

1 **Inherent length scales of periodic mesoscale density**
2 **structures in the solar wind over two solar cycles**

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6 **Key Points:**

- 7 • 25 years of Wind solar wind data are analyzed for periodic mesoscale structures
8 in the proton density
- 9 • Periodic density structures recur with particular length scales, suggesting solar for-
10 mation
- 11 • The observed length scales are solar cycle dependent

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Abstract

It is now well-established through multiple event and statistical studies that the solar wind at 1 AU contains periodic, mesoscale ($L \sim 100 - 1000$ Mm) structures in the proton density. Composition variations observed at 1 AU within periodic density structures and remote sensing observations of similar structures in the young solar wind indicate that at least some of these periodic structures originate in the solar atmosphere as a part of solar wind formation. Viall et al. (2008) analyzed 11 years of data from the Wind spacecraft near L1 and demonstrated a recurrence to the length scales of periodic structures in the solar wind proton density observed, and a potential solar cycle dependence. In the time since that study, Wind has collected 14 additional years of solar wind data, allowing a more thorough examination of the dependence of these structures as a function of solar cycle. In addition, the Wind plasma data have been reprocessed in the interim, and new methods for spectral background approximation have been developed, allowing a reevaluation of the precision and accuracy of the initial study. In this study, we analyze 25 years of Wind data collected near L1, and produce occurrence distributions of statistically significant periodic length scales in proton density. The results confirm the Viall et al. (2008) study and significantly extend those results to show a solar cycle dependence of the length scales, and a possible relation to solar “termination” events.

Plain Language Summary

The plasma and magnetic field in the solar atmosphere flows away from the Sun, filling interplanetary space. This plasma is called the solar wind, and it constantly bombards all of the planets in the solar system. The solar wind is comprised of mesoscale structures - larger than scales where particle dynamics are important, but smaller than global scales - of increased density, and therefore pressure. They are of order the size of Earth’s magnetosphere, and often quasi-periodic. These periodic density structures are an important driver of dynamics in Earth’s space environment. In this study, we examine the statistics of the size scales of these structures using 25 years, or approximately two solar cycles, of Wind spacecraft solar wind data. We confirm earlier work showing a persistence of particular length scales to the periodicities, and find that the periodicities are a function of solar cycle. In addition to their driving of magnetospheric dynamics, periodic density structures are a tracer of solar wind formation. Their lengths scales and evolution are an important constraint of solar wind formation.

1 Introduction

The solar wind often contains intervals of mesoscale ($L \sim 100\text{--}1000$ Mm), quasi-periodic proton density enhancements, termed periodic density structures (PDSs). They were initially discovered through event studies that showed a direct correspondence between magnetospheric pulsations in the mHz range and a one-to-one correlation with discrete frequencies in the solar wind density observed in the upstream solar wind (Kepko et al., 2002; Kepko & Spence, 2003). Numerous event studies have observed similar direct links between the periodicities in solar wind density and periodicities in radar (Stephenson & Walker, 2002; Fenrich & Waters, 2008), ionospheric (Dyrud et al., 2008), and ground magnetometer (Villante et al., 2007; Villante & Tiberi, 2016) observations.

At frequencies $< \sim 4$ mHz the magnetosphere is generally incapable of supporting standing oscillations, such as via cavity mode or fieldline resonances (e.g., Hartinger et al. (2013)). Meanwhile, PDSs have been observed to directly drive magnetospheric pulsations from ~ 4 mHz down to ~ 0.2 mHz. Hence, there is a general split between directly driven oscillations at $f < 4$ mHz, and internally supported oscillations at around $f > 4$ mHz. This divide straddles the traditional Pc5 range of 1.7-6.7 mHz. We further note that solar wind driven pulsations span the Pc5 range but extend to lower frequencies, into the $f < 1.7$ mHz ($T > 600$ s) Pc6 range (Saito, 1978). Pc6 oscillations are rarely included in magnetospheric ULF waves studies, yet these lower frequency solar wind driven oscillations may have important consequences for magnetospheric dynamics (Kepko & Viall, 2019).

While there have been many statistical studies of magnetospheric pulsations in the few mHz range (Mathie et al., 1999; Francia et al., 2005; Takahashi & Ukhorskiy, 2007), to date only two studies have examined the statistical occurrence rate of the solar wind periodic density structures that drive them. Viall, Kepko, and Spence (2009) identified statistically significant frequencies observed in 11 years of Wind proton density data near L1 and 10 years of dayside GOES magnetospheric B_z data. They showed that both the solar wind and dayside magnetosphere contained recurrent, similar sets of observed frequencies. The apparent frequency of a periodic density structures as it flows past Earth and an in situ spacecraft is related to the radial length scale of the structure as $f_{pds} = V_{sw}/L_{pds}$. Viall et al. (2008) used the same solar wind data to demonstrate that L_{pds} are on the order of the dayside magnetosphere and larger. They therefore quasi-statically

76 drive magnetospheric pulsations at a frequency f_{pds} . At the typically observed length
 77 scales of ~ 80 to several 100 Mm, this equates to oscillations of a few mHz for nomi-
 78 nal solar wind conditions, which are observed as Pc5-6 pulsations in Earth’s magneto-
 79 sphere.

80 Since the initial papers describing the existence of periodic density structures in
 81 the solar wind, there have been several attempts to identify their source. A key measure-
 82 ment are the occurrence distributions of statistically significant frequencies and length
 83 scales measured by Viall et al. (2008) and Viall, Kepko, and Spence (2009). These dis-
 84 tributions of statistically significant spectral peaks in time series of solar wind density
 85 consists of 3 sources: in situ generated structures (e.g., via turbulence); ‘false positives’
 86 at a rate determined by the chosen confidence thresholds and appropriateness of the back-
 87 ground spectral fit; and periodic density structures injected through the process of so-
 88 lar wind formation. The first two of these sources would generate a smoothly varying
 89 distribution of observed periodicities, while the third could produce localized occurrence
 90 distribution peaks. While it is possible within any segment of solar wind to generate a
 91 single occurrence of a PDS by turbulent processes during transit from the Sun to L1, such
 92 turbulence-driven generation of structures would show a smooth occurrence distribution
 93 of found length scales and frequencies, rather than the recurrent sets of enhancements
 94 found by Viall et al. (2008) and Viall, Kepko, and Spence (2009). Although it is theo-
 95 retically possible that there exists an MHD instability that could generate periodic struc-
 96 tures in transit to 1 AU, for example a slow mode wave (Hollweg et al., 2014), to date
 97 there has been no published observations of such instabilities on mesoscales. Further-
 98 more, Viall, Spence, and Kasper (2009) found a lower occurrence rate of recurrent so-
 99 lar wind periodicities analyzing frequencies than Viall et al. (2008) did analyzing length
 100 scales. This suggests advecting structures, rather than locally generated oscillations or
 101 waves at particular frequencies.

102 Multiple lines of evidence suggest that periodic solar wind density structures are
 103 tracers of solar wind formation. In situ observations show composition, magnetic field,
 104 and electron strahl changes that indicate magnetic reconnection effects that could only
 105 have occurred during solar wind release and acceleration (Viall, Spence, & Kasper, 2009;
 106 Kepko et al., 2016; Matteo et al., 2019). Matteo et al. (2019), using Helios data, found
 107 anisotropic temperature changes within PDSs that are not observed near L1, consistent
 108 with solar formation followed by temperature isotropization while in transit. Remote imag-

109 ing studies using the Solar Terrestrial Relations Observatory (STEREO)/Sun Earth Con-
 110 nection Coronal and Heliospheric Investigation (SECCHI) white light instruments have
 111 identified periodic density structures in the solar corona as close as 2.5 solar radii, ob-
 112 served as they accelerate with the surrounding solar wind (Viall et al., 2010; Viall & Vourl-
 113 idas, 2015; DeForest et al., 2016, 2018). In short, it is now clear that the solar wind is
 114 often formed of quasi-periodic mesoscale plasma density structures released as a part of
 115 solar wind formation.

116 Three factors motivate this investigation. First, while previous studies used only
 117 11 years of data, 25 years of Wind solar wind data are now available, which allows an
 118 examination of evolution of the recurrent length scales as a function of two complete so-
 119 lar cycles. Second, the Wind SWE data that the Viall et al. (2008) and Viall, Kepko,
 120 and Spence (2009) statistical studies analyzed has recently been reprocessed, providing
 121 an opportunity to test the accuracy and precision of those previous results. Third, re-
 122 cent progress on techniques used to identify statistically significant spectral peaks has
 123 shown that there are limitations to using the AR(1) background assumption, and sug-
 124 gests a different background assumption may be more appropriate (Vaughan et al., 2011).

125 2 Methods

126 2.1 Data processing and quality checks

127 We follow the general process of data preparation and spectral analysis as the Viall
 128 et al. (2008) study. We used the proton number density and proton velocity measured
 129 by the Solar Wind Experiment (SWE) onboard the Wind spacecraft (Ogilvie et al., 1995)
 130 to examine the characteristics of mesoscale periodic density structures observed over the
 131 full lifetime of Wind to this point, from 1995-2019. In the time since the Viall et al. (2008)
 132 study, Wind SWE data have been reprocessed, leading to more accurate measurements
 133 of the proton number density and velocity (Kasper et al., 2006). The primary impact
 134 of that reprocessing on this study is that the velocity increased on average by a few per-
 135 cent, which increases the length-scales by a few percent.

136 For continuity with and comparison to the Viall et al. (2008) study, we follow the
 137 same processing steps prior to the spectral analysis. We first converted the time series
 138 of solar wind proton density, $n(t)$, to a length series, $L(t)$, by multiplying each time step
 139 by the radial velocity, $v_x(t)$. Then we separated the length series into overlapping seg-

140 ments $L = 9072$ Mm in length, shifting by 252 Mm each segment, and then separated
 141 segments into two categories, “fast” ($\overline{v_x} \geq 550$ km/s) or “slow” ($\overline{v_x} < 550$ km/s), based
 142 on the average proton velocity for each segment. Since the conversion of the time series
 143 to a length series produces an irregularly sampled series that is not compatible with Fourier
 144 analysis, we then resampled to common length steps of $\Delta L_s = 35.4$ Mm for slow, and
 145 $\Delta L_f = 56.7$ Mm for fast. For slow wind, 9072 Mm is approximately 6 hours of data
 146 at the median slow solar wind speed, and the 35.4 Mm ΔL is approximately equivalent
 147 to the SWE instrument sampling rate (typically 90-100 seconds) converted to length. Sim-
 148 ilarly, for the fast wind 9072 Mm is approximately 4 hours, and 56.7 Mm is the equiv-
 149 alent sampling rate multiplied by the median fast speed. Note that the categorization
 150 of fast and slow data segments is not an attempt at a physics-based classification of so-
 151 lar wind type, for which speed is not the best measure (Zurbuchen et al., 2002; Roberts
 152 et al., 2020; Borovsky, 2012). Rather, these two categories are the result of the effective
 153 sampling rate of the data segment.

154 Figure 1 shows both a slow (panels a-c) and fast (panels d-f) segment of solar wind
 155 data comparing the original (blue) and reprocessed (red) SWE data as a time series, and
 156 both datasets converted into a length series (panels c and f). These segments are typ-
 157 ical of other intervals in that they exhibit the very slight increase of a few percent in ve-
 158 locity in the reprocessed data. The reprocessed data also show differences in higher fre-
 159 quency variations, particularly for the fast wind (see Figure 1d).

160 For each data segment, we imposed data quality requirements to minimize spuri-
 161 ous spectral signals, and do not analyze segments that failed the data quality check. We
 162 required that the Wind spacecraft be located at least 50 Earth radii (R_E) upstream of
 163 Earth, to exclude any solar wind collected within or near Earth’s magnetosphere, or that
 164 could be contaminated with foreshock activity. This reduced the number of segments dur-
 165 ing the early part of the Wind mission, when it occasionally enters Earth’s magnetosphere.
 166 We remove single point data spikes and interpolated over them. We excluded any seg-
 167 ment that contained more than 10% flagged or missing data over the entire segment, or
 168 3% consecutive flagged or missing data. Finally, we excluded segments that contained
 169 discontinuous jumps (e.g., shocks) in the number density, since this would introduce “ring-
 170 ing” in the spectra. To determine a discontinuous jump, we subtracted a third order poly-
 171 nomial fit to the data segment, and discarded segments that contained changes in 5-point
 172 running averages that exceeded 3.7 standard deviations of the detrended median. The

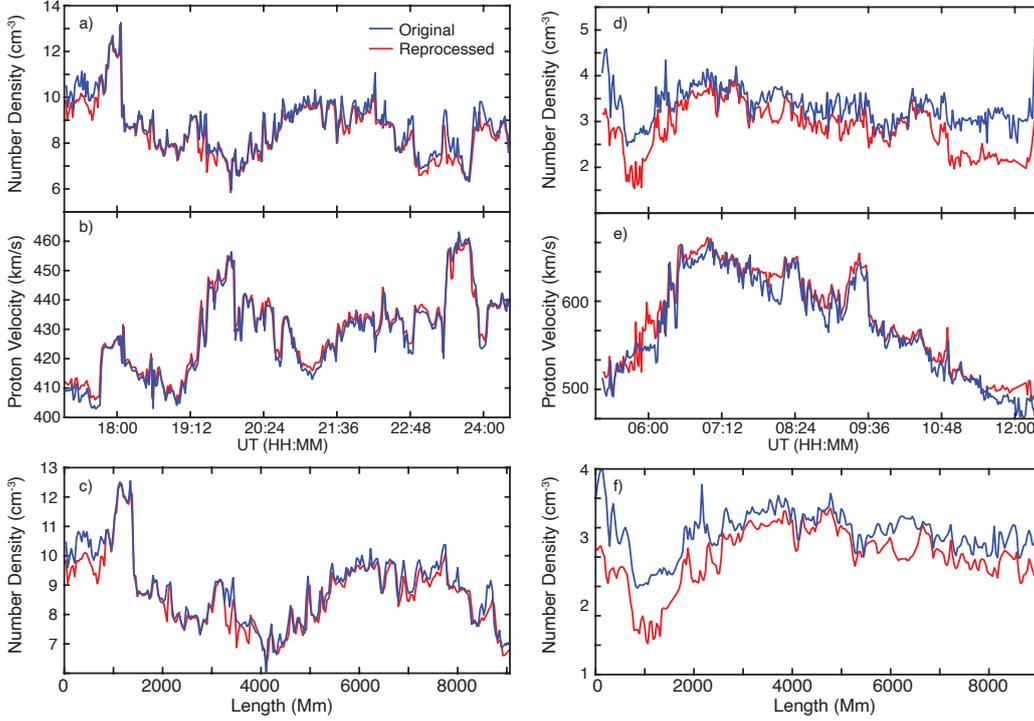


Figure 1. Comparison of the original (blue) and reprocessed (red) solar wind data from Wind SWE for a representative segment, for both slow (left) and fast (right). Reprocessed data show slightly lower density (a and d), slightly higher velocity (b and e), and lower amplitude high frequency variations compared to the original data.

173 fraction of segments that passed these quality control checks is shown in Figure 2. There
 174 is a slight decrease in the number of segments that passed these checks using the repro-
 175 cessed SWE data for the slow wind compared to the original data used by Viall et al.
 176 (2008).

177 2.2 Spectral analysis and peak detection

178 We perform spectral analysis on each segment that passed the quality checks. We
 179 identify statistically significant spectral speaks using an amplitude test and a harmonic
 180 F-test. For the amplitude test, we calculate the spectra, estimate the background fit, then
 181 identifying statistically significant peaks above this background. We use the segments
 182 in Figure 1c and 1f to demonstrate the process, and present the results in Figure 3. Es-
 183 timation of the spectra relies on the multitaper method (MTM), in which multiple, or-
 184 thogonal Slepian tapers are convolved with the data segment to provide multiple, inde-

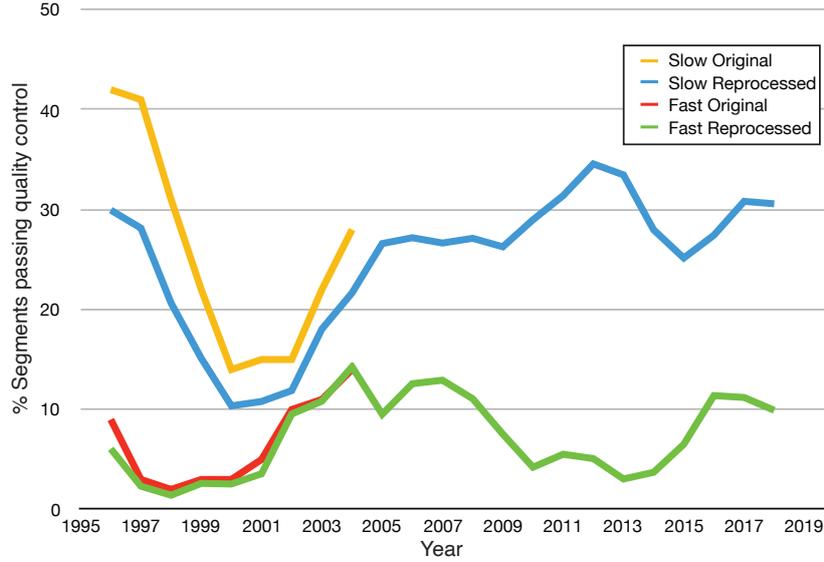


Figure 2. The percentage of the slow and fast solar wind length series segments that passed the quality control checks, and that were analyzed for periodic density structures. We also include the percentage of segments that passed these same quality checks in the original Viall et al. (2008) study. The differences are due to the reprocessed Wind SWE data.

185 pendent estimates of the spectra (Thomson, 1982). While producing a robust spectral
 186 estimate, this technique reduces the effective frequency resolution of the data as a func-
 187 tion of the number of tapers chosen, K , to $2pf_R$, where $f_R = 1/(N\Delta L)$ is the Rayleigh
 188 frequency, and $p = (K + 1)/2$. In this study we used 5 Slepian tapers, leading to an
 189 effective resolution of $6f_R$. We zeropad the data segments by a factor of 10 prior to cal-
 190 culating the spectral estimates. In Figure 3a and c we plot MTM spectra for the fast and
 191 slow length series segments shown in Figure 1, for both the original and reprocessed data.
 192 Note that the X-axis is in units of wavenumber Mm^{-1} , and we also list the equivalent
 193 length scale. Both the original and reprocessed data sets show similar spectral charac-
 194 teristics at the longer length scales (lower wavenumbers), but differ slightly at the smaller
 195 length scales (higher wavenumbers); the differences are more pronounced in the fast wind
 196 spectra. These trends are generally persistent across all segments, and is consistent with
 197 the reprocessed data having lower noise.

198 Viall et al. (2008), following Mann and Lees (1996), modeled the spectral background
 199 under the assumption that the observations x_i , at point t_i , followed an auto-regressive

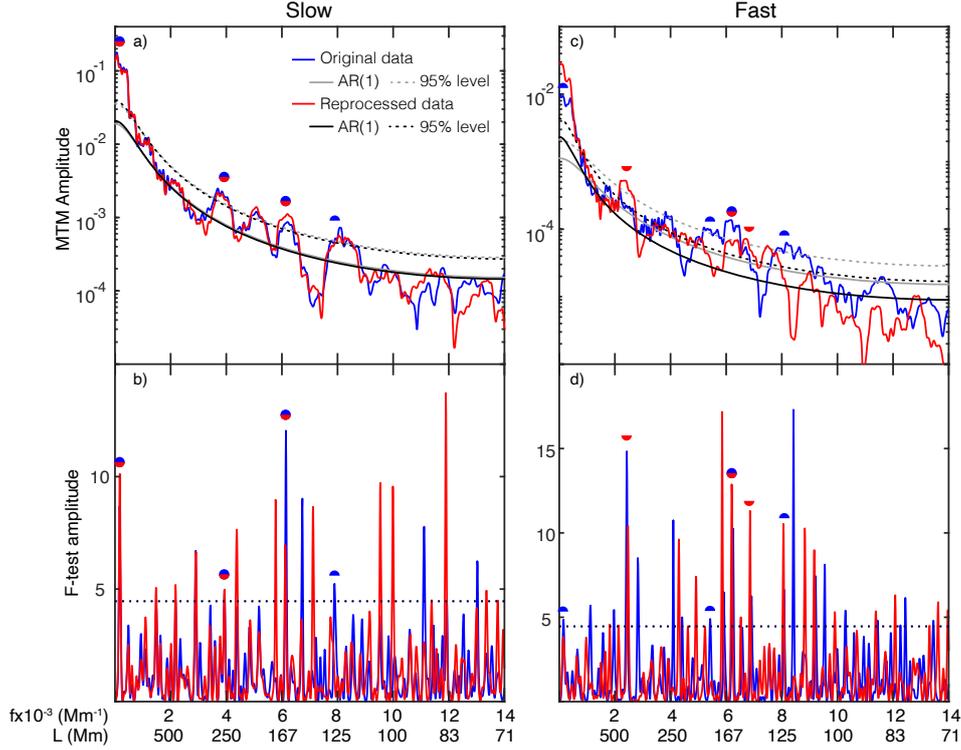


Figure 3. MTM spectra and f-test for both the slow (left) and fast (right) segment shown in Figure 1, and for both the original (blue) and reprocessed (red) Wind data. We have plotted an AR(1) background fit for both datasets, with the 95% confidence level. Peaks that simultaneously pass the amplitude and f -test are marked with half circles for both original (blue) and reprocessed (red) data.

200 AR(1) process, such that

$$x(t_i) = ax(t_{i-1}) + \epsilon_i \quad (1)$$

201 where a is the degree of correlation between sequential data points, and ϵ is random noise
 202 with zero mean (white noise). The limit of $a = 0$ produces a purely white noise spec-
 203 trum, while larger values of a produce more strongly red-noise data series. The analyt-
 204 ical spectrum of (1) is

$$S_{AR1}(f) = S_0 \frac{1 - a^2}{1 - 2a \cos(\pi f / f_N) + a^2} \quad (2)$$

205 where $S_0 = \sigma^2 / (1 - a^2)$ is the average value of the power spectrum, and σ^2 is the vari-
 206 ance of the white-noise. We fit (2) to the spectra computed using the MTM to produce
 207 an estimation of the background under the assumption of red+white noise, and confi-
 208 dence levels are determined relative to that background. AR(1) background fits and 95%

209 confidence levels for the original and reprocessed datasets, for the fast and slow segments,
 210 are shown in Figure 3a and c, overlaid on the MTM amplitude spectra. The background
 211 AR(1) fit for both the original and reprocessed data are quite similar for the slow wind,
 212 with $a = 0.836$ and $a = 0.846$, respectively. For the fast wind, however, the spectra
 213 and AR(1) fits are quite different, due to reduced high frequency power in the reprocessed
 214 data, with $a = 0.792$ and $a = 0.883$ for the original and reprocessed data, respectively.
 215 For both fast and slow wind, the AR(1) background fits lie well above the background
 216 at shorter scales (higher wavenumber), suggesting AR(1) may not be a good background
 217 assumption. We return to this in the next section.

218 The determination of a significant spectral peak, in this example frequencies that
 219 have spectral power that exceed the 95% confidence threshold, is complicated by two is-
 220 sues. First, by definition power spectrum and confidence levels produce false positives
 221 at the rate determined by the confidence thresholds (Thomson, 1982; Mann & Lees, 1996).
 222 That is, for each frequency tested for significance, for a 95% test, e.g., there is a 5% prob-
 223 ability of exceeding the threshold. These false positives would be randomly distributed
 224 in frequency, and therefore could not produce the types of preferential occurrence dis-
 225 tributions identified by Viall et al. (2008). To minimize these “false positives”, in addi-
 226 tion to the amplitude test, we apply a second type of spectral test, the harmonic F -test,
 227 which is independent of the background fit (Mann & Lees, 1996). The amplitude test
 228 requires a signal to have strong power, but does not explicitly test the discrete nature
 229 of the power enhancement. On the other hand, the harmonic F -test tests for phase co-
 230 herent signals, but does not test the power contained in those signals. As in Viall et al.
 231 (2008) we require that a spectral peak pass both the narrowband (amplitude) and F -
 232 test simultaneously to be considered significant and counted in our statistics. The pre-
 233 cise value of the peak we identify is fixed to the maximum F -test frequency within the
 234 spectral amplitude band that exceeds the threshold. Because a peak has to pass both,
 235 independent, tests simultaneously at the 95% level, our confidence threshold in appli-
 236 cation is significantly higher than 95%. Assuming that the false positives from the two
 237 tests are uncorrelated, requiring that a signal pass both tests is analogous to testing at
 238 a 99.75% confidence threshold. The second issue in identifying significant spectral peaks
 239 is that the choice of the background noise model, while not affecting the F -test, clearly
 240 affects the narrowband (amplitude) test, an issues we discuss below.

241 In Figure 3b and d we show the F -test for the representative segments, and we in-
 242 dicate peaks that pass both the narrowband and F -test at the 95% level with dots. Note
 243 that many peaks pass the harmonic F -test with little power, and are therefore not iden-
 244 tified as significant in this combined test. Similarly, there are several amplitude peaks
 245 that exceed the amplitude threshold, but not the F -test. For example, the amplitude
 246 peak at $L = 200$ Mm in the slow wind, while significant in terms of spectral amplitude,
 247 was not considered phase coherent by the F -test, and therefore was not considered sig-
 248 nificant. Since the F -test is a test for phase coherence, our study likely undercounts so-
 249 lar wind signals that have significant power but are not precisely phase coherent. As such,
 250 results that use this technique should be considered a lower bound.

251 2.3 Background estimation

252 The narrowband (amplitude) spectral test is a measure of the power of a discrete
 253 signal relative to a background spectra. The AR(1) process assumption (Equation 1) is
 254 widely used, since it is reasonable to expect a physical system to have memory. How-
 255 ever, whether that memory takes the precise form of the AR(1) in any particular seg-
 256 ment of solar wind data is impossible to know *a priori*. Indeed, Figures 3a and c shows
 257 that the AR(1) does not fit the highest and lowest wavenumbers well. We find this to
 258 be a persistent characteristic of the AR(1) fit when applied to the solar wind number
 259 density data. In effect, this bias imposes a higher confidence threshold in order to pass,
 260 and indicates that the solar wind cannot be modeled as an AR(1) process over the ~ 6
 261 hour windows we consider here.

262 The paleoclimatology community has studied the AR(1) background assumption
 263 extensively, where the choice of noise model impacts the ability to detect cycles in the
 264 stratigraphic record. In response to these concerns, Vaughan et al. (2011) suggest a bend-
 265 ing power law (BPL) background spectrum fit

$$S_{BPL}(f) = \frac{Nf^{-\beta}}{1 + (f/f_b)^{\gamma-\beta}} \quad (3)$$

266 which has the AR(1) as a special case, and performs well in mixed noise spectra. Here
 267 N is the normalization, β is the spectral slope index at low frequencies, γ is the spec-
 268 tral slope index at high frequencies, and f_b is the frequency at which the bend occurs.
 269 For low values of f_b , the BPL reduces to a straight power law with spectral slope $-\gamma$.

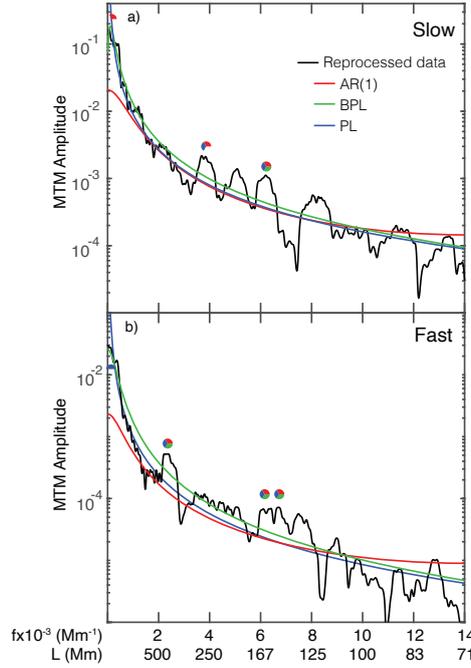


Figure 4. A comparison of three different background assumptions for the solar wind intervals shown in Figure 1. Shown are an AR(1) (red), a BPL (green) and power law (blue). We have not plotted the 95% confidence levels for clarity. Peaks that simultaneously pass the amplitude and f -test at 95 % are marked for the different fits. The spectral background model parameters are $N = 24.33$, $\beta = -0.51$, $\gamma = 1.87$, $f_b = 1.8 \times 10^{-4} \text{ Mm}^{-1}$ for BPL slow wind; $p = -1.74$ for PL slow wind; $N = .02$, $\beta = .02$, $\gamma = 2.26$, $f_b = 3.2 \times 10^{-4}$ for BPL fast wind; $p = -2.1$ for PL fast wind.

270 The BPL fit, and the 95% confidence level, is shown in Figure 4a and b in green
 271 for the same segments shown in Figure 3. Note how the BPL is a better representation
 272 of the background at both the higher and lower wavenumbers compared to AR(1) (red).
 273 We also plot a straight power law (blue) with spectral slope, p , for both slow and fast
 274 segments for reference. There is consistency in the identified peaks using the different
 275 background assumptions, with the BPL assumption producing fewer peaks in the slow
 276 wind segment. This tendency for BPL to identify fewer significant peaks than AR(1),
 277 particularly at lower frequencies, is a consistent feature across the entire 25-year study.

278 The BPL is flexible in that it allows for an AR(1) solution, a single power law, and
 279 a host of solutions in between. Since the BPL approximates the solar wind background
 280 spectra better than AR(1), and because it is more versatile than a straight power law,

281 we utilize BPL as one of the two background assumptions we use for our statistical study.
 282 For consistency with Viall et al. (2008) we also run the analysis with an AR(1) background
 283 estimate.

284 **2.4 Occurrence Distributions**

285 We applied the data processing and spectral analysis methods described above to
 286 the reprocessed solar wind measured by the Wind spacecraft from 1995-2019. For each
 287 segment we determine statistically significant peaks that pass the amplitude and F -tests
 288 simultaneously, for both BPL and AR(1) background assumptions. We create separate
 289 occurrence distributions (ODs) of the statistically significant lengths (inverse wavenum-
 290 bers) identified using the AR(1)+ F -test and BPL+ F -test criteria. For each set, we com-
 291 pute occurrence distributions over overlapping, three-year intervals, with bins of width
 292 $6\lambda_R$, the effective resolution of the MTM with our choice of $K = 5$, stepping by $3\lambda_R$
 293 for each subsequent bin. For each 3-year window, we applied the bootstrap technique
 294 ($N = 500$) to estimate the uncertainty of local peaks on the histogram, and calculated
 295 a median histogram, median fit, and standard deviation from these 500 instantiations.

296 To demonstrate this process we show the median histograms, representing an oc-
 297 currence distribution, for 1995-1998 for both the fast and slow solar wind in Figure 5a
 298 and b, with 2σ standard deviation bars determined via the bootstrap method (Efron &
 299 Tibshirani, 1993). Visually, these histograms exhibit locally enhanced counts for partic-
 300 ular lengthscale bands, with strong correlation between the occurrence enhancements us-
 301 ing the AR(1) and BPL background fits. The residuals (Figure 5b and d) highlight the
 302 similarity in local occurrence enhancements between the AR(1) and BPL histograms,
 303 despite the differences in the overall shape of the occurrence distributions. We use the
 304 bootstrapped occurrence distributions to determine statistically significant occurrence
 305 enhancements as those points that are $> 2\sigma$ above the background fit. These are high-
 306 lighted with circle in Figures 5b and d, and with thick lines in Figure 5a and c.

307 Importantly, although the AR(1) and BPL background models produce different
 308 overall shapes of the occurrence distributions, they produce similar residuals, and sim-
 309 ilar occurrence enhancements are identified as statistically significant with the bootstrap
 310 method for each. For the slow wind, the OD determined with the AR(1) assumption ex-
 311 hibits a steep slope on the short length scale (higher wavenumber) end, consistent with

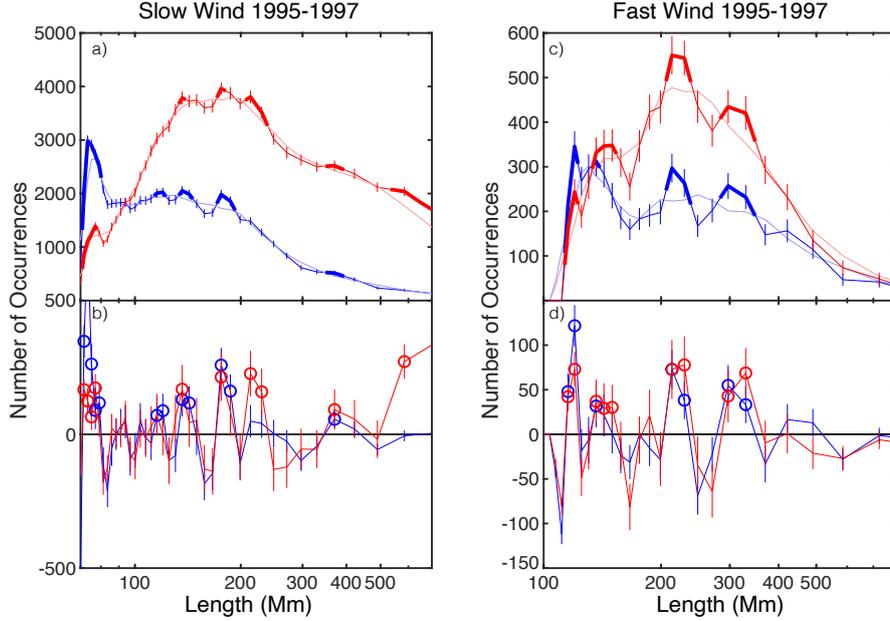


Figure 5. 3-year occurrence distributions for 1995-1997 for the slow and fast solar wind calculated for both the AR(1) (red) and BPL (blue) spectral background assumptions. Vertical bars represent $\pm 2\sigma$ standard deviation. Length scales that are greater than 2σ above the median fit (dashed lines) are shown in thick lines in (a) and (c), where we have extended the significant length scale range by $f_R/2$ in either direction. The residual distributions, obtained by subtracting the median fits from the occurrence distributions, are shown in (b) for the slow and (d) for the fast wind. Circles denote points that exceed 2σ .

312 the examples shown in Figures 3 and 4. The BPL assumption does not exhibit this bias,
 313 which provides confidence for local occurrence enhancements within this region (between
 314 $\sim 90 - 150$ Mm). For example, there is a local occurrence enhancement identified in
 315 the BPL OD near 110 Mm as $> 2\sigma$ significant, on top of a relatively flat part of the dis-
 316 tribution. In the AR(1) OD, this shows up as a relatively small local enhancement, and
 317 appears in the residual histogram as well, but is not significant at the 2σ level. In ad-
 318 dition, the ODs produced with the BPL assumption identify $\sim 50\%$ fewer significant
 319 peaks than those with the AR(1) assumption. This trend is consistent throughout the
 320 25-year interval, and indicates that the BPL is likely a better approximation for the so-
 321 lar wind background spectra, with fewer false positive detections. Despite the difference
 322 between the AR(1) and BPL results in absolute counts, the relative amplitude of the en-

323 hancements in the occurrence distribution are similar between the two background model
 324 assumptions.

325 **3 Results**

326 We ran the entire 25 year Wind SWE dataset through the analysis process described
 327 in Section 2. Figure 6 shows the percentage of analyzed segments that contained at least
 328 1 statistically significant peak that simultaneously passed the amplitude and F-test at
 329 the 95% confidence levels, for each of the AR(1) and BPL background assumptions, com-
 330 pared to the Viall et al. (2008) study. Viall et al. (2008), using the original Wind data,
 331 showed an increasing trend with time of the fraction of segments containing ≥ 1 sta-
 332 tistically significant frequency. This trend does not appear in the reprocessed data. In-
 333 stead, there is a relatively consistent number of significant radial-length peaks identified
 334 in segments during the 25-year interval, with the BPL background assumption produc-
 335 ing consistently fewer statistically significant peaks than AR(1).

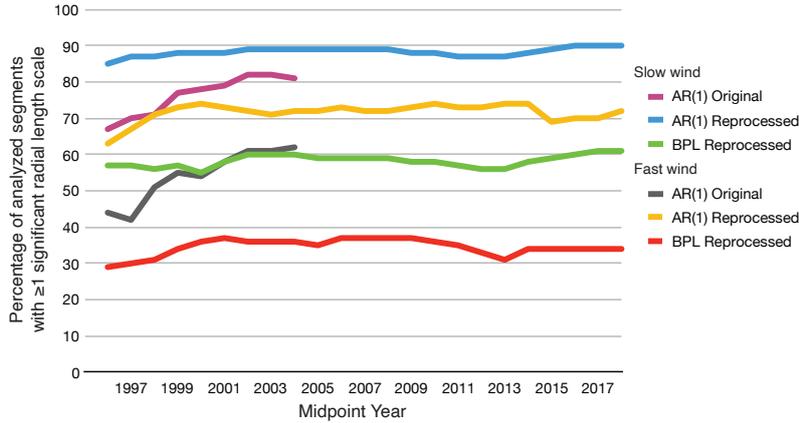


Figure 6. The percentage number of segments containing ≥ 1 statistically significant length scale for the two different fits, for both fast and slow wind, compared to the results of Viall et al. (2008)

336 We show in Figure 7 the normalized occurrence distributions of statistically sig-
 337 nificant radial length scales for slow and fast wind, and for both the AR(1) and BPL back-
 338 ground assumptions, for all 25 years of Wind data. We computed the histograms in 3-
 339 year intervals, shifting by 1-year for each new histogram. We mark the occurrence en-
 340 hancements (i.e. the persistent length scales) that are $> 2\sigma$ above the occurrence dis-

341 tribution with thick lines in Figure 7. For example, the histograms for 2017-2019 slow
 342 wind in Figure 7 show in the BPL histogram 3 clear peaks below 100 Mm, and two broad
 343 peaks near 130 and 160 Mm. The histogram derived from the AR(1) assumption show
 344 the first 2 peaks below 100 Mm and the two broad peaks near 130 and 160 Mm, but at
 345 a reduced relative amplitude compared to the BPL histogram.

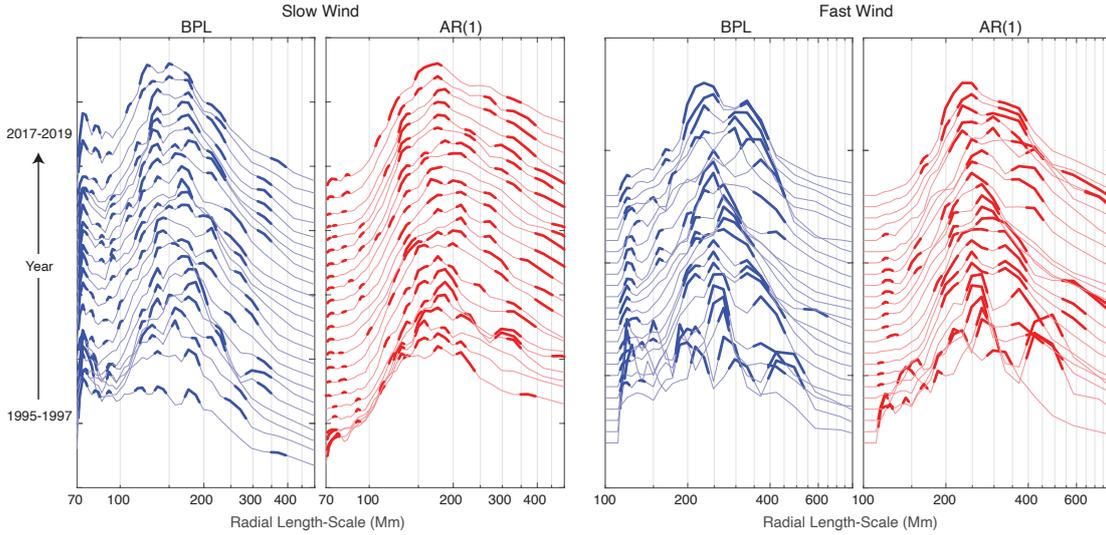


Figure 7. Bootstrapped occurrence distributions of length scales identified as significant by passing both the amplitude and F -test at the 95% level, for both slow (left) and fast (right) wind, and both background assumptions. Local peaks that exceed the background by 2σ are considered significant and are marked with thick lines.

346 To compare between the two background assumptions, we plot the significant length
 347 scales identified in both the AR(1) and BPL derived occurrence distributions as signif-
 348 icant at the 2σ level as horizontal bars in Figure 8. This comparison shows that the method
 349 applied with both background model assumptions resulted in the very similar length-
 350 scales identified as occurrence enhancements in the distributions. The primary differ-
 351 ences are at the ends of the spectral range analyzed, and follow the general pattern iden-
 352 tified in the example shown in Figure 4. At the long length-scale end (low wavenumber),
 353 fewer significant peaks were identified with the BPL (blue) background assumption, while
 354 at the short-length-scale end (high wavenumber), fewer peaks were identified with the
 355 AR(1) (red) background assumption. Many of the occurrence distributions exhibit lo-
 356 cal enhancements at the smallest length scales, very near the Nyquist, and therefore we

357 shade those particular length scales lighter to emphasize they may not be significant. Lengths
 358 that were identified concurrently in the occurrence distributions of both model fits are
 359 shown in Figure 8 as solid black bars.

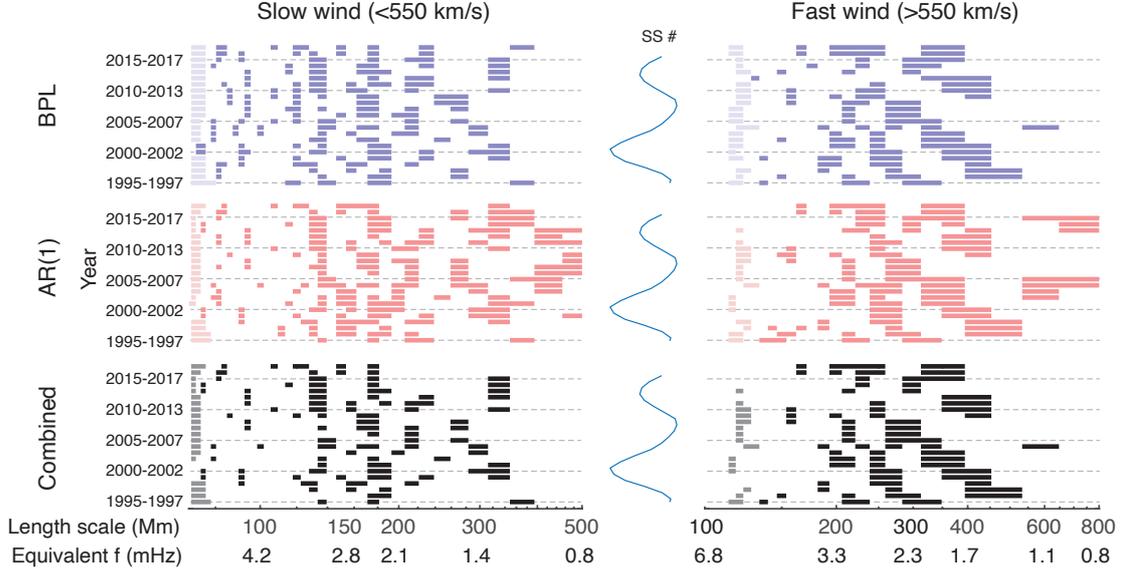


Figure 8. Bars represent statistically significant length scales identified in the occurrence distributions of Figure 7 as above the background at the 2σ level. Length scales at edge of the OD that may be affected by the Nyquist are lightly shaded. Black bars represent length scales identified simultaneously in both BPL and AR(1) distributions. We also show the equivalent frequencies using the median solar wind speed of 420 km/s for slow, and 675 for fast. For Earth’s magnetosphere, or an in situ spacecraft, these length scales would appear as periodicities at these frequencies. The sunspot number cycle is shown in the middle for reference.

360 The new results are consistent with the previous results of Viall et al. (2008) that
 361 covered the years 1995-2005 using the original Wind data. Figure 9 shows the concur-
 362 rently identified significant length scales from Figure 8 with the AR(1) derived results
 363 from Viall et al. (2008). For the slow wind (Figure 9a), both studies identified signifi-
 364 cant lengths near 130 and 170 Mm, and an additional set near 330 Mm. The differences
 365 between the original and reprocessed data occur primarily in the first 3 rows, covering
 366 years 1995-2001, during the earliest portion of the Wind mission. The fast wind results
 367 compare very well to the previous Viall et al. (2008) results, with 3 sets of length scales
 368 near 100, 300, and 400-500 Mm detected in both the original and reprocessed Wind data.
 369 The slight shift to shorter length scales in the 80-500 Mm bands in the reprocessed data

370 results is due to a reduced central peak in the OD in the reprocessed data compared to
 371 the original data.

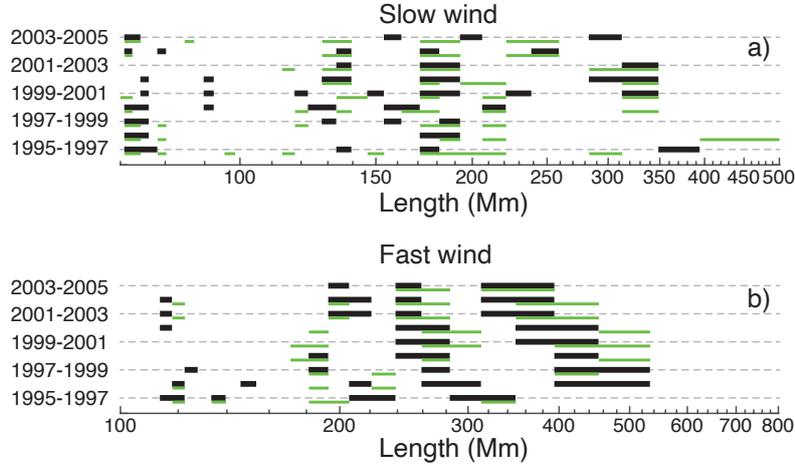


Figure 9. Comparison between the statistically significant length scales identified by Viall et al. (2008) (green) and this study (black).

372 4 Discussion

373 The histograms shown in Figure 7 represent occurrence distributions of significant
 374 length scales observed in the solar wind near L1 over 2 solar cycles. The overall shape
 375 of these distributions exhibits a consistent pattern across the full 25 years of Wind data
 376 (Figure 7). For the slow wind, the statistically significant length scales identified using
 377 the BPL background assumption exhibit comparatively few counts at the longer length
 378 scales (> 300 Mm), and a broad peak near the center of the distribution (100-200 Mm).
 379 The AR(1)-derived histograms exhibit a steep slope at the smaller length scales, followed
 380 by a slow decline at the longer length scales. The histograms for the fast wind length
 381 scales show a similar, although less pronounced, trend. Future work examining longer
 382 data segments is required to understand the nature of the shape of the occurrence dis-
 383 tribution over these length scales.

384 In addition to these overall trends, the occurrence distributions exhibit local en-
 385 hancements of length scales identified as significant. These are highlighted in the occur-
 386 rence distribution histograms in Figure 7 and pulled out separately in Figure 8 as bars.
 387 Figures 7 and 8 together provide evidence for persistent bands of significant periodic length

388 scales. To highlight these trends we have plotted colored contour plots, along with the
 389 normalized residuals from which these length scales were determined, in Figure 10a and
 390 10b. The residuals here are the addition of the normalized occurrence distribution resid-
 391 uals from the BPL and AR(1) background assumptions. The plotted values are $(OD_{BPL} -$
 392 $fit_{BPL}) + (OD_{AR(1)} - fit_{AR(1)})$, where OD is the 3-year occurrence distribution and
 393 fit is the occurrence distribution fit for the two spectral background assumptions. Length
 394 scale occurrence enhancements that were detected in both occurrence distribution resid-
 395 uals would add together (red), while parts of the distributions that are less correlated
 396 would tend to zero (green).

397 Figure 10a and 10b shows clear patterns of periodic length scales that evolve over
 398 the full 25 years of Wind SWE data. In the slow wind, $L \sim 90$ Mm (VI), $L \sim 130 -$
 399 140 Mm (III), and $L \sim 170-190$ Mm (II), are all observed for the majority of the 25-
 400 year dataset, with some noticeable variations we discuss below. There are two smaller
 401 bands near $L \sim 210$ Mm (IV) in the middle years and between 310 and 350 Mm in the
 402 later years, and a sloped band between 250 and 400 Mm (I) for the first half of the in-
 403 terval. An additional band appears near $L \sim 120$ Mm in the BPL-derived histograms
 404 in Figure 7a, but is not apparent in the AR(1)-derived histograms, likely because this
 405 region has a very strong slope; there is a similar effect with the $L \sim 90$ Mm band (see
 406 Figure 7). For the fast wind there is an intermittent band between $L \sim 200-220$ Mm
 407 (IV), and a series of sloped bands that appear to decrease in L as functions of year.

408 Figure 10c shows a pictorial summary of the significant length scale bands, derived
 409 by examining the combined bar plots and residual contours of Figure 10a and 10b, and
 410 using the additional information of the histograms in Figure 7 to provide visual guid-
 411 ance on persistence. We have also combined the significant length scales observed in the
 412 slow and fast wind together. We note that bands I, IV and V, observed in both fast and
 413 slow wind analysis, clearly overlap, suggesting that these periodic density structures are
 414 not a strictly “slow” (< 550 km/s) wind phenomenon.

415 Many characteristics of the Sun, solar corona, and solar wind are correlated with
 416 solar cycle, so unraveling the specific nature of the correlation of periodic density struc-
 417 tures with solar cycle is a topic for future work. Here we speculate on a likely connec-
 418 tion. In general, the solar corona is hotter, and its magnetic topology increases in com-
 419 plexity, at solar maximum, as manifested in active regions and their underlying magnetic

420 concentrations, sunspots. To the right of Figure 10c we show the gradual solar cycle change
421 as measured by 3-year averages of the sunspot number, along with the more abrupt “ter-
422 minator” events that occur during the rising phase of the solar cycle (McIntosh et al.,
423 2015, 2019). The terminator events are observed as abrupt changes in the distribution
424 of solar EUV bright points, and are correlated with the appearance of active regions (and
425 their associated sunspots) that have the polarity of the next solar cycle. Active regions
426 with the past-cycle polarity can occur even after solar minimum, hence these termina-
427 tor events are offset in time from sunspot minimum. Related, Schonfeld et al. (2017) showed
428 that the amount of hot plasma (plasma greater than $10^{6.1}\text{K}$) in the solar corona abruptly
429 increases at the terminator, due to an increased amount of hot plasma in active regions.

430 The length scale bands that we find in this paper exhibit a solar cycle dependence,
431 but they exhibit breaks that are associated more closely with terminators than with sunspot
432 minimum. For example, bands I and II at both ends, and band III for the termination
433 event of solar cycle 22. Additionally, there is a gradual evolution of the characteristic
434 length scales between termination events, most pronounced in bands I and V. With data
435 from only two, very different, solar cycles, we cannot draw definitive conclusions about
436 the exact relationship between solar wind periodic length scales and the solar cycle, but
437 the result suggests that a relationship exists.

438 As reviewed in Section 1, there is strong evidence that periodic density structures
439 originate from the sun and are associated with magnetic reconnection of plasma from
440 closed-field regions. The association of periodic length scales with solar cycle could be
441 the result of changes in the nature of the interchange reconnection that releases the plasma
442 into the solar wind, due to the increase in complexity of the global magnetic topology
443 (Antiochos et al., 2011) and the local nature of the magnetic field concentrations (Mason
444 et al., 2019). Coronal temperature is correlated with solar wind speed, so it could also
445 be that the hotter active regions accelerate solar wind, and any embedded periodic den-
446 sity structures, differently.

447 While this study focused specifically on mesoscale structures measured at L1 that
448 exhibit periodicity in density, many other studies have observed mesoscale structures in
449 the solar wind that form at the sun and advect to 1 AU. A general picture is emerging
450 in which mesoscale structures that form through spatial structures that rotate (Borovsky,
451 2008, 2020) or time dynamics such as reconnection in the corona (Sanchez-Diaz et al.,

452 2016, 2017, 2019; M. J. Owens et al., 2018; M. Owens et al., 2020; Stansby & Horbury,
 453 2018), are an inherent part of solar wind formation (see review by Viall and Borovsky,
 454 2010).

455 In a series of papers, Rouillard, Davies, et al. (2010), Rouillard, Lavraud, et al. (2010),
 456 and Rouillard et al. (2011) tracked larger mesoscale structures from their formation in
 457 the corona through the inner heliosphere using SECCHI HI images, all the way to their
 458 impact at the Earth. They identified the corresponding compositional and magnetic field
 459 variations inherent to the structures, which were retained out to 1 AU. This set of stud-
 460 dies unequivocally demonstrated that mesoscale structures created at the Sun survive to
 461 1 AU with identifiable in situ signatures. More recently, Rouillard et al. (2020) tracked
 462 density structures through the STEREO COR2 and HI1 FOVs to their impact at Parker
 463 Solar Probe, where they observed a one-to-one correlation between the $\sim 3 - 4$ hour
 464 density structures observed remotely and the in situ Parker measurements. They showed
 465 that Parker measured additional sequences of small density peaks separated in time by
 466 approximately 90-120 minutes, suggestive of the types of periodic density enhancements
 467 at 90 minute timescales that have been observed in situ at L1 (Viall et al., 2008; Kepko
 468 & Spence, 2003), near Mercury’s orbit with Helios (Matteo et al., 2019) and remotely
 469 with STEREO (Viall & Vourlidas, 2015). Many of these event studies exhibited still smaller
 470 substructures at tens of minutes (Matteo et al., 2019; Kepko & Viall, 2019; Stansby &
 471 Horbury, 2018; Kepko & Spence, 2003). Several studies also found composition signa-
 472 tures which could only have come from formation at the sun (Viall, Spence, & Kasper,
 473 2009; Kepko & Viall, 2019). These studies together demonstrate that the solar wind is
 474 often composed of mesoscale density structures, and provide ample evidence that struc-
 475 tures of order tens of minutes timescales and longer form with the solar wind and sur-
 476 vive through the inner heliosphere, out to 1 AU. This current study further demonstrated
 477 that at least some of those structures are quasi periodic, and occur at repeatable sets
 478 of frequencies and/or length scales.

479 We emphasize that these length scales represent periodic density structures that
 480 advect with the solar wind. In the rest frame of a spacecraft or planet, they would ap-
 481 pear as a periodic density variations at a frequency determined by $f_{PDS} = V_{sw}/L_{PDS}$.
 482 Statistically, for any particular year the magnetosphere or a spacecraft would see a spec-
 483 trum of equivalent frequencies determined by convolving the distribution of solar wind
 484 V_x with the length scales identified in Figures 8 and 10c from that year. To zeroth or-

485 der, we can estimate these frequencies using the median solar wind speed for “fast” and
 486 “slow” solar wind. These equivalent frequencies are listed at the bottom of Figure 10c.
 487 The equivalent frequencies of these structures fall in the few mHz range, which for the
 488 magnetosphere is considered the Pc5 band and higher. Previously, Viall, Kepko, and Spence
 489 (2009) studied 11 years of Wind SWE data covering 1995-2005 for evidence of discrete
 490 frequency periodicities in the solar wind number density. They found that $f = 0.7, 1.3$ –
 491 $1.5, 2.0 - 2.3,$ and $4.7 - 4.8$ mHz occurred most often over that 11-year interval. Fig-
 492 ure 10c demonstrates that $f = 1.4$ mHz corresponds to Band I in the slow wind, $f =$
 493 $2.0 - 2.3$ mHz corresponds to Band IV in the slow and I in the fast, and $f = 4.7 - 4.8$
 494 mHz corresponds to Band VI in the slow wind.

495 Since these are periodic structures in solar wind density, they would periodically
 496 compress the magnetosphere via periodic dynamic pressure changes, and we would ex-
 497 pect the magnetosphere to show these same sets of frequencies. In the same Viall, Kepko,
 498 and Spence (2009) study, they also examined GOES magnetospheric magnetic field data
 499 for intervals when GOES was near the dayside magnetopause, and found in the GOES
 500 data a similar set of frequencies to those found in the solar wind. In a direct compar-
 501 ison between Wind and GOES, they found when a spectral peak was observed in the so-
 502 lar wind, that same peak was observed at GOES 54% of the time. Other statistical stud-
 503 ies have similarly identified persistent bands of significant mHz frequencies (e.g., Francia
 504 and Villante (1997); Chisham and Orr (1997); Ziesolleck and McDiarmid (1995)). While
 505 originally attributed to global cavity modes (e.g., Harrold and Samson (1992)), we now
 506 know these $< \sim 4$ mHz oscillations are largely driven by solar wind periodic density struc-
 507 tures. Since these periodic length scales directly drive the magnetosphere, we would ex-
 508 pect the spectrum of discrete mHz oscillations in the magnetosphere to vary year-to-year
 509 as the L_{PDS} vary. Since the L_{PDS} have a solar cycle dependence, this would mean the
 510 spectrum of discrete mHz waves in the magnetosphere would also have a solar cycle de-
 511 pendence, although the variability of the solar wind speed would produce broad, rather
 512 than narrow, enhancements. This slow year-to-year variability, and the distribution of
 513 solar wind speeds, can explain year-to-year changes in measured frequencies. In addi-
 514 tion, Kepko and Viall (2019), showed that ambient periodic density structures in the slow
 515 solar wind were sometimes compressed and amplified by a faster solar wind stream from
 516 behind, and that these amplified PDSs had an observable impact on radiation belt par-

517 ticles. These particular PDSs were observed with stream interaction regions, which are
518 known to be important drivers of radiation belt flux enhancements.

519 **5 Conclusions**

520 Using 25 years of Wind solar wind number density data observed near L1 we have
521 identified bands of periodic length scales that occur more often than others. Each oc-
522 currence of periodic length scales passed two independent spectral tests at the 95% level,
523 and we tested each occurrence with two different background spectral models. We iden-
524 tify bands of occurrence enhancements that are persistent in time, and regardless of back-
525 ground spectral model (Figure 8c). These bands, particularly at larger length scales, have
526 a clear solar cycle dependence, and their evolution may be related to “terminator” events
527 (Figure 10). This study provides further evidence that large portions of the solar wind
528 plasma consist of mesoscale structures that are released via magnetic reconnection. Fi-
529 nally, in the rest frame of a spacecraft or Earth, these periodic mesoscale density struc-
530 tures would appear as Pc5-6 pulsations, which are known to be important in processes
531 leading to radiation belt particle loss, diffusion, and acceleration. Given the statistical
532 bands of recurrent length scales in the solar wind, and a solar cycle dependence, it may
533 be possible in the future to produce a statistical model for these solar-wind driven dis-
534 crete oscillations.

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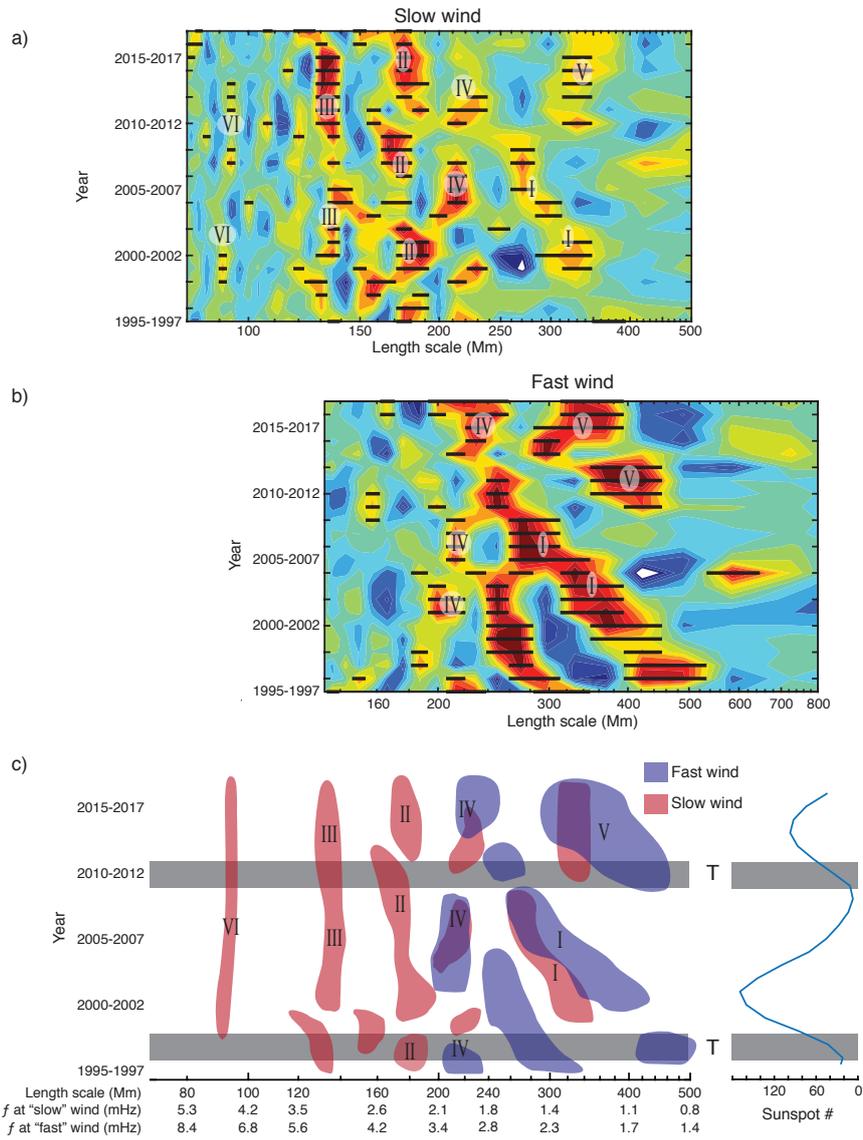


Figure 10. The contour plots (a) and (b) are the addition of the normalized (to the peak value) residuals for both the BPL and AR(1) derived occurrence distributions for the slow and fast wind. Red indicates areas of enhancement observed in both OD residuals, blue indicates areas where both found length scales significantly below the background fit, and green indicates regions near the background or areas where BPL and AR(1) were in disagreement. The bars superimposed on (a) and (b) are from Figure 8(e) and (f), and indicate length scales that exceeded the background by 2σ . The schematic (c) is a pictorial representation of (a) and (b) combined, and includes the 3-year running average of sunspot number, and locations of the terminator. slow wind is 420, fast 675 ** Need to get last 2 years of sunspot data

Figure 1.

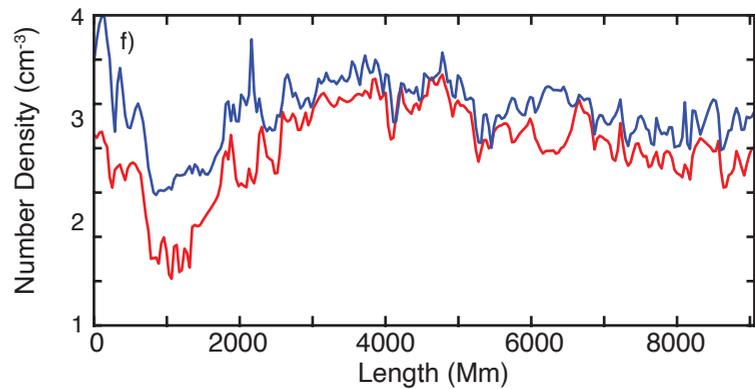
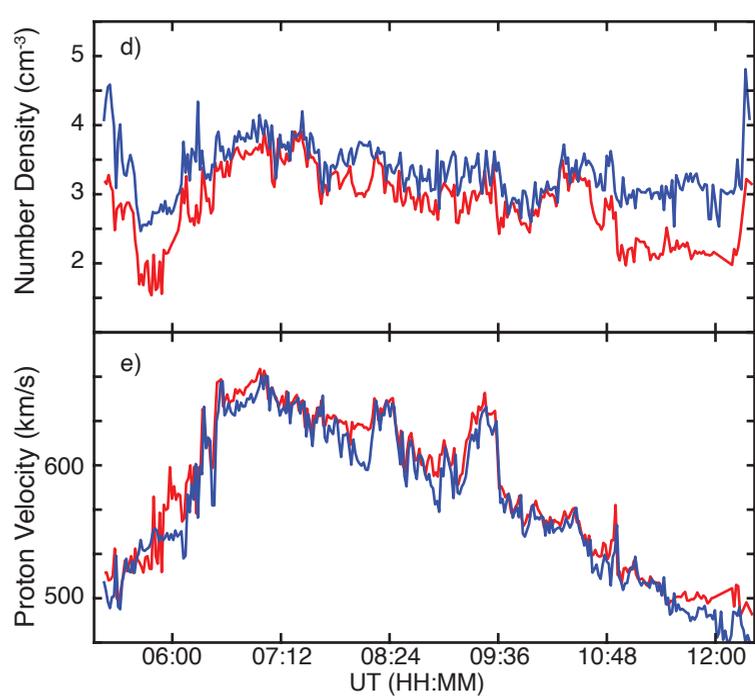
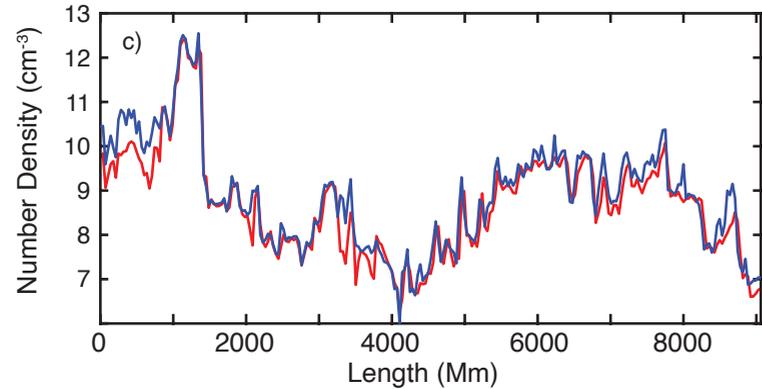
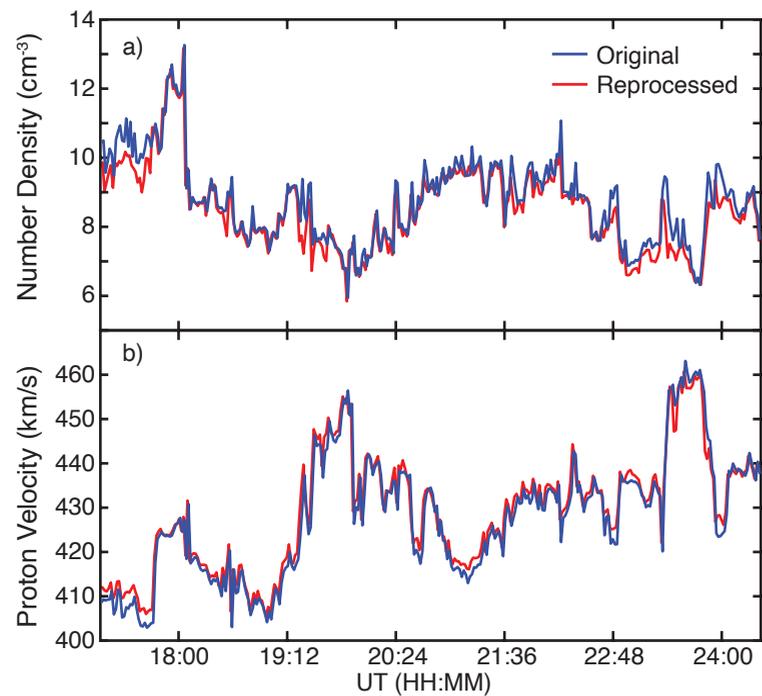


Figure 2.

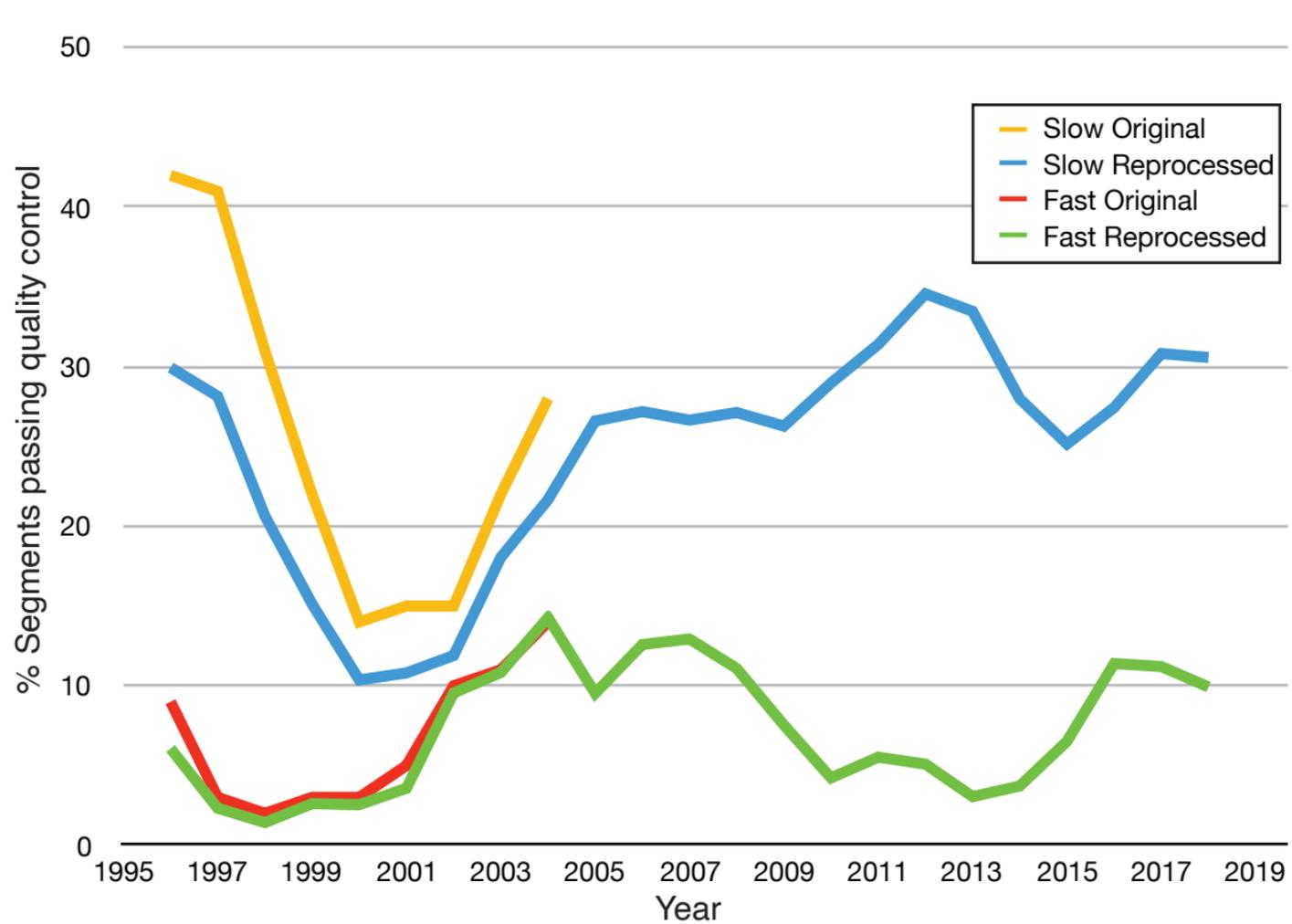
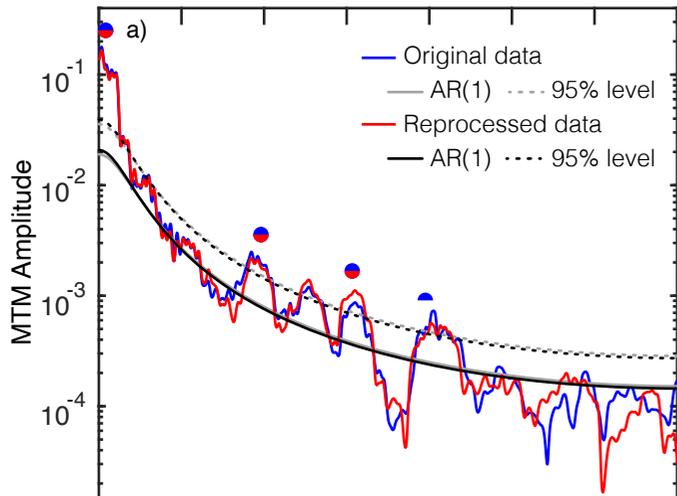
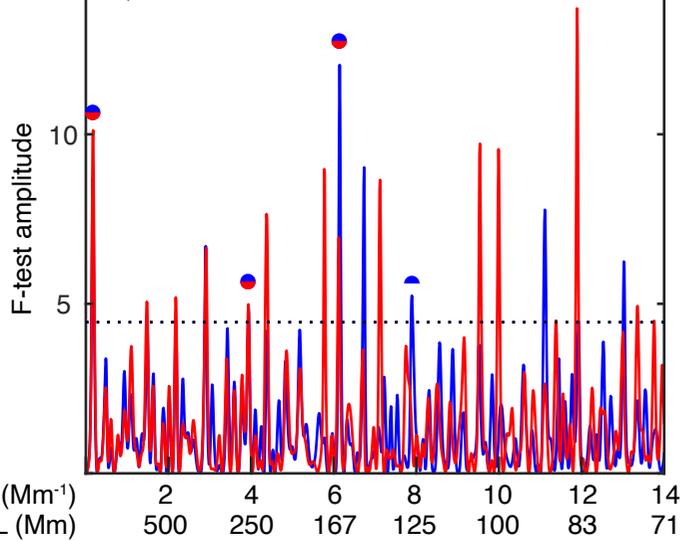


Figure 3.

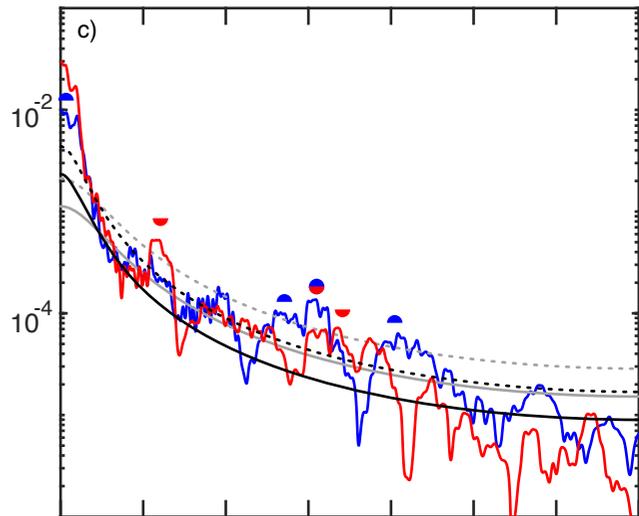
Slow



b)



Fast



d)

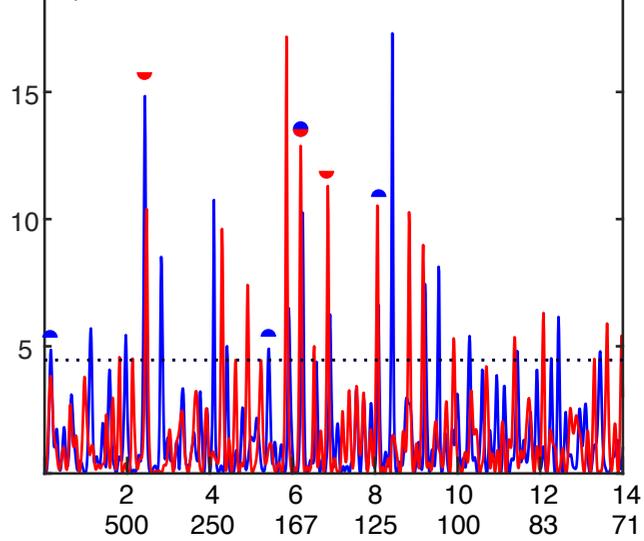
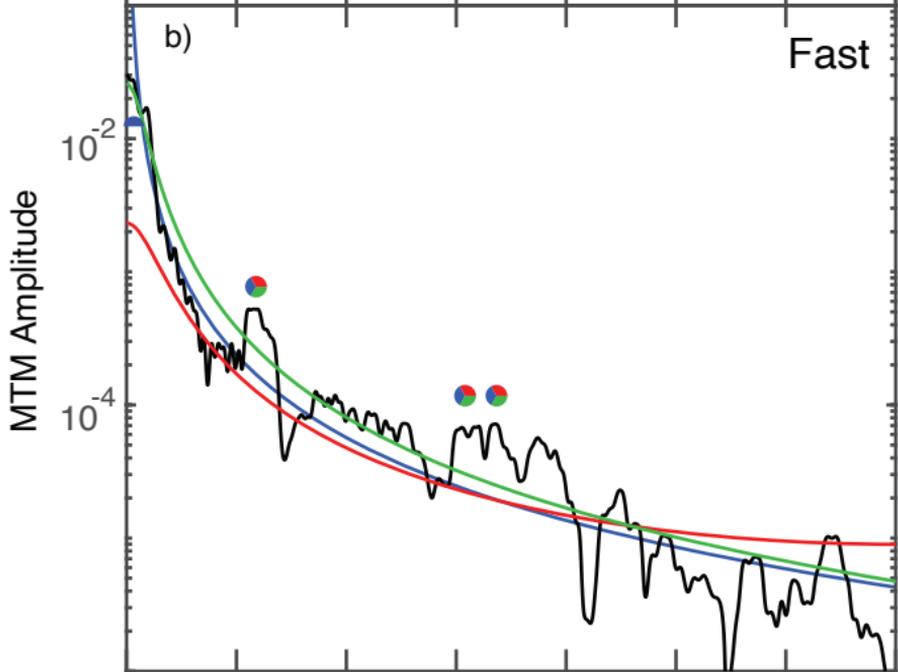
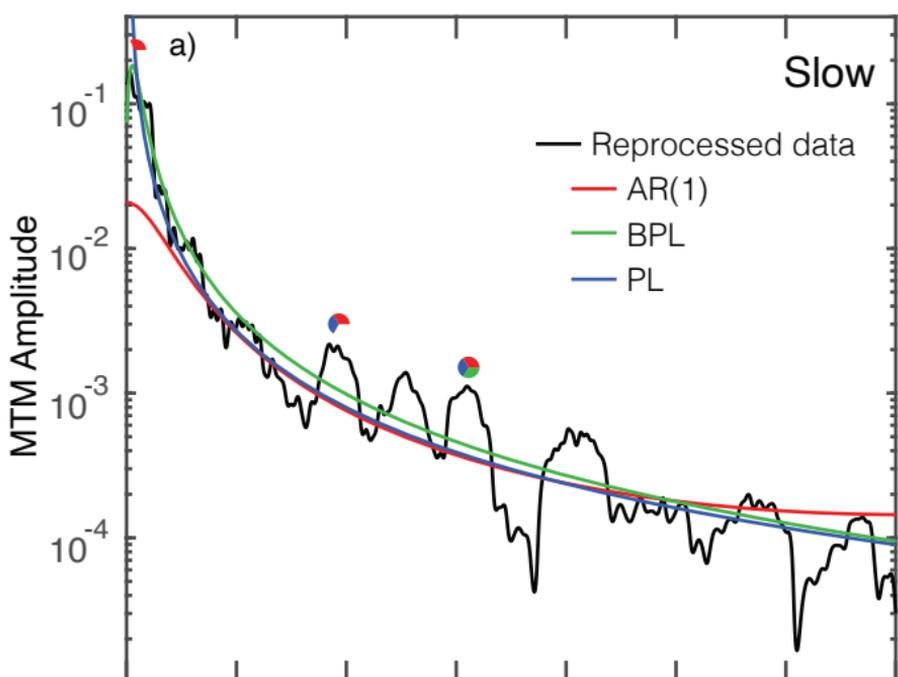


Figure 4.



$\times 10^{-3}$ (Mm $^{-1}$)
L (Mm)

500

250

167

125

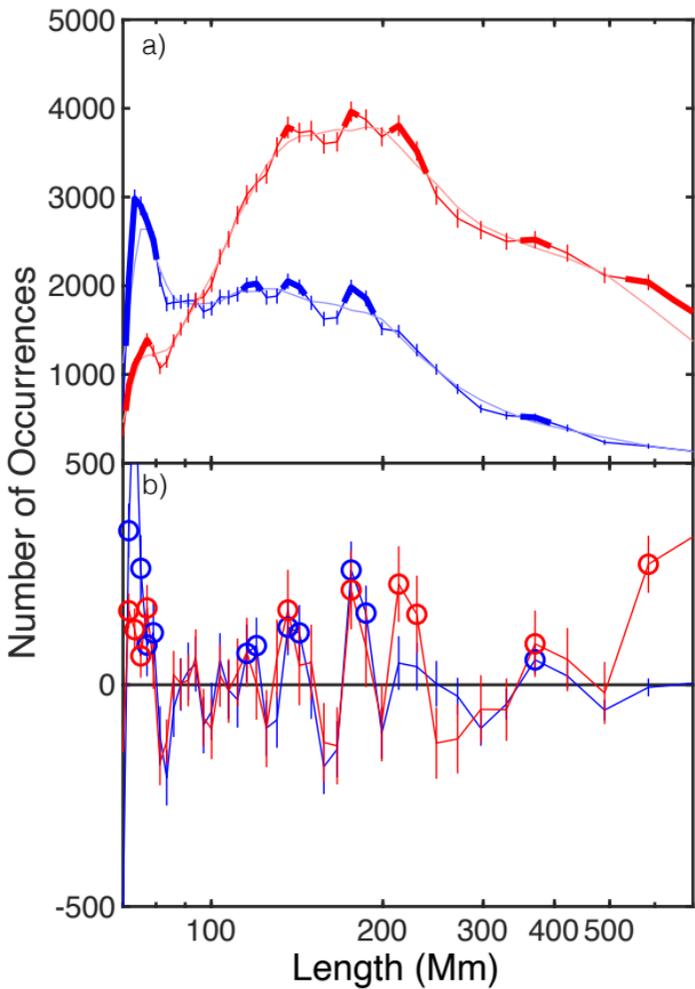
100

83

71

Figure 5.

Slow Wind 1995-1997



Fast Wind 1995-1997

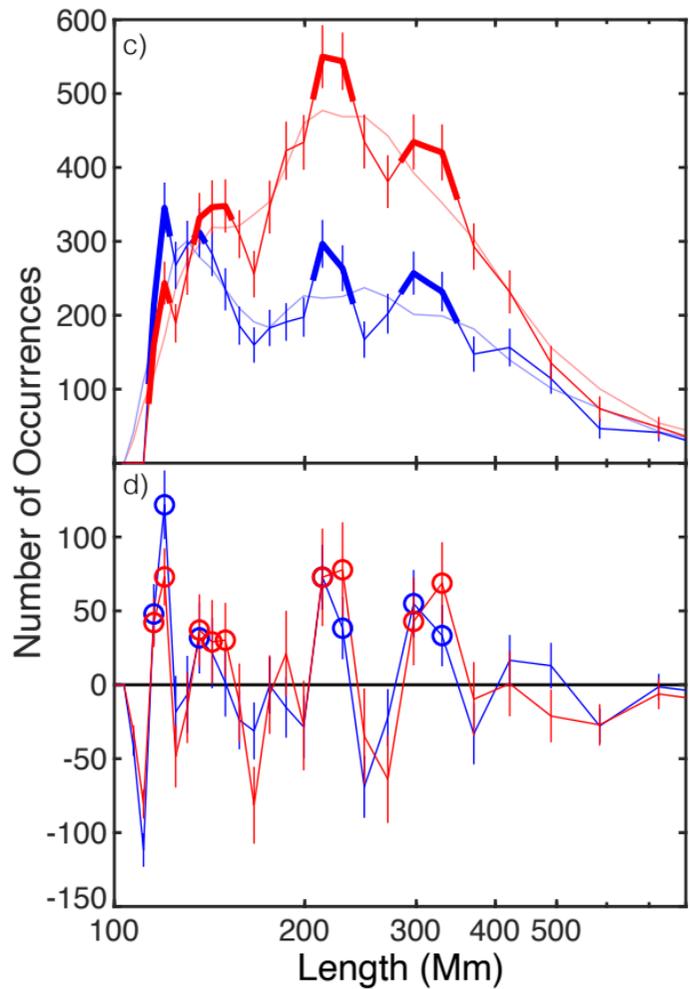


Figure 6.

Percentage of analyzed segments
with ≥ 1 significant radial length scale

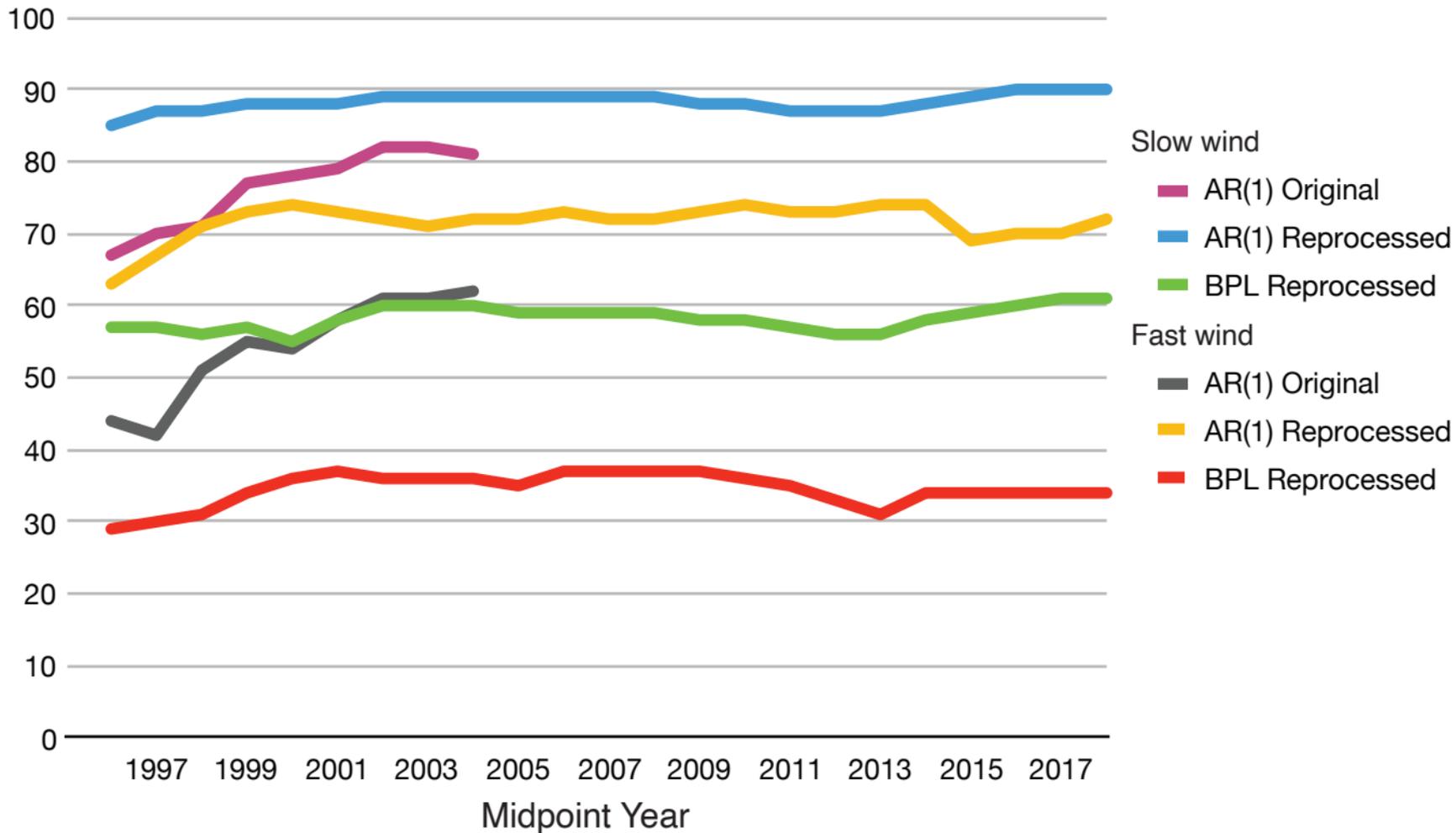


Figure 7.

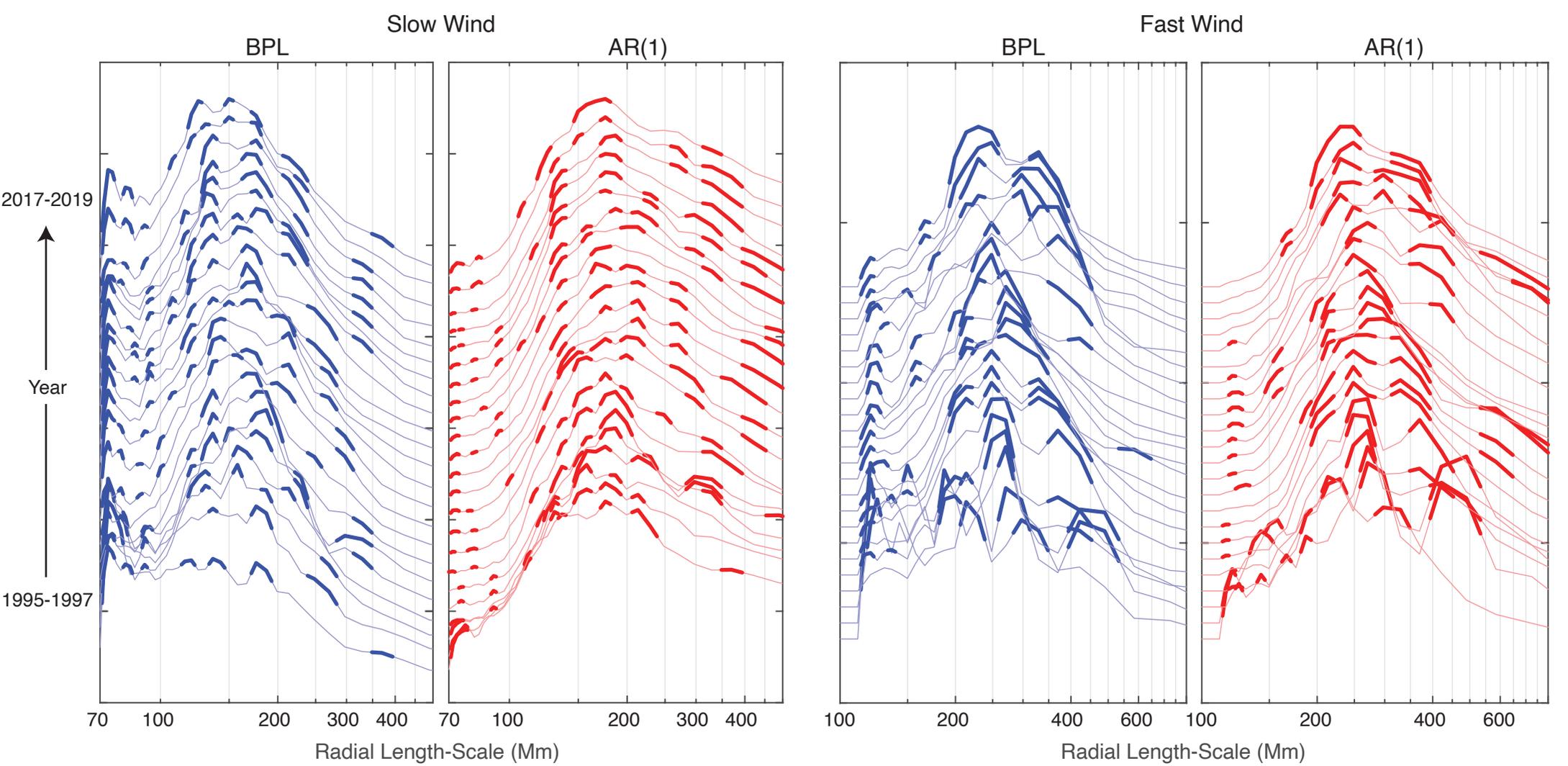


Figure 8.

Slow wind (<550 km/s)

Fast wind (>550 km/s)

BPL

AR(1)

Combined

Year

2015-2017
2010-2013
2005-2007
2000-2002
1995-1997

SS #

Length scale (Mm) 100 150 200 300 500
Equivalent f (mHz) 4.2 2.8 2.1 1.4 0.8

100 200 300 400 600 800
6.8 3.3 2.3 1.7 1.1 0.8

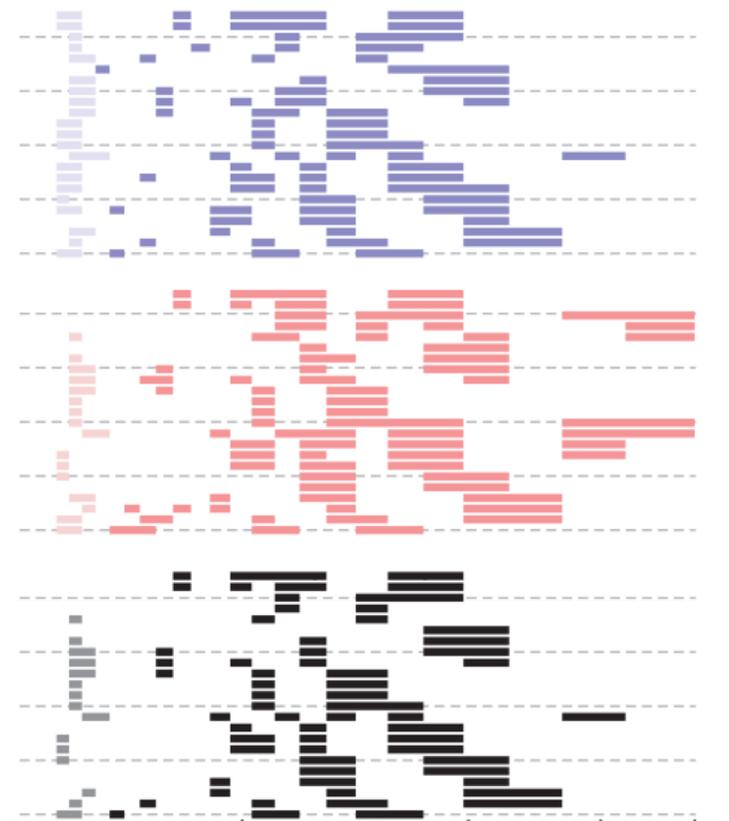
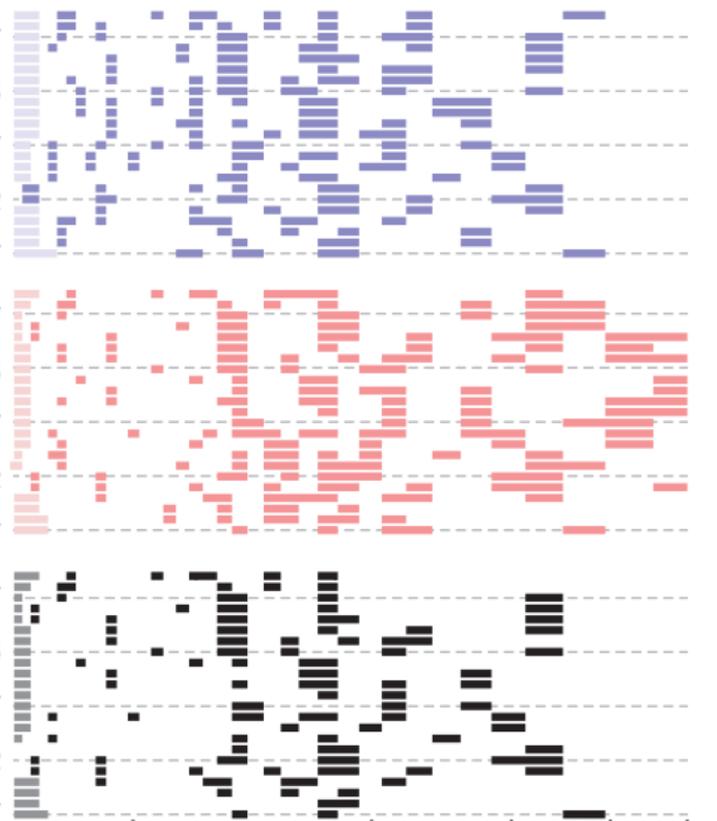
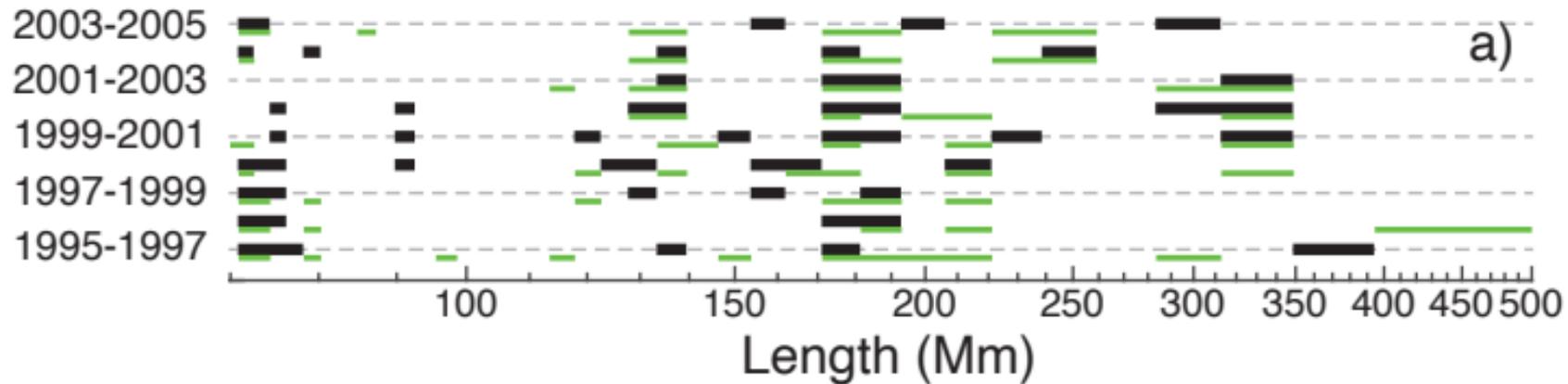


Figure 9.

Slow wind



Fast wind

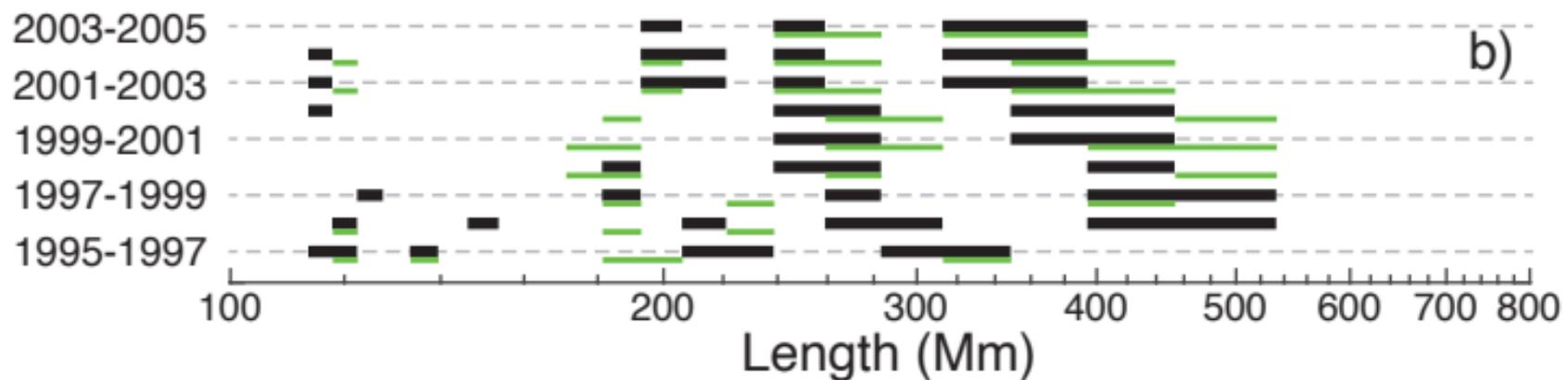


Figure 10.

