

1  
2  
3  
4  
5  
6  
7  
8  
9  
10  
11  
12

# Prolonged Lifetime of the Transient Ionized Layer in the Martian Atmosphere Caused by Comet Siding Spring

Z. A. Luppen<sup>1</sup>, Z. Girazian<sup>1</sup>, D. D. Morgan<sup>1</sup>, A. J. Kopf<sup>1</sup>, F. Chu<sup>1</sup>, J. S. Halekas<sup>1</sup>, D. A. Gurnett<sup>1</sup>

<sup>1</sup>Department of Physics and Astronomy, University of Iowa, Iowa City, Iowa, USA

## Key Points:

- The transient ionospheric layer at Mars caused by comet Siding Spring's flyby may have lasted at least 7 days, and up to 32 days.
- All detected transient layer measurements were located between 20°N-60°N latitude, and most were confined to one longitudinal hemisphere.
- We discuss how solar flares may have contributed to the prolonged lifetime of the transient layer.

---

Corresponding author: Zachary Luppen, zaluppen@iastate.edu

## Abstract

In October 2014, the close encounter between Mars and comet Siding Spring produced a transient ionized layer in the upper atmosphere composed primarily of  $\text{Mg}^+$  ions. The layer was detected by instruments on three spacecraft, including the Mars Advanced Radar for Subsurface and Ionosphere Sounding (MARSIS) on Mars Express. Analyses of the MARSIS data indicated the transient layer persisted up to  $\sim 19$  hours after the comet's closest approach. We report MARSIS observations that suggest the transient layer lasted at least 7 days – and potentially as long as 32 days – after closest approach. During this period, the transient layer was mostly confined to a narrow latitude range between  $20^\circ\text{N}$ - $60^\circ\text{N}$  and a longitude range spanning  $275^\circ\text{E}$  to  $95^\circ\text{E}$ . Since this period coincided with a highly active Sun, we discuss how solar flares may have contributed to the layer's prolonged lifetime.

## 1 Introduction

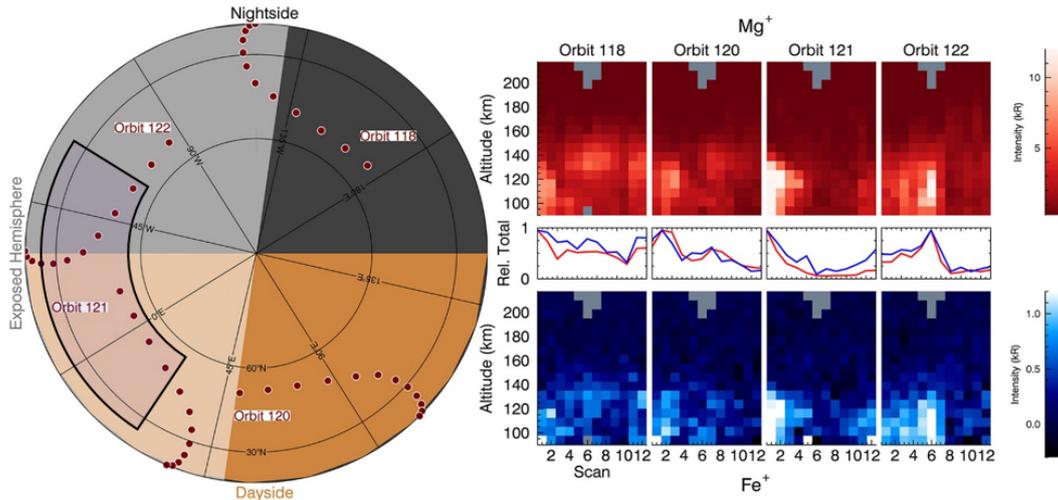
On 19 October 2014, at 18:29 UT, comet Siding Spring (C/2013 A1) flew close to Mars, passing within 40 Mars radii along a highly-inclined orbital path. The event presented a rare opportunity to understand how cometary material is deposited and distributed in the Martian upper atmosphere, and to test models of metal ion chemistry. During the close flyby,  $82 (\pm 25)$  tons of cometary material was deposited into the upper atmosphere of Mars with an approach velocity of  $56 \text{ km s}^{-1}$  [Crismani *et al.*, 2018]. Upon entry, the ablated material deposited an abundance of metallic species such as Mg, Na, and Fe into the upper atmosphere [Benna *et al.*, 2015; Schneider *et al.*, 2015; Crismani *et al.*, 2018]. The material formed a transient ionized layer with a peak altitude near 115 km that was detected by instruments on multiple spacecraft. This layer, composed primarily of  $\text{Mg}^+$  ions, was produced by the ionization of ablated material through collisions, photoionization, and charge exchange with existing ionospheric ions [Whalley and Plane, 2010; Crismani *et al.*, 2018; Plane *et al.*, 2018].

Within 10 hours after the comet's closest approach ( $\Delta t_{ca} = 10$  hours), the Mars SHARAD sounder (SHARAD) on the Mars Reconnaissance orbiter (MRO) detected the transient ionized layer during two MRO orbits. Both of the observations, from the nightside of the planet, showed an exceptional increase in the nightside total electron content (TEC) that was attributed to ablated cometary material [Restano *et al.*, 2015].

The Mars Atmosphere and Volatile Evolution (MAVEN) spacecraft, which had just reached Mars in late September 2014, was in a limited powered-on state to collect data during the event. In situ measurements by MAVEN's Neutral Gas and Ion Mass Spectrometer (NGIMS) showed an abrupt increase in  $\text{Mg}^+$ ,  $\text{Na}^+$ ,  $\text{Fe}^+$ , and other metallic ion species during its first observations on 20 October ( $\Delta t_{ca} \approx 10$  hours) [Benna *et al.*, 2015]. The metal ions were continuously detected until periapse observations ceased on 22 October ( $\Delta t_{ca} \approx 2.5$  days). Benna *et al.* [2015] also reported that the  $\text{Mg}^+$  abundance at 185 km decayed with an e-folding timescale of  $\sim 8$  days, much longer than the 1.8 days predicted by a 1-D chemical model [Whalley and Plane, 2010]. This discrepancy was attributed to the rapid transport of metallic ions by thermospheric winds, which were not included in the 1-D model.

During the flyby, periapse limb scans from MAVEN's Imaging Ultraviolet Spectrograph (IUVS) between  $50^\circ\text{N}$  and  $5^\circ\text{S}$  measured  $\text{Mg}^+$  and  $\text{Fe}^+$  transient layer altitude profiles [Crismani *et al.*, 2018]. The observations showed that the transient layer was primarily composed of  $\text{Mg}^+$ , with a peak density of as much as  $2 \times 10^5 \text{ cm}^{-3}$ , and a peak altitude near 115 km [Schneider *et al.*, 2015; Crismani *et al.*, 2018]. These results also indicated that ablation occurred on a timescale less than 4.5 hr, primarily impacting only one hemisphere of Mars between  $\sim 50^\circ\text{E}$  and  $\sim 230^\circ\text{E}$  longitude (see Figure 1) [Crismani *et al.*, 2018]. Hours after closest approach, the metallic species were observed globally, including on the unexposed hemisphere, once again implying that they were transported across the planet by thermospheric winds.

62 Figure 1, adapted from Figure 9 in *Crismani et al.* [2018], shows the IUVS observa-  
 63 tions from 20-21 October ( $\Delta t_{ca} \approx 1.0\text{-}2.5$  days). These IUVS limb scans measure ultraviolet  
 64 emissions, which can be converted into altitude profiles of the  $\text{Mg}^+$  and  $\text{Fe}^+$  densities. Dur-  
 65 ing orbits 121-122, the transient layer was confined to a narrow latitude range between  $20^\circ\text{N}$ -  
 66  $60^\circ\text{N}$ . The highest densities were observed near a region of the exposed hemisphere spanning  
 67  $\sim 90^\circ$  between  $300^\circ\text{E}$  and  $30^\circ\text{E}$  longitude during MAVEN orbits 121 and 122 (shaded region,  
 68 Figure 1). These measurements taken several days after closest approach are peculiar, and  
 69 hint at a transport mechanism that confines metal ions to specific locations. Such a mecha-  
 70 nism has yet to be determined, although *Crismani et al.* [2017] suggested global circulation  
 71 patterns or crustal magnetic fields may play a role.



72 **Figure 1.** MAVEN IUVS observations of the transient ionized layer caused by comet Siding Spring. The  
 73 observations range from 20-21 October, covering 1.0-2.5 days after the comet's closest approach. The obser-  
 74 vational geometry on the left side of the figure shows the location of each IUVS limb scan. For each orbit,  
 75 the scans move consecutively from north to south and west to east. The  $\text{Mg}^+$  and  $\text{Fe}^+$  emissions observed  
 76 during the limb scans are shown on the right side. Localized, bright emissions indicate the metal ions were  
 77 patchy and variable throughout the upper atmosphere. The locations of these localized emissions during orbits  
 78 121-122 are marked with a shaded region. This figure is adapted from Figure 9 of *Crismani et al.* [2018].

79 The Mars Advanced Radar for Subsurface and Ionosphere Sounding (MARSIS) on  
 80 Mars Express (MEX) also detected the transient layer on both the dayside and nightside up  
 81 to  $\Delta t_{ca} \approx 19$  hours [*Gurnett et al.*, 2015; *Venkateswara Rao et al.*, 2016]. The maximum  
 82 observed peak electron density of  $2 \times 10^5 \text{ cm}^{-3}$  was consistent with the IUVS observations  
 83 [*Gurnett et al.*, 2015; *Venkateswara Rao et al.*, 2016; *Crismani et al.*, 2018]. Although the  
 84 reported peak altitude of the transient layer by MARSIS was  $100 \pm 7$  km [*Gurnett et al.*,  
 85 2015], improved data processing techniques by the MARSIS team later showed that the ob-  
 86 served peak altitude was  $113 \pm 7$  km, much more consistent with the IUVS observations.

87 Despite these observations and analyses, there are still unanswered questions in regards  
 88 to the chemistry and dynamics of the metal ions in the upper atmosphere of Mars. To date,  
 89 models have been unable to reproduce the behavior of metal ions following Siding Spring's  
 90 closest approach [*Benna et al.*, 2015; *Crismani et al.*, 2018]. One reason for this is that the  
 91 models thus far do not include winds, which can transport the metal species to locations far  
 92 from where they were created. Moreover, although it is generally accepted that metal ions  
 93 have long chemical lifetimes of days or even weeks, their chemical lifetimes are poorly con-  
 94 strained by direct observations [*Crismani et al.*, 2017; *Grebowsky et al.*, 2017].

In this paper, we further analyze MARSIS data to look for transient layer detections up to a month after Siding Spring’s closest approach. About 2.5 days after closest approach, MAVEN ceased normal science operations for several weeks, making MARSIS one of the only instruments able to monitor the ionosphere during this time. MARSIS is a topside sounder that can only measure electron density and cannot discriminate between different ions [Gurnett *et al.*, 2015]. Thus, we cannot definitively prove that we observe the  $\text{Mg}^+$  layer. Nonetheless, we present peculiar MARSIS observations after the comet’s closest approach that suggest the transient  $\text{Mg}^+$  layer lasted for at least 7 days.

It is also worth noting that Siding Spring’s encounter with Mars coincided with a period of strong space weather. Two days prior to closest approach, Mars encountered an interplanetary coronal mass ejection (ICME) which was detected by MEX, MAVEN, Mars Odyssey, and the Mars Science Laboratory [Witasse *et al.*, 2017]. These instruments indicated an increase in the solar wind density by a factor of  $\sim 2$  as well as an increase in solar wind velocity, from  $\sim 400$  km/s to  $\sim 700$  km/s [Witasse *et al.*, 2017]. Sánchez-Cano *et al.* [2020] discussed how, before and during closest approach, the ionosphere was unusually variable, likely due to combined effects of the comet flyby and strong space weather [Espley *et al.*, 2015]. Sánchez-Cano *et al.* [2018] noted an increase in solar wind speed at Mars lasted from October 17-22. While the comet’s closest approach to Mars was during this period, our key follow-up measurements on October 26 that we discuss are not. In Section 4, we discuss the potential effects of space weather on our measurements.

## 2 MARSIS Observations of the Transient Layer

### 2.1 The MARSIS Instrument

MEX is in a  $\sim 7$  hour, near-polar orbit with an inclination of  $86^\circ$ . The orbit is highly elliptical, with an apoapsis altitude near 10,000 km, and a periapsis altitude near 375 km. During the periapsis portion of the orbit, when MEX is below  $\sim 1500$  km, MARSIS conducts radar soundings using two 20 meter antennas, operating in one of two different observing modes: one for sounding the subsurface of Mars, and one for sounding the ionosphere [Pircardi *et al.*, 2004]. In the Active Ionospheric Sounding (AIS) mode, the MARSIS transmitter emits radio pulses at frequencies between 0.1-5.4 MHz, and then measures the intensities and time delays of the return echoes – radio pulses that have been reflected off the ionosphere below  $\sim 250$  km. Each frequency sweep takes 1.26 seconds and the entire process is repeated every 7.54 seconds.

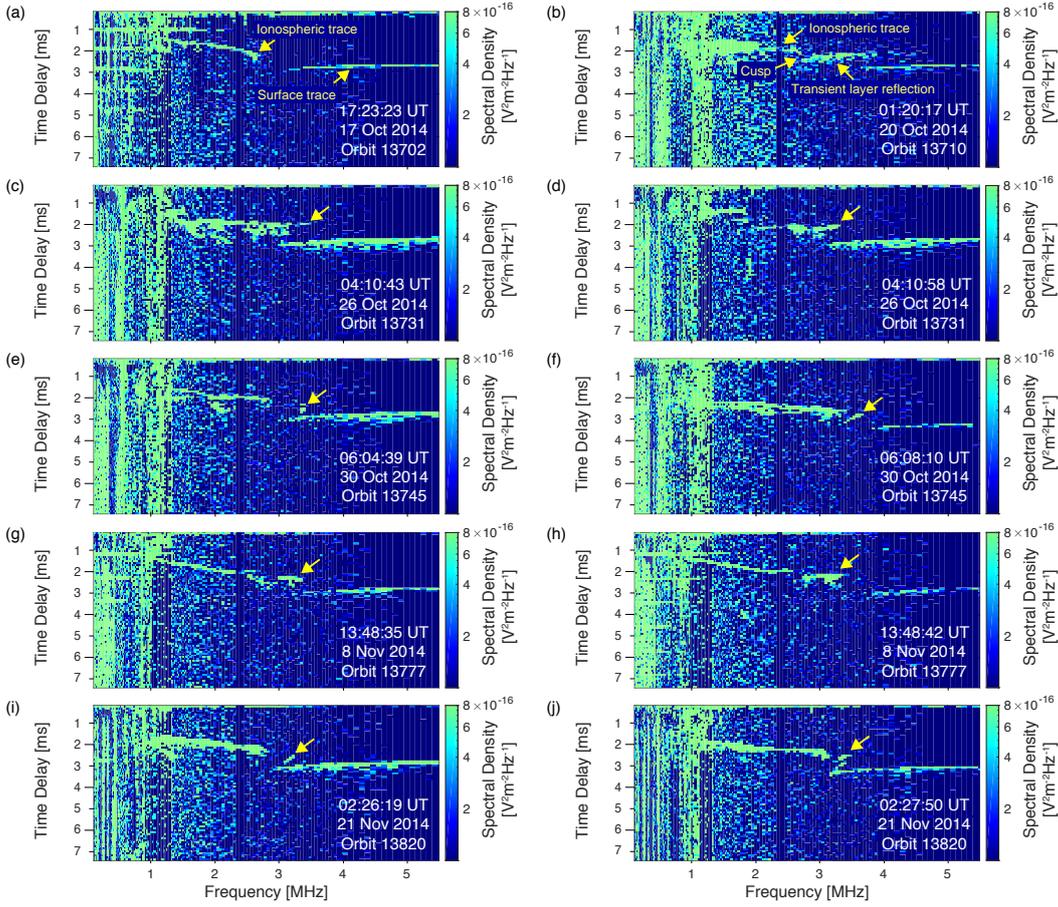
The main data products produced by MARSIS soundings are ionograms – color-coded plots of the return echo intensities as a function of frequency and time delay [Gurnett *et al.*, 2005, 2008]. The time delay of each return echo provides information about its reflection altitude, with longer time delays indicating a lower altitude of reflection. The frequency of each return echo is a measure of the electron density at the reflection altitude, because radio waves are reflected when the local electron plasma frequency is equivalent to that of the MARSIS signal. For an extensive discussion of the MARSIS ionospheric sounding in context of the Siding Spring encounter, we refer the reader to Gurnett *et al.* [2015] and its supplementary material.

### 2.2 Identifying the Transient Layer in MARSIS Ionograms

#### 2.2.1 Observations shortly before and after closest approach

Figure 2a shows an ionogram from 17 October 2014, two days before Siding Spring’s closest approach. The ionogram has a clear ionospheric trace that is typical for observations of the dayside ionosphere far from appreciable crustal magnetic fields [Gurnett *et al.*, 2005, 2008]. The ionospheric trace is the thin line extending from 1.0-2.8 MHz whose time delay gradually increases with frequency. The time delay increases with frequency because, in the

143 topside ionosphere, the electron density generally increases exponentially towards lower alti-  
 144 tudes until the altitude of the main peak. The abrupt end of the ionospheric trace at  $\sim 2.8$   
 145 MHz represents the reflection from the main peak of the ionosphere. Given the equation  
 146 for the electron plasma frequency,  $f_p = 8980 \sqrt{N_e}$  Hz, where  $N_e$  is the electron density in  
 147  $\text{cm}^{-3}$ , this indicates a peak electron density of  $9.7 \times 10^4 \text{ cm}^{-3}$ . Note that because MARSIS  
 148 is a topside sounder, it cannot observe below the main peak of the ionosphere unless there  
 149 is an additional layer there whose peak density exceeds that of the main peak. In addition to  
 150 the ionospheric trace, the surface reflection trace is also present in Figure 2a, appearing as an  
 151 approximately flat line between 3.2-5.5 MHz.



152 **Figure 2.** (a) A dayside ionogram with a typical ionospheric trace and a typical surface trace. (b) An  
 153 ionogram from shortly after Siding Spring’s closest approach ( $\Delta t_{ca} \approx 7$  hours) showing the reflection of the  
 154 transient  $\text{Mg}^+$  layer produced by the deposition of cometary material. (c-j) Eight ionograms with transient  
 155 layer reflections (arrows) that we identified using the method explained in the text. The dates of the transient  
 156 layer detections range from 26 October - 21 November ( $\Delta t_{ca} \approx 7$ -32 days).

157 Figure 2b shows an ionogram from 20 October 2014, shortly after Siding Spring’s  
 158 closest approach ( $\Delta t_{ca} \approx 5$  hours). The ionospheric trace is present, and there is an addi-  
 159 tional distinct feature between 2.8-3.8 MHz that is not present in the typical ionogram shown  
 160 in Figure 2a. Based on its unique shape and extension to high frequencies, this feature has  
 161 been identified as the  $\text{Mg}^+$  transient layer reflection [Gurnett et al., 2015; Venkateswara  
 162 Rao et al., 2016; Crismani et al., 2018]. Unlike a typical ionospheric trace, the time delay  
 163 of the transient layer reflection decreases with frequency, suggesting that there is a “cusp”  
 164 between the ionospheric trace and the transient layer reflection. A cusp is a discontinuity in

165 the trace that occurs when there is a local maximum in the electron density, such as at the  
 166 main peak of the ionosphere [Gurnett *et al.*, 2008; Kopf *et al.*, 2008]. The presence of the  
 167 cusp implies a non-monotonic variation in the electron density with altitude, i.e., the tran-  
 168 sient layer lies below the main peak of the ionosphere, and has a larger peak density than the  
 169 main peak [Gurnett *et al.*, 2008; Kopf *et al.*, 2008; Wang *et al.*, 2009]. This is consistent with  
 170 the MAVEN IUVS observations around the same time which observed the Mg<sup>+</sup> peak around  
 171 115 km [Crismani *et al.*, 2018]. The maximum frequency of the transient layer reflection is  
 172  $\sim 3.8$  MHz, which corresponds to an electron density of  $\sim 1.8 \times 10^5 \text{ cm}^{-3}$ , one of the highest  
 173 densities ever observed by the MARSIS instrument. The transient layer reflection appears  
 174 in many ionograms throughout MEX orbits 13,710 and 13,711 which cover  $\sim 15$  hours after  
 175 closest approach [Gurnett *et al.*, 2015; Venkateswara Rao *et al.*, 2016].

176 The transient layer in Figure 2b overlaps the surface trace. The overlap indicates that  
 177 at least some of the transient layer reflections are off-nadir [Duru *et al.*, 2010]. The presence  
 178 of such off-nadir or oblique echoes implies strong horizontal gradients, or patchiness, in the  
 179 ionosphere [Duru *et al.*, 2010; Gurnett *et al.*, 2015]. This patchiness of the transient layer  
 180 is expected, given the localized regions of enhanced Mg<sup>+</sup> observed by MAVEN IUVS (Fig-  
 181 ure 1).

182 We must be cautious when interpreting oblique echoes, as they are sometimes observed  
 183 near the terminator between  $85^\circ$ - $95^\circ$  SZA due to a strong day-to-night gradient in the iono-  
 184 spheric plasma density [Duru *et al.*, 2010]. However, since most of our observations have  
 185 SZA less than  $85^\circ$ , this is unlikely to be a cause of the oblique echoes we observe. Addition-  
 186 ally, oblique echoes and abnormally high peak densities are commonly observed near strong,  
 187 vertical crustal magnetic fields [Gurnett *et al.*, 2008; Fallows *et al.*, 2019]. However, we  
 188 show in Section 3 that all of our observations are obtained at northern mid-latitudes ( $45^\circ\text{N}$ )  
 189 far from any substantial crustal magnetism. Thus, it is unlikely that crustal magnetic fields  
 190 are the primary cause of the oblique echoes we observe.

## 191 2.2.2 Observations more than one day after closest approach

192 Studies of MARSIS ionograms to date have detected the transient layer only up to  
 193 MEX orbit 13,712 ( $\Delta t_{ca} \approx 19$  hours) [Gurnett *et al.*, 2015; Venkateswara Rao *et al.*, 2016].  
 194 To determine if the layer lasted longer, we search for transient layer reflections in dayside  
 195 (solar zenith angle (SZA)  $< 90^\circ$ ) ionograms that were obtained up to a month after closest  
 196 approach.

197 To identify transient layer reflections, we search by eye for ionospheric traces that ei-  
 198 ther overlap the ground trace, or appear similar to the transient reflection shown in 2b. The  
 199 former would indicate an oblique echo as discussed above [Duru *et al.*, 2008] and the lat-  
 200 ter serves as a good benchmark for interpreting distinct ionograms associated with Siding  
 201 Spring. This simple visual analysis is typical for studying MARSIS ionograms [Gurnett  
 202 *et al.*, 2008; Morgan *et al.*, 2008]. We exclude ionograms that are excessively noisy or that  
 203 have obvious features caused by crustal magnetic fields [Gurnett *et al.*, 2008]. We extract  
 204 the peak density from each transient layer using in-house software that identifies the highest  
 205 frequency of a trace, and converts it into an electron density as described in Morgan *et al.*  
 206 [2008]. Eight examples of transient layer reflections we identified are shown in Figures 2c-j.

207 Figures 2c-d show ionograms from 26 October 2014, seven days after closest approach.  
 208 In both examples there is no discernible cusp at the end of the ionospheric trace. Instead,  
 209 the trace extends to high frequencies that overlap the ground trace. The highest frequencies  
 210 of the trace in Figure 2c, 3.5 MHz, correspond to a peak electron density of  $1.5 \times 10^5 \text{ cm}^{-3}$ .  
 211 These peak densities are much larger than expected for the main peak of the ionosphere at a  
 212 SZA near  $77^\circ$ , where the peak density is typically on the order of  $1.0 \times 10^5 \text{ cm}^{-3}$  [Morgan  
 213 *et al.*, 2008; Withers, 2009]. Furthermore, the trace shown in Figure 2d looks similar to the  
 214 transient layer reflection observed shortly after closest approach (Figure 2b), with the time  
 215 delay decreasing with frequency. Note, however, that both of these examples lack a cusp that

216 is expected to be visible at the frequency corresponding to the main peak density as in Fig-  
 217 ure 2a-b. The lack of any cusp may suggest that the transient layer peak has moved upward  
 218 and nearly merged with the main peak of the ionosphere.

219 Figures 2e-f show ionograms from 30 October 2014, eleven days after closest ap-  
 220 proach. Figure 2e has a small plasma feature at 3.4 MHz that overlaps the ground trace. This  
 221 feature is well-separated from the ionospheric trace. Figure 2f shows a plasma feature at the  
 222 high-frequency portion of the ionospheric trace. This feature appears similar to Figure 2d,  
 223 with no discernible cusp and a time delay that decreases with frequency.

224 Ionograms shown in Figures 2g-j demonstrate that these unique trace features are still  
 225 detectable until 21 November 2014, 32 days after Siding Spring’s closest approach. After  
 226 this date, we do not find any more ionograms with convincing features as those explained  
 227 above. In total, we identify 86 ionograms with unusual features such as these. Characteris-  
 228 tics of the 86 ionograms containing these features are listed in Table 1. We interpret these  
 229 unusual features as being detections of the transient layer produced by Siding Spring’s close  
 230 flyby. In the next section, we show that the locations and peak densities of the transient layer  
 231 measurements are consistent with the IUVS observations obtained shortly after closest ap-  
 232 proach.

### 233 3 Transient Layer Locations and Peak Densities

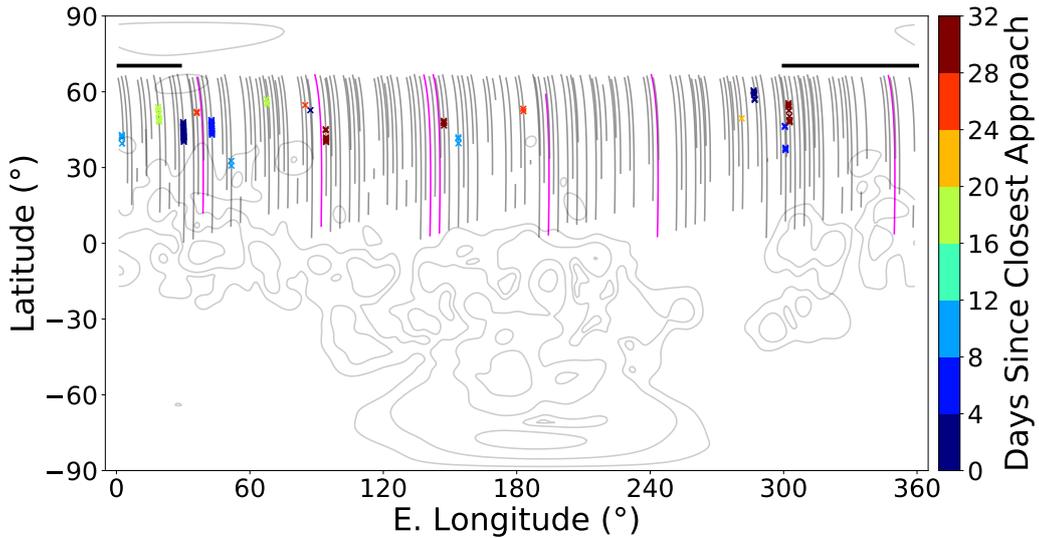
234 Figure 3 shows the latitudes and longitudes of all dayside MARSIS ionograms ob-  
 235 tained between 19 October and 21 November in 2014. Ionograms with transient layer de-  
 236 tections are colored according to  $\Delta t_{ca}$ , the number of days since closest approach. A map  
 237 of the Martian crustal magnetic field strength at 400 km [Morschhauser *et al.*, 2014] is also  
 238 shown. Most of the transient layer detections are located in weak crustal field regions where  
 239 the field strength at 400 km is less than 10 nT (Table 1). Oblique echoes and unusually high  
 240 peak electron densities, are much more commonly observed near radial crustal fields with an  
 241 elevation angle (the angle between the magnetic field and the vertical) of less than 30° [Gur-  
 242 nett *et al.*, 2008]. Of the 86 transient layer detections used in this study, 55 have an elevation  
 243 angle greater than 30° (Table 1), so crustal magnetic fields likely have minor effects on the  
 244 measurements.

257 Although the MARSIS observations cover latitudes between 0°N-65°N, the detections  
 258 are confined to a narrow range between 20°N-60°N. They are also non-uniformly distributed  
 259 in longitude, with only 9 of the 86 transient layer detections between 95°E-275°E. This sug-  
 260 gests the transient layer is mostly confined to one longitudinal hemisphere. Additionally,  
 261 63 of the 86 transient layer detections are between 280°E-45°E, which is near the locations  
 262 where IUVS observed large amounts of Mg<sup>+</sup>. In particular, the localized, bright Mg<sup>+</sup> mea-  
 263 surements in scans 1-2 of orbit 121, and scans 5-6 of orbit 122, nearly coincide with the loca-  
 264 tions of our transient layer detections, even weeks after closest approach.

265 Figure 4 shows peak densities as a function of SZA using the same data. Each panel  
 266 shows a different time range after closest approach. The gray points are the peak densities  
 267 from ionograms that had normal ionospheric traces. The gray shaded regions encompass the  
 268 3 $\sigma$  standard deviations of the peak densities calculated using 2° SZA bins. For comparison,  
 269 the peak densities of the transient layer detections are shown using red crosses.

270 The peak densities from October 19-22 ( $\Delta t_{ca} \approx 0-3$  days) are shown in Figure 4a. The  
 271 transient layer peak densities lie well above the 3 $\sigma$  envelope, indicating they are significantly  
 272 larger than expected for the main peak of the ionosphere. The two groups of elevated tran-  
 273 sient layer peak densities near SZA=77° and SZA=85° are consistent with the transient layer  
 274 peak densities reported in Gurnett *et al.* [2015].

275 The peak density measurements from October 23-25 ( $\Delta t_{ca} \approx 4-6$  days) are shown in  
 276 Figure 4b. The MARSIS coverage during this time period is highlighted in Figure 3 with



245 **Figure 3.** MARSIS observational coverage during 20 October - 21 November. The contours map the  
 246 crustal magnetic field strength at 400 km [Morschhauser *et al.*, 2014]. The straight gray lines show all nom-  
 247 inal MARSIS ionospheric measurements. The crosses mark the locations of transient layer detections we  
 248 identified, colored according to days since closest approach. The pink lines show the coverage between 23-25  
 249 October, when we did not detect the transient layer (Fig. 4b). The black line above the data indicates the lon-  
 250 gitudes at which MAVEN’s IUVS instrument detected localized, bright emissions 1.0-2.5 days after closest  
 251 approach (see Figure 1).

277 pink lines. Despite MEX having similar orbital coverage (Fig. 3), we do not identify any  
 278 transient layer reflections during this period. One orbit in particular, orbit 13,724, even cover-  
 279 ed nearly the same longitudes (E. Lon  $\approx 40^\circ$ ) as orbit 13,710 (E. Lon  $\approx 30^\circ$ ), when MAR-  
 280 SIS readily observed the transient layer shortly after closest approach (Fig. 2b).

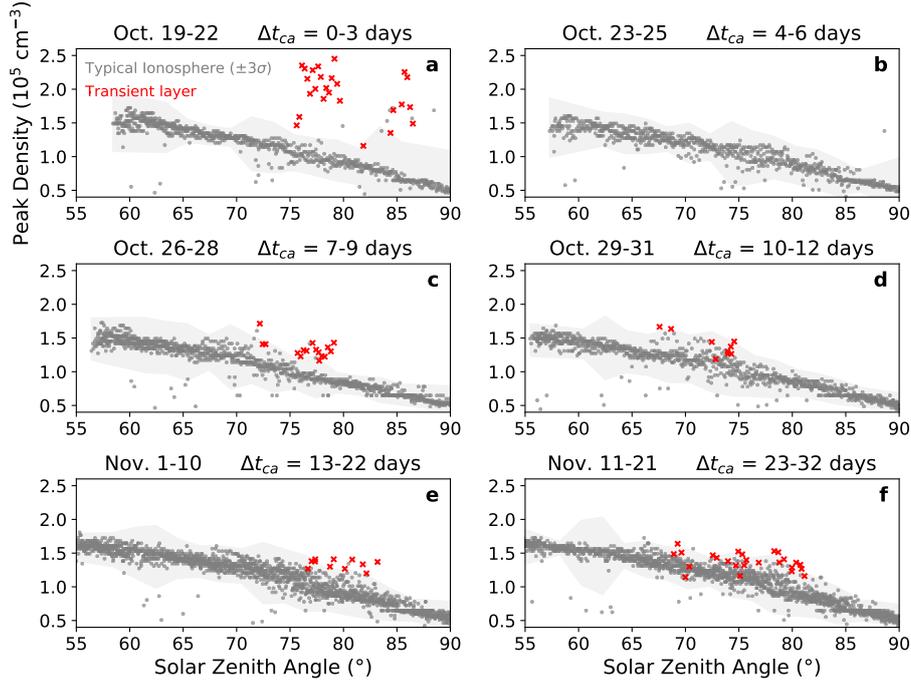
281 We again identify the transient layer beginning on 26 October ( $\Delta t_{ca} \approx 7$  days), as  
 282 shown in Figure 4c. The SZAs of the transient layer measurements are similar to those de-  
 283 tected shortly after closest approach (Fig. 4a). The transient layer peak densities are smaller  
 284 than before, but mostly still lie above the  $3\sigma$  envelope. The smaller peak densities suggest  
 285 that the transient layer peak density decayed over time.

286 The peak densities from 29 October - 21 November ( $\Delta t_{ca} \approx 10$ -32 days) are shown  
 287 in Figures 4d-f. We identify the transient layer during these three periods, at similar SZA  
 288 ranges between  $65^\circ$ - $85^\circ$ . During all three of these periods, the transient layer peak densities  
 289 are less significant outliers than before, suggesting that the transient layer continues to decay.

## 290 4 Discussion

291 Since MARSIS is a topside sounder that is unable to discriminate between different  
 292 ions, we cannot be certain that the unusual features we identified in MARSIS ionograms are  
 293 produced by reflections from the transient  $\text{Mg}^+$  layer. Nevertheless, as described below, sev-  
 294 eral aspects of our transient layer detections are consistent with expectations for this layer.

295 As described in Section 2, the transient layer reflections we identified between 26-28  
 296 October ( $\Delta t_{ca} = 7$ -9 days) have characteristics similar to those reported immediately after  
 297 comet Siding Spring’s closest approach [Gurnett *et al.*, 2015], such as the one shown in Fig-



252 **Figure 4.** Each panel shows dayside peak densities as a function of solar zenith angle for a specified time  
 253 period. The gray points are peak densities derived from normal ionospheric traces and the gray envelopes  
 254 encompass detections the  $3\sigma$  standard deviations in  $2^\circ$  SZA bins. The red crosses are the transient layer peak  
 255 densities. The panel titles display the date ranges of the measurements and the time since comet Siding  
 256 Spring's closest approach ( $\Delta t_{ca}$ ).

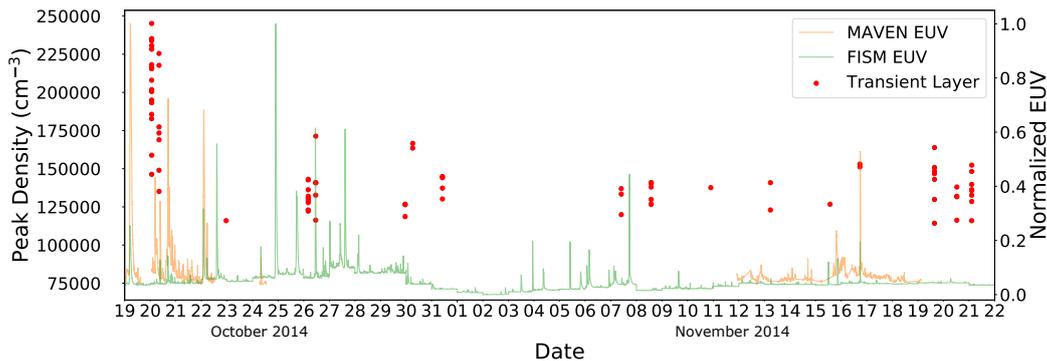
298 ure 2b. These characteristics include an unusual reflection shape, a ground echo overlap, and  
 299 extension to abnormally high frequencies. Most of the transient layer detections are located  
 300 between  $20^\circ\text{N}$ - $60^\circ\text{N}$  latitude and  $275^\circ\text{E}$  to  $95^\circ\text{E}$  longitude, near the MAVEN IUVS observa-  
 301 tions obtained at  $\Delta t_{ca} = 1.0$ - $2.5$  days (Fig. 1). Nearly all of the transient layer peak electron  
 302 densities are exceptionally high ( $>3\sigma$ ) compared to the peak densities derived from typical  
 303 ionospheric traces during the same time period (Fig. 4c). Together, these results make a compelling  
 304 case for the transient layer lasting at least 7 days after closest approach, which is  $\sim 3.5$   
 305 times longer than previously reported.

306 The transient layer reflections we identified between 29 October - 21 November ( $\Delta t_{ca}$   
 307 = 10-32 days) are from similar locations, but most of the peak densities are not  $3\sigma$  outliers  
 308 over the average peak densities of the normal ionosphere (Fig. 4d-f). However, during all  
 309 three periods, the transient layer peak densities are mostly  $1\sigma$  higher. This suggests that the  
 310 transient layer may have lasted up to 32 days after closest approach, and the peak density decayed  
 311 with time. The prolonged lifetime of the transient  $\text{Mg}^+$  layer is perhaps not surprising  
 312 given that metallic ion species can have long chemical lifetimes of a week or longer [Whalley  
 313 and Plane, 2010; Grebowsky et al., 2017]. However, there is currently no comprehensive  
 314 time-dependent model of the Siding Spring event that allows us to compare our observed  
 315 transient layer lifetime with predictions.

316 Despite detecting the transient layer up to 32 days after closest approach, we did not  
 317 detect it continuously; no transient layer could be identified between 4-6 days after closest  
 318 approach (23-25 October). This lack of detection is perplexing because the MARSIS in-

319 strument had similar observational coverage during the entirety of the 32 days we examined.  
 320 One possible explanation is that the transient layer peak moved upwards and merged with the  
 321 main peak of the ionosphere. This would cause the transient layer reflection and main iono-  
 322 spheric trace to be indistinguishable in ionograms. A second possible explanation is that the  
 323 transient layer was so localized in longitude that MARSIS was unable to observe the right lo-  
 324 cation during this period. A third possible explanation is that, by this time, the transient layer  
 325 peak density had decayed to a value smaller than the peak density of the main ionosphere,  
 326 preventing MARSIS from detecting the transient layer.

327 If the third explanation is true and the transient layer did decay between 23-25 October  
 328 ber, then what caused it to come back? The flyby of comet Siding Spring coincidentally took  
 329 place during a time of intense space weather [Sánchez-Cano *et al.*, 2018] when there were  
 330 multiple active flare regions on the Sun. One of these regions, AR 12192, produced multi-  
 331 ple X-class flares directed towards Mars as shown in Figure 5 [Xu *et al.*, 2018; Fang *et al.*,  
 332 2019]. It is known that solar flares increase the X-ray and extreme ultraviolet (EUV) photon  
 333 flux, which leads to increased  $O_2^+$  and electron densities, especially below the main peak of  
 334 the ionosphere [Mendillo *et al.*, 2006; Lollo *et al.*, 2012; Fallows *et al.*, 2015; Xu *et al.*, 2018;  
 335 Thirupathiah *et al.*, 2019]. The increased densities below the main peak could potentially  
 336 lead to ionogram features similar to those we identified as the transient layer. However, these  
 337 false detections are unlikely because, at these altitudes, ionospheric plasma is composed pri-  
 338 marily of  $NO^+$  and  $O_2^+$  [Fox, 2004], which rapidly recombine on timescales of minutes after a  
 339 flare ends [Fang *et al.*, 2019].



340 **Figure 5.** Transient layer detections during 20 October - 22 November, plotted with two solar extreme  
 341 ultraviolet (EUV) irradiance datasets for reference. The irradiances are normalized relative to their maximum  
 342 values during the period considered. The red data points indicate the transient layer detections. Orange data  
 343 are MAVEN Extreme Ultraviolet Monitor measurements in the 0.1-7 nm wavelength range [Eparvier *et al.*,  
 344 2015]. Green data are from the Earth-based Flare Irradiance Spectral Model (FISM) [Chamberlin *et al.*,  
 345 2008]. During this period the Mars-Sun-Earth angle was  $\sim 90^\circ$ .

346 Alternatively, if neutral Mg from the comet was still present in the upper atmosphere,  
 347 then  $Mg^+$  would be produced through charge exchange between Mg and  $O_2^+$  ( $Mg + O_2^+ \rightarrow$   
 348  $Mg^+ + O_2$ ), which is the major chemical pathway for  $Mg^+$  production [Whalley and Plane,  
 349 2010; Plane *et al.*, 2018]. The long lifetime of  $Mg^+$  would allow MARSIS to detect the trans-  
 350 sient layer well after a flare occurred. This hypothesis requires Mg to be present in the upper  
 351 atmosphere long after Siding Spring's closest approach. This may be possible, although cur-  
 352 rent  $Mg^+$  chemistry suggests that the buildup of Mg is limited [Plane *et al.*, 2018]. However,  
 353 without detailed modeling that is out of the scope of this work, it is unclear how long Mg  
 354 was present in the upper atmosphere after closest approach. Thus, this hypothesis is poten-  
 355 tially possible.

356 Another possibility is that the transient layer we detect is created by a different mech-  
357 anism proposed in *Crismani et al.* [2019]. In this mechanism, a sporadic layer is produced  
358 when the interplanetary magnetic field (IMF) is radial, allowing direct access of solar wind  
359 protons to the upper atmosphere. Ionization by precipitating solar wind protons can create a  
360 sporadic ionospheric layer below the main peak of the ionosphere.

361 This mechanism is also not global, but acts to enhance the electron density in a small  
362 region of the planet. Thus, this mechanism would also create localized enhancements of  
363 electron density, similar to what we observe. However, nearly all observations of a sporadic  
364 layer so far have a peak density much smaller than the main peak density and thus could not  
365 be observed by MARSIS [*Pätzold et al.*, 2005; *Withers et al.*, 2008]. Furthermore, it would  
366 be quite a coincidence if this mechanism was causing enhanced ionization in the same loca-  
367 tions as the Siding Spring  $\text{Mg}^+$  layer.

368 A final possibility is that a meteor shower unrelated to Siding Spring happened some  
369 time after closest approach. In general, this seems like a very unlikely possibility given that  
370 no meteor showers at Mars were predicted to occur during the month following closest ap-  
371 proach [*Christou*, 2010; *Vaubailon et al.*, 2014].

## 372 5 Conclusions

373 We have presented MARSIS observations that suggest the transient  $\text{Mg}^+$  layer pro-  
374 duced by comet Siding Spring's encounter with Mars lasted at least 7 days, and potentially  
375 up to 32 days. This is longer than reported by previous studies, which suggested the layer  
376 lasted  $\sim 1$ -2 days. The intense space weather during and after closest approach may have  
377 played a role in the prolonged lifetime of the transient layer. However, global simulations are  
378 needed to fully untangle the mechanisms that caused the transient layer to be localized and  
379 patchy for a prolonged period of time.

Table 1: Characteristics of the 86 transient layer measurements we identified in MARSIS ionograms. The columns list the date, time, MEX orbit number, East longitude (LON, °), latitude (LAT, °), solar zenith angle (SZA, °), peak electron density ( $N_{max}$ ,  $10^5 \text{ cm}^{-3}$ ), the crustal magnetic field strength at 400 km ( $B_{crust}$ , nT), and the angle between the magnetic field and the vertical direction ( $\theta_{elev}$ , degrees).

Date	UT	Orbit	LON	LAT	SZA	$N_{max}$	$B_{crust}$	$\theta_{elev}$
2014-10-20	1:19:32	13710	30.0	48.0	79.7	1.83	14.1	33.4
2014-10-20	1:19:40	13710	30.0	47.5	79.4	2.08	13.8	35.5
2014-10-20	1:19:47	13710	30.1	47.0	79.2	2.45	13.0	34.8
2014-10-20	1:19:55	13710	30.1	46.5	78.9	2.16	11.9	30.6
2014-10-20	1:20:02	13710	30.1	46.0	78.6	1.95	10.7	21.5
2014-10-20	1:20:10	13710	30.1	45.6	78.4	2.02	9.5	7.9
2014-10-20	1:20:17	13710	30.2	45.1	78.1	1.86	8.9	8.6
2014-10-20	1:20:25	13710	30.2	44.6	77.9	2.18	8.9	24.4
2014-10-20	1:20:32	13710	30.2	44.1	77.6	2.34	9.3	34.5
2014-10-20	1:20:40	13710	30.2	43.6	77.4	2.01	10.1	37.5
2014-10-20	1:20:47	13710	30.2	43.1	77.1	2.28	10.9	35.6
2014-10-20	1:20:55	13710	30.3	42.6	76.9	1.93	11.7	31.8
2014-10-20	1:21:03	13710	30.3	42.2	76.6	2.15	12.9	28.2
2014-10-20	1:21:10	13710	30.3	41.7	76.4	2.31	14.4	25.9
2014-10-20	1:21:18	13710	30.3	41.2	76.1	2.35	16.2	25.7
2014-10-20	1:21:25	13710	30.3	40.7	75.9	1.59	18.1	26.6
2014-10-20	1:21:33	13710	30.3	40.2	75.6	1.46	19.8	28.6
2014-10-20	8:15:50	13711	286.5	60.6	86.5	1.49	5.5	26.3
2014-10-20	8:15:57	13711	286.6	60.1	86.2	1.73	5.2	19.5
2014-10-20	8:16:05	13711	286.7	59.6	86.0	2.18	5.0	12.3
2014-10-20	8:16:13	13711	286.7	59.2	85.7	2.25	4.7	4.6
2014-10-20	8:16:20	13711	286.8	58.7	85.4	1.77	4.5	3.9
2014-10-20	8:16:43	13711	287.0	57.3	84.7	1.69	4.2	29.4
2014-10-20	8:16:50	13711	287.0	56.8	84.4	1.35	4.2	37.3
2014-10-22	23:13:39	13720	87.2	52.7	81.9	1.16	5.2	22.1
2014-10-26	4:09:42	13731	42.7	48.8	79.1	1.43	36.4	68.3
2014-10-26	4:09:50	13731	42.7	48.3	78.8	1.3	39.1	72.8
2014-10-26	4:09:57	13731	42.8	47.8	78.5	1.36	41.1	77.9
2014-10-26	4:10:05	13731	42.8	47.4	78.2	1.22	41.8	82.3
2014-10-26	4:10:12	13731	42.8	46.9	78.0	1.23	41.1	82.4
2014-10-26	4:10:20	13731	42.8	46.4	77.7	1.29	38.9	77.3
2014-10-26	4:10:35	13731	42.9	45.4	77.1	1.43	31.5	63.1
2014-10-26	4:10:50	13731	42.9	44.4	76.6	1.31	23.0	48.1
2014-10-26	4:10:58	13731	42.9	44.0	76.3	1.32	19.3	40.8
2014-10-26	4:11:05	13731	43.0	43.5	76.0	1.23	16.3	34.2
2014-10-26	4:11:13	13731	43.0	43.0	75.7	1.28	13.9	28.6
2014-10-26	11:09:53	13732	300.6	46.6	77.7	1.16	4.6	39.3
2014-10-26	11:10:01	13732	300.6	46.1	77.4	1.33	5.2	43.4
2014-10-26	11:12:09	13732	300.9	37.9	72.7	1.41	9.2	54.9
2014-10-26	11:12:17	13732	300.9	37.4	72.4	1.41	10.1	56.4
2014-10-26	11:12:24	13732	300.9	36.9	72.2	1.71	11.4	58.3
2014-10-29	23:05:41	13744	153.7	42.0	74.3	1.26	2.7	78.8
2014-10-29	23:05:48	13744	153.7	41.5	74.0	1.27	2.7	80.4
2014-10-29	23:06:18	13744	153.8	39.6	72.8	1.19	2.8	81.2
2014-10-30	6:07:40	13745	51.7	32.6	68.6	1.63	37.3	31.3
2014-10-30	6:08:10	13745	51.7	30.8	67.6	1.67	50.4	28.0
2014-10-31	10:03:06	13749	2.4	42.9	74.6	1.45	31.9	1.9
2014-10-31	10:03:13	13749	2.5	42.4	74.2	1.37	36.2	8.6

Table 1: Characteristics of the 86 transient layer measurements we identified in MARSIS ionograms. The columns list the date, time, MEX orbit number, East longitude (LON, °), latitude (LAT, °), solar zenith angle (SZA, °), peak electron density ( $N_{max}$ ,  $10^5 \text{ cm}^{-3}$ ), the crustal magnetic field strength at 400 km ( $B_{crust}$ , nT), and the angle between the magnetic field and the vertical direction ( $\theta_{elev}$ , degrees).

Date	UT	Orbit	LON	LAT	SZA	$N_{max}$	$B_{crust}$	$\theta_{elev}$
2014-10-31	10:03:21	13749	2.5	41.9	74.0	1.3	41.1	15.9
2014-10-31	10:03:58	13749	2.6	39.5	72.5	1.44	65.6	54.6
2014-11-07	9:48:23	13773	67.5	57.1	83.2	1.37	15.0	24.8
2014-11-07	9:48:45	13773	67.6	55.6	82.2	1.2	17.0	43.0
2014-11-07	9:48:53	13773	67.7	55.2	81.8	1.33	17.9	48.1
2014-11-08	13:47:19	13777	18.8	53.9	80.8	1.4	8.4	10.3
2014-11-08	13:47:34	13777	18.9	53.0	80.1	1.27	9.5	12.8
2014-11-08	13:47:57	13777	19.0	51.5	79.1	1.41	10.4	1.9
2014-11-08	13:48:05	13777	19.0	51.0	78.7	1.3	10.5	10.8
2014-11-08	13:48:35	13777	19.1	49.1	77.4	1.41	11.7	51.7
2014-11-08	13:48:42	13777	19.2	48.6	77.0	1.38	12.1	58.8
2014-11-08	13:48:50	13777	19.2	48.2	76.7	1.27	12.4	63.4
2014-11-10	21:44:48	13785	281.1	49.5	77.3	1.38	2.6	43.9
2014-11-13	5:39:47	13793	182.9	53.3	80.0	1.23	10.1	9.4
2014-11-13	5:40:02	13793	183.0	52.3	79.2	1.41	12.3	23.2
2014-11-15	13:35:26	13801	84.9	54.7	80.9	1.27	8.1	55.4
2014-11-16	17:34:07	13805	36.1	52.1	78.7	1.51	37.3	72.1
2014-11-16	17:34:15	13805	36.2	51.6	78.3	1.53	39.6	76.8
2014-11-19	15:31:05	13815	94.1	45.2	72.9	1.43	3.7	66.4
2014-11-19	15:31:13	13815	94.1	44.7	72.6	1.47	3.7	68.7
2014-11-19	15:31:58	13815	94.2	41.9	70.4	1.3	3.8	74.6
2014-11-19	15:32:05	13815	94.2	41.5	70.0	1.14	3.7	74.4
2014-11-19	15:32:13	13815	94.2	41.0	69.6	1.51	3.7	74.1
2014-11-19	15:32:20	13815	94.3	40.6	69.3	1.64	3.6	73.7
2014-11-19	15:32:28	13815	94.3	40.1	68.9	1.49	3.5	73.0
2014-11-20	12:28:45	13818	147.2	48.5	75.5	1.32	2.4	35.1
2014-11-20	12:28:52	13818	147.2	48.0	75.1	1.16	2.2	38.7
2014-11-20	12:29:00	13818	147.2	47.6	74.7	1.32	2.1	43.0
2014-11-20	12:29:15	13818	147.3	46.6	74.0	1.38	2.0	53.3
2014-11-21	2:25:57	13820	302.2	55.5	81.1	1.16	2.2	28.9
2014-11-21	2:26:04	13820	302.2	55.0	80.8	1.33	2.4	36.4
2014-11-21	2:26:12	13820	302.3	54.5	80.4	1.36	2.5	41.4
2014-11-21	2:26:19	13820	302.3	54.1	80.0	1.28	2.6	44.0
2014-11-21	2:26:42	13820	302.4	52.6	78.8	1.36	2.4	38.5
2014-11-21	2:27:20	13820	302.6	50.3	76.9	1.36	1.5	28.9
2014-11-21	2:27:42	13820	302.7	48.8	75.7	1.4	3.2	37.9
2014-11-21	2:27:50	13820	302.7	48.4	75.3	1.48	4.0	38.5
2014-11-21	2:27:57	13820	302.8	47.9	74.9	1.52	4.8	39.6

## Acknowledgments

All the MARSIS data used in this study are publicly available at <https://space.physics.uiowa.edu/marsx/public-data.html>. MAVEN EUV data can be found at <https://lasp.colorado.edu/maven/sdc/public/pages/datasets/euv.html>. MAVEN FISM EUV data, produced by the Laboratory for Atmospheric and Space Physics (LASP) can be found at <http://lasp.colorado.edu/lisird/>. The research at the University of

386 Iowa was supported by NASA through contract 1224107 with the Jet Propulsion Laboratory.  
 387 We thank the many members of the scientific, technical, and management teams at NASA  
 388 Headquarters, the Jet Propulsion Laboratory, the European Space Agency, and the Italian  
 389 Space Agency for their efforts in planning the spacecraft operations required to successfully  
 390 obtain these data. We thank Bill Kurth, Hadi Madanian, John Plane, and Matteo Crismani  
 391 for their insightful discussions on the topic. We give special thanks to Chris Piker for helping  
 392 with data acquisition using the das2py software package (<https://das2.org/das2py/>)  
 393 and to Larry Granroth who provided useful programs for analyzing MARSIS data. Addi-  
 394 tionally, we thank Thomas Koeppl for his dedicated administrative work during Z. Luppen's  
 395 time at the University of Iowa.

## 396 References

- 397 Benna, M., P. R. Mahaffy, J. M. Grebowsky, J. M. C. Plane, R. V. Yelle, and B. M. Jakosky  
 398 (2015), Metallic ions in the upper atmosphere of Mars from the passage of comet C/2013  
 399 A1 (Siding Spring), *Geophys. Res. Lett.*, *42*, 4670–4675, doi:10.1002/2015GL064159.
- 400 Chamberlin, P. C., T. N. Woods, and F. G. Eparvier (2008), Flare Irradiance Spectral  
 401 Model (FISM): Flare component algorithms and results, *Space Weather*, *6*, S05001,  
 402 10.1029/2007SW000372, doi:10.1029/2007SW000372.
- 403 Christou, A. A. (2010), Annual meteor showers at venus and mars: lessons from the earth,  
 404 *Mon. Not. R. Astron. Soc.*, *402*(4), 2759–2770, doi:10.1111/j.1365-2966.2009.16097.x.
- 405 Crismani, M. M. J., N. M. Schneider, J. M. C. Plane, J. S. Evans, S. K. Jain, M. S. Chaf-  
 406 fin, J. D. Carrillo-Sanchez, J. I. Deighan, R. V. Yelle, A. I. F. Stewart, W. McClintock,  
 407 J. Clarke, G. M. Holsclaw, A. Stiepen, F. Montmessin, and B. M. Jakosky (2017), De-  
 408 tection of a persistent meteoric metal layer in the Martian atmosphere, *Nat. Geosci.*, *10*(6),  
 409 401–404, doi:10.1038/ngeo2958.
- 410 Crismani, M. M. J., N. M. Schneider, J. S. Evans, J. M. C. Plane, J. D. Carrillo-Sánchez,  
 411 S. Jain, J. Deighan, and R. Yelle (2018), The Impact of Comet Siding Spring's Meteors  
 412 on the Martian Atmosphere and Ionosphere, *J. Geophys. Res.*, *123*(10), 2613–2627, doi:  
 413 10.1029/2018JE005750.
- 414 Crismani, M. M. J., J. Deighan, N. M. Schneider, J. M. C. Plane, P. Withers, J. Halekas,  
 415 M. Chaffin, and S. Jain (2019), Localized ionization hypothesis for transient ionospheric  
 416 layers, *J. Geophys. Res.*, *124*(6), 4870–4880, doi:10.1029/2018JA026251.
- 417 Duru, F., D. A. Gurnett, D. D. Morgan, R. Modolo, A. F. Nagy, and D. Najib (2008), Elec-  
 418 tron densities in the upper ionosphere of mars from the excitation of electron plasma oscil-  
 419 lations, *J. Geophys. Res.*, *113*, A07302, doi:10.1029/2008JA013073.
- 420 Duru, F., D. D. Morgan, and D. A. Gurnett (2010), Overlapping ionospheric and surface  
 421 echoes observed by the Mars Express radar sounder near the Martian terminator, *Geophys.*  
 422 *Res. Lett.*, *37*(23), L23102, doi:10.1029/2010GL045859.
- 423 Eparvier, F. G., P. C. Chamberlin, T. N. Woods, and E. M. B. Thiemann (2015), The Solar  
 424 Extreme Ultraviolet Monitor for MAVEN, *Space Sci. Rev.*, *195*, 293–301, doi:10.1007/  
 425 s11214-015-0195-2.
- 426 Espley, J. R., G. A. DiBraccio, J. E. P. Connerney, D. Brain, J. Gruesbeck, Y. Soobiah,  
 427 J. Halekas, M. Combi, J. Luhmann, Y. Ma, Y. Jia, and B. Jakosky (2015), A comet engulfs  
 428 mars: Maven observations of comet siding spring's influence on the martian magneto-  
 429 sphere, *Geophys. Res. Lett.*, *42*(21), 8810–8818, doi:10.1002/2015GL066300.
- 430 Fallows, K., P. Withers, and G. Gonzalez (2015), Response of the Mars ionosphere to solar  
 431 flares: Analysis of MGS radio occultation data, *J. Geophys. Res.*, *120*(11), 9805–9825,  
 432 doi:10.1002/2015JA021108.
- 433 Fallows, K., P. Withers, D. Morgan, and A. Kopf (2019), Extremely high plasma densities in  
 434 the mars ionosphere associated with cusp-like magnetic fields, *J. Geophys. Res.*, *124*(7),  
 435 6029–6046, doi:10.1029/2019JA026690.
- 436 Fang, X., D. Pawlowski, Y. Ma, S. Bougher, E. Thiemann, F. Eparvier, W. Wang, C. Dong,  
 437 C. O. Lee, Y. Dong, M. Benna, M. Elrod, P. Chamberlin, P. Mahaffy, and B. Jakosky

- 438 (2019), Mars Upper Atmospheric Responses to the 10 September 2017 Solar Flare:  
439 A Global, Time-Dependent Simulation, *Geophys. Res. Lett.*, *46*(16), 9334–9343, doi:  
440 10.1029/2019GL084515.
- 441 Fox, J. L. (2004), Response of the martian thermosphere/ionosphere to enhanced fluxes of  
442 solar soft x rays, *J. Geophys. Res.*, *109*, A11310, doi:10.1029/2004JA010380.
- 443 Grebowsky, J. M., M. Benna, J. M. C. Plane, G. A. Collinson, P. R. Mahaffy, and B. M.  
444 Jakosky (2017), Unique, non-earthlike, meteoritic ion behavior in upper atmosphere of  
445 mars, *Geophys. Res. Lett.*, *44*(7), 3066–3072, doi:10.1002/2017GL072635.
- 446 Gurnett, D. A., D. L. Kirchner, R. L. Huff, D. D. Morgan, A. M. Persoon, T. F. Averkamp,  
447 F. Duru, E. Nielsen, A. Safaeinili, J. J. Plaut, and G. Picardi (2005), Radar soundings of  
448 the ionosphere of Mars, *Science*, *310*, 1929–1933, doi:10.1126/science.1121868.
- 449 Gurnett, D. A., R. L. Huff, D. D. Morgan, A. M. Persoon, T. F. Averkamp, D. L. Kirchner,  
450 F. Duru, F. Akalin, A. J. Kopf, E. Nielsen, A. Safaeinili, J. J. Plaut, and G. Picardi (2008),  
451 An overview of radar soundings of the martian ionosphere from the Mars Express space-  
452 craft, *Adv. Space Res.*, *41*, 1335–1346, doi:10.1016/j.asr.2007.01.062.
- 453 Gurnett, D. A., D. D. Morgan, A. M. Persoon, L. J. Granroth, A. J. Kopf, J. J. Plaut, and  
454 J. L. Green (2015), An ionized layer in the upper atmosphere of Mars caused by dust im-  
455 pacts from comet Siding Spring, *Geophys. Res. Lett.*, *42*(12), 4745–4751, doi:10.1002/  
456 2015GL063726.
- 457 Kopf, A. J., D. A. Gurnett, D. D. Morgan, and D. L. Kirchner (2008), Transient layers in the  
458 topside ionosphere of mars, *Geophys. Res. Lett.*, *35*, L17102, doi:10.1029/2008GL034948.
- 459 Lollo, A., P. Withers, K. Fallows, Z. Girazian, M. Matta, and P. C. Chamberlin (2012), Nu-  
460 merical simulations of the ionosphere of Mars during a solar flare, *J. Geophys. Res.*,  
461 *117*(A5), doi:10.1029/2011JA017399.
- 462 Mendillo, M., P. Withers, D. Hinson, H. Rishbeth, and B. Reinisch (2006), Effects of Solar  
463 Flares on the Ionosphere of Mars, *Science*, *311*(5764), 1135–1138, doi:10.1126/science.  
464 1122099.
- 465 Morgan, D. D., D. A. Gurnett, D. L. Kirchner, J. L. Fox, E. Nielsen, and J. J. Plaut (2008),  
466 Variation of the martian ionospheric electron density from Mars Express radar soundings,  
467 *J. Geophys. Res.*, *113*, A09303, doi:10.1029/2008JA013313.
- 468 Morschhauser, A., V. Lesur, and M. Grott (2014), A spherical harmonic model of the  
469 lithospheric magnetic field of Mars, *J. Geophys. Res.*, *119*, 1162–1188, doi:10.1002/  
470 2013JE004555.
- 471 Pätzold, M., S. Tellmann, B. Häusler, D. Hinson, R. Schaa, and G. L. Tyler (2005), A spo-  
472 radic third layer in the ionosphere of Mars, *Science*, *310*, 837–839, doi:10.1126/science.  
473 1117755.
- 474 Picardi, G., D. Biccari, R. Seu, J. Plaut, W. T. K. Johnson, R. L. Jordan, A. Safaeinili, D. A.  
475 Gurnett, R. Huff, R. Orosei, O. Bombaci, D. Calabrese, and E. Zampolini (2004), *MAR-*  
476 *SIS: Mars Advanced Radar for Subsurface and Ionosphere Sounding*, pp. 51–69, ESA SP-  
477 1240: Mars Express: the Scientific Payload, available online at <http://sci.esa.int/science->  
478 [e/www/object/index.cfm?fobjectid=34885](http://sci.esa.int/science-e/www/object/index.cfm?fobjectid=34885).
- 479 Plane, J. M. C., J. D. Carrillo-Sanchez, T. P. Mangan, M. M. J. Crismani, N. M. Schneider,  
480 and A. Määttänen (2018), Meteoric metal chemistry in the martian atmosphere, *Journal of*  
481 *Geophysical Research: Planets*, *123*(3), 695–707, doi:10.1002/2017JE005510.
- 482 Restano, M., J. J. Plaut, B. A. Campbell, Y. Gim, D. Nunes, F. Bernardini, A. Egan, R. Seu,  
483 and R. J. Phillips (2015), Effects of the passage of Comet C/2013 A1 (Siding Spring) ob-  
484 served by the Shallow Radar (SHARAD) on Mars Reconnaissance Orbiter, *Geophys. Res.*  
485 *Lett.*, *42*(12), 4663–4669, doi:10.1002/2015GL064150.
- 486 Sánchez-Cano, B., O. Witasse, M. Lester, A. Rahmati, R. Ambrosi, R. Lillis, F. Leblanc, P.-  
487 L. Blelly, M. Costa, S. W. H. Cowley, J. R. Espley, S. E. Milan, J. J. Plaut, C. Lee, and  
488 D. Larson (2018), Energetic particle showers over mars from comet c/2013 a1 siding  
489 spring, *J. Geophys. Res.*, *123*(10), 8778–8796, doi:10.1029/2018JA025454.
- 490 Sánchez-Cano, B., O. Witasse, M. Lester, A. Rahmati, R. Ambrosi, R. Lillis, F. Leblanc, P.-  
491 L. Blelly, M. Costa, S. W. H. Cowley, J. R. Espley, S. E. Milan, J. J. Plaut, C. Lee, and

- 492 D. Larson (2018), Energetic Particle Showers Over Mars from Comet C/2013 A1 Siding  
493 Spring, *J. Geophys. Res.*, *123*(10), 8778–8796, doi:10.1029/2018JA025454.
- 494 Sánchez-Cano, B., M. Lester, O. Witasse, D. D. Morgan, H. Opgenoorth, D. J. Andrews, P.-  
495 L. Blelly, S. W. H. Cowley, A. J. Kopf, F. Leblanc, J. R. Espley, and A. Cardesín-Moinelo  
496 (2020), Mars' ionospheric interaction with comet c/2013 a1 siding spring's coma at their  
497 closest approach as seen by mars express, *J. Geophys. Res.*, *125*(1), e2019JA027,344, doi:  
498 10.1029/2019JA027344.
- 499 Schneider, N. M., J. I. Deighan, A. I. F. Stewart, W. E. McClintock, S. K. Jain, M. S. Chaf-  
500 fin, A. Stiepen, M. Crismani, J. M. C. Plane, J. D. Carrillo-Sánchez, J. S. Evans, M. H.  
501 Stevens, R. V. Yelle, J. T. Clarke, G. M. Holsclaw, F. Montmessin, and B. M. Jakosky  
502 (2015), MAVEN IUVS observations of the aftermath of the Comet Siding Spring meteor  
503 shower on Mars, *Geophys. Res. Lett.*, *42*(12), 4755–4761, doi:10.1002/2015GL063863.
- 504 Thirupathaiah, P., S. Y. Shah, and S. Haider (2019), Characteristics of solar x-ray flares and  
505 their effects on the ionosphere and human exploration to mars: Mgs radio science observa-  
506 tions, *Icarus*, *330*, 60 – 74, doi:https://doi.org/10.1016/j.icarus.2019.04.015.
- 507 Vaubaillon, J., L. Maquet, and R. Soja (2014), Meteor hurricane at Mars on 2014 October  
508 19 from comet C/2013 A1, *Mon. Not. R. Astron. Soc.*, *439*(4), 3294–3299, doi:10.1093/  
509 mnras/stu160.
- 510 Venkateswara Rao, N., P. ManasaMohana, A. Jayaraman, and S. V. B. Rao (2016), Some  
511 new aspects of the transient ionization layer of comet Siding Spring origin in the Martian  
512 upper atmosphere, *J. Geophys. Res.*, *121*(4), 3592–3602, doi:10.1002/2015JA022189.
- 513 Wang, X. D., J. S. Wang, E. Nielsen, and H. Zou (2009), “Hook” structure in MARSIS  
514 ionogram and its interpretation, *Geophys. Res. Lett.*, *36*(13), L13103, doi:10.1029/  
515 2009GL038844.
- 516 Whalley, C. L., and J. M. C. Plane (2010), Meteoric ion layers in the Martian atmosphere,  
517 *Faraday Discuss.*, *147*, 349, doi:10.1039/c003726e.
- 518 Witasse, O., B. Sánchez-Cano, M. L. Mays, P. KajdiÄN, H. Opgenoorth, H. A. Elliott, I. G.  
519 Richardson, I. Zouganelis, J. Zender, R. F. Wimmer-Schweingruber, L. Turc, M. G. G. T.  
520 Taylor, E. Roussos, A. Rouillard, I. Richter, J. D. Richardson, R. Ramstad, G. Provan,  
521 A. Posner, J. J. Plaut, D. Odstrcil, H. Nilsson, P. Niemenen, S. E. Milan, K. Mandt,  
522 H. Lohf, M. Lester, J.-P. Lebreton, E. Kuulkers, N. Krupp, C. Koenders, M. K. James,  
523 D. Intzekara, M. Holmstrom, D. M. Hassler, B. E. S. Hall, J. Guo, R. Goldstein, C. Goetz,  
524 K. H. Glassmeier, V. GÄhnot, H. Evans, J. Espley, N. J. T. Edberg, M. Dougherty, S. W. H.  
525 Cowley, J. Burch, E. Behar, S. Barabash, D. J. Andrews, and N. Altobelli (2017), Inter-  
526 planetary coronal mass ejection observed at STEREO-A, Mars, comet 67P/Churyumov-  
527 Gerasimenko, Saturn, and New Horizons en route to Pluto: Comparison of its Forbush  
528 decreases at 1.4, 3.1, and 9.9 AU, *J. Geophys. Res.*, *122*(8), 7865–7890, doi:10.1002/  
529 2017JA023884.
- 530 Withers, P. (2009), A review of observed variability in the dayside ionosphere of Mars, *Adv.*  
531 *Space Res.*, *44*, 277–307, doi:10.1016/j.asr.2009.04.027.
- 532 Withers, P., M. Mendillo, D. P. Hinson, and K. Cahoy (2008), Physical characteristics and  
533 occurrence rates of meteoric plasma layers detected in the martian ionosphere by the mars  
534 global surveyor radio science experiment, *J. Geophys. Res.*, *113*, A12314, doi:10.1029/  
535 2008JA013636.
- 536 Xu, S., E. Thiemann, D. Mitchell, F. Eparvier, D. Pawlowski, M. Benna, L. Andersson,  
537 M. W. Liemohn, S. Bougher, and C. Mazelle (2018), Observations and Modeling of the  
538 Mars Low-Altitude Ionospheric Response to the 10 September 2017 X-Class Solar Flare,  
539 *Geophys. Res. Lett.*, *45*(15), 7382–7390, doi:10.1029/2018GL078524.