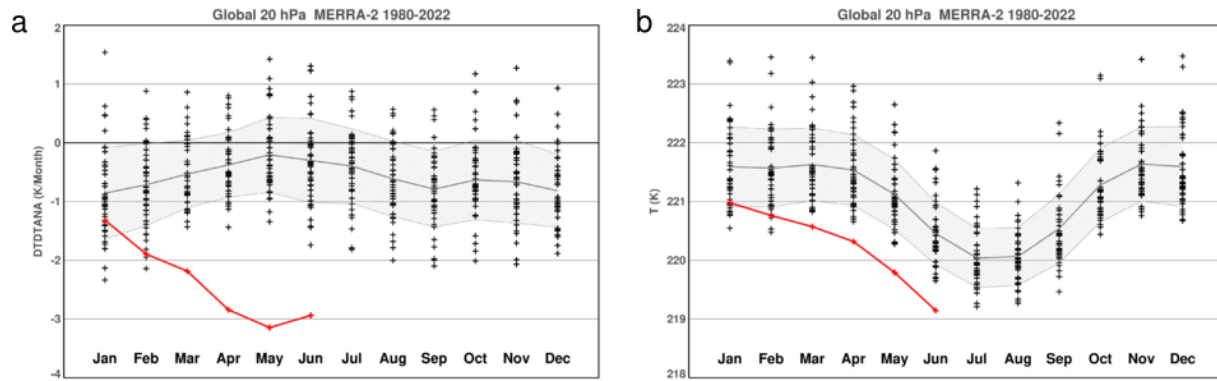
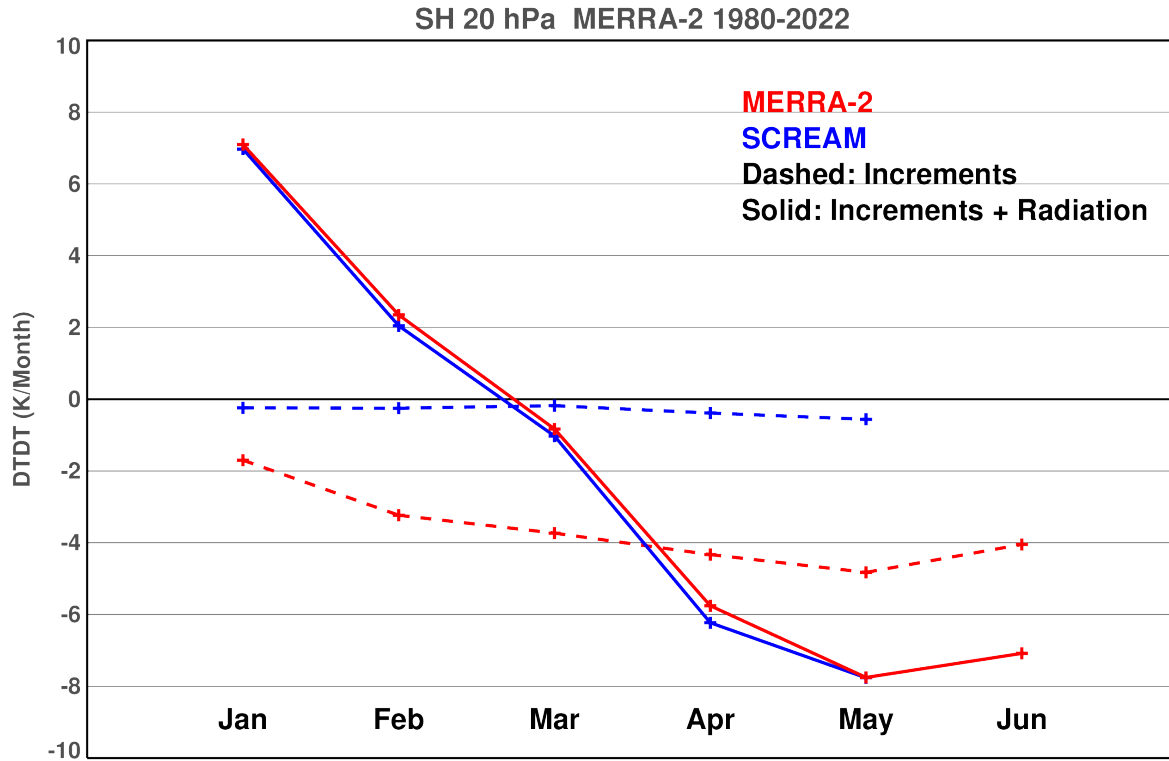


Figure 1: The monthly and globally averaged a) temperature increments and b) temperature at 20 hPa. The gray curve denotes the multi-year mean, the gray shading denotes the standard deviation for each month, and the red curve denotes the year 2022.

105 anomaly isolated in Millán et al. (2022), are the largest in MERRA-2. The increments show
 106 weaker additional cooling at 10hPa and weak additional warming at 30hPa (not shown).



107
 108 *Figure 2: The 20 hPa monthly and Southern Hemisphere averaged MERRA-2 temperature*
 109 *increments (red dashed), MERRA-2 sum of temperature increments and radiative tendency (red*
 110 *solid), M2-SCREAM temperature increments (blue dashed), and SCREAM sum of temperature*
 111 *increments and radiative tendency (blue solid).*

112 The global temperature at 20 hPa (Fig. 1b) depicts the cooler than average global
 113 temperatures associated with the record cold increments during January through June 2022. The
 114 January temperatures are approximately one standard deviation below average, the February
 115 temperatures continue to cool, while the March through June temperatures are setting new
 116 records for cold global temperatures at 20 hPa. Note that the temperature increments can suggest
 117 temperatures larger than the temperature departures. For example, in June 2022 the temperature
 118 increments are nearly 3 K/month below the mean, while the temperatures for the month are only
 119 ~1.25 K below average indicating the importance of the other terms, such as dynamical forcing
 120 and radiative transfer, in determining temperature. Nevertheless, these record cold increments are
 121 associated with the record low temperatures.

122 More evidence that the anomalous increments correspond to expected anomalous water vapor is
 123 provided by M2-SCREAM temperature tendency fields (Fig. 2). These are shown averaged over
 124 the Southern Hemisphere (SH) rather than globally as the global averaged summed temperature
 125 tendencies are much smaller. The M2-SCREAM system's smaller magnitude temperature
 126 increments than MERRA-2 (dashed curves) show that M2-SCREAM can respond radiatively to

the water vapor anomaly without a strong reliance on the temperature observations. However, when the analysis and radiative tendencies are summed (solid curves), both systems show similar temperature tendencies indicating that the MERRA-2 increments are replacing the additional water vapor portion of the radiative tendency.

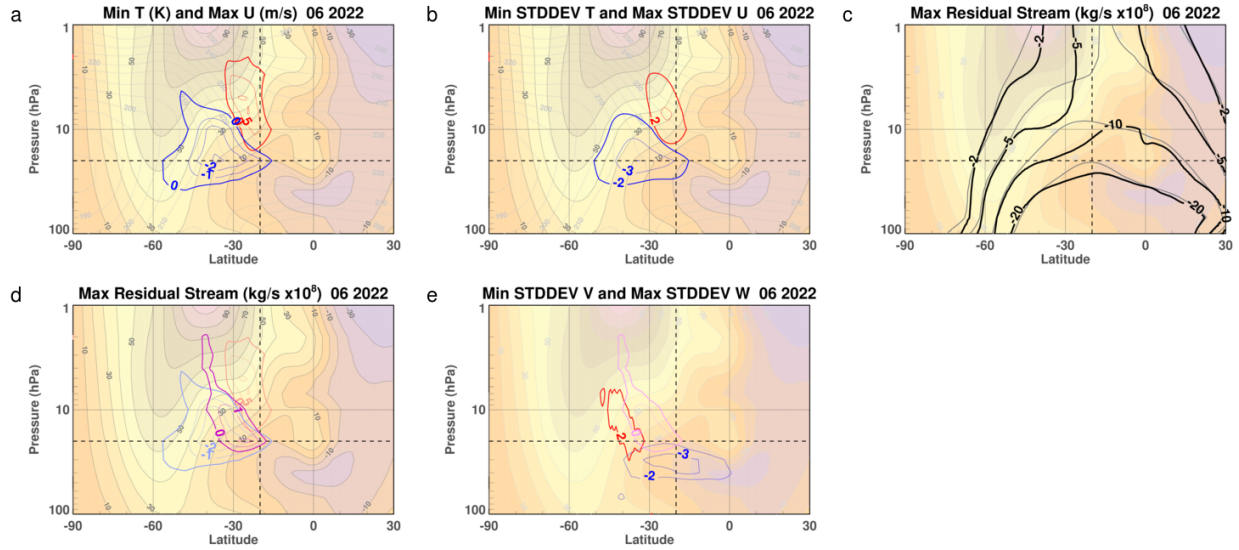


Figure 3 Cross sections for June showing zonal mean zonal winds (10 m/s, filled contours) and temperatures (5K, gray contours (a and b): a) Blue contours: record cold regions where the contours (1 K) denote how much June 2022 is below the previous record low temperature. Red contours: record strong zonal mean winds where the contours (5 m/s) denote how much June 2022 is above the previous record strong wind; b) Blue contours: standard deviations of temperature below the mean (contour interval of 1, starting at -2) and red contours: standard deviations of zonal wind above the mean (contour interval of 1, starting at 2); c) Residual mean stream function ($-2, -5, -10, -20 \times 10^8$ kg/s) for 1980-2021 average (gray) and 2022 (black); d) residual mean circulation stream function greater than the past maximum value (magenta, contour interval 1×10^8 kg/s) with wind and temperature record as in a); e) Blue contours: standard deviations of residual mean meridional wind below the mean (contour levels of -3 and -2) and red contours: standard deviations of residual mean vertical velocity above the mean (contour level of 2). Dashed lines denote 20°S and 20 hPa.

This cooling is not uniform over the globe but is strongest near 30°S and 20 hPa. In June 2022 record low temperatures for the month stretch from 55°S to 15°S (Fig. 3a). These temperatures break the previous low temperature record by as much as 3K. In addition, the zonal mean winds are breaking records by as much as 10 m/s. The location of these record strong winds near the low temperatures is consistent with the geostrophic relation where increased cooling toward the pole requires increased vertical wind shear. In addition to setting records for the month of June, these 2022 low temperatures and strong winds were outside the standard deviation of the year-to-year variability (Fig. 3b) with values greater than double the standard deviation.

These wind and temperature anomalies are likely associated with changes in the mean circulation as the atmosphere adjusts to the temperature perturbation. The counter-clockwise

flow of the residual mean stream function for 2022 (Fig. 3c) shows large distortions in the region near the wind and temperature anomalies compared to the 1980-2022 averaged June residual mean stream function. In particular, the strong vertical gradient in the stream function at 30°S and 30 hPa denotes a stronger than average poleward (negative) flow in 2022. This can be represented as a clockwise anomaly in the residual mean stream function (Fig. 3d). In Fig. 2d, the stream function plotted is greater than any of the previous years in that region. This means that the distortion of the stream function from the mean seen in Fig. 3c is larger than in previous years.

The residual mean circulation can also be expressed in terms of residual mean meridional and vertical velocities (Fig. 3e). The residual mean meridional velocity is particularly striking with negative (poleward) values over three standard deviations below the mean from 10°S-30°S near 30 hPa. The upward mean vertical wind anomalies are over two standard deviations above the mean on the poleward side of the stream function anomaly. This upward anomaly does not correspond to an actual upward circulation but expresses the weaker downward circulation than average as seen in the nearly horizontal stream function regions in Fig. 3c.

4 Conclusions

Anomalous temperatures and circulation patterns analyzed by MERRA-2 in the southern hemisphere during June 2022 can be forensically attributed to the stratospheric water vapor injection from the January 2022 eruption of the Hunga Tonga-Hunga Ha'apai underwater volcano. These anomalies can be traced back to March 2022. Their consistency in space and time suggests a realistic response to a geophysical event rather than a yearly random dynamical fluctuation. In June the record winds are part of an unusual secondary jet maximum at 10 hPa, 30°S-20°S.

These wind and temperature anomalies (Fig. 3a, b) develop from the assimilation of data, mainly routine, satellite based, nadir viewing radiometers and geostrophic balance and are likely to be very realistic. Note that MERRA-2 does assimilate MLS temperatures, but only at pressures of 5hPa and lower. The residual mean circulation might be more difficult to interpret. If the cooling analysis temperature increments mainly reflect the missing water vapor cooling then they can be considered to be the missing cooling term from the lack of stratospheric water vapor in the assimilation system. Then the calculated residual circulation (Fig. 3c, d, e) should realistically capture the perturbed residual circulation. If, however, the analysis temperature increments also contain cooling induced by circulations in the atmosphere's response to the water vapor perturbation, then it is possible that the residual circulation may adjust in an unphysical manner. However, the good SH 20 hPa agreement between MERRA-2 and M2-SCREAM seen in the sum of the analysis and radiative temperature tendencies (Fig. 2) suggests that the MERRA-2 analysis tendencies are representative of the missing water vapor cooling. Future work is planned for model simulations that include a realistic representation of the stratospheric water perturbations. These calculations should provide a more complete picture of the atmospheric response to the volcanic perturbation.

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Open Research

The MERRA-2 products are available on the NASA Goddard Earth Sciences Data and Information Services Center (GES DISC). Specific MERRA-2 product collections used here are cited appropriately in the references. The 2022 temperature tendencies from M2-SCREAM are available from <https://gmao.gsfc.nasa.gov/gmaoftp/STRATOSPHERE/M2-SCREAM/T-tendencies/>. Historical M2-SCREAM output can be accessed from <https://disc.gsfc.nasa.gov/datasets?keywords=m2-scream&page=1>.

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