

1                   **Adiabatic temperature profile in the mantle, revised**

2                   **Tomoo Katsura**

3                   Bayerisches Geoinstitut, University of Bayreuth, Universitaetsstrasse 30, 95447 Bayreuth,  
4                   Germany.

5                   Corresponding author: Tomoo Katsura ([tomo.katsura@uni-bayreuth.de](mailto:tomo.katsura@uni-bayreuth.de))

6                   **Key Points:**

- 7                   • The previous olivine-wadsleyite phase relations and P-V-T relations of mantle  
8                   minerals are corrected using the pressure effect on EMF.
- 9                   • Comparing the 410-km discontinuity depth with the olivine-wadsleyite phase  
10                  relations suggests 1805(18) K at the discontinuity.
- 11                  • The adiabatic temperatures are 1617(16) at 50-km depth and 2530(30) K at 2800-km  
12                  depth.

13

14 **Abstract**

15 This study re-evaluates the adiabatic temperature profile of the Earth's mantle. The global  
16 average temperature at the 410-km discontinuity is estimated to be 1805(18) K by comparing  
17 the globally averaged depths of the 410-km discontinuity with the previously determined  
18 phase diagram of the olivine-wadsleyite transition in the system  $(\text{Mg},\text{Fe})_2\text{SiO}_4$  at two  
19 temperatures. The temperature at the 410-km discontinuity is extrapolated to shallower and  
20 deeper regions using the adiabatic temperature gradient, which is estimated from the  
21 pressure-volume-temperature relations and heat capacities of the major mantle minerals,  
22 namely, olivine, wadselyite, ringwoodite, and bridgmanite. The experimental temperatures  
23 and pressures in the original studies used in these evaluations are re-calculated using the  
24 recently proposed pressure correction on EMF of the  $\text{W}_{97}\text{Re}_3$ - $\text{W}_{75}\text{Re}_{25}$  thermocouple. The  
25 uncertainties are evaluated by the Monte Carlo simulation. The temperatures on the adiabatic  
26 geotherm are found to be 1617(16), 1959(19), 1925(19), and 2530(30) K, respectively, at a  
27 50-km depth, just above the 660-km discontinuity, just below the 660-km discontinuity, and a  
28 2800-km depth. These temperatures are higher than those given by Katsura *et al.* (2010). The  
29 50-km depth temperature is slightly higher but generally agrees to that estimated from the  
30 melting of depleted peridotite.

31 **Plain Language Summary**

32 This study estimates the temperature profile of the Earth's mantle by generally following the  
33 approach described in Katsura *et al.* (2010). The estimation consists of two steps. First, the  
34 temperature at the 410-km seismic discontinuity (D410), at which the seismic wave velocities  
35 abruptly increase almost everywhere in the mantle, is evaluated. The D410 is usually  
36 attributed to the olivine-wadsleyite transition in peridotite. Comparing the globally averaged  
37 D410 depth with the phase diagram of the olivine-wadsleyite transition yields the D410  
38 temperature of 1805(18) K. Second, this temperature is extrapolated to shallower and deeper  
39 regions by assuming that the heat is mainly transferred by convection in the mantle. The  
40 temperature gradient in such cases is the product of the thermal expansion coefficient and the  
41 temperature divided by the density and the heat capacity. The thermal expansion coefficients  
42 of the major mantle mineral are obtained by recalculating our previous experimental data. We  
43 found that the temperatures at 50-km depth, the bottom of the mantle transition zone, the top  
44 of the lower mantle, and 2800-km depth are found 1617(16), 1959(19), 1925(19), and  
45 2530(30) K, respectively. The 50-km depth temperature is slightly higher but generally  
46 agrees to that estimated from the melting of depleted peridotite.

47 **1 Introduction**

48 The temperature is one of the essential parameters for modeling the dynamics of the  
49 Earth's interior. Therefore, estimating the temperature distribution in the mantle is a vital task  
50 in solid geophysics. However, we cannot directly measure temperatures in the Earth's deep  
51 interior. We can only estimate the temperature distribution by combining various pieces of  
52 information obtained by indirect methods.

53 For this reason, Katsura *et al.* (2010) estimated the adiabatic temperature distribution  
54 in the mantle. They first estimated the temperature at the 410-km seismic discontinuity  
55 (D410) by comparing the global average depth of the D410 with the olivine-wadsleyite  
56 transition pressure as a function of temperature (Katsura *et al.*, 2004a). Then, they estimated  
57 the temperatures above and below the D410 using the formula of the adiabatic temperature  
58 gradient with depth as:

$$59 \left(\frac{dT}{dz}\right)_S = \frac{\alpha g T}{c_p} \quad (1)$$

60 where T is the temperature, z is the depth,  $\alpha$  is the thermal expansivity of the constituent, g is  
61 the gravitational acceleration, and  $C_p$  is the isobaric heat capacity per weight of the  
62 constituent (Turcotte and Schubert, 2014). In their calculation, the mantle rocks were  
63 approximated by the Mg endmembers of olivine, wadsleyite, ringwoodite, and bridgmanite.  
64 The P-V-T relations of these minerals were determined using the multianvil *in situ* X-ray  
65 diffraction experiments by Katsura *et al.* (2004b; 2009a; 2009b; 2009c). The  $C_p$  for olivine,  
66 wadsleyite, and ringwoodite were taken from Saxena *et al.* (1993), and that for bridgmanite  
67 was calculated using the Debye model.

68 However, the adiabatic temperature profile by Katsura *et al.* (2010) has to be revised  
69 for the following reasons. First, the temperatures in Katsura *et al.* (2004a; 2004b; 2009a;  
70 2009b; 2009c) were measured using W<sub>97</sub>Re<sub>3</sub>-W<sub>75</sub>Re<sub>25</sub> thermocouples without any pressure  
71 correction. Nishihara *et al.* (2020) provided the pressure correction of the EMF-temperature  
72 relations of the W<sub>97</sub>Re<sub>3</sub>-W<sub>75</sub>Re<sub>25</sub> thermocouple. Therefore, the experimental data in Katsura  
73 *et al.* (2004a; 2004b; 2009a; 2009b; 2009c) should be recalculated using Nishihara *et al.*'s  
74 (2020) correction. Second, Tange *et al.* (2012) obtained more reliable P-V-T data of MgSiO<sub>3</sub>  
75 bridgmanite than Katsura *et al.* (2009c). These data should be included in estimating the  
76 lower-mantle adiabatic profile. Third, the calculation program used in Katsura *et al.* (2010)  
77 was found to have errors leading to incorrect thermal expansivity evaluation, as shown later.

78 This paper presents a revised average adiabatic temperature profile in the mantle by  
79 integrating the above-mentioned new data and using a newly made calculation program.  
80

## 81 **2 Methods**

82 The current study estimates an adiabatic temperature profile in the mantle by the  
83 procedure very similar to Katsura *et al.* (2010). Namely, it first estimates the temperature at  
84 D410, then estimates the temperature gradient, finally calculates the geotherm profile from  
85 the D410 temperature using the estimated temperature gradient. Details of the procedure are  
86 explained below.

### 87     2.1 Temperature at the 410-km discontinuity

88 The current study first considers the most probable global average of the D410 depth.  
89 Flanagan and Shearer (1998; 1999), Chambers *et al.* (2005), Houser *et al.* (2008; 2016)  
90 mapped the D410 depths globally and suggested the averaged depths of 418, 418, 409, 410,  
91 and 411 km, respectively. Although the older studies (Flanagan and Shearer, 1998; 1999)  
92 should have contained some uncertainties in absolute D410 depths due to the velocity  
93 structure variation of the crust, the recent two studies (Houser *et al.*, 2008; 2016) corrected  
94 them using the CRUST model (Mooney *et al.*, 1998). Therefore, the current study adopts the  
95 global average D410 depth of 410±1 km for the following calculation.

96 The current study employs the data given by Katsura *et al.* (2004a) for the phase  
97 relations of the olivine-wadsleyite transition in (Mg,Fe)<sub>2</sub>SiO<sub>4</sub>. They synthesized coexisting  
98 olivine and wadsleyite in a multi-anvil press with temperatures measured by W<sub>97</sub>Re<sub>3</sub>-W<sub>75</sub>Re<sub>25</sub>  
99 thermocouples and pressures estimated from MgO volumes by *in situ* X-ray diffraction using  
100 the P-V-T relations of MgO suggested by Matsui *et al.* (2000). Then, Katsura *et al.* (2004a)  
101 measured the compositions of recovered olivine and wadsleyite grains using an electron  
102 probe microanalyzer (EPMA). In the current study, the temperatures in Katsura *et al.* (2004a)  
103 are corrected using Nishihara *et al.*'s (2020) pressure correction of the EMF-temperature  
104 relations. The pressures in Katsura *et al.* (2004a) are recalculated using these new

105 temperatures and the two MgO equations of state (EOS) based on the 3rd-order Birch-  
106 Murnaghan EOS and Vinet EOS by Tange *et al.* (2009).

107 The recalculated olivine-wadsleyite phase-relation data are fitted to Stixrude's (1997)  
108 equations to express the binary loops at the above two temperatures. One essential parameter  
109 in his equations is the Fe-Mg partition coefficients between olivine and wadsleyite,  $K_D^{\text{Mg-Fe}}$ ,

$$110 \quad K_D^{\text{Mg-Fe}} = \frac{x_{\text{Fe}_2\text{SiO}_4}^{\text{Ol}} x_{\text{Mg}_2\text{SiO}_4}^{\text{Wd}}}{x_{\text{Fe}_2\text{SiO}_4}^{\text{Wd}} x_{\text{Mg}_2\text{SiO}_4}^{\text{Ol}}} \quad (2)$$

111 where  $X_i^j$  is the mole fraction of component  $i$  in phase  $j$ , and  $\Pi$  is the reduced pressure using  
112 the transition pressures of the endmembers of  $\text{Mg}_2\text{SiO}_4$  and  $\text{Fe}_2\text{SiO}_4$ ,  $P_{\text{Mg}_2\text{SiO}_4}$  and  $P_{\text{Fe}_2\text{SiO}_4}$ ,  
113 respectively as:

$$114 \quad \Pi = \frac{P - P_{\text{Fe}_2\text{SiO}_4}}{P_{\text{Mg}_2\text{SiO}_4} - P_{\text{Fe}_2\text{SiO}_4}} \quad (3)$$

115 where  $P$  is the experimental pressure. The mole fraction of wadsleyite,  $f_{\text{wd}}$ , is expressed as a  
116 function of  $\Pi$  with the parameter  $K_D^{\text{Mg-Fe}}$  as:

$$117 \quad f_{\text{wd}}(\Pi) = \frac{x_{\text{bulk}}(1 - K_D^{\text{Mg-Fe}}) - K_D^{\text{Mg-Fe}}\Pi + K_D^{\text{Mg-Fe}}}{1 - K_D^{\text{Mg-Fe}}(1 - \Pi) - K_D^{\text{Mg-Fe}}\Pi + K_D^{\text{Mg-Fe}}} \quad (4)$$

118 where  $x_{\text{bulk}}$  is the bulk mole fraction of the  $\text{Fe}_2\text{SiO}_4$  component. Fitting the experimental  
119 data to Eqs. (2) – (4) yields  $K_D^{\text{Mg-Fe}}$ ,  $P_{\text{Mg}_2\text{SiO}_4}$  and  $P_{\text{Fe}_2\text{SiO}_4}$ . Although Katsura *et al.* (2004a;  
120 2010) assumed that the difference in the endmember transition pressures,  $P_{\text{Mg}_2\text{SiO}_4}$  –  
121  $P_{\text{Fe}_2\text{SiO}_4}$ , is independent of the temperature, this assumption is not adopted in the current  
122 study.

123 Similarly to Katsura *et al.* (2010), the current study assumes that the upper mantle is  
124 pyrolytic, namely,  $X_{\text{Mg}} = 0.89$  ( $X_{\text{Mg}} = \text{Mg}/(\text{Mg}+\text{Fe})$ ) (Green and Falloon, 2009). Stixrude  
125 (1997) suggested that the D410 depth corresponds to the pressure where the olivine-to-  
126 wadsleyite ratio is 1:2, namely,  $f_{\text{wd}} = 2/3$ . The current study follows this idea to obtain the  
127 olivine-wadsleyite transition pressure in the mantle peridotite ( $P_{\text{Ol-Wd}}$ ) as a function of  
128 temperature. Comparing the pressure corresponding to the D410 depth with  $P_{\text{Ol-Wd}}$  allows  
129 estimating the D410 temperature.

130 The uncertainties in the above estimations are evaluated using Monte Carlo simulation by  
131 producing 1000 replica sets of the phase relation data. This procedure produces the replica  
132 sets' data points, including pressure and compositions of olivine and wadsleyite, as:

$$133 \quad x_i^j = x_i^0 + p \cdot \sigma_{x_i^0} \quad (5)$$

134 where  $x_i^j$  is the  $i$ -th data values in the  $j$ -th replica data set,  $x_i^0$  and  $\sigma_{x_i^0}$  are the average and  
135 standard deviation of the  $i$ -th datum in the original data set, and  $p$  is the normally distributed  
136 random number. Assuming the D410 depths of 409 and 411 km, the D410 temperatures are  
137 calculated for each replica set. The average and uncertainty of the D410 temperature are  
138 obtained from the mean value and standard deviation of the 1000 replica sets at the two D410  
139 depths.

## 140 2.2 Evaluation of the parameters for the adiabatic temperature gradient

141 The adiabatic temperature gradient with pressure is given as:

$$142 \quad \left( \frac{dT}{dP} \right)_S = \frac{\alpha T}{\rho_c C_p} \quad (6)$$

143 where  $\rho_c$  is the density of the constituent (Turcotte and Schubert, 2014). Therefore, evaluating  
 144 the adiabatic temperature gradient requires the thermal expansivity, density, and isobaric heat  
 145 capacity of the constituents as a function of pressure and temperature.

146 Following Katsura *et al.* (2010), the current study approximates the mantle  
 147 constituents by the Mg-endmembers of olivine, wadsleyite, ringwoodite, and bridgmanite at  
 148 depth ranges of 50-410, 410-520, 520-660, and 660-2800 km, respectively. Wolf *et al.* (2015)  
 149 found that the adiabatic temperature gradients of  $\text{MgSiO}_3$  and  $(\text{Mg}_{0.83}\text{Fe}_{0.17})\text{SiO}_3$  bridgmanite  
 150 are very similar, which supports the above approximation. On the other hand, although some  
 151 studies suggested that the mantle transition zone contains weight percent levels of  $\text{H}_2\text{O}$   
 152 (Pearson *et al.*, 2014; Fei *et al.*, 2017), Houser (2016) predicted that the mantle transition  
 153 zone is generally dry by combining the global analysis of long-period seismic data and the  
 154 mineral physics data. Furthermore, the mantle transition zone is only 250-km thick and  
 155 notably smaller than the whole mantle (2800 km thickness), and therefore, it should not  
 156 change the temperature profile of the whole mantle significantly even if the transition zone  
 157 were  $\text{H}_2\text{O}$ -rich.

158 The data sets for the density and thermal expansivity of olivine, wadsleyite, and  
 159 ringwoodite are taken from Katsura *et al.* (2009a), Katsura *et al.* (2009b), and Katsura *et al.*  
 160 (2004b), respectively. Those of bridgmanite are taken from Katsura *et al.* (2009c) and Tange  
 161 *et al.* (2012). The temperatures of these data sets are recalculated using Nishihara *et al.*'s  
 162 (2020) pressure correction on the thermocouple EMF. The pressures are then recalculated  
 163 based on these new temperatures using the two MgO EOS's given by Tange *et al.* (2009).

164 The  $P$ - $V$ - $T$  data set of each mineral is fitted to the Mie-Grüneisen-Debye EOS  
 165 (Jackson and Rigden, 1996) with the 3<sup>rd</sup>-order Birch-Murnaghan EOS (Katsura *et al.*, 2019).  
 166 At the standard temperature of  $T_0 = 300$  K, the 3<sup>rd</sup>-order Birch-Murnaghan EOS is expressed  
 167 as:

$$168 P(V, T_0) = \frac{3}{2} K_{T_0} \left[ \left( \frac{V_{P_0, T_0}}{V_{P, T_0}} \right)^{\frac{7}{3}} - \left( \frac{V_{P_0, T_0}}{V_{P, T_0}} \right)^{\frac{5}{3}} \right] \times \left\{ 1 - \frac{3}{4} (4 - K_{T_0}') \left[ \left( \frac{V_{P_0, T_0}}{V_{P, T_0}} \right)^{\frac{2}{3}} - 1 \right] \right\} \quad (7)$$

169 where  $P(V, T_0)$  is the pressure at volume  $V$  and the standard temperature of  $T_0$ ,  $K_{T_0}$  is the  
 170 isothermal bulk modulus at the temperature of  $T_0$ ,  $K_{T_0}'$  is its pressure derivative, and  $V_{P_0, T_0}$  and  
 171  $V_{P, T_0}$  are the volume at the standard pressure of  $P_0 = 0$  and at the pressure of  $P$ , respectively,  
 172 under the temperature condition of  $T_0$ . The Mie-Grüneisen-Debye EOS is:

$$173 P(V, T) = P(V, T_0) + \frac{\gamma}{V} [E_{\text{th}}(V_P, T) - E_{\text{th}}(V_P, T_0)] \quad (8)$$

174 where  $P(V, T)$  is the pressure at the volume  $V$  and temperature  $T$ ,  $\gamma$  is the Grüneisen  
 175 parameter, and  $E_{\text{th}}$  is the thermal energy.

176 The volume dependence of the Grüneisen parameter is expressed using the constant  $q$  as:

$$177 \gamma = \gamma_0 \left( \frac{V_{P, T}}{V_{P_0, T_0}} \right)^q \quad (9)$$

178 where  $\gamma_0$  is the Grüneisen parameter at the standard pressure  $P_0$  and temperature  $T_0$ . The  
 179 thermal energy,  $E_{\text{th}}$ , is given by:

$$180 E_{\text{th}}(V, T) = \frac{9nRT}{(\theta/T)^3} \int_0^{\theta} \left( \frac{\xi^3}{e^{\xi} - 1} \right) d\xi \quad (10)$$

181 where  $\theta$  is the Debye temperature at the volume  $V$ ,  $n$  is the number of atoms per formula unit,  
 182 and  $R$  is the gas constant. The Debye temperature at the volume  $V$  is expressed as:

$$183 \quad \theta = \theta_0 \exp \left[ \frac{\gamma_0 - \gamma(V)}{q} \right] \quad (11)$$

184 where  $\theta_0$  is the Debye temperature at the standard volume  $V_0$ .

185 With the fixed values of  $K_{T_0}$  and  $\theta_0$  given by literature (Isaak *et al.*, 1989; Mao *et al.*,  
186 2008; Higo *et al.*, 2006; Tange *et al.*, 2012; Watanabe, 1982; Akaogi and Ito, 1993), the  $K_{T_0}'$ ,  
187  $\gamma_0$ , and  $q$  are fitted to minimize the sum of squared differences of the pressures based on each  
188 mineral's EOS, Pmineral, from those based on the MgO pressure marker using Tange *et al.*'s  
189 (2009) two EOS's,  $P_{\text{MgO}}$ . Namely,  $\chi^2 = \sum(P_{\text{mineral}} - P_{\text{MgO}})^2$  is minimized.

190 The uncertainties of the fitting results of  $K_{T_0}'$ ,  $\gamma_0$ , and  $q$  are evaluated using the  
191 Bootstrap method as follows. Firstly, 1000 bootstrap data sets are produced by randomly  
192 choosing data points from the original data set with allowing duplication. Then, uncertainties  
193 multiplied by the normally distributed random numbers are added to the individual data  
194 points using Eq. (5). The three parameters of  $K_{T_0}'$ ,  $\gamma_0$ , and  $q$  are obtained by fitting the  
195 individual bootstrap data sets to the Eqs. (7) – (11). The averages and standard deviations of  
196 the three parameters are obtained from the mean values and standard deviations of the 1000  
197 bootstrap data sets. Note that the Bootstrap method is not adopted in modeling the phase  
198 relations because the number of the data points for the phase relations is too small for the  
199 bootstrap method (3 pairs of olivine and wadsleyite compositions at each temperature).

200 The density at given pressure and temperature is evaluated from the ambient density  
201 divided by the relative volumes given by Eqs. (7) – (10). The thermal expansivity is obtained  
202 from the differentiation of the relative volumes with respect to the temperature.

203 It is assumed that the isochoric heat capacity,  $C_V$ , of olivine is independent of the pressure  
204 because of the high temperatures. Therefore, the isobaric heat capacity at high pressure,  $C_P$ ,  
205 can be obtained using the isobaric heat capacity of olivine at ambient pressure,  $C_P^0$  as:

$$206 \quad C_P = (1 + \alpha\gamma T)C_V = (1 + \alpha\gamma T) \frac{C_P^0}{1 + \alpha_0\gamma_0 T} = \frac{1 + \alpha\gamma T}{1 + \alpha_0\gamma_0 T} C_P^0 \quad (12)$$

207 where  $\alpha$  and  $\alpha_0$  are the temperature-dependent thermal expansivities of olivine at high and  
208 ambient pressures, respectively. The isobaric heat capacity of olivine at ambient pressure is  
209 taken from Saxena *et al.* (1993), which was also adopted in Katsura *et al.* (2010).

210 The isobaric heat capacities of wadsleyite, ringwoodite and bridgmanite are obtained by  
211 multiplying the Debye heat capacity by  $1 + \alpha\gamma T$ .

### 212 2.3 Evaluation of adiabatic temperature profile in the Earth's mantle

213 The adiabatic temperature gradient to the depth ( $z$ ) in the mantle is written as:

$$214 \quad \left( \frac{dT}{dz} \right)_S = \rho_m g \left( \frac{dT}{dP} \right)_S \quad (13)$$

215 where  $\rho_m$  and  $g$  are the density and gravity acceleration at a certain depth of the Earth's  
216 mantle given by PREM (Dziewonski and Anderson, 1981). Note that the  $\rho_m$  in this equation  
217 is different from  $\rho_c$  in Eq. (6). The former is the density of the Earth's mantle, which consists  
218 of Fe-bearing peridotite. On the other hand, the latter is that of the Fe-free endmembers of the  
219 major mantle minerals. Katsura *et al.* (2010) did not consider this difference and simply  
220 obtained the adiabatic temperature gradient to the depth using Eq. (1), which is the major  
221 error source of Katsura *et al.* (2010).

222 Using Eq. (13), the temperatures at depths are obtained by 10-km increments and  
223 decrements from the D410 depth to deeper and shallower depths, respectively. It is well  
224 known that the latent heat associated with the phase transition of olivine to wadsleyite,

wadsleyite to ringwoodite, and ringwoodite to bridgmanite + ferropericlase abruptly changes the geotherm (Ito and Katsura, 1989, Katsura *et al.*, 2010). Following Katsura *et al.* (2010), the current study assumes the geotherm change of +60, +43, and -34 K by these transitions. Since the D410 is taken as the depth where the olivine and wadsleyite have a volume ratio of 1:2 (Stixrude, 1997), the temperature changes of +20 and -40 K are delivered above and below the D410, as Katsura *et al.* (2010) did. The wadsleyite to ringwoodite and ringwoodite to bridgmanite + ferropericlase transitions are assumed to occur at fixed depths of 520 and 660 km, respectively.

### 3 Results and discussion

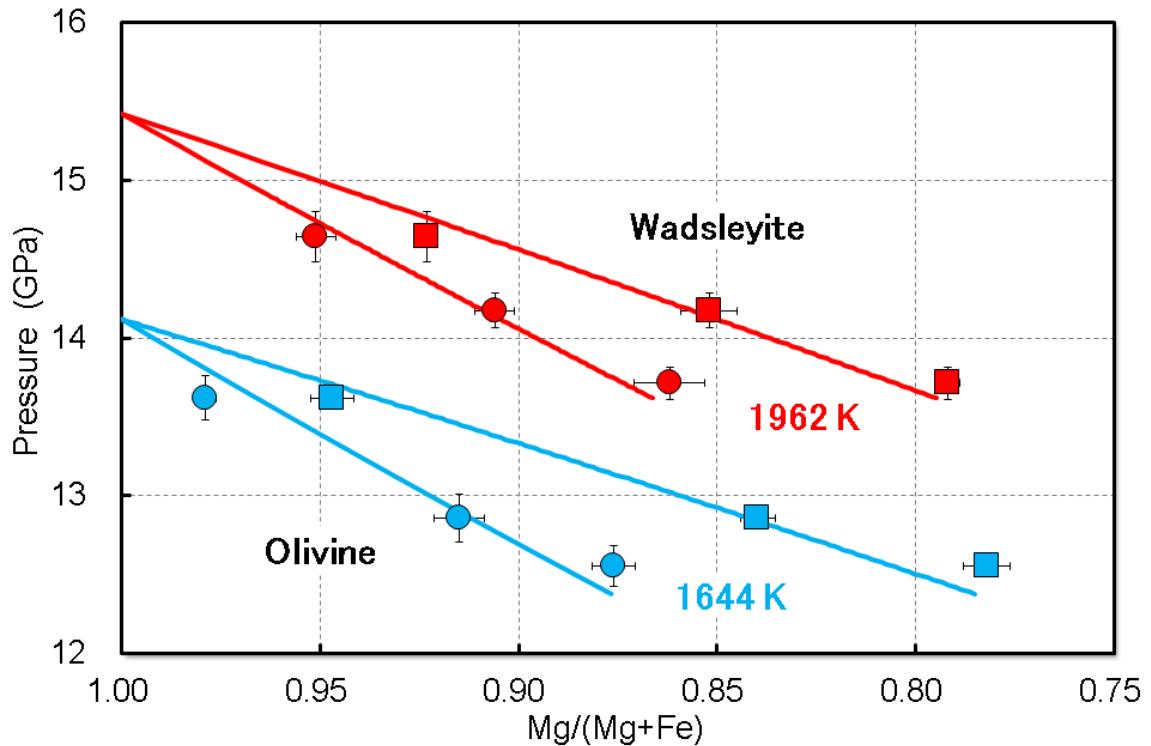
#### 3.1 Olivine-wadsleyite transition in $(\text{Mg},\text{Fe})_2\text{SiO}_4$

Table 1 presents the recalculated pressures and temperatures with the olivine and wadsleyite compositions from Katsura *et al.*'s (2004a) experiments. Figure 1 shows the recalculated phase relations of the olivine-wadsleyite transitions in  $(\text{Mg},\text{Fe})_2\text{SiO}_4$ . The Nishihara *et al.*'s (2020) pressure correction of thermocouple EMF has increased 1600 and 1900 K to 1644 and 1962 K, respectively. This correction has accordingly increased the pressures by 0.22 and 0.32 GPa at these temperatures, respectively. The fitting yields  $K_D^{\text{Mg}-\text{Fe}} = 0.514 \pm 0.035$ ,  $PM_{\text{Mg}_2\text{SiO}_4} = 14.12 \pm 0.13$  GPa, and  $P_{\text{Fe}_2\text{SiO}_4} = 3.6 \pm 1.0$  GPa at 1962 K and  $K_D^{\text{Mg}-\text{Fe}} = 0.598 \pm 0.036$ ,  $PM_{\text{Mg}_2\text{SiO}_4} = 15.42 \pm 0.17$  GPa, and  $P_{\text{Fe}_2\text{SiO}_4} = 4.7 \pm 1.0$  GPa at 1962 K. Consequently, the Clapeyron slopes of the Mg and Fe endmember transitions are  $4.1 \pm 0.5$  and  $3.4 \pm 3.6$  MPa/K, respectively. The uncertainties of  $P_{\text{Fe}_2\text{SiO}_4}$  estimation (1.0 and 1.0 GPa) are 6~8 times larger than those of  $PM_{\text{Mg}_2\text{SiO}_4}$  (0.13 and 0.17 GPa). This is because Katsura *et al.* (2004a) conducted the experiments with the bulk compositions around  $X_{\text{Mg}} = 0.9$ . The estimated  $PM_{\text{Mg}_2\text{SiO}_4}$  and  $P_{\text{Fe}_2\text{SiO}_4}$  are anticorrelated, as expected (Fig. S1).

**Table 1.** Temperature, Pressure, and Phase Compositions of Coexisting Olivine and Wadsleyite.

Run #	$V/V_0^{\text{MgO}}$	Tempera-ture before Correc-tion (K)	Pressure before Correc-tion (GPa)	Tempera-ture after Correc-tion (K)	Pressure after Correc-tion (GPa)	$X_{\text{Mg}}^{\text{Ol}}$	$X_{\text{Mg}}^{\text{Wd}}$
733	0.9738(6)	1900	14.24(16)	1964	14.64(16)	0.951(5)	0.923(1)
734	0.9763(3)	1900	13.78(11)	1962	14.18(11)	0.906(5)	0.852(7)
735	0.9788(3)	1900	13.33(10)	1961	13.71(10)	0.862(9)	0.792(3)
763	0.9741(5)	1600	12.29(13)	1643	12.56(13)	0.876(6)	0.782(6)
779	0.9785(5)	1600	13.33(14)	1645	13.62(14)	0.979(3)	0.947(6)
780	0.9725(6)	1600	12.58(15)	1643	12.86(15)	0.915(6)	0.840(4)

The initial data are from Katsura *et al.* (2004a).



**Figure 1.** Binary phase relations of the olivine-wadsleyite transition in  $(\text{Mg},\text{Fe})_2\text{SiO}_4$ . The original data are from Katsura *et al.* (2004a). The temperatures were recalculated using Nishihara *et al.*'s (2020) pressure correction of the EMF-temperature relations of  $\text{W}_{97}\text{Re}_{3}\text{-We}_{75}\text{Re}_{25}$  thermocouple. The temperatures of 1600 and 1900 K in Katsura *et al.* (2004a) became 1644 (blue) and 1962 (red) K. The circle and square symbols denote the compositions and pressures of coexisting olivine and wadsleyite, respectively.

Yagi *et al.* (1987) determined the fayalite-ahrensite transition in  $\text{Fe}_2\text{SiO}_4$  using *in situ* X-ray diffraction with a multianvil press. Their data suggested  $P_{\text{Fe}_2\text{SiO}_4}$ 's of  $7.0 \pm 0.1$  and  $6.2 \pm 0.1$  GPa at 1644 and 1962 K, respectively, which are significantly higher than those estimated in the current study. Hence, the binary loop should have an old-crescent shape. However, these curvatures should be insignificant for the current study because the compositional range of Katsura *et al.* (2004a) already covers the mantle composition ( $X_{\text{Mg}} = 0.89$  (Green and Falloon, 1998)) (Fig. 1).

### 3.2 Temperature at the 410-km discontinuity

The olivine-wadsleyite transition pressures in peridotite,  $P_{\text{Ol-Wd}}$ , are found to be  $13.09 \pm 0.07$  and  $14.34 \pm 0.08$  GPa at temperatures of 1644 and 1962 K, respectively. The smaller uncertainties in  $P_{\text{Ol-Wd}}$  than  $P_{\text{Mg}_2\text{SiO}_4}$  are because Katsura *et al.* (2004a) conducted the experiments with the bulk compositions around  $X_{\text{Mg}} = 0.9$  and not at  $X_{\text{Mg}} = 1.0$ , or equivalently because the  $P_{\text{Mg}_2\text{SiO}_4}$  and  $P_{\text{Fe}_2\text{SiO}_4}$  are anticorrelated.

The temperature dependence of  $P_{\text{Ol-Wd}}$ , is  $dP_{\text{Ol-Wd}}/dT = 3.9 \pm 0.3$  MPa/K. This temperature dependence is significant because the temperature increase by 100 K causes the olivine-wadsleyite transition pressure to increase by  $9.8 \pm 0.7$  km. Comparing  $P_{\text{Ol-Wd}}$  with the pressures of the D410 depths suggests the temperature at D410 to be  $1805 \pm 18$  K. That estimated by Katsura *et al.* (2010),  $1830 \pm 48$  K, was slightly higher than the present

estimation. The reason for this difference is the thermocouple correction, which raises not only temperatures but also pressures of the experimental data points. The ratio of the pressure increase to the temperature increase is equal to the thermal pressure of MgO, which is 6.4 MPa/K at around 13 GPa and 1900~1600 K based on Tange *et al.* (2009). On the other hand, the temperature dependence of  $P_{\text{Ol-Wd}}$  is smaller, *i.e.*,  $3.9 \pm 0.3$  MPa/K. As a result, the required temperature for the olivine-wadsleyite transition becomes lower by the thermocouple correction. Also, note that the uncertainty in the temperature estimation is smaller in the current study than in Katsura *et al.* (2010). This smaller uncertainty is because Katsura *et al.* (2010) did not consider the anticorrelation of  $P_{\text{Mg}_2\text{SiO}_4}$  and  $P_{\text{Fe}_2\text{SiO}_4}$ .

### 3.3 $P$ - $V$ - $T$ relations of major mantle minerals

Tables S1-S4 show the recalculated  $P$ - $V$ - $T$  data of the four minerals. Table 2 presents the optimized values together with the assumed parameters. Figures S2-S5 show the comparisons of pressures obtained using the Tange *et al.*'s (2009) MgO EOSs and the EOS of each mineral in the current study.

The obtained  $K'$  of the four minerals are identical within the uncertainties (Table 2). A higher-pressure mineral has a larger  $\gamma_0$ . This is reasonable because, according to the definition,  $\gamma_0$  is the rate of the pressure increase to the thermal energy increase at constant volume and ambient pressure. Since the thermal expansivity of minerals vastly increases with decreasing pressure below their stability fields, the pressure increase rate caused by the energy increase should be more significant at ambient pressure in higher-pressure minerals. However, the relation between the stability field and  $q$  among different minerals is unclear due to the significant uncertainties. Although ringwoodite apparently has a larger  $q$  than wadsleyite ( $2.9 \pm 0.7$  and  $1.5 \pm 1.1$ ), this difference is within the uncertainties. As expected, the  $\gamma_0$  and  $q$  are strongly correlated, especially for wadsleyite and ringwoodite due to their narrow experimental pressure ranges. (Fig. S6). Consequently, the uncertainty in  $q$  has a relatively small effect on the estimations of density and thermal expansivity.

**Table 2.** Thermoelastic Parameters of the Major Mantle Minerals

Mineral	$K_{T_0}$ (GPa)	$K_{T_0}'$	$\theta$ (K)	$\gamma_0$	$q$
Olivine	127.4 <sup>a</sup>	4.2(4)	768 <sup>a</sup>	1.00(2)	2.4(5)
Wadsleyite	169.2 <sup>b</sup>	4.2(2)	814 <sup>c</sup>	1.23(6)	1.5(11)
Ringwoodite	182.0 <sup>d</sup>	4.8(4)	830 <sup>c</sup>	1.34(6)	2.9(7)
Bridgmanite	256.7 <sup>e</sup>	4.09(4)	1030 <sup>f</sup>	1.53(3)	1.6(4)

*a*: Isaak *et al.* (1989); *b*: Mao *et al.* (2008); *c*: Watanabe *et al.* (1982); *d*: converted from  $K_S$  given by Higo *et al.* (2006) *e*: Tange *et al.* (2012); *f*: Akaogi and Ito (1993)

The obtained  $K'$  of the four minerals are identical within the uncertainties (Table 2). A higher-pressure mineral has a larger  $\gamma_0$ . This is reasonable because, according to the definition,  $\gamma_0$  is the rate of the pressure increase to the thermal energy increase at constant volume and ambient pressure. Since the thermal expansivity of minerals vastly increases with decreasing pressure below their stability fields, the pressure increase rate caused by the energy increase should be more significant at ambient pressure in higher-pressure minerals.

However, the relation between the stability field and  $q$  among different minerals is unclear due to the significant uncertainties. Although ringwoodite apparently has a larger  $q$  than wadsleyite ( $2.9 \pm 0.7$  and  $1.5 \pm 1.1$ ), this difference is within the uncertainties. As

320 expected, the  $\gamma_0$  and  $q$  are strongly correlated, especially for wadsleyite and ringwoodite due  
321 to their narrow experimental pressure ranges. (Fig. S6). Consequently, the uncertainty in  $q$   
322 has a relatively small effect on the estimations of density and thermal expansivity.

323

324        3.4 Adiabatic temperature profile in the Earth's mantle

325        Table 3 shows the adiabatic temperature profile in the mantle at depths from 50 to  
326 2800 km obtained in the current study. Although the temperature profile in the real mantle  
327 should not be adiabatic in the lower-most region due to the thermal boundary layer, the  
328 present profile is constructed by ignoring the presence of the thermal boundary layer.  
329 Supplementary Table S5 shows various thermoelastic parameters at depths from 50 to 2800  
330 km with a 10-km step. Among the shown parameters in Table S5, the density, gravitational  
331 acceleration, and pressure are taken from PREM (Dziewonski and Anderson, 1981). The  
332 other parameters are obtained through the procedure described above. The thermal  
333 expansivity, adiabatic temperature gradient, and adiabatic geotherm are plotted in Fig. 2, 3,  
334 and 4.

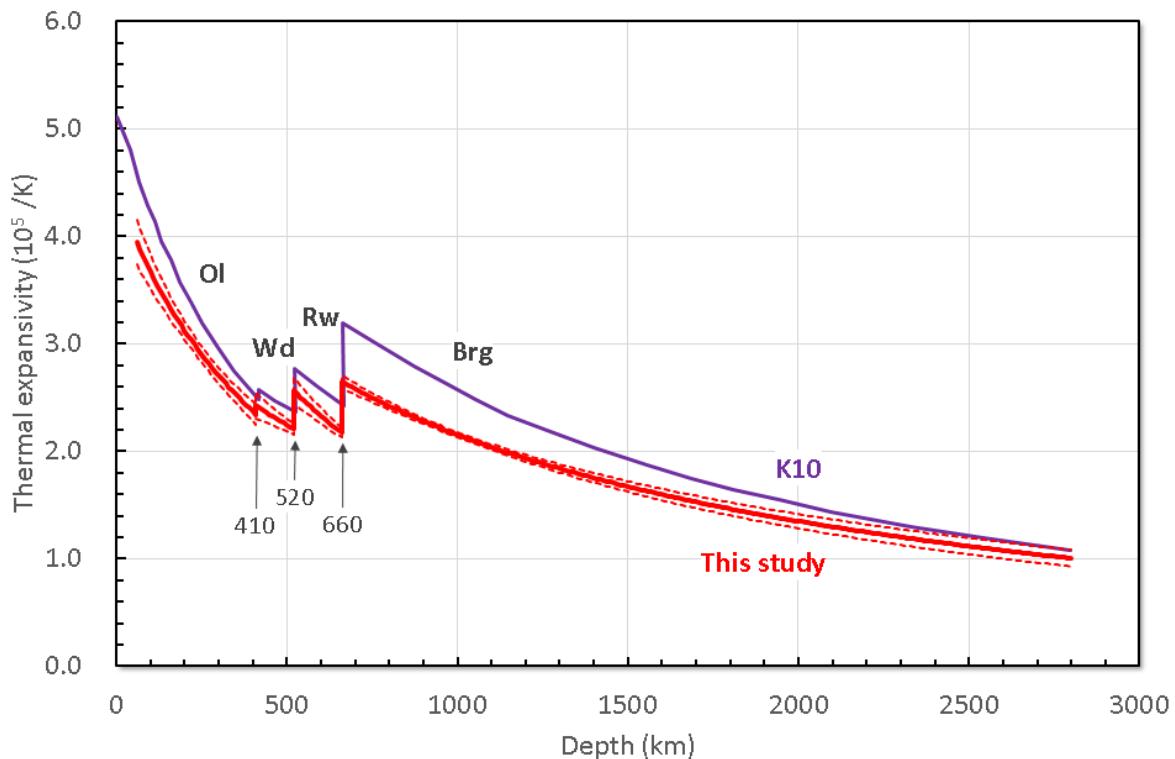
335

336 **Table 3.** Adiabatic Temperature Profile in the Earth's Mantle

Depth (km)	Pressure (GPa)	Temperature (K)	dT/dz (K/km)
50	1.5	1616(17)	0.54
70	2.1	1627(17)	0.52
90	2.8	1637(17)	0.50
100	3.1	1642(17)	0.50
120	3.8	1651(17)	0.48
140	4.4	1660(17)	0.46
160	5.1	1669(17)	0.45
180	5.8	1678(17)	0.44
200	6.4	1686(17)	0.43
220	7.1	1695(17)	0.42
240	7.8	1703(18)	0.41
260	8.5	1711(18)	0.40
280	9.2	1718(18)	0.39
300	9.9	1726(18)	0.38
320	10.6	1734(18)	0.38
340	11.2	1741(18)	0.37
360	11.9	1748(18)	0.36
380	12.6	1755(18)	0.35
400	13.4	1762(18)	0.34
410	13.7	1765(18)	0.34
410	13.7	1825(18)	0.35
420	14.1	1829(18)	0.36
440	14.9	1836(18)	0.35
460	15.6	1843(18)	0.35
480	16.4	1850(18)	0.35
500	17.1	1857(18)	0.34

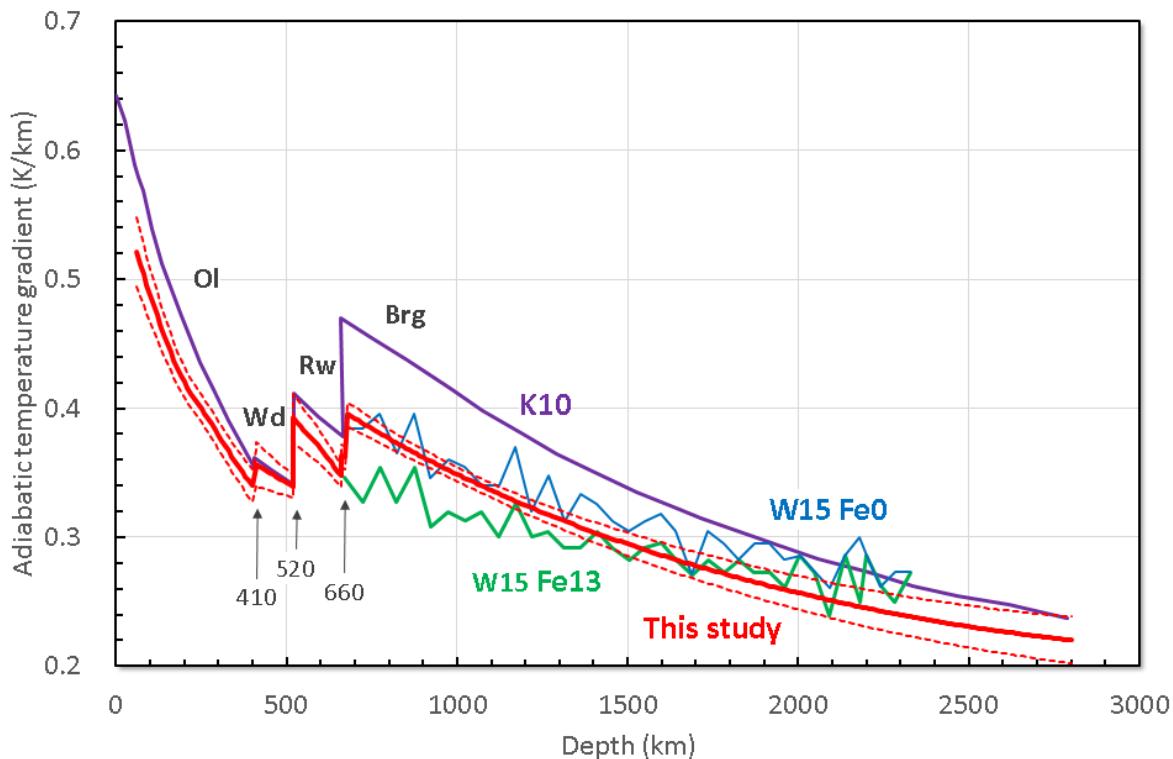
520	17.9	1864(18)	0.34
520	17.9	1907(18)	0.34
540	18.7	1915(18)	0.39
560	19.5	1922(18)	0.38
580	20.3	1930(18)	0.38
600	21.0	1937(19)	0.37
620	21.8	1945(19)	0.37
640	22.6	1952(19)	0.36
660	23.4	1959(19)	0.35
660	23.4	1925(19)	0.35
700	25.2	1940(19)	0.39
800	29.6	1979(20)	0.38
1000	38.6	2051(21)	0.35
1200	47.8	2119(21)	0.33
1400	57.3	2182(22)	0.31
1600	66.9	2241(23)	0.29
1800	76.8	2296(24)	0.27
2000	86.9	2349(25)	0.26
2200	97.3	2399(26)	0.25
2400	108.0	2447(28)	0.24
2600	119.0	2494(30)	0.23
2800	130.4	2538(32)	0.22

337

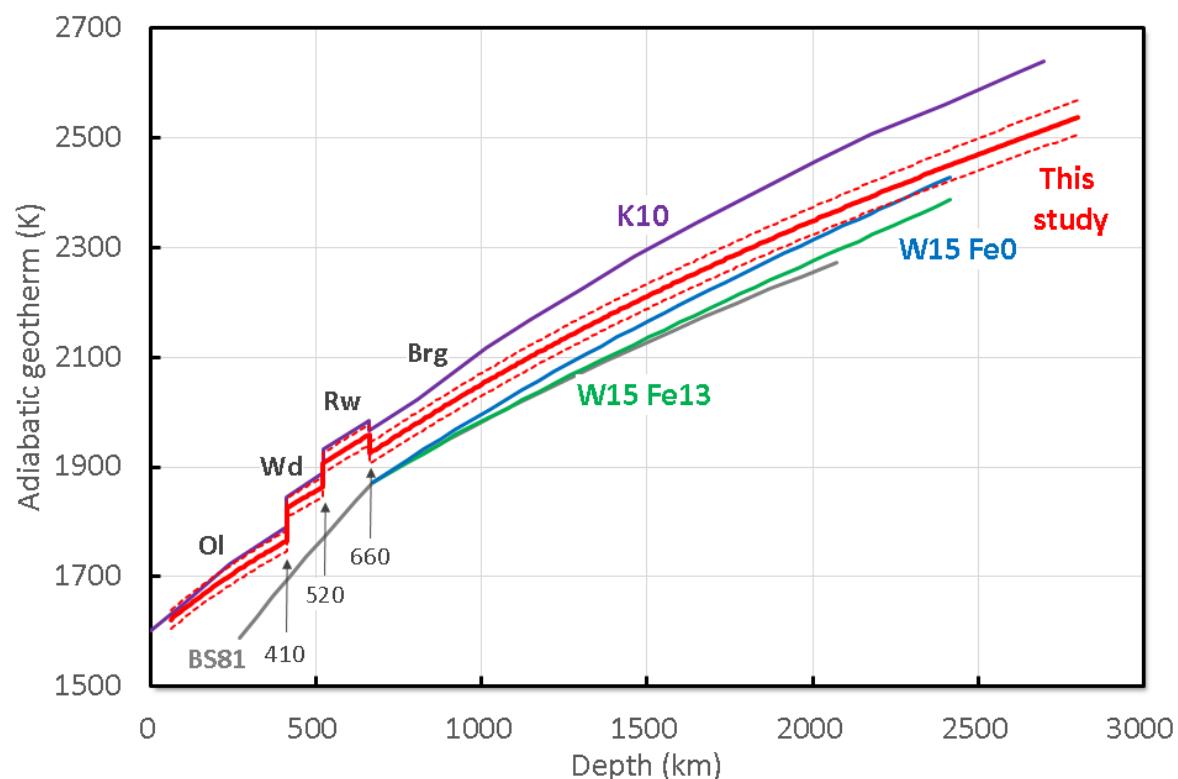


338

339 **Figure 2.** Thermal expansivity in the mantle. Red: the current study,  
340 violet: Katsura *et al.* (2010). The solid and dashed curves show the most probably and 68% confidence intervals,  
341 respectively.



344 **Figure 3.** Adiabatic temperature gradients in the mantle. Red: the current study; violet:  
 345 Katsura *et al.* (2010); blue:  $\text{MgSiO}_3$  bridgemanite lower mantle by Wolf *et al.* (2015); green:  
 346  $\text{Mg}_{0.87}\text{Fe}_{0.13}\text{SiO}_3$  bridgemanite lower mantle by Wolf *et al.* (2015). The solid and dashed  
 347 curves show the most probably and 68% confidence intervals, respectively.  
 348



350 **Figure 4.** Adiabatic temperature profiles in the mantle. Red: the current study; violet: Katsura  
351 *et al.* (2010); blue: MgSiO<sub>3</sub> bridgemanite lower mantle by Wolf *et al.* (2015); green:  
352 Mg<sub>0.87</sub>Fe<sub>0.13</sub>SiO<sub>3</sub> bridgemanite lower mantle by Wolf *et al.* (2015); grey: Brown and  
353 Shankland (1981). The solid and dashed curves show the most probable and 68% confidence  
354 intervals, respectively.  
355

356 As Katsura *et al.* (2010) argued, the thermal expansivity continuously decreases with  
357 increasing depth due to the negative temperature dependence when no phase transition  
358 occurs, whereas the phase transitions increase the thermal expansivity (Fig. 3). In the olivine  
359 stability field, the thermal expansivity decreases from  $(4.0 \pm 0.2) \times 10^{-5}$  to  $(2.3 \pm 0.1) \times 10^{-5} \text{ K}^{-1}$   
360 down to 410-km depth. In the wadsleyite and ringwoodite stability fields, it decreases from  
361  $(2.4 \pm 0.1) \times 10^{-5}$  to  $(2.2 \pm 0.5) \times 10^{-5}$  and from  $(2.6 \pm 0.1) \times 10^{-5}$  to  $(2.2 \pm 0.5) \times 10^{-5} \text{ K}^{-1}$ , respectively,  
362 at 410-520 and 520-660 km depths. In the bridgemanite stability field, it decreases from  
363  $(2.7 \pm 0.5) \times 10^{-5}$  to  $(1.0 \pm 0.1) \times 10^{-5} \text{ K}^{-1}$  to 2800-km depth. These values are about 10 % smaller  
364 than those of Katsura *et al.* (2010). Since the data sources are very similar between the  
365 current study and Katsura *et al.* (2010), Katsura *et al.* (2010) could have miscalculated the  
366 thermal expansivity.

367 Since the other parameters for the adiabatic temperature gradient given by Eq. (13)  
368 other than the thermal expansivity do not vary significantly, the adiabatic temperature  
369 gradient varies similarly to the thermal expansivity (Figs. 2 and 3). The adiabatic temperature  
370 gradient continuously decreases with depth without a phase transition, whereas it increases  
371 when a phase transition occurs. In the olivine stability field, the adiabatic gradient decreases  
372 from  $(0.53 \pm 0.03)$  to  $(0.35 \pm 0.01) \text{ K km}^{-1}$  down to 410-km depth. In the wadsleyite and  
373 ringwoodite stability fields, it decreases from  $0.36 \pm 0.02$  to  $0.34 \pm 0.01$  and from  $0.39 \pm 0.02$  to  
374  $0.35 \pm 0.01 \text{ K km}^{-1}$ , respectively, at 410-520 and 520-660 km depths. In the bridgemanite  
375 stability field, it decreases from  $0.36 \pm 0.01$  to  $0.22 \pm 0.02 \text{ K km}^{-1}$  to 2800-km depth. As  
376 expected, the current adiabatic gradient is smaller than Katsura *et al.* (2010), similarly to the  
377 thermal expansivity. On the other hand, the adiabatic gradient in the lower mantle from the  
378 current study agrees very well to those estimated by Wolf *et al.* (2015), which has gradients  
379 of  $0.35\text{--}0.39 \text{ K/km}$  at the top of the lower mantle and decrease to  $0.25\text{--}0.27 \text{ K/km}$  at 2300-km  
380 depth. Turcotte and Schubert (2014) suggested an adiabatic temperature gradient of  $0.3 \text{ K}$   
381  $\text{km}^{-1}$  in the mantle. This value may be a good approximation in the regions from the bottom  
382 of the upper mantle to the mid-mantle, but it may be underestimated for the uppermost mantle  
383 and overestimated for the deeper regions of the lower mantle.

384 According to the above argument, the temperatures are  $1765 \pm 18$  and  $1825 \pm 18 \text{ K}$ ,  
385 respectively, just above and below the D410,  $1864 \pm 18$  and  $1907 \pm 18 \text{ K}$ , respectively, just  
386 above and below 520-km depth, and  $1959 \pm 19$  and  $1925 \pm 19 \text{ K}$ , respectively, just above and  
387 below 660-km depth (Table 2, Fig. 4). The temperature at 50-km depth is  $1616 \pm 17 \text{ K}$ ,  
388 whereas that at 2800-km depth is  $2538 \pm 32 \text{ K}$ . An extrapolation of the temperature at 50-km  
389 depth to the surface using the temperature gradient at 50-km depth ( $0.53 \text{ K km}^{-1}$ ) yields the  
390 mantle potential temperature of  $1643 \text{ K}$  ( $1370^\circ\text{C}$ ). These temperatures are lower than those of  
391 Katsura *et al.* (2010). One factor causing the lower temperature profile is that the temperature  
392 at D410 obtained in the current study is by 25 K lower than that of Katsura *et al.* (2010).  
393 More importantly, the temperature gradients of the current study are smaller than Katsura *et*  
394 *al.* (2010) (Fig. 3). As a result, the current study indicates *ca.* 120 K lower temperature at  
395 2700 km depth than Katsura *et al.* (2010) (Fig. 4). Wolf *et al.* (2015) reported an adiabatic  
396 temperature profile in the lower mantle down to 2400-km depth. He suggested a temperature  
397 of 2390 K at 2400-km depth, which is 60 K lower than the current study (2450 K). This is  
398 because Wolf *et al.* (2015) arbitrarily assumed the temperature at 670-km depth of 1873 K,

which is ca. 60 K higher than the present estimation ( $1929 \pm 29$  K) (Fig. 4). Brown and Shankland (1981) estimated the adiabatic temperature profile and reported 90 K lower temperature at 2070 km depth ( $2359 \pm 26$  K) than the current study (Fig. 5), which is also caused by the assumption of 1873 K at the top of the lower mantle in Brown and Shankland (1981).

Sarafian *et al.* (2017) determined the solidus temperature of mantle peridotite with 140 wt. ppm of H<sub>2</sub>O at a pressure of 1.5 GPa, corresponding to 50 km depth. Their solidus temperature was  $1590 \pm 10$  K, which is slightly lower than that of the present profile,  $1617 \pm 16$  K. It is possible that the peridotite melting may start in the regions beneath mid-oceanic ridges slightly deeper than 50-km depth, and the melt finally separates from the source rocks at around 50 km depth.

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414 787527). There is no conflict of interest.  
415

#### 416 Open Research

417 The P-V-T data used in this study after the EMF correction using Nishihara *et al.* (2020) and  
418 the pressure calculation using Tange *et al.* (2009) are given at  
419 <https://doi.org/10.5281/zenodo.5644426>. The Matlab scripts to fit the P-V-T data to the  
420 equations of state, the temperature at the 410-km discontinuity, and calculate the adiabatic  
421 temperature profile are given at <https://doi.org/10.5281/zenodo.5644430>.

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536 **Supplementary information**

537 **Table S1.**  $P$ - $V$ - $T$  Data of Olivine

538 **Table S2.**  $P$ - $V$ - $T$  Data of Wadsleyite

539 **Table S3.**  $P$ - $V$ - $T$  Data of Ringwoodite

540 **Table S4.**  $P$ - $V$ - $T$  Data of Bridgmanite

541 **Table S5.** Thermoelastic Parameters Of The Adiabatic Mantle

542

543 **Figure S1.** Correlation of the Endmember Transition Pressures of the Olivine-Wadsleyite  
544 Transition.

545 **Figure S2.** Comparison of the Pressures using Tange *et al.*'s [2009] Mgo and the Current  
546 Study's Olivine EOS's.

547 **Figure S3.** Comparison of the Pressures using Tange *et al.*'s [2009] Mgo and the Current  
548 Study's Wadsleyite EOS's.

549 **Figure S4.** Comparison of the Pressures using Tange *et al.*'s [2009] Mgo and the Current  
550 Study's Ringwoodite EOS's.

551 **Figure S5.** Comparison of the Pressures using Tange *et al.*'s [2009] Mgo and the Current  
552 Study's Bridgmanite EOS's.

553 **Figure S6.** The Correlations of  $\gamma_0$  and  $q$ .

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556 **Table S1.**  $P$ - $V$ - $T$  data of olivine. The initial data is from Katsura et al. [2009a], and the pressure and tempeatures are recalculated in this study.

Data #	T before correction (K)	$V/V_0^{\text{MgO}}$	$P^{\text{MgO}}$ before (GPa)	T after correction (K)	$V/V_0^{\text{Ol}}$	$P^{\text{MgO}}$ after correction (GPa)	$P^{\text{Ol}}$ (GPa)	$P^{\text{MgO}} -$ $P^{\text{Ol}}$ (GPa)
M371008	1300	1.0371(2)	0.56(2)	1302	1.0300(3)	0.57(2)	0.37	-0.08
M371011	1100	1.0299(2)	0.27(3)	1101	1.0255(3)	0.27(3)	0.02	0.25
M371014	900	1.0227(1)	0.04(2)	900	1.0193(3)	0.04(2)	-0.10	0.14
M371017	700	1.0148(1)	-0.05(2)	700	1.0125(3)	-0.05(2)	-0.11	0.06
M371020	500	1.0069(1)	-0.02(2)	500	1.0051(3)	-0.02(2)	0.03	-0.05
M371023	304	0.9998(1)	0.06(2)	304	0.9996(3)	0.06(2)	0.07	-0.01
M371029	1300	1.0363(2)	0.66(3)	1303	1.0291(4)	0.67(3)	0.47	0.21
M371033	1300	1.0334(2)	1.05(3)	1304	1.0265(5)	1.07(3)	0.74	0.33
M371036	1500	1.0404(4)	1.39(6)	1507	1.0297(4)	1.43(6)	1.25	0.18
M371039	1700	1.0496(1)	1.49(2)	1709	1.0368(4)	1.54(2)	1.37	0.18
M375006	1300	1.0148(3)	3.66(6)	1312	0.9925(6)	3.74(6)	4.71	-0.97
M375009	1500	1.0250(2)	3.46(4)	1515	1.0047(4)	3.55(4)	4.01	-0.46
M375012	1300	1.0178(1)	3.23(3)	1311	1.0027(5)	3.30(3)	3.45	-0.15
M375015	1500	1.0236(2)	3.64(3)	1515	1.0061(5)	3.74(3)	3.85	-0.11
M375018	1300	1.0183(1)	3.15(3)	1311	1.0031(5)	3.22(3)	3.39	-0.18
M375021	1100	1.0134(1)	2.62(3)	1107	1.0013(4)	2.66(3)	2.81	-0.15
M375024	900	1.0090(1)	2.05(3)	904	0.9995(4)	2.07(3)	2.24	-0.17
M375027	700	1.0048(1)	1.47(2)	702	0.9985(4)	1.48(3)	1.60	-0.12
M375030	500	1.0011(1)	0.89(3)	501	0.9976(4)	0.90(3)	0.97	-0.07

M375033	305	0.9975(1)	0.42(2)	305	0.9967(4)	0.42(2)	0.43	-0.01
M375037	1500	1.0009(3)	7.06(7)	1526	0.9798(5)	7.22(7)	7.19	0.03
M375040	1700	1.0107(2)	6.80(4)	1731	0.9881(6)	6.99(4)	6.89	0.11
M375043	1500	1.0058(2)	6.28(4)	1524	0.9854(5)	6.42(4)	6.44	-0.01
M375046	1300	1.0008(2)	5.82(4)	1318	0.9825(5)	5.93(4)	6.02	-0.10
M375049	1100	0.9963(1)	5.30(3)	1112	0.9812(5)	5.37(3)	5.42	-0.05
M375052	900	0.9923(1)	4.71(3)	908	0.9798(5)	4.76(3)	4.84	-0.08
M375055	700	0.9883(1)	4.17(3)	705	0.9787(5)	4.20(3)	4.25	-0.05
M375058	500	0.9848(1)	3.59(2)	503	0.9777(5)	3.61(2)	3.66	-0.05
M375061	308	0.9815(1)	3.15(3)	309	0.9771(5)	3.16(3)	3.12	0.03
M375067	1700	0.9838(2)	11.18(7)	1745	0.9570(5)	11.46(7)	11.27	0.19
M375070	1900	0.9923(2)	10.99(7)	1954	0.9636(6)	11.33(7)	11.07	0.26
M375073	1700	0.9867(3)	10.66(8)	1743	0.9596(5)	10.93(8)	10.87	0.07
M375076	1500	0.9830(2)	10.05(7)	1533	0.9578(5)	10.26(7)	10.35	-0.10
M375079	1300	0.9786(2)	9.57(6)	1325	0.9554(5)	9.73(6)	9.96	-0.23
M375082	1100	0.9747(2)	9.04(6)	1119	0.9541(5)	9.15(6)	9.40	-0.25
M375085	900	0.9699(2)	8.68(5)	913	0.9528(5)	8.76(5)	8.88	-0.12
M375088	700	0.9663(2)	8.14(5)	709	0.9519(5)	8.19(5)	8.30	-0.10
M375091	500	0.9628(2)	7.65(4)	505	0.9508(4)	7.68(4)	7.77	-0.09
M375094	309	0.9598(2)	7.23(4)	312	0.9502(4)	7.24(4)	7.27	-0.03
M375100	1500	0.9679(3)	12.81(10)	1539	0.9417(5)	13.06(10)	12.94	0.11
M375103	1700	0.9729(3)	13.15(11)	1750	0.9448(6)	13.47(11)	13.19	0.28
M375106	1900	0.9840(2)	12.40(8)	1958	0.9534(6)	12.77(8)	12.60	0.16
M375109	1700	0.9801(2)	11.83(8)	1746	0.9514(5)	12.12(8)	12.14	-0.01

M375112	1500	0.9756(2)	11.38(7)	1536	0.9491(5)	11.61(8)	11.73	-0.12
M375115	1300	0.9710(2)	10.98(7)	1328	0.9476(5)	11.16(7)	11.20	-0.04
M375118	1100	0.9668(2)	10.51(6)	1121	0.9457(5)	10.64(7)	10.76	-0.12
M375121	900	0.9627(2)	10.07(6)	915	0.9443(5)	10.16(6)	10.27	-0.11
M375124	700	0.9590(2)	9.58(5)	710	0.9433(5)	9.63(5)	9.71	-0.07
M375127	500	0.9553(1)	9.15(4)	506	0.9422(4)	9.18(4)	9.22	-0.03
M375130	309	0.9520(1)	8.81(3)	312	0.9402(4)	8.83(3)	8.97	-0.14
M375135	1700	0.9716(4)	13.37(12)	1750	0.9444(5)	13.69(13)	13.26	0.43

557 The original data are from Katsura *et al.* [2009a].  
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560**Table S2.**  $P$ - $V$ - $T$  data of wadsleyite. . The initial data is from Katsura et al. [2009b], and the pressure and tempeatures are recalculated in this study.

Data #	T before correction (K)	$V/V_0^{\text{MgO}}$	$P^{\text{MgO}}$ before correction (GPa)	T after correction (K)	$V/V_0^{\text{Wd}}$	$P^{\text{MgO}}$ after correction (GPa)	$P^{\text{Wd}}$ (GPa)	$P^{\text{MgO}} - P^{\text{Wd}}$ (GPa)
M306028	1500	0.9650(1)	13.38(8)	1540	0.9566(2)	13.64(9)	13.72	-0.08
M306030	1500	0.9661(1)	13.17(8)	1540	0.9572(1)	13.42(9)	13.58	-0.16
M306033	1600	0.9697(1)	13.08(8)	1645	0.9602(2)	13.37(8)	13.46	-0.09
M306041	1700	0.9681(1)	14.05(9)	1752	0.9578(1)	14.38(9)	14.46	-0.08
M306044	1600	0.9649(1)	14.03(9)	1647	0.9556(1)	14.32(9)	14.41	-0.09
M306047	1500	0.9623(1)	13.90(9)	1541	0.9539(1)	14.17(9)	14.30	-0.13
M306050	1400	0.9599(1)	13.75(9)	1437	0.9523(1)	13.98(9)	14.15	-0.17
M306053	1300	0.9577(1)	13.56(8)	1332	0.9509(1)	13.76(9)	13.97	-0.21
M306056	1100	0.9531(1)	13.23(8)	1124	0.9483(1)	13.38(8)	13.57	-0.19
M306059	900	0.9486(1)	12.92(7)	917	0.9459(1)	13.03(7)	13.17	-0.14
M306062	700	0.9445(1)	12.57(5)	712	0.9437(1)	12.64(6)	12.75	-0.11
M306065	500	0.9407(1)	12.23(4)	507	0.9419(1)	12.27(4)	12.29	-0.02
M306068	309	0.9376(1)	11.92(4)	313	0.9405(1)	11.94(4)	11.87	0.07
M306071	308	0.9349(10)	12.53(26)	312	0.9383(1)	12.54(26)	12.38	0.17
M306073	1700	0.9622(1)	15.19(10)	1754	0.9524(2)	15.54(11)	15.61	-0.07
M306078	316	0.9313(1)	13.38(4)	320	0.9336(1)	13.40(4)	13.53	-0.13
M307010	1500	0.9493(0)	16.55(11)	1546	0.9476(1)	16.84(11)	16.66	0.03
M307013	1700	0.9565(1)	16.33(11)	1757	0.9441(1)	16.69(12)	16.45	-0.20
M307016	1500	0.9521(1)	15.96(11)	1545	0.9418(1)	16.24(11)	15.99	-0.17

M307019	1300	0.9478(1)	15.60(10)	1335	0.9395(1)	15.82(10)	15.57	-0.16
M307022	1100	0.9436(1)	15.25(9)	1126	0.9374(1)	15.41(9)	15.11	-0.12
M307025	900	0.9396(1)	14.87(7)	919	0.9353(1)	14.99(8)	14.70	-0.11
M307028	700	0.9358(1)	14.51(6)	713	0.9336(1)	14.58(6)	14.23	-0.05
M307031	500	0.9323(1)	14.13(5)	508	0.9324(1)	14.17(5)	13.77	0.08
M307034	306	0.9292(1)	13.83(3)	310	0.9524(1)	13.85(3)	14.61	-0.16
M307039	1500	0.9609(0)	14.19(9)	1542	0.9566(1)	14.45(9)	14.22	-0.02
M307042	1600	0.9656(1)	13.90(9)	1646	0.9556(1)	14.19(9)	13.92	-0.08
M307045	1500	0.9639(1)	13.58(9)	1541	0.9530(1)	13.84(9)	13.50	-0.15
M307048	1300	0.9597(1)	13.16(8)	1331	0.9506(1)	13.36(8)	13.06	-0.16
M307051	1100	0.9555(1)	12.75(7)	1123	0.9486(1)	12.90(7)	12.57	-0.17
M307054	900	0.9516(1)	12.30(6)	917	0.9467(1)	12.40(6)	12.06	-0.03
M307057	700	0.9474(1)	11.96(5)	711	0.9451(1)	12.03(5)	11.54	0.00
M307060	500	0.9440(1)	11.50(4)	507	0.9439(1)	11.54(4)	11.08	0.08
M307063	305	0.9411(1)	11.14(3)	309	0.9388(1)	11.16(3)	17.67	-0.13
M308012	1500	0.9461(1)	17.24(12)	1547	0.9328(1)	17.54(12)	14.42	-0.10
M308014	1500	0.9468(1)	17.08(12)	1547	0.9322(1)	17.38(12)	13.84	0.06
M308017	1600	0.9500(1)	17.04(12)	1653	0.9408(1)	17.38(12)	18.19	-0.04
M308020	1700	0.9533(1)	16.98(12)	1758	0.9388(2)	17.35(12)	17.66	0.07
M308025	1700	0.9543(1)	16.78(12)	1758	0.9363(1)	17.14(12)	17.27	-0.07
M308028	1500	0.9499(1)	16.42(11)	1546	0.9341(1)	16.71(11)	16.83	-0.10
M308031	1300	0.9458(1)	16.04(10)	1336	0.9322(1)	16.26(10)	16.34	-0.20
M308034	1100	0.9419(1)	15.62(9)	1127	0.9304(1)	15.78(9)	15.86	-0.21
M308037	900	0.9385(0)	15.11(7)	919	0.9293(1)	15.23(8)	15.28	-0.06

M308040	700	0.9351(0)	14.66(6)	713	0.9278(1)	14.73(6)	14.91	-0.09
M308043	500	0.9316(1)	14.28(5)	508	0.9382(1)	14.33(5)	18.80	0.25
M308046	307	0.9290(1)	13.88(4)	311	0.9360(2)	13.90(4)	18.32	0.13
M308050	1700	0.9495(1)	17.77(13)	1760	0.9334(1)	18.15(13)	17.96	-0.03
M308054	1500	0.9451(1)	17.43(12)	1548	0.9315(1)	17.74(12)	17.46	-0.10
M308057	1300	0.9415(1)	16.97(11)	1337	0.9296(1)	17.20(11)	16.97	-0.11

561 The original data are from Katsura *et al.* [2009b].

562

563 **Table S3.**  $P$ - $V$ - $T$  data of ringwoodite. . The initial data is from Katsura et al. [2004b], and the pressure and tempeatures are recalculated in this  
 564 study.

Data #	T before correction (K)	$V/V_0^{\text{MgO}}$	$P^{\text{MgO}}$ before correction (GPa)	T after correction (K)	$V/V_0^{\text{Rw}}$	$P^{\text{MgO}}$ after correction (GPa)	$P^{\text{Rw}}$ (GPa)	$P^{\text{MgO}} - P^{\text{Rw}}$ (GPa)
M038090	1500	0.9263(4)	21.71(18)	1554	0.9300(3)	22.05(19)	21.69	0.37
M038092	1550	0.9278(3)	21.67(18)	1607	0.9311(3)	22.03(19)	21.68	0.35
M038094	1600	0.9291(3)	21.69(18)	1661	0.9316(3)	22.08(19)	21.83	0.25
M038096	1650	0.9306(2)	21.66(17)	1714	0.9321(2)	22.06(18)	21.96	0.10
M038098	1700	0.9317(3)	21.71(18)	1767	0.9328(3)	22.14(19)	22.06	0.07
M038100	1600	0.9301(3)	21.46(18)	1660	0.9318(2)	21.84(18)	21.75	0.09
M038102	1500	0.9283(3)	21.24(17)	1554	0.9306(2)	21.58(17)	21.53	0.05
M038104	1400	0.9265(3)	21.03(16)	1447	0.9296(2)	21.33(17)	21.24	0.10
M038106	1300	0.9247(3)	20.83(16)	1342	0.9282(3)	21.10(16)	21.08	0.02
M038108	1200	0.9227(2)	20.66(14)	1236	0.9271(2)	20.89(15)	20.85	0.04
M038111	1100	0.9211(2)	20.43(13)	1131	0.9256(3)	20.63(13)	20.74	-0.11
M038113	1000	0.9191(2)	20.29(13)	1027	0.9243(3)	20.46(13)	20.55	-0.09
M038115	900	0.9171(2)	20.16(12)	923	0.9236(2)	20.30(12)	20.24	0.05
M038117	800	0.9156(3)	19.93(12)	819	0.9223(2)	20.04(13)	20.11	-0.07
M038119	700	0.9138(3)	19.76(11)	715	0.9212(1)	19.86(11)	19.94	-0.08
M038121	600	0.9120(3)	19.62(11)	612	0.9202(2)	19.70(11)	19.74	-0.05
M038123	500	0.9104(3)	19.45(10)	510	0.9190(1)	19.51(11)	19.62	-0.11
M038125	400	0.9091(4)	19.25(12)	407	0.9181(2)	19.29(12)	19.46	-0.17
M038127	300	0.9084(4)	18.96(11)	305	0.9180(2)	18.98(11)	19.12	-0.14

M038129	1700	0.9322(3)	21.59(19)	1767	0.9331(2)	22.02(19)	21.97	0.05
M038130	1750	0.9325(3)	21.85(18)	1821	0.9337(2)	22.31(19)	22.10	0.21
M038132	1800	0.9348(4)	21.65(19)	1874	0.9348(2)	22.13(20)	22.11	0.02
M038134	1850	0.9359(2)	21.71(18)	1928	0.9359(2)	22.21(19)	22.11	0.10
M038136	1900	0.9376(2)	21.66(18)	1982	0.9369(2)	22.18(19)	22.15	0.03
M038138	1800	0.9354(3)	21.50(18)	1874	0.9365(3)	21.97(19)	21.68	0.29
M038140	1700	0.9337(3)	21.25(17)	1766	0.9344(2)	21.68(18)	21.64	0.04
M038142	1600	0.9318(2)	21.05(17)	1660	0.9332(3)	21.43(17)	21.39	0.04
M038144	1500	0.9300(2)	20.84(16)	1553	0.9321(3)	21.18(16)	21.14	0.04
M038146	1900	0.9379(3)	21.59(18)	1981	0.9374(2)	22.11(19)	22.04	0.07
M038148	1950	0.9397(2)	21.50(18)	2035	0.9387(2)	22.05(18)	22.00	0.05
M038150	2000	0.9412(2)	21.48(18)	2089	0.9398(2)	22.05(18)	22.00	0.05
M038154	1900	0.9433(2)	20.39(16)	1979	0.9421(3)	20.90(17)	20.88	0.01
M038156	1800	0.9408(3)	20.29(17)	1872	0.9410(3)	20.75(18)	20.59	0.16
M087012	1600	0.9341(1)	20.53(15)	1659	0.9326(2)	20.90(16)	21.54	-0.64
M087016	300	0.9072(1)	19.26(5)	305	0.9177(1)	19.28(5)	19.20	0.09
M087019	1100	0.9203(1)	20.63(13)	1131	0.9250(1)	20.83(13)	20.89	-0.06
M087021	1300	0.9234(1)	21.13(14)	1342	0.9270(2)	21.40(15)	21.40	0.00
M087023	1500	0.9272(1)	21.48(16)	1554	0.9295(1)	21.83(16)	21.82	0.01
M087025	1600	0.9295(2)	21.59(17)	1659	0.9311(2)	21.98(18)	21.94	0.03
M087027	1700	0.9326(2)	21.51(17)	1767	0.9330(1)	21.94(18)	22.00	-0.06
M087029	1800	0.9365(2)	21.25(17)	1873	0.9362(2)	21.72(18)	21.77	-0.05
M087032	1900	0.9404(3)	21.01(18)	1980	0.9393(2)	21.53(19)	21.55	-0.02
M087033	1800	0.9389(1)	20.71(16)	1872	0.9377(2)	21.17(17)	21.38	-0.21

M087035	1700	0.9370(1)	20.51(16)	1765	0.9367(2)	20.93(16)	21.06	-0.14
M087037	1600	0.9349(1)	20.35(15)	1658	0.9355(1)	20.72(16)	20.83	-0.10
M087039	1500	0.9327(1)	20.20(15)	1552	0.9342(2)	20.53(15)	20.60	-0.07
M087041	1400	0.9308(1)	20.01(14)	1446	0.9329(1)	20.30(14)	20.39	-0.09
M087043	1300	0.9290(1)	19.79(13)	1340	0.9317(2)	20.05(14)	20.17	-0.12
M087045	1200	0.9270(1)	19.63(12)	1235	0.9306(1)	19.85(13)	19.92	-0.07
M087047	1100	0.9251(2)	19.46(13)	1130	0.9290(5)	19.65(13)	19.81	-0.16
M087049	1000	0.9234(1)	19.24(11)	1026	0.9282(1)	19.41(11)	19.49	-0.09
M087051	900	0.9217(2)	19.04(11)	922	0.9271(1)	19.17(11)	19.27	-0.10
M087053	800	0.9197(1)	18.90(9)	818	0.9262(1)	19.01(9)	19.04	-0.02
M087055	700	0.9180(2)	18.72(10)	715	0.9250(1)	18.81(10)	18.86	-0.06
M087057	600	0.9164(1)	18.50(8)	612	0.9240(2)	18.57(8)	18.65	-0.08
M087059	500	0.9147(1)	18.35(7)	509	0.9231(1)	18.40(7)	18.46	-0.06
M087061	400	0.9133(2)	18.18(7)	407	0.9223(2)	18.22(7)	18.25	-0.03
M087063	305	0.9123(0)	17.97(4)	310	0.9219(2)	17.99(4)	18.00	-0.01
M087065	302	0.9108(2)	18.34(6)	307	0.9212(2)	18.36(6)	18.19	0.18
M087067	302	0.9099(1)	18.56(5)	307	0.9201(2)	18.59(5)	18.51	0.07
M087069	500	0.9128(1)	18.84(7)	509	0.9211(2)	18.90(7)	19.02	-0.12
M087071	700	0.9156(1)	19.32(9)	715	0.9227(1)	19.41(9)	19.49	-0.09
M087073	900	0.9188(1)	19.73(10)	922	0.9246(2)	19.87(10)	19.98	-0.11
M087075	304	0.9096(1)	18.65(6)	309	0.9194(2)	18.67(6)	18.72	-0.04
M087077	300	0.9044(2)	20.01(7)	305	0.9148(2)	20.03(8)	20.06	-0.03
M087083	700	0.9054(1)	21.95(10)	716	0.9136(2)	22.05(10)	22.14	-0.10
M087085	1100	0.9124(1)	22.60(14)	1133	0.9173(2)	22.81(14)	23.05	-0.24

M087087	1300	0.9161(2)	22.94(16)	1344	0.9198(1)	23.22(17)	23.37	-0.15
M087089	1400	0.9179(2)	23.13(17)	1450	0.9209(2)	23.45(18)	23.59	-0.14
M087091	1500	0.9198(1)	23.29(17)	1556	0.9221(1)	23.65(18)	23.79	-0.13
M087093	1300	0.9155(1)	23.09(16)	1344	0.9196(2)	23.37(16)	23.44	-0.07
M087096	1100	0.9121(1)	22.69(14)	1133	0.9177(1)	22.90(14)	22.94	-0.04
M087098	900	0.9084(1)	22.37(12)	924	0.9154(2)	22.52(13)	22.57	-0.05
M087100	700	0.9052(1)	22.01(10)	716	0.9135(2)	22.11(10)	22.16	-0.05
M087102	500	0.9021(1)	21.66(7)	510	0.9116(2)	21.71(8)	21.82	-0.10
M087104	300	0.8995(1)	21.33(6)	306	0.9106(2)	21.35(6)	21.33	0.02
M087106	1500	0.9198(2)	23.30(18)	1556	0.9223(1)	23.66(18)	23.74	-0.08
M087109	1500	0.9222(3)	22.70(18)	1556	0.9253(1)	23.05(19)	22.93	0.12
M087111	1700	0.9265(1)	22.95(18)	1769	0.9278(2)	23.40(19)	23.36	0.04
M087113	1700	0.9277(1)	22.67(18)	1769	0.9290(1)	23.11(18)	23.04	0.07
M087115	1700	0.9288(1)	22.39(17)	1768	0.9299(1)	22.83(18)	22.80	0.03
M087117	1800	0.9312(2)	22.47(18)	1876	0.9314(1)	22.96(19)	22.96	-0.01
M087119	1900	0.9341(1)	22.44(18)	1983	0.9342(1)	22.98(19)	22.82	0.15
M087121	1700	0.9308(1)	21.93(17)	1768	0.9317(1)	22.36(18)	22.34	0.02
M087123	1500	0.9267(1)	21.62(16)	1554	0.9295(2)	21.97(16)	21.82	0.14
M087125	1300	0.9228(1)	21.28(14)	1342	0.9271(1)	21.54(15)	21.38	0.16
M087127	1100	0.9195(1)	20.83(13)	1132	0.9248(1)	21.03(13)	20.93	0.10
M087129	900	0.9160(1)	20.44(11)	923	0.9228(1)	20.58(11)	20.48	0.11
M087131	700	0.9125(1)	20.10(9)	715	0.9210(1)	20.19(9)	19.99	0.20
M087133	500	0.9095(2)	19.68(8)	510	0.9195(1)	19.74(8)	19.49	0.25
M087135	300	0.9072(1)	19.26(5)	305	0.9182(1)	19.29(5)	19.07	0.22

M087137	1700	0.9331(1)	21.39(16)	1767	0.9338(2)	21.82(17)	21.80	0.02
M087139	1700	0.9363(1)	20.67(16)	1765	0.9369(2)	21.09(16)	21.03	0.06
M087141	1700	0.9397(1)	19.90(15)	1764	0.9392(2)	20.31(16)	20.46	-0.15
M087143	1500	0.9356(1)	19.55(14)	1551	0.9365(1)	19.88(14)	20.02	-0.14
M087153	1800	0.9354(3)	21.51(19)	1874	0.9348(3)	21.98(19)	22.11	-0.13
M087155	1900	0.9375(2)	21.66(18)	1982	0.9365(2)	22.18(18)	22.25	-0.07
M087158	1500	0.9327(1)	20.20(15)	1552	0.9344(2)	20.53(15)	20.55	-0.02
M087160	1500	0.9367(3)	19.30(16)	1551	0.9383(1)	19.62(16)	19.57	0.05
M087162	1500	0.9398(1)	18.61(13)	1550	0.9409(1)	18.92(14)	18.93	-0.01
M087164	1300	0.9362(2)	18.14(12)	1338	0.9380(2)	18.38(13)	18.56	-0.17
M087166	1100	0.9325(1)	17.72(11)	1129	0.9361(2)	17.90(11)	17.96	-0.06
M087168	900	0.9285(2)	17.40(10)	921	0.9339(2)	17.53(10)	17.47	0.06
M087170	700	0.9255(1)	16.88(7)	714	0.9320(2)	16.97(7)	16.96	0.01
M087172	500	0.9222(1)	16.50(6)	509	0.9301(1)	16.55(6)	16.51	0.04
M087174	300	0.9197(0)	16.09(4)	305	0.9287(1)	16.11(4)	16.09	0.02
M087176	1100	0.9400(2)	16.04(10)	1127	0.9434(2)	16.21(11)	16.12	0.09
M087178	900	0.9363(2)	15.60(10)	919	0.9414(2)	15.72(10)	15.56	0.16
M087180	700	0.9333(1)	15.07(7)	713	0.9392(2)	15.15(7)	15.08	0.08
M087182	500	0.9293(1)	14.82(6)	508	0.9377(1)	14.86(6)	14.51	0.36
M087184	300	0.9256(3)	14.64(9)	304	0.9360(2)	14.66(9)	14.11	0.55
M092019	1700	0.9329(1)	21.44(17)	1767	0.9328(4)	21.87(17)	22.07	-0.20
M092021	1800	0.9364(2)	21.28(17)	1873	0.9367(3)	21.75(18)	21.65	0.10
M092023	1900	0.9402(2)	21.06(17)	1980	0.9397(3)	21.57(18)	21.48	0.10
M092025	2000	0.9433(4)	21.01(19)	2088	0.9424(3)	21.57(20)	21.39	0.19

M092028	1900	0.9418(2)	20.72(17)	1980	0.9409(4)	21.23(18)	21.17	0.06
M092030	1900	0.9424(2)	20.57(17)	1979	0.9402(4)	21.08(18)	21.35	-0.27
M092032	1900	0.9432(2)	20.40(17)	1979	0.9412(6)	20.91(17)	21.09	-0.19
M092034	1900	0.9446(1)	20.11(16)	1978	0.9420(5)	20.61(17)	20.91	-0.30
M092036	1900	0.9456(3)	19.87(17)	1978	0.9429(5)	20.37(17)	20.68	-0.31
M092038	1900	0.9467(1)	19.64(15)	1977	0.9443(4)	20.13(16)	20.36	-0.23

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567**Table S4.**  $P$ - $V$ - $T$  data of bridgmanite. . The initial data is from Katsura et al. [2009c] and Tange et al. [2012], and the pressure and tempeatures are recalculated in this study.

Data #	T before correction (K)	$V/V_0^{\text{MgO}}$	$P^{\text{MgO}}$ before correction (GPa)	T after correction (K)	$V/V_0^{\text{Brg}}$	$P^{\text{MgO}}$ after correction (GPa)	$P^{\text{Brg}}$ (GPa)	$P^{\text{MgO}} - P^{\text{Brg}}$ (GPa)
M1	1300	0.8891(1)	30.25(21)	1351	0.9254(7)	30.58(22)	30.28	0.30
M1	1100	0.8854(1)	30.07(18)	1139	0.9215(9)	30.31(19)	30.18	0.14
M1	900	0.8829(1)	29.54(15)	928	0.9184(9)	29.72(16)	29.84	-0.13
M1	700	0.8810(3)	28.87(15)	719	0.9162(9)	28.98(15)	29.26	-0.28
M1	500	0.8780(1)	28.60(10)	512	0.9141(6)	28.67(11)	28.80	-0.13
M1	300	0.8769(0)	27.91(6)	306	0.9131(4)	27.94(7)	28.14	-0.20
M2	1500	0.8548(4)	42.56(35)	1577	0.8987(4)	43.06(36)	42.53	0.53
M2	1300	0.8529(3)	41.91(29)	1360	0.8967(4)	42.30(30)	41.80	0.50
M2	1100	0.8513(2)	41.17(23)	1145	0.8941(5)	41.47(24)	41.41	0.06
M2	900	0.8486(3)	40.86(22)	933	0.8920(4)	41.07(22)	40.84	0.22
M2	700	0.8470(1)	40.21(16)	722	0.8899(5)	40.34(16)	40.34	0.00
M2	500	0.8454(1)	39.58(12)	514	0.8886(4)	39.66(12)	39.67	-0.01
M2	300	0.8443(4)	38.96(19)	308	0.8877(5)	39.00(19)	39.05	-0.05
M3	1500	0.8297(3)	52.00(36)	1584	0.8771(3)	52.56(38)	52.29	0.28
M3	1300	0.8277(3)	51.48(31)	1365	0.8753(4)	51.92(33)	51.58	0.34
M3	1100	0.8259(3)	50.94(27)	1149	0.8731(4)	51.26(28)	51.14	0.12
M3	900	0.8239(3)	50.47(23)	936	0.8712(5)	50.70(23)	50.62	0.08
M3	700	0.8218(5)	50.04(29)	724	0.8695(5)	50.19(30)	50.05	0.13
M3	500	0.8206(5)	49.35(28)	515	0.8682(4)	49.43(28)	49.48	-0.04

M3	300	0.8198(4)	48.67(20)	308	0.8673(3)	48.71(20)	48.97	-0.26
M4	1500	0.8067(3)	61.93(37)	1590	0.8581(7)	62.54(39)	61.89	0.64
M4	1300	0.8030(4)	62.32(35)	1371	0.8548(7)	62.80(37)	62.13	0.66
M4	1100	0.8003(8)	62.24(49)	1154	0.8527(6)	62.60(50)	61.77	0.83
M4	900	0.7987(7)	61.68(40)	939	0.8507(8)	61.94(41)	61.41	0.53
M4	700	0.7965(7)	61.43(38)	727	0.8487(12)	61.60(39)	61.17	0.43
M4	500	0.7952(7)	60.87(38)	517	0.8472(10)	60.97(38)	60.74	0.23
M4	300	0.7937(7)	60.58(38)	309	0.8461(10)	60.62(38)	60.46	0.16
M252031	1900	0.9313(1)	23.09(19)	1985	0.9614(1)	23.63(20)	23.05	0.57
M252029	1800	0.9283(0)	23.15(18)	1877	0.9583(1)	23.64(19)	23.16	0.48
M252027	1700	0.9253(0)	23.24(18)	1770	0.9551(1)	23.69(19)	23.31	0.37
M252034	1700	0.9270(1)	22.82(18)	1769	0.9559(1)	23.26(19)	23.06	0.20
M252025	1600	0.9226(1)	23.24(18)	1663	0.9522(1)	23.64(18)	23.44	0.21
M252023	1500	0.9195(2)	23.36(18)	1556	0.9489(1)	23.72(18)	23.66	0.06
M252036	1500	0.9235(1)	22.39(16)	1555	0.9523(1)	22.74(17)	22.58	0.16
M252038	1300	0.9203(1)	21.90(15)	1343	0.9490(1)	22.17(15)	22.02	0.16
M252040	1100	0.9167(1)	21.51(13)	1132	0.9458(1)	21.72(13)	21.52	0.20
M252042	900	0.9142(1)	20.90(11)	923	0.9429(1)	21.04(11)	20.96	0.08
M252044	700	0.9113(1)	20.42(9)	716	0.9402(1)	20.51(9)	20.43	0.08
M252046	500	0.9086(1)	19.93(7)	510	0.9378(1)	19.99(7)	19.91	0.08
M252048	306	0.9064(1)	19.50(5)	311	0.9358(1)	19.52(5)	19.56	-0.03
M252050	300	0.9076(1)	19.15(5)	305	0.9385(1)	19.18(5)	18.59	0.58
M253054	2500	0.9160(1)	30.69(31)	2661	0.9480(1)	31.74(33)	32.43	-0.70
M253052	2300	0.9127(1)	30.23(29)	2437	0.9455(1)	31.12(31)	31.53	-0.41

M253050	2100	0.9096(1)	29.75(28)	2216	0.9432(1)	30.50(29)	30.56	-0.06
M253032	1900	0.9071(0)	29.11(25)	1996	0.9402(1)	29.73(27)	29.91	-0.19
M253013	1900	0.9173(1)	26.46(22)	1991	0.9482(1)	27.05(24)	27.22	-0.17
M253034	1700	0.9046(1)	28.49(23)	1778	0.9380(1)	29.00(25)	28.99	0.01
M253015	1700	0.9144(1)	25.94(21)	1774	0.9462(1)	26.41(22)	26.22	0.19
M253036	1500	0.9028(1)	27.70(21)	1562	0.9359(1)	28.10(22)	28.08	0.02
M253017	1500	0.9124(1)	25.17(19)	1559	0.9438(1)	25.55(20)	25.36	0.19
M253038	1300	0.9001(1)	27.15(19)	1348	0.9335(1)	27.46(19)	27.32	0.14
M253019	1300	0.9095(1)	24.62(17)	1346	0.9414(1)	24.91(18)	24.59	0.32
M253040	1100	0.8982(1)	26.40(16)	1136	0.9311(1)	26.63(17)	26.64	-0.01
M253021	1100	0.9074(1)	23.91(14)	1134	0.9388(1)	24.12(15)	23.90	0.23
M253042	900	0.8953(1)	25.93(14)	926	0.9287(1)	26.09(14)	25.98	0.11
M253023	900	0.9044(1)	23.44(12)	925	0.9363(1)	23.59(13)	23.23	0.37
M253044	700	0.8934(1)	25.25(11)	718	0.9263(1)	25.36(11)	25.41	-0.05
M253025	700	0.9023(1)	22.79(10)	717	0.9340(1)	22.89(10)	22.58	0.30
M253046	500	0.8904(1)	24.93(9)	511	0.9243(1)	24.99(9)	24.84	0.15
M253027	500	0.8993(1)	22.41(8)	510	0.9314(1)	22.47(8)	22.19	0.28
M253031	300	0.8886(1)	24.41(7)	306	0.9234(1)	24.44(7)	24.11	0.33
M253048	300	0.8890(1)	24.28(7)	306	0.9229(1)	24.31(7)	24.29	0.02
M253029	300	0.8978(1)	21.81(6)	306	0.9297(1)	21.84(6)	21.74	0.09
M253059	300	1.0000(0)	0.00(0)	300	1.0000(1)	0.00(0)	0.00	0.00
M526017	1600	0.9015(5)	28.70(27)	1670	0.9368(3)	29.15(28)	28.59	0.56
M526020	1700	0.8998(5)	29.81(30)	1780	0.9358(3)	30.33(31)	29.78	0.55
M526021	1800	0.9033(6)	29.50(31)	1888	0.9369(3)	30.07(32)	30.20	-0.13

M526024	1900	0.9021(5)	30.46(32)	1998	0.9379(3)	31.09(33)	30.68	0.41
M526025	1700	0.9013(6)	29.39(31)	1779	0.9348(3)	29.90(32)	30.10	-0.20
M526028	1500	0.8975(6)	29.15(29)	1564	0.9311(4)	29.56(30)	29.79	-0.23
M526029	1300	0.8946(6)	28.68(27)	1350	0.9280(4)	29.00(27)	29.34	-0.33
M526032	1100	0.8919(5)	28.16(25)	1137	0.9251(3)	28.40(25)	28.81	-0.41
M526033	900	0.8894(6)	27.63(24)	927	0.9221(3)	27.80(24)	28.43	-0.63
M526045	1700	0.8812(7)	35.21(38)	1787	0.9170(4)	35.78(40)	36.67	-0.90
M526046	1800	0.8808(5)	35.98(36)	1897	0.9176(4)	36.61(38)	37.27	-0.66
M526047	1700	0.8796(5)	35.71(35)	1787	0.9160(4)	36.28(37)	37.08	-0.80
M526048	1500	0.8771(5)	35.17(33)	1570	0.9131(4)	35.63(34)	36.57	-0.94
M526049	1300	0.8745(5)	34.69(31)	1355	0.9101(4)	35.04(31)	36.19	-1.14
M526050	1100	0.8719(5)	34.23(28)	1141	0.9073(4)	34.49(29)	35.79	-1.30
M526051	900	0.8694(5)	33.73(26)	930	0.9050(4)	33.92(26)	35.23	-1.31
M526052	700	0.8671(5)	33.25(24)	720	0.9026(4)	33.37(24)	34.82	-1.45
M554011	1700	0.9211(4)	24.25(22)	1772	0.9513(4)	24.71(23)	24.55	0.16
M554012	1700	0.9191(3)	24.74(21)	1772	0.9497(3)	25.21(22)	25.07	0.13
M554013	1500	0.9150(4)	24.49(21)	1558	0.9457(4)	24.87(22)	24.73	0.14
M554014	1300	0.9116(5)	24.07(21)	1345	0.9435(4)	24.36(22)	23.86	0.50
M554015	1100	0.9085(4)	23.61(19)	1134	0.9398(4)	23.82(20)	23.56	0.26
M554016	900	0.9055(4)	23.14(18)	924	0.9368(4)	23.29(18)	23.06	0.24
M554017	700	0.9030(4)	22.59(16)	716	0.9342(4)	22.69(16)	22.53	0.17
M554018	500	0.9005(4)	22.09(15)	510	0.9316(4)	22.15(15)	22.14	0.01
M554019	304	0.8979(3)	21.78(11)	310	0.9292(3)	21.81(11)	21.94	-0.13
M554021	1700	0.9176(5)	25.12(25)	1773	0.9486(4)	25.59(26)	25.41	0.18

M554022	1700	0.9172(5)	25.21(25)	1773	0.9482(4)	25.68(25)	25.55	0.13
M554023	1800	0.9188(5)	25.45(25)	1881	0.9490(4)	25.98(26)	26.11	-0.14
M554024	1800	0.9189(5)	25.44(25)	1881	0.9502(5)	25.96(26)	25.74	0.22
M554025	1900	0.9220(5)	25.33(26)	1989	0.9515(3)	25.90(27)	26.15	-0.25
M557019	1900	0.8685(5)	40.52(41)	2013	0.9087(5)	41.27(43)	41.62	-0.35
M554020	1700	0.8658(5)	40.09(38)	1792	0.9065(4)	40.70(40)	40.88	-0.18
M554021	1500	0.8638(5)	39.45(34)	1574	0.9043(4)	39.94(35)	40.17	-0.24
M645023	1700	0.8302(5)	53.15(44)	1805	0.8778(5)	53.85(47)	53.55	0.30
M645024	1500	0.8270(4)	53.11(39)	1584	0.8736(5)	53.68(41)	54.00	-0.32
M645025	1300	0.8250(4)	52.62(35)	1366	0.8719(4)	53.06(36)	53.26	-0.20
M645026	1100	0.8233(4)	51.98(31)	1150	0.8704(4)	52.31(32)	52.50	-0.19
M645027	900	0.8221(4)	51.20(27)	936	0.8692(4)	51.44(28)	51.63	-0.19
M645028	700	0.8207(4)	50.53(24)	725	0.8679(4)	50.69(24)	50.89	-0.20
M645029	500	0.8194(4)	49.86(21)	515	0.8664(4)	49.94(21)	50.38	-0.43
M645030	304	0.8184(4)	49.29(21)	312	0.8658(3)	49.32(21)	49.75	-0.42
M654032	1900	0.8303(5)	54.45(49)	2029	0.8781(3)	55.32(52)	55.02	0.30
M659028	1700	0.8508(5)	45.29(41)	1798	0.8947(3)	45.94(43)	45.85	0.09
M659029	1500	0.8486(4)	44.77(35)	1578	0.8933(3)	45.28(37)	44.85	0.44
M659030	1300	0.8470(4)	44.04(32)	1361	0.8907(3)	44.44(33)	44.42	0.02
M659031	1100	0.8444(4)	43.70(28)	1146	0.8880(3)	44.00(29)	44.10	-0.10
M659032	900	0.8432(4)	42.85(26)	933	0.8866(3)	43.07(27)	43.27	-0.21
M659033	700	0.8417(4)	42.14(22)	723	0.8850(3)	42.28(22)	42.61	-0.33
M659034	500	0.8401(4)	41.55(19)	514	0.8834(2)	41.64(19)	42.07	-0.43
M659035	308	0.8389(4)	41.06(18)	316	0.8821(2)	41.09(18)	41.71	-0.61

M659036	1700	0.8501(4)	45.54(40)	1798	0.8943(3)	46.19(42)	46.02	0.18
M697010	305	0.8724(5)	29.35(19)	312	0.9104(2)	29.38(19)	29.25	0.14

568 Original data labeled as ‘M1’, ‘M2’, ‘M3, and ‘M4’ are from Tange *et al.* [2014]. The other data are from Katsura *et al.* [2009c].

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**Table S5.** Thermoelastic parameters of the adiabatic mantle

Depth (km)	$\rho$ (kg m <sup>-3</sup> ) <sup>#</sup>	g (m s <sup>-2</sup> ) <sup>#</sup>	P (GPa) <sup>#</sup>	T (K)	dT/dP (K GPa <sup>-1</sup> )	dT/dP (K GPa <sup>-1</sup> )	$\alpha$ (K <sup>-1</sup> )	Cp (J kg <sup>-1</sup> )	Cp (kJ m <sup>-3</sup> )	Cv/3nR	V/V <sub>0</sub>	$\gamma$	1+ $\alpha\gamma T$	K <sub>T</sub> (GPa)	K <sub>T</sub> (GPa)
50	3127(4)	9.85	1.46	1616(17)	15.9(9)	0.530(29)	4.02(23)	1305(5)	4079(10)	0.981(1)	1.032(1)	1.102(4)	1.072(4)	102.5(4)	109.8(6)
60	3136(4)	9.85	1.79	1622(17)	15.7(8)	0.521(27)	3.95(21)	1303(4)	4086(9)	0.981(1)	1.029(1)	1.094(3)	1.070(4)	103.9(4)	111.2(6)
70	3145(4)	9.86	2.12	1627(17)	15.4(8)	0.512(25)	3.87(20)	1301(4)	4093(8)	0.981(1)	1.026(1)	1.086(3)	1.068(4)	105.3(4)	112.5(6)
80	3155(3)	9.86	2.45	1632(17)	15.1(7)	0.504(24)	3.80(18)	1300(4)	4100(8)	0.982(1)	1.023(1)	1.077(3)	1.067(3)	106.7(4)	113.8(6)
90	3164(3)	9.86	2.79	1637(17)	14.9(7)	0.495(22)	3.74(17)	1298(4)	4108(7)	0.982(1)	1.020(1)	1.069(3)	1.065(3)	108.1(4)	115.2(5)
100	3173(3)	9.87	3.12	1642(17)	14.6(6)	0.487(21)	3.67(16)	1297(3)	4115(7)	0.982(1)	1.017(1)	1.062(3)	1.064(3)	109.5(4)	116.5(5)
110	3182(3)	9.87	3.45	1646(17)	14.4(6)	0.480(20)	3.61(15)	1296(3)	4122(6)	0.982(1)	1.014(1)	1.054(3)	1.063(3)	110.9(4)	117.8(5)
120	3191(3)	9.87	3.78	1651(17)	14.2(6)	0.472(18)	3.55(14)	1294(3)	4129(6)	0.983(1)	1.011(1)	1.046(2)	1.061(3)	112.3(4)	119.2(5)
130	3199(3)	9.87	4.12	1656(17)	14.0(5)	0.465(17)	3.49(13)	1293(3)	4137(5)	0.983(1)	1.009(1)	1.039(2)	1.060(2)	113.7(4)	120.5(5)
140	3208(3)	9.88	4.45	1660(17)	13.8(5)	0.457(16)	3.43(12)	1292(3)	4144(5)	0.983(1)	1.006(1)	1.032(2)	1.059(2)	115.0(4)	121.8(5)
150	3217(3)	9.88	4.78	1665(17)	13.5(5)	0.450(15)	3.38(11)	1291(3)	4151(5)	0.983(1)	1.003(1)	1.025(2)	1.058(2)	116.4(4)	123.1(5)
160	3225(3)	9.88	5.12	1669(17)	13.3(4)	0.444(15)	3.32(11)	1289(2)	4159(5)	0.983(1)	1.001(1)	1.018(2)	1.056(2)	117.8(4)	124.4(4)
170	3234(3)	9.89	5.45	1674(17)	13.1(4)	0.437(14)	3.27(10)	1288(2)	4166(4)	0.984(1)	0.998(1)	1.011(2)	1.055(2)	119.2(4)	125.8(4)
180	3242(3)	9.89	5.78	1678(17)	12.9(4)	0.431(13)	3.22(10)	1287(2)	4173(4)	0.984(1)	0.995(1)	1.004(2)	1.054(2)	120.5(4)	127.1(4)
190	3251(2)	9.89	6.11	1682(17)	12.8(4)	0.425(13)	3.17(9)	1286(2)	4181(4)	0.984(1)	0.993(1)	0.997(2)	1.053(2)	121.9(4)	128.4(4)
200	3259(2)	9.89	6.45	1686(17)	12.6(4)	0.420(12)	3.12(9)	1285(2)	4188(4)	0.984(1)	0.990(1)	0.991(2)	1.052(2)	123.3(4)	129.7(4)
210	3267(2)	9.90	6.78	1690(17)	12.4(3)	0.416(12)	3.08(8)	1284(2)	4195(4)	0.984(1)	0.988(1)	0.984(2)	1.051(2)	124.6(4)	131.0(4)
220	3275(2)	9.90	7.11	1695(17)	12.2(3)	0.411(11)	3.03(8)	1283(2)	4203(4)	0.984(1)	0.985(1)	0.978(2)	1.050(1)	126.0(4)	132.3(4)
230	3284(2)	9.90	7.45	1699(18)	12.1(3)	0.407(11)	2.99(8)	1282(2)	4210(4)	0.985(1)	0.983(1)	0.972(2)	1.049(1)	127.4(4)	133.6(4)
240	3292(2)	9.91	7.79	1703(18)	11.9(3)	0.403(11)	2.94(8)	1281(2)	4218(5)	0.985(1)	0.980(1)	0.965(2)	1.048(1)	128.7(4)	135.0(4)
250	3300(2)	9.91	8.14	1707(18)	11.7(3)	0.400(11)	2.90(7)	1280(2)	4225(5)	0.985(1)	0.978(1)	0.959(2)	1.047(1)	130.1(4)	136.3(4)
260	3308(2)	9.92	8.48	1711(18)	11.6(3)	0.396(11)	2.86(7)	1279(2)	4233(5)	0.985(1)	0.975(1)	0.953(2)	1.047(1)	131.5(4)	137.6(4)
270	3316(2)	9.92	8.82	1715(18)	11.4(3)	0.392(11)	2.82(7)	1279(2)	4240(5)	0.985(1)	0.973(1)	0.947(2)	1.046(1)	132.9(4)	139.0(4)
280	3324(2)	9.93	9.17	1718(18)	11.2(3)	0.388(11)	2.78(7)	1278(2)	4248(5)	0.985(1)	0.971(1)	0.941(2)	1.045(1)	134.3(4)	140.3(4)
290	3332(2)	9.94	9.51	1722(18)	11.1(3)	0.383(11)	2.74(7)	1277(2)	4255(6)	0.985(1)	0.968(1)	0.935(2)	1.044(1)	135.7(4)	141.7(4)
300	3340(2)	9.94	9.86	1726(18)	10.9(3)	0.379(11)	2.70(7)	1276(2)	4263(6)	0.985(1)	0.966(1)	0.929(2)	1.043(1)	137.0(4)	143.0(4)
310	3348(2)	9.95	10.20	1730(18)	10.8(3)	0.375(11)	2.67(8)	1275(2)	4270(6)	0.986(1)	0.964(1)	0.924(2)	1.043(1)	138.4(4)	144.3(4)

320	3356(2)	9.95	10.55	1734(18)	10.7(3)	0.371(11)	2.63(8)	1275(3)	4278(6)	0.986(1)	0.961(1)	0.918(2)	1.042(1)	139.8(4)	145.7(4)
330	3364(2)	9.95	10.90	1737(18)	10.5(3)	0.367(11)	2.60(8)	1274(3)	4285(7)	0.986(1)	0.959(1)	0.913(2)	1.041(1)	141.2(4)	147.0(4)
340	3372(2)	9.95	11.25	1741(18)	10.4(3)	0.362(11)	2.56(8)	1273(3)	4293(7)	0.986(1)	0.957(1)	0.907(2)	1.040(1)	142.6(4)	148.4(4)
350	3380(2)	9.95	11.60	1744(18)	10.3(3)	0.358(12)	2.53(8)	1272(3)	4300(7)	0.986(1)	0.955(1)	0.902(2)	1.040(1)	144.0(4)	149.7(4)
360	3388(2)	9.95	11.95	1748(18)	10.1(3)	0.355(12)	2.49(8)	1272(3)	4308(7)	0.986(1)	0.953(1)	0.897(2)	1.039(1)	145.3(4)	151.0(4)
370	3395(2)	9.96	12.30	1751(18)	10.0(3)	0.351(12)	2.46(8)	1271(3)	4315(8)	0.986(1)	0.950(1)	0.891(2)	1.038(1)	146.7(4)	152.4(4)
380	3403(2)	9.96	12.65	1755(18)	9.9(4)	0.347(12)	2.43(9)	1270(3)	4323(8)	0.986(1)	0.948(1)	0.886(2)	1.038(1)	148.1(4)	153.7(4)
390	3411(3)	9.97	13.00	1758(18)	9.7(4)	0.344(13)	2.40(9)	1270(3)	4330(8)	0.986(1)	0.946(1)	0.881(2)	1.037(1)	149.5(4)	155.1(4)
400	3418(3)	9.97	13.35	1762(18)	9.6(4)	0.340(13)	2.37(9)	1269(3)	4338(8)	0.986(1)	0.944(1)	0.876(2)	1.037(1)	150.9(4)	156.4(4)
410	3426(3)	9.97	13.73	1765(18)	9.5(4)	0.354(14)	2.34(9)	1268(3)	4346(9)	0.987(1)	0.942(1)	0.871(2)	1.036(2)	152.4(4)	157.8(4)
410	3606(3)	9.97	13.73	1825(18)	9.6(5)	0.357(17)	2.42(12)	1282(2)	4623(5)	0.992(1)	0.963(1)	0.927(1)	1.041(2)	193.0(3)	200.9(4)
420	3613(2)	9.97	14.10	1829(18)	9.5(4)	0.355(16)	2.40(11)	1281(2)	4630(5)	0.992(1)	0.961(1)	0.925(1)	1.041(2)	194.3(3)	202.2(4)
430	3620(2)	9.98	14.48	1833(18)	9.4(4)	0.353(15)	2.38(10)	1281(2)	4636(5)	0.992(1)	0.959(1)	0.922(1)	1.040(2)	195.6(3)	203.5(4)
440	3626(2)	9.98	14.85	1836(18)	9.3(4)	0.352(15)	2.36(9)	1280(2)	4643(5)	0.992(1)	0.958(1)	0.919(1)	1.040(2)	196.9(3)	204.7(4)
450	3633(2)	9.98	15.23	1840(18)	9.3(4)	0.350(14)	2.34(9)	1280(2)	4650(5)	0.992(1)	0.956(1)	0.917(1)	1.039(2)	198.1(3)	206.0(3)
460	3639(2)	9.98	15.61	1843(18)	9.2(3)	0.348(13)	2.32(8)	1279(2)	4656(5)	0.992(1)	0.954(1)	0.914(1)	1.039(1)	199.4(3)	207.2(3)
470	3646(2)	9.98	15.99	1847(18)	9.1(3)	0.347(12)	2.30(7)	1279(2)	4663(5)	0.992(2)	0.953(1)	0.912(1)	1.039(1)	200.7(3)	208.5(3)
480	3652(2)	9.99	16.37	1850(18)	9.0(3)	0.345(11)	2.28(7)	1279(2)	4670(6)	0.992(2)	0.951(1)	0.909(1)	1.038(1)	202.0(3)	209.8(3)
490	3659(2)	9.99	16.75	1854(18)	9.0(3)	0.344(11)	2.26(6)	1278(2)	4677(6)	0.992(2)	0.949(1)	0.906(1)	1.038(1)	203.3(3)	211.1(3)
500	3665(2)	9.99	17.13	1857(18)	8.9(3)	0.342(10)	2.24(6)	1278(2)	4683(7)	0.992(2)	0.948(0)	0.904(1)	1.038(1)	204.6(3)	212.3(3)
510	3672(2)	9.99	17.52	1860(18)	8.8(2)	0.341(10)	2.23(6)	1277(2)	4690(7)	0.992(2)	0.946(0)	0.901(1)	1.037(1)	205.9(3)	213.6(3)
520	3678(2)	9.99	17.91	1864(18)	8.8(2)	0.339(9)	2.21(5)	1277(2)	4697(7)	0.992(2)	0.944(0)	0.899(1)	1.037(1)	207.2(3)	214.9(3)
520	3737(3)	9.99	17.91	1907(18)	10.1(5)	0.392(19)	2.56(12)	1290(3)	4822(6)	0.996(1)	0.953(1)	0.898(1)	1.044(2)	224.1(4)	233.9(5)
530	3743(3)	10.00	18.29	1911(18)	10.0(5)	0.389(18)	2.53(12)	1290(2)	4827(6)	0.996(1)	0.952(1)	0.895(1)	1.043(2)	225.7(4)	235.4(5)
540	3750(2)	10.00	18.68	1915(18)	9.9(4)	0.386(17)	2.50(11)	1289(2)	4832(5)	0.996(1)	0.950(1)	0.892(1)	1.043(2)	227.2(4)	236.9(5)
550	3756(2)	10.00	19.07	1919(18)	9.8(4)	0.383(17)	2.47(10)	1288(2)	4837(5)	0.996(1)	0.949(1)	0.888(1)	1.042(2)	228.8(4)	238.4(5)
560	3763(2)	10.00	19.46	1922(18)	9.7(4)	0.380(16)	2.44(10)	1287(2)	4842(4)	0.996(1)	0.947(1)	0.885(1)	1.042(2)	230.4(4)	239.9(5)
570	3769(2)	10.00	19.86	1926(18)	9.6(4)	0.377(15)	2.41(9)	1286(2)	4848(4)	0.995(1)	0.945(1)	0.882(1)	1.041(2)	232.0(4)	241.5(4)
580	3776(2)	10.00	20.25	1930(18)	9.5(4)	0.374(14)	2.38(9)	1285(2)	4853(4)	0.995(1)	0.944(1)	0.879(1)	1.040(2)	233.5(4)	243.0(4)
590	3782(2)	10.00	20.65	1934(18)	9.4(3)	0.372(13)	2.36(8)	1284(2)	4858(3)	0.995(1)	0.942(0)	0.877(1)	1.040(2)	235.1(4)	244.5(4)

600	3788(2)	10.00	21.04	1937(19)	9.3(3)	0.369(13)	2.33(7)	1284(1)	4863(3)	0.995(1)	0.940(0)	0.874(1)	1.039(1)	236.7(4)	246.0(4)
610	3795(2)	10.00	21.44	1941(19)	9.2(3)	0.365(12)	2.30(7)	1283(1)	4868(3)	0.995(1)	0.939(0)	0.871(1)	1.039(1)	238.3(4)	247.5(4)
620	3801(2)	10.01	21.84	1945(19)	9.1(3)	0.362(11)	2.28(6)	1282(1)	4874(3)	0.995(1)	0.937(0)	0.868(1)	1.038(1)	239.9(4)	249.1(4)
630	3807(2)	10.01	22.24	1948(19)	9.0(3)	0.358(11)	2.25(6)	1281(1)	4879(3)	0.995(1)	0.936(0)	0.865(1)	1.038(1)	241.4(4)	250.6(3)
640	3814(2)	10.01	22.64	1952(19)	8.9(3)	0.355(10)	2.22(5)	1281(1)	4884(3)	0.995(1)	0.934(0)	0.862(1)	1.037(1)	243.0(4)	252.1(3)
650	3820(2)	10.01	23.04	1955(19)	8.8(2)	0.351(9)	2.20(5)	1280(1)	4890(3)	0.995(1)	0.933(0)	0.859(1)	1.037(1)	244.6(4)	253.7(3)
660	3826(2)	10.01	23.44	1959(19)	8.7(2)	0.348(9)	2.17(5)	1279(1)	4895(3)	0.995(1)	0.931(0)	0.857(1)	1.036(1)	246.2(4)	255.2(3)
660	4286(3)	10.01	23.44	1925(19)	9.1(2)	0.363(9)	2.65(6)	1306(2)	5600(7)	1.007(1)	0.958(1)	0.872(2)	1.044(1)	302.7(5)	316.1(5)
670	4292(3)	10.01	23.84	1929(19)	9.1(2)	0.362(9)	2.63(6)	1306(2)	5605(7)	1.007(1)	0.957(1)	0.869(2)	1.044(1)	304.2(5)	317.6(5)
680	4298(3)	10.01	24.27	1932(19)	9.0(2)	0.395(9)	2.61(6)	1305(2)	5610(7)	1.007(1)	0.956(1)	0.865(2)	1.044(1)	305.8(5)	319.2(4)
690	4304(3)	10.01	24.71	1936(19)	9.0(2)	0.394(9)	2.60(5)	1305(2)	5616(6)	1.007(1)	0.954(1)	0.861(2)	1.043(1)	307.5(5)	320.8(4)
700	4310(3)	10.01	25.15	1940(19)	8.9(2)	0.392(9)	2.58(5)	1304(2)	5621(6)	1.007(1)	0.953(1)	0.857(2)	1.043(1)	309.1(5)	322.4(4)
710	4316(3)	10.01	25.59	1944(19)	8.9(2)	0.391(9)	2.56(5)	1304(2)	5627(6)	1.007(1)	0.952(1)	0.854(2)	1.043(1)	310.8(5)	324.0(4)
720	4322(3)	10.01	26.03	1948(19)	8.8(2)	0.389(9)	2.55(5)	1303(2)	5632(6)	1.007(1)	0.950(1)	0.850(2)	1.042(1)	312.4(5)	325.6(4)
730	4328(3)	10.01	26.48	1952(19)	8.8(2)	0.387(8)	2.53(5)	1303(2)	5638(6)	1.007(1)	0.949(1)	0.846(2)	1.042(1)	314.1(5)	327.2(4)
740	4333(2)	10.01	26.92	1956(19)	8.7(2)	0.386(8)	2.51(4)	1302(2)	5643(6)	1.006(1)	0.948(1)	0.843(2)	1.041(1)	315.8(5)	328.9(4)
750	4339(2)	10.00	27.36	1960(19)	8.7(2)	0.384(8)	2.50(4)	1302(2)	5649(6)	1.006(1)	0.947(1)	0.839(2)	1.041(1)	317.4(5)	330.5(4)
760	4345(2)	10.00	27.81	1963(19)	8.6(2)	0.383(8)	2.48(4)	1301(2)	5654(6)	1.006(1)	0.945(1)	0.836(1)	1.041(1)	319.1(5)	332.1(4)
770	4351(2)	10.00	28.25	1967(19)	8.6(2)	0.381(8)	2.47(4)	1301(2)	5659(6)	1.006(1)	0.944(1)	0.832(1)	1.040(1)	320.7(5)	333.7(4)
780	4357(2)	10.00	28.70	1971(19)	8.5(2)	0.380(8)	2.45(4)	1300(2)	5665(6)	1.006(1)	0.943(1)	0.829(1)	1.040(1)	322.4(5)	335.3(4)
790	4363(2)	10.00	29.14	1975(20)	8.5(2)	0.378(7)	2.44(4)	1300(2)	5670(6)	1.006(1)	0.941(1)	0.825(1)	1.040(1)	324.1(5)	336.9(4)
800	4369(2)	9.99	29.59	1979(20)	8.4(2)	0.376(7)	2.42(4)	1299(2)	5676(6)	1.006(1)	0.940(1)	0.822(1)	1.039(1)	325.7(5)	338.6(4)
810	4374(2)	9.99	30.04	1982(20)	8.4(2)	0.375(7)	2.41(3)	1299(2)	5681(6)	1.006(1)	0.939(1)	0.818(1)	1.039(1)	327.4(5)	340.2(4)
820	4380(2)	9.99	30.48	1986(20)	8.4(2)	0.373(7)	2.39(3)	1298(2)	5687(6)	1.006(1)	0.938(1)	0.815(1)	1.039(1)	329.1(5)	341.8(4)
830	4386(2)	9.99	30.93	1990(20)	8.3(2)	0.372(7)	2.38(3)	1298(2)	5692(6)	1.006(1)	0.936(1)	0.811(1)	1.038(1)	330.7(5)	343.4(4)
840	4392(2)	9.99	31.38	1994(20)	8.3(1)	0.371(7)	2.36(3)	1297(2)	5698(6)	1.006(1)	0.935(1)	0.808(1)	1.038(1)	332.4(5)	345.0(4)
850	4397(2)	9.98	31.82	1997(20)	8.2(1)	0.369(7)	2.35(3)	1297(2)	5703(6)	1.006(1)	0.934(1)	0.805(1)	1.038(1)	334.0(5)	346.6(4)
860	4403(2)	9.98	32.27	2001(20)	8.2(1)	0.368(6)	2.34(3)	1297(2)	5709(6)	1.006(1)	0.933(0)	0.801(1)	1.037(1)	335.7(5)	348.3(4)
870	4409(2)	9.98	32.72	2005(20)	8.1(1)	0.366(6)	2.32(3)	1296(2)	5714(6)	1.006(1)	0.932(0)	0.798(1)	1.037(1)	337.3(5)	349.9(4)
880	4414(2)	9.98	33.17	2008(20)	8.1(1)	0.365(6)	2.31(3)	1296(2)	5719(6)	1.006(1)	0.930(0)	0.795(1)	1.037(1)	339.0(5)	351.5(4)

890	4420(2)	9.98	33.62	2012(20)	8.1(1)	0.363(6)	2.30(3)	1295(2)	5725(6)	1.006(1)	0.929(0)	0.792(1)	1.037(1)	340.7(5)	353.1(4)
900	4426(2)	9.98	34.07	2016(20)	8.0(1)	0.362(6)	2.28(3)	1295(2)	5730(6)	1.006(1)	0.928(0)	0.788(1)	1.036(1)	342.3(5)	354.7(4)
910	4432(2)	9.98	34.53	2019(20)	8.0(1)	0.361(6)	2.27(3)	1294(2)	5736(7)	1.006(1)	0.927(0)	0.785(1)	1.036(1)	344.0(5)	356.4(4)
920	4437(2)	9.98	34.98	2023(20)	7.9(1)	0.359(6)	2.26(2)	1294(2)	5741(7)	1.006(1)	0.926(0)	0.782(1)	1.036(1)	345.7(5)	358.0(4)
930	4443(2)	9.97	35.43	2026(20)	7.9(1)	0.358(6)	2.24(2)	1293(2)	5747(7)	1.006(1)	0.924(0)	0.779(1)	1.035(1)	347.3(5)	359.6(4)
940	4448(2)	9.97	35.88	2030(20)	7.9(1)	0.357(6)	2.23(2)	1293(2)	5752(7)	1.006(1)	0.923(0)	0.776(1)	1.035(1)	349.0(5)	361.2(4)
950	4454(2)	9.97	36.34	2034(20)	7.8(1)	0.355(6)	2.22(2)	1293(2)	5758(7)	1.005(1)	0.922(0)	0.773(1)	1.035(1)	350.6(5)	362.9(4)
960	4460(2)	9.97	36.79	2037(20)	7.8(1)	0.354(6)	2.20(2)	1292(2)	5763(7)	1.005(1)	0.921(0)	0.770(1)	1.035(1)	352.3(5)	364.5(4)
970	4465(2)	9.97	37.24	2041(20)	7.8(1)	0.353(6)	2.19(2)	1292(2)	5768(7)	1.005(1)	0.920(0)	0.767(1)	1.034(1)	354.0(5)	366.1(4)
980	4471(2)	9.97	37.70	2044(20)	7.7(1)	0.351(6)	2.18(2)	1291(2)	5774(7)	1.005(1)	0.919(1)	0.764(1)	1.034(1)	355.6(5)	367.7(4)
990	4476(2)	9.97	38.15	2048(20)	7.7(1)	0.350(6)	2.17(2)	1291(2)	5779(7)	1.005(1)	0.918(1)	0.761(1)	1.034(1)	357.3(5)	369.4(4)
1000	4482(2)	9.97	38.61	2051(21)	7.6(1)	0.349(6)	2.15(2)	1291(2)	5785(8)	1.005(1)	0.916(1)	0.758(1)	1.033(1)	359.0(5)	371.0(4)
1010	4488(2)	9.97	39.07	2055(21)	7.6(1)	0.348(5)	2.14(2)	1290(2)	5790(8)	1.005(1)	0.915(1)	0.755(1)	1.033(1)	360.6(5)	372.6(4)
1020	4493(2)	9.97	39.53	2058(21)	7.6(1)	0.346(5)	2.13(2)	1290(2)	5795(8)	1.005(1)	0.914(1)	0.752(1)	1.033(1)	362.3(5)	374.3(4)
1030	4499(3)	9.96	39.98	2062(21)	7.5(1)	0.345(5)	2.12(2)	1289(2)	5801(8)	1.005(1)	0.913(1)	0.749(1)	1.033(1)	364.0(5)	375.9(4)
1040	4504(3)	9.96	40.44	2065(21)	7.5(1)	0.344(5)	2.11(2)	1289(2)	5806(8)	1.005(1)	0.912(1)	0.746(1)	1.032(1)	365.6(5)	377.5(4)
1050	4510(3)	9.96	40.90	2069(21)	7.5(1)	0.343(6)	2.10(2)	1289(2)	5812(8)	1.005(1)	0.911(1)	0.743(1)	1.032(1)	367.3(5)	379.1(4)
1060	4515(3)	9.96	41.36	2072(21)	7.4(1)	0.341(6)	2.08(2)	1288(2)	5817(8)	1.005(1)	0.910(1)	0.740(1)	1.032(1)	369.0(5)	380.8(4)
1070	4520(3)	9.96	41.81	2075(21)	7.4(1)	0.340(6)	2.07(2)	1288(3)	5822(8)	1.005(1)	0.909(1)	0.738(1)	1.032(1)	370.6(5)	382.4(4)
1080	4526(3)	9.96	42.28	2079(21)	7.4(1)	0.339(6)	2.06(3)	1288(3)	5828(9)	1.005(2)	0.907(1)	0.735(1)	1.031(1)	372.3(5)	384.0(4)
1090	4531(3)	9.96	42.74	2082(21)	7.3(1)	0.338(6)	2.05(3)	1287(3)	5833(9)	1.005(2)	0.906(1)	0.732(1)	1.031(1)	374.0(5)	385.7(4)
1100	4537(3)	9.95	43.20	2086(21)	7.3(1)	0.336(6)	2.04(3)	1287(3)	5839(9)	1.005(2)	0.905(1)	0.729(1)	1.031(1)	375.7(5)	387.3(4)
1110	4542(3)	9.95	43.67	2089(21)	7.3(1)	0.335(6)	2.03(3)	1287(3)	5844(9)	1.005(2)	0.904(1)	0.726(1)	1.031(1)	377.3(5)	389.0(4)
1120	4548(3)	9.95	44.13	2092(21)	7.2(1)	0.334(6)	2.02(3)	1286(3)	5849(9)	1.005(2)	0.903(1)	0.724(1)	1.031(1)	379.0(5)	390.6(4)
1130	4553(3)	9.95	44.59	2096(21)	7.2(1)	0.333(6)	2.01(3)	1286(3)	5855(9)	1.005(2)	0.902(1)	0.721(1)	1.030(1)	380.7(5)	392.2(4)
1140	4559(3)	9.95	45.05	2099(21)	7.2(1)	0.332(6)	2.00(3)	1286(3)	5860(9)	1.005(2)	0.901(1)	0.718(1)	1.030(1)	382.4(5)	393.9(4)
1150	4564(3)	9.94	45.52	2102(21)	7.1(1)	0.330(6)	1.99(3)	1285(3)	5865(10)	1.004(2)	0.900(1)	0.716(1)	1.030(1)	384.0(5)	395.5(4)
1160	4569(3)	9.94	45.98	2106(21)	7.1(1)	0.329(6)	1.98(3)	1285(3)	5871(10)	1.004(2)	0.899(1)	0.713(1)	1.030(1)	385.7(5)	397.1(4)
1170	4575(3)	9.94	46.44	2109(21)	7.1(1)	0.328(6)	1.97(3)	1285(3)	5876(10)	1.004(2)	0.898(1)	0.710(1)	1.029(1)	387.4(5)	398.8(4)
1180	4580(3)	9.94	46.91	2112(21)	7.0(1)	0.327(6)	1.96(3)	1284(3)	5882(10)	1.004(2)	0.897(1)	0.708(1)	1.029(1)	389.0(5)	400.4(4)

1190	4585(3)	9.94	47.38	2115(21)	7.0(1)	0.326(6)	1.95(3)	1284(3)	5887(10)	1.004(2)	0.896(1)	0.705(1)	1.029(1)	390.7(5)	402.1(4)
1200	4591(3)	9.94	47.85	2119(21)	7.0(1)	0.325(6)	1.94(3)	1284(3)	5892(10)	1.004(2)	0.895(1)	0.703(1)	1.029(1)	392.4(5)	403.7(4)
1210	4596(3)	9.94	48.31	2122(21)	6.9(1)	0.324(6)	1.93(3)	1283(3)	5898(10)	1.004(2)	0.894(1)	0.700(1)	1.029(1)	394.1(5)	405.3(4)
1220	4601(3)	9.94	48.78	2125(21)	6.9(1)	0.323(6)	1.92(3)	1283(3)	5903(10)	1.004(2)	0.893(1)	0.697(1)	1.028(1)	395.8(5)	407.0(4)
1230	4607(3)	9.94	49.25	2128(21)	6.9(1)	0.321(6)	1.91(3)	1283(3)	5908(11)	1.004(2)	0.892(1)	0.695(1)	1.028(1)	397.4(5)	408.6(4)
1240	4612(3)	9.94	49.72	2132(22)	6.8(1)	0.320(7)	1.90(3)	1282(3)	5914(11)	1.004(2)	0.891(1)	0.692(1)	1.028(1)	399.1(5)	410.3(4)
1250	4617(3)	9.94	50.18	2135(22)	6.8(1)	0.319(7)	1.89(3)	1282(3)	5919(11)	1.004(2)	0.890(1)	0.690(1)	1.028(1)	400.8(5)	411.9(4)
1260	4623(3)	9.94	50.65	2138(22)	6.8(1)	0.318(7)	1.88(4)	1282(3)	5924(11)	1.004(2)	0.889(1)	0.687(1)	1.028(1)	402.5(5)	413.6(4)
1270	4628(3)	9.94	51.12	2141(22)	6.7(1)	0.317(7)	1.87(4)	1281(3)	5930(11)	1.004(2)	0.888(1)	0.685(1)	1.027(1)	404.1(5)	415.2(4)
1280	4633(3)	9.94	51.59	2144(22)	6.7(1)	0.316(7)	1.86(4)	1281(3)	5935(11)	1.004(2)	0.887(1)	0.682(1)	1.027(1)	405.8(5)	416.8(4)
1290	4638(3)	9.94	52.07	2147(22)	6.7(1)	0.315(7)	1.85(4)	1281(3)	5940(11)	1.004(2)	0.885(1)	0.680(1)	1.027(1)	407.5(5)	418.5(4)
1300	4644(3)	9.94	52.54	2151(22)	6.7(1)	0.314(7)	1.84(4)	1280(3)	5946(11)	1.004(2)	0.884(1)	0.678(1)	1.027(1)	409.2(5)	420.2(4)
1310	4649(3)	9.94	53.01	2154(22)	6.6(2)	0.313(7)	1.83(4)	1280(3)	5951(11)	1.004(2)	0.884(1)	0.675(1)	1.027(1)	410.9(5)	421.8(4)
1320	4654(3)	9.94	53.49	2157(22)	6.6(2)	0.312(7)	1.82(4)	1280(3)	5956(12)	1.004(2)	0.883(1)	0.673(1)	1.026(1)	412.6(5)	423.5(4)
1330	4659(3)	9.93	53.96	2160(22)	6.6(2)	0.311(7)	1.81(4)	1280(3)	5962(12)	1.004(2)	0.882(1)	0.670(1)	1.026(1)	414.2(5)	425.1(4)
1340	4664(3)	9.93	54.43	2163(22)	6.5(2)	0.310(7)	1.80(4)	1279(3)	5967(12)	1.004(2)	0.881(1)	0.668(1)	1.026(1)	415.9(5)	426.8(4)
1350	4670(3)	9.93	54.91	2166(22)	6.5(2)	0.309(7)	1.79(4)	1279(3)	5972(12)	1.003(2)	0.880(1)	0.666(1)	1.026(1)	417.6(5)	428.4(4)
1360	4675(3)	9.93	55.38	2169(22)	6.5(2)	0.308(8)	1.79(4)	1279(3)	5977(12)	1.003(2)	0.879(1)	0.663(1)	1.026(1)	419.3(5)	430.1(4)
1370	4680(3)	9.93	55.85	2172(22)	6.5(2)	0.307(8)	1.78(4)	1278(3)	5983(12)	1.003(2)	0.878(1)	0.661(1)	1.026(1)	420.9(5)	431.7(4)
1380	4685(3)	9.93	56.33	2175(22)	6.4(2)	0.306(8)	1.77(4)	1278(3)	5988(12)	1.003(2)	0.877(1)	0.659(1)	1.025(1)	422.6(5)	433.4(4)
1390	4690(3)	9.93	56.81	2179(22)	6.4(2)	0.305(8)	1.76(4)	1278(3)	5993(12)	1.003(2)	0.876(1)	0.657(2)	1.025(1)	424.3(5)	435.0(4)
1400	4695(3)	9.93	57.29	2182(22)	6.4(2)	0.304(8)	1.75(4)	1278(3)	5999(12)	1.003(2)	0.875(1)	0.654(2)	1.025(1)	426.0(5)	436.7(4)
1410	4701(3)	9.93	57.76	2185(22)	6.3(2)	0.303(8)	1.74(4)	1277(4)	6004(13)	1.003(2)	0.874(1)	0.652(2)	1.025(1)	427.7(5)	438.3(4)
1420	4706(3)	9.93	58.24	2188(22)	6.3(2)	0.302(8)	1.74(4)	1277(4)	6009(13)	1.003(2)	0.873(1)	0.650(2)	1.025(1)	429.4(5)	440.0(4)
1430	4711(4)	9.93	58.72	2191(22)	6.3(2)	0.301(8)	1.73(5)	1277(4)	6014(13)	1.003(2)	0.872(1)	0.647(2)	1.024(1)	431.1(5)	441.7(4)
1440	4716(4)	9.93	59.20	2194(22)	6.3(2)	0.300(8)	1.72(5)	1276(4)	6020(13)	1.003(2)	0.871(1)	0.645(2)	1.024(1)	432.8(5)	443.3(4)
1450	4721(4)	9.93	59.68	2197(22)	6.2(2)	0.299(8)	1.71(5)	1276(4)	6025(13)	1.003(2)	0.870(1)	0.643(2)	1.024(1)	434.5(5)	445.0(4)
1460	4726(4)	9.93	60.16	2200(22)	6.2(2)	0.298(8)	1.70(5)	1276(4)	6030(13)	1.003(2)	0.869(1)	0.641(2)	1.024(1)	436.2(5)	446.6(4)
1470	4731(4)	9.93	60.64	2203(22)	6.2(2)	0.298(9)	1.69(5)	1276(4)	6035(13)	1.003(2)	0.868(1)	0.639(2)	1.024(1)	437.8(5)	448.3(4)
1480	4736(4)	9.93	61.12	2206(22)	6.2(2)	0.297(9)	1.69(5)	1275(4)	6041(13)	1.003(2)	0.867(1)	0.637(2)	1.024(1)	439.5(5)	449.9(4)

1490	4741(4)	9.93	61.60	2209(23)	6.1(2)	0.296(9)	1.68(5)	1275(4)	6046(13)	1.003(2)	0.866(1)	0.634(2)	1.024(1)	441.2(5)	451.6(4)
1500	4747(4)	9.93	62.09	2212(23)	6.1(2)	0.295(9)	1.67(5)	1275(4)	6051(13)	1.003(2)	0.865(1)	0.632(2)	1.023(1)	442.9(5)	453.3(4)
1510	4752(4)	9.93	62.57	2215(23)	6.1(2)	0.294(9)	1.66(5)	1275(4)	6056(14)	1.003(2)	0.864(1)	0.630(2)	1.023(1)	444.6(5)	455.0(4)
1520	4757(4)	9.93	63.05	2217(23)	6.1(2)	0.293(9)	1.66(5)	1274(4)	6062(14)	1.003(2)	0.863(1)	0.628(2)	1.023(1)	446.3(5)	456.6(5)
1530	4762(4)	9.93	63.54	2220(23)	6.0(2)	0.292(9)	1.65(5)	1274(4)	6067(14)	1.003(2)	0.863(1)	0.626(2)	1.023(1)	448.0(5)	458.3(5)
1540	4767(4)	9.93	64.02	2223(23)	6.0(2)	0.291(9)	1.64(5)	1274(4)	6072(14)	1.003(2)	0.862(1)	0.624(2)	1.023(1)	449.7(5)	460.0(5)
1550	4772(4)	9.93	64.50	2226(23)	6.0(2)	0.290(9)	1.63(5)	1274(4)	6077(14)	1.003(2)	0.861(1)	0.622(2)	1.023(1)	451.4(5)	461.6(5)
1560	4777(4)	9.93	64.99	2229(23)	6.0(2)	0.289(9)	1.63(5)	1273(4)	6083(14)	1.002(2)	0.860(1)	0.620(2)	1.022(1)	453.1(5)	463.3(5)
1570	4782(4)	9.93	65.47	2232(23)	5.9(2)	0.289(9)	1.62(5)	1273(4)	6088(14)	1.002(2)	0.859(1)	0.618(2)	1.022(1)	454.8(5)	464.9(5)
1580	4787(4)	9.93	65.96	2235(23)	5.9(2)	0.288(10)	1.61(5)	1273(4)	6093(14)	1.002(2)	0.858(1)	0.616(2)	1.022(1)	456.5(5)	466.6(5)
1590	4792(4)	9.93	66.45	2238(23)	5.9(2)	0.287(10)	1.60(5)	1273(4)	6098(14)	1.002(2)	0.857(1)	0.614(2)	1.022(1)	458.2(5)	468.3(5)
1600	4797(4)	9.93	66.94	2241(23)	5.9(2)	0.286(10)	1.60(5)	1272(4)	6103(14)	1.002(2)	0.856(1)	0.612(2)	1.022(1)	459.9(5)	470.0(5)
1610	4802(4)	9.93	67.43	2243(23)	5.8(2)	0.285(10)	1.59(5)	1272(4)	6109(15)	1.002(2)	0.855(1)	0.610(2)	1.022(1)	461.6(5)	471.7(5)
1620	4807(4)	9.93	67.92	2246(23)	5.8(2)	0.284(10)	1.58(5)	1272(4)	6114(15)	1.002(2)	0.854(1)	0.608(2)	1.022(1)	463.3(5)	473.3(5)
1630	4812(4)	9.93	68.41	2249(23)	5.8(2)	0.283(10)	1.58(5)	1272(4)	6119(15)	1.002(2)	0.854(1)	0.606(2)	1.021(1)	465.0(5)	475.0(5)
1640	4817(4)	9.93	68.90	2252(23)	5.8(2)	0.283(10)	1.57(5)	1271(4)	6124(15)	1.002(2)	0.853(1)	0.604(2)	1.021(1)	466.7(5)	476.7(5)
1650	4822(4)	9.93	69.38	2255(23)	5.7(2)	0.282(10)	1.56(6)	1271(4)	6129(15)	1.002(2)	0.852(1)	0.602(2)	1.021(1)	468.4(5)	478.4(5)
1660	4827(4)	9.93	69.87	2258(23)	5.7(2)	0.281(10)	1.55(6)	1271(4)	6135(15)	1.002(2)	0.851(1)	0.600(2)	1.021(1)	470.1(5)	480.0(5)
1670	4831(4)	9.93	70.36	2260(23)	5.7(2)	0.280(10)	1.55(6)	1271(4)	6140(15)	1.002(2)	0.850(1)	0.598(2)	1.021(1)	471.8(5)	481.7(5)
1680	4836(4)	9.93	70.86	2263(23)	5.7(2)	0.279(10)	1.54(6)	1271(4)	6145(15)	1.002(2)	0.849(1)	0.596(2)	1.021(1)	473.5(5)	483.4(5)
1690	4841(4)	9.93	71.35	2266(23)	5.7(2)	0.279(11)	1.53(6)	1270(4)	6150(15)	1.002(2)	0.848(1)	0.594(2)	1.021(1)	475.3(5)	485.1(5)
1700	4846(4)	9.94	71.85	2269(23)	5.6(2)	0.278(11)	1.53(6)	1270(4)	6155(15)	1.002(2)	0.848(1)	0.592(2)	1.021(1)	477.0(5)	486.8(5)
1710	4851(4)	9.94	72.34	2272(23)	5.6(2)	0.277(11)	1.52(6)	1270(4)	6161(15)	1.002(2)	0.847(1)	0.590(2)	1.020(1)	478.7(5)	488.4(5)
1720	4856(4)	9.94	72.84	2274(24)	5.6(2)	0.276(11)	1.51(6)	1270(4)	6166(15)	1.002(2)	0.846(1)	0.588(2)	1.020(1)	480.4(5)	490.1(5)
1730	4861(4)	9.94	73.33	2277(24)	5.6(2)	0.276(11)	1.51(6)	1269(4)	6171(15)	1.002(2)	0.845(1)	0.586(2)	1.020(1)	482.1(5)	491.8(5)
1740	4866(4)	9.94	73.83	2280(24)	5.5(2)	0.275(11)	1.50(6)	1269(4)	6176(16)	1.002(2)	0.844(1)	0.585(2)	1.020(1)	483.8(5)	493.5(5)
1750	4871(4)	9.95	74.32	2283(24)	5.5(2)	0.274(11)	1.50(6)	1269(4)	6181(16)	1.002(2)	0.843(1)	0.583(2)	1.020(1)	485.5(5)	495.2(5)
1760	4876(4)	9.95	74.82	2285(24)	5.5(2)	0.273(11)	1.49(6)	1269(4)	6186(16)	1.002(2)	0.842(1)	0.581(2)	1.020(1)	487.2(5)	496.9(5)
1770	4880(4)	9.95	75.31	2288(24)	5.5(2)	0.273(11)	1.48(6)	1269(4)	6191(16)	1.001(2)	0.842(1)	0.579(2)	1.020(1)	488.9(5)	498.5(5)
1780	4885(4)	9.95	75.81	2291(24)	5.5(2)	0.272(11)	1.48(6)	1268(4)	6197(16)	1.001(2)	0.841(1)	0.577(2)	1.020(1)	490.7(5)	500.2(5)

1790	4890(4)	9.95	76.31	2294(24)	5.4(2)	0.271(11)	1.47(6)	1268(4)	6202(16)	1.001(2)	0.840(1)	0.575(2)	1.019(1)	492.4(5)	501.9(5)
1800	4895(4)	9.95	76.81	2296(24)	5.4(2)	0.270(11)	1.46(6)	1268(4)	6207(16)	1.001(2)	0.839(1)	0.574(2)	1.019(1)	494.1(5)	503.6(5)
1810	4900(4)	9.95	77.31	2299(24)	5.4(2)	0.270(12)	1.46(6)	1268(4)	6212(16)	1.001(2)	0.838(1)	0.572(2)	1.019(1)	495.8(5)	505.3(5)
1820	4905(4)	9.95	77.81	2302(24)	5.4(2)	0.269(12)	1.45(6)	1268(4)	6217(16)	1.001(2)	0.837(1)	0.570(2)	1.019(1)	497.6(5)	507.0(5)
1830	4910(4)	9.96	78.31	2304(24)	5.4(2)	0.268(12)	1.45(6)	1267(4)	6222(16)	1.001(2)	0.837(1)	0.568(2)	1.019(1)	499.3(5)	508.7(5)
1840	4914(5)	9.96	78.81	2307(24)	5.3(2)	0.267(12)	1.44(6)	1267(4)	6228(16)	1.001(3)	0.836(1)	0.567(2)	1.019(1)	501.0(5)	510.4(5)
1850	4919(5)	9.96	79.31	2310(24)	5.3(2)	0.267(12)	1.43(6)	1267(4)	6233(16)	1.001(3)	0.835(1)	0.565(2)	1.019(1)	502.7(5)	512.1(5)
1860	4924(5)	9.96	79.82	2312(24)	5.3(2)	0.266(12)	1.43(6)	1267(4)	6238(16)	1.001(3)	0.834(1)	0.563(2)	1.019(1)	504.4(5)	513.8(5)
1870	4929(5)	9.96	80.32	2315(24)	5.3(2)	0.265(12)	1.42(6)	1267(4)	6243(16)	1.001(3)	0.833(1)	0.561(2)	1.018(1)	506.1(5)	515.5(5)
1880	4934(5)	9.96	80.82	2318(24)	5.3(2)	0.265(12)	1.42(6)	1266(4)	6248(17)	1.001(3)	0.832(1)	0.560(2)	1.018(1)	507.9(5)	517.2(5)
1890	4939(5)	9.97	81.33	2320(24)	5.2(2)	0.264(12)	1.41(6)	1266(4)	6253(17)	1.001(3)	0.832(1)	0.558(2)	1.018(1)	509.6(5)	518.9(5)
1900	4943(5)	9.97	81.84	2323(24)	5.2(2)	0.263(12)	1.40(6)	1266(4)	6258(17)	1.001(3)	0.831(1)	0.556(2)	1.018(1)	511.4(5)	520.6(5)
1910	4948(5)	9.97	82.34	2326(24)	5.2(2)	0.263(12)	1.40(6)	1266(4)	6263(17)	1.001(3)	0.830(1)	0.554(2)	1.018(1)	513.1(5)	522.3(5)
1920	4953(5)	9.97	82.85	2328(25)	5.2(2)	0.262(12)	1.39(6)	1266(4)	6269(17)	1.001(3)	0.829(1)	0.553(2)	1.018(1)	514.8(5)	524.0(5)
1930	4958(5)	9.98	83.36	2331(25)	5.2(2)	0.261(12)	1.39(6)	1265(4)	6274(17)	1.001(3)	0.828(1)	0.551(2)	1.018(1)	516.6(5)	525.8(5)
1940	4962(5)	9.98	83.86	2334(25)	5.1(2)	0.261(13)	1.38(6)	1265(4)	6279(17)	1.001(3)	0.828(1)	0.549(2)	1.018(1)	518.3(5)	527.5(5)
1950	4967(5)	9.98	84.37	2336(25)	5.1(2)	0.260(13)	1.38(7)	1265(4)	6284(17)	1.001(3)	0.827(1)	0.548(2)	1.018(1)	520.0(5)	529.2(5)
1960	4972(5)	9.99	84.88	2339(25)	5.1(2)	0.260(13)	1.37(7)	1265(4)	6289(17)	1.001(3)	0.826(1)	0.546(2)	1.018(1)	521.7(5)	530.9(5)
1970	4977(5)	9.99	85.38	2341(25)	5.1(3)	0.259(13)	1.36(7)	1265(4)	6294(17)	1.001(3)	0.825(1)	0.544(2)	1.017(1)	523.5(5)	532.6(5)
1980	4981(5)	9.99	85.90	2344(25)	5.1(3)	0.258(13)	1.36(7)	1265(4)	6299(17)	1.001(3)	0.825(1)	0.543(2)	1.017(1)	525.2(5)	534.3(5)
1990	4986(5)	10.00	86.41	2346(25)	5.0(3)	0.258(13)	1.35(7)	1264(4)	6304(17)	1.001(3)	0.824(1)	0.541(2)	1.017(1)	527.0(5)	536.0(5)
2000	4991(5)	10.00	86.92	2349(25)	5.0(3)	0.257(13)	1.35(7)	1264(4)	6309(17)	1.0(3)	0.823(1)	0.540(2)	1.017(1)	528.7(5)	537.7(5)
2010	4996(5)	10.00	87.43	2352(25)	5.0(3)	0.256(13)	1.34(7)	1264(5)	6315(17)	1.0(3)	0.822(1)	0.538(2)	1.017(1)	530.4(5)	539.5(5)
2020	5000(5)	10.00	87.95	2354(25)	5.0(3)	0.256(13)	1.34(7)	1264(5)	6320(17)	1.0(3)	0.821(1)	0.536(2)	1.017(1)	532.2(5)	541.2(5)
2030	5005(5)	10.01	88.46	2357(25)	5.0(3)	0.255(13)	1.33(7)	1264(5)	6325(17)	1.0(3)	0.821(1)	0.535(2)	1.017(1)	533.9(5)	542.9(5)
2040	5010(5)	10.01	88.97	2359(25)	4.9(3)	0.255(13)	1.33(7)	1263(5)	6330(18)	1.0(3)	0.820(1)	0.533(2)	1.017(1)	535.7(5)	544.6(5)
2050	5015(5)	10.01	89.49	2362(25)	4.9(3)	0.254(13)	1.32(7)	1263(5)	6335(18)	1.0(3)	0.819(1)	0.532(2)	1.017(1)	537.4(5)	546.3(5)
2060	5019(5)	10.02	90.00	2364(25)	4.9(3)	0.253(13)	1.32(7)	1263(5)	6340(18)	1.0(3)	0.818(1)	0.530(2)	1.017(1)	539.2(5)	548.1(5)
2070	5024(5)	10.02	90.51	2367(25)	4.9(3)	0.253(13)	1.31(7)	1263(5)	6346(18)	1.0(3)	0.818(1)	0.528(2)	1.016(1)	540.9(5)	549.8(5)
2080	5029(5)	10.02	91.03	2369(25)	4.9(3)	0.252(14)	1.31(7)	1263(5)	6352(18)	1.0(3)	0.817(1)	0.527(2)	1.016(1)	542.7(5)	551.5(5)

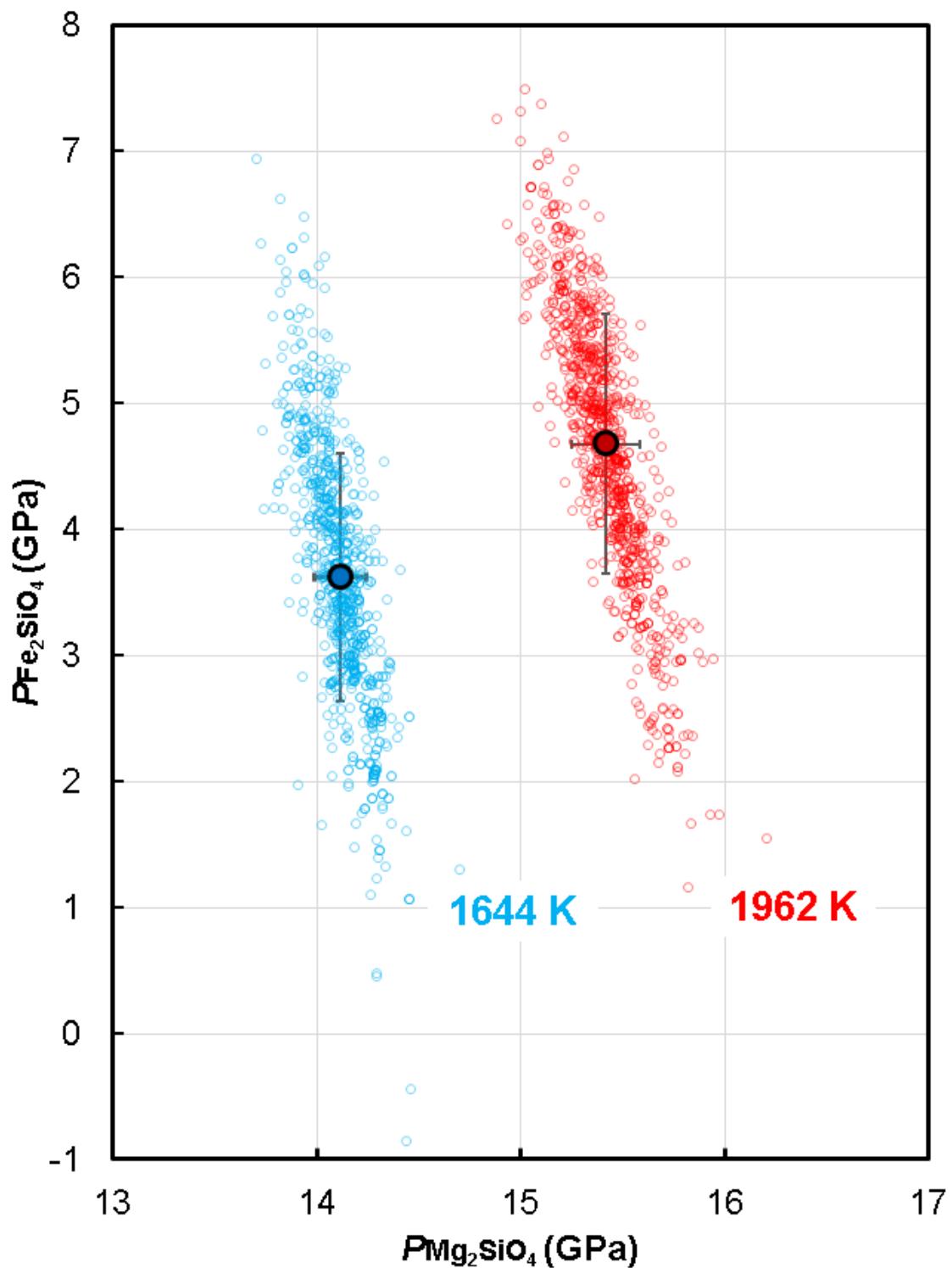
2090	5033(5)	10.03	91.55	2372(26)	4.9(3)	0.251(14)	1.30(7)	1263(5)	6357(18)	1.0(3)	0.816(1)	0.525(2)	1.016(1)	544.4(5)	553.2(5)
2100	5038(5)	10.03	92.07	2374(26)	4.8(3)	0.251(14)	1.30(7)	1263(5)	6363(18)	1.0(3)	0.815(1)	0.524(2)	1.016(1)	546.2(5)	555.0(5)
2110	5043(5)	10.03	92.59	2377(26)	4.8(3)	0.250(14)	1.29(7)	1263(5)	6368(18)	1.0(3)	0.814(1)	0.522(2)	1.016(1)	547.9(5)	556.7(5)
2120	5048(5)	10.03	93.11	2379(26)	4.8(3)	0.250(14)	1.29(7)	1263(5)	6374(18)	1.001(3)	0.814(1)	0.521(2)	1.016(1)	549.7(5)	558.4(5)
2130	5052(5)	10.04	93.63	2382(26)	4.8(3)	0.249(14)	1.28(7)	1263(5)	6379(18)	1.001(3)	0.813(1)	0.519(2)	1.016(1)	551.4(5)	560.2(5)
2140	5057(5)	10.04	94.15	2384(26)	4.8(3)	0.248(14)	1.28(7)	1263(5)	6385(18)	1.001(3)	0.812(1)	0.518(2)	1.016(1)	553.2(5)	561.9(5)
2150	5062(5)	10.04	94.67	2387(26)	4.7(3)	0.248(14)	1.27(7)	1263(5)	6390(18)	1.001(3)	0.811(1)	0.516(2)	1.016(1)	555.0(5)	563.7(5)
2160	5066(5)	10.05	95.19	2389(26)	4.7(3)	0.247(14)	1.27(7)	1262(5)	6396(18)	1.001(3)	0.811(1)	0.515(2)	1.016(1)	556.7(5)	565.4(5)
2170	5071(5)	10.05	95.71	2392(26)	4.7(3)	0.247(14)	1.26(7)	1262(5)	6401(19)	1.001(3)	0.810(1)	0.513(2)	1.015(1)	558.5(5)	567.1(5)
2180	5076(5)	10.05	96.24	2394(26)	4.7(3)	0.246(14)	1.26(7)	1262(5)	6407(19)	1.001(3)	0.809(1)	0.512(2)	1.015(1)	560.2(5)	568.9(5)
2190	5080(5)	10.06	96.77	2397(26)	4.7(3)	0.246(14)	1.25(7)	1262(5)	6413(19)	1.001(3)	0.808(1)	0.510(2)	1.015(1)	562.0(5)	570.6(5)
2200	5085(5)	10.06	97.29	2399(26)	4.7(3)	0.245(14)	1.25(7)	1262(5)	6418(19)	1.001(3)	0.808(1)	0.509(2)	1.015(1)	563.8(5)	572.4(5)
2210	5090(5)	10.07	97.82	2402(26)	4.6(3)	0.244(14)	1.24(7)	1262(5)	6424(19)	1.001(3)	0.807(1)	0.507(2)	1.015(1)	565.6(5)	574.1(5)
2220	5094(5)	10.07	98.35	2404(26)	4.6(3)	0.244(14)	1.24(7)	1262(5)	6429(19)	1.001(3)	0.806(1)	0.506(2)	1.015(1)	567.3(5)	575.9(5)
2230	5099(5)	10.08	98.87	2407(27)	4.6(3)	0.243(15)	1.23(7)	1262(5)	6435(19)	1.001(3)	0.806(1)	0.504(2)	1.015(1)	569.1(5)	577.6(5)
2240	5103(5)	10.08	99.40	2409(27)	4.6(3)	0.243(15)	1.23(7)	1262(5)	6440(19)	1.001(3)	0.805(1)	0.503(2)	1.015(1)	570.9(5)	579.4(5)
2250	5108(5)	10.09	99.93	2411(27)	4.6(3)	0.242(15)	1.22(7)	1262(5)	6446(19)	1.001(3)	0.804(1)	0.501(2)	1.015(1)	572.6(5)	581.1(5)
2260	5113(5)	10.09	100.46	2414(27)	4.6(3)	0.242(15)	1.22(7)	1262(5)	6451(19)	1.001(3)	0.803(1)	0.500(2)	1.015(1)	574.4(5)	582.9(5)
2270	5117(5)	10.10	100.98	2416(27)	4.5(3)	0.241(15)	1.21(7)	1262(5)	6457(19)	1.001(3)	0.803(1)	0.499(2)	1.015(1)	576.2(5)	584.6(5)
2280	5122(5)	10.11	101.52	2419(27)	4.5(3)	0.241(15)	1.21(7)	1262(5)	6462(19)	1.001(3)	0.802(1)	0.497(2)	1.015(1)	578.0(5)	586.4(5)
2290	5127(5)	10.11	102.05	2421(27)	4.5(3)	0.240(15)	1.21(7)	1262(5)	6468(19)	1.001(3)	0.801(1)	0.496(2)	1.014(1)	579.8(5)	588.2(5)
2300	5131(5)	10.12	102.59	2424(27)	4.5(3)	0.240(15)	1.20(7)	1262(5)	6473(19)	1.001(3)	0.800(1)	0.494(2)	1.014(1)	581.6(5)	589.9(5)
2310	5136(5)	10.12	103.12	2426(27)	4.5(3)	0.239(15)	1.20(7)	1261(5)	6479(19)	1.001(3)	0.800(1)	0.493(2)	1.014(1)	583.3(5)	591.7(5)
2320	5141(5)	10.13	103.66	2428(27)	4.5(3)	0.239(15)	1.19(7)	1261(5)	6484(19)	1.001(3)	0.799(1)	0.492(2)	1.014(1)	585.1(5)	593.5(5)
2330	5145(5)	10.14	104.19	2431(27)	4.4(3)	0.238(15)	1.19(7)	1261(5)	6490(19)	1.001(3)	0.798(1)	0.490(2)	1.014(1)	586.9(5)	595.2(5)
2340	5150(6)	10.14	104.73	2433(27)	4.4(3)	0.238(15)	1.18(7)	1261(5)	6495(19)	1.001(3)	0.798(1)	0.489(2)	1.014(1)	588.7(5)	597.0(5)
2350	5154(6)	10.15	105.26	2435(28)	4.4(3)	0.237(15)	1.18(7)	1261(5)	6501(19)	1.001(3)	0.797(1)	0.487(2)	1.014(1)	590.5(5)	598.8(5)
2360	5159(6)	10.15	105.80	2438(28)	4.4(3)	0.237(15)	1.17(7)	1261(5)	6506(19)	1.001(3)	0.796(1)	0.486(2)	1.014(1)	592.3(5)	600.5(5)
2370	5163(6)	10.16	106.33	2440(28)	4.4(3)	0.236(15)	1.17(7)	1261(5)	6511(19)	1.001(3)	0.795(1)	0.485(2)	1.014(1)	594.1(5)	602.3(5)
2380	5168(6)	10.17	106.88	2443(28)	4.4(3)	0.236(15)	1.17(7)	1261(5)	6517(19)	1.001(3)	0.795(1)	0.483(2)	1.014(1)	595.9(5)	604.1(5)

2390	5173(6)	10.17	107.42	2445(28)	4.4(3)	0.235(16)	1.16(7)	1261(5)	6522(19)	1.001(3)	0.794(1)	0.482(2)	1.014(1)	597.7(5)	605.9(5)
2400	5177(6)	10.18	107.96	2447(28)	4.3(3)	0.235(16)	1.16(7)	1261(5)	6528(20)	1.001(3)	0.793(1)	0.481(2)	1.014(1)	599.5(5)	607.7(5)
2410	5182(6)	10.19	108.51	2450(28)	4.3(3)	0.235(16)	1.15(7)	1261(5)	6533(20)	1.001(3)	0.793(1)	0.479(2)	1.014(1)	601.3(5)	609.5(5)
2420	5187(6)	10.19	109.05	2452(28)	4.3(3)	0.234(16)	1.15(7)	1261(5)	6539(20)	1.001(3)	0.792(1)	0.478(2)	1.013(1)	603.1(5)	611.2(5)
2430	5191(6)	10.20	109.59	2454(28)	4.3(3)	0.234(16)	1.14(7)	1261(5)	6544(20)	1.001(3)	0.791(1)	0.477(2)	1.013(1)	604.9(5)	613.0(5)
2440	5196(6)	10.21	110.14	2457(28)	4.3(3)	0.233(16)	1.14(7)	1261(5)	6550(20)	1.001(3)	0.791(1)	0.475(2)	1.013(1)	606.7(5)	614.8(5)
2450	5200(6)	10.22	110.68	2459(28)	4.3(3)	0.233(16)	1.14(7)	1261(5)	6555(20)	1.001(3)	0.790(1)	0.474(2)	1.013(1)	608.5(5)	616.6(5)
2460	5205(6)	10.22	111.22	2461(29)	4.2(3)	0.232(16)	1.13(7)	1261(5)	6561(20)	1.001(3)	0.789(1)	0.473(2)	1.013(1)	610.3(5)	618.4(5)
2470	5209(6)	10.23	111.77	2464(29)	4.2(3)	0.232(16)	1.13(7)	1260(5)	6566(20)	1.001(3)	0.788(1)	0.471(2)	1.013(1)	612.1(5)	620.2(5)
2480	5214(6)	10.24	112.32	2466(29)	4.2(3)	0.232(16)	1.12(7)	1260(5)	6572(20)	1.001(3)	0.788(1)	0.470(2)	1.013(1)	614.0(5)	622.0(5)
2490	5219(6)	10.25	112.87	2468(29)	4.2(3)	0.231(16)	1.12(7)	1260(5)	6577(20)	1.001(3)	0.787(1)	0.469(2)	1.013(1)	615.8(5)	623.8(5)
2500	5223(6)	10.25	113.42	2471(29)	4.2(3)	0.231(16)	1.12(7)	1260(5)	6582(20)	1.002(3)	0.786(1)	0.467(2)	1.013(1)	617.6(5)	625.6(5)
2510	5228(6)	10.26	113.98	2473(29)	4.2(3)	0.230(16)	1.11(7)	1260(5)	6588(20)	1.002(3)	0.786(1)	0.466(2)	1.013(1)	619.5(5)	627.4(5)
2520	5232(6)	10.27	114.53	2475(29)	4.2(3)	0.230(16)	1.11(7)	1260(5)	6593(20)	1.002(3)	0.785(1)	0.465(2)	1.013(1)	621.3(5)	629.2(5)
2530	5237(6)	10.28	115.08	2478(29)	4.1(3)	0.229(16)	1.10(8)	1260(5)	6599(20)	1.002(3)	0.784(1)	0.464(2)	1.013(1)	623.1(5)	631.0(5)
2540	5241(6)	10.29	115.63	2480(29)	4.1(3)	0.229(16)	1.10(8)	1260(5)	6604(20)	1.002(3)	0.784(1)	0.462(2)	1.013(1)	625.0(5)	632.9(5)
2550	5246(6)	10.29	116.19	2482(29)	4.1(3)	0.229(16)	1.10(8)	1260(5)	6610(20)	1.002(3)	0.783(1)	0.461(2)	1.013(1)	626.8(5)	634.7(5)
2560	5251(6)	10.30	116.74	2484(29)	4.1(3)	0.228(17)	1.09(8)	1260(5)	6615(20)	1.002(3)	0.782(1)	0.460(2)	1.012(1)	628.6(5)	636.5(5)
2570	5255(6)	10.31	117.29	2487(30)	4.1(3)	0.228(17)	1.09(8)	1260(5)	6620(20)	1.002(3)	0.782(1)	0.459(2)	1.012(1)	630.4(5)	638.3(5)
2580	5260(6)	10.32	117.85	2489(30)	4.1(3)	0.227(17)	1.08(8)	1260(5)	6626(20)	1.002(3)	0.781(1)	0.457(2)	1.012(1)	632.3(5)	640.1(5)
2590	5264(6)	10.33	118.42	2491(30)	4.1(3)	0.227(17)	1.08(8)	1260(5)	6631(20)	1.002(3)	0.780(1)	0.456(2)	1.012(1)	634.2(5)	641.9(5)
2600	5269(6)	10.34	118.98	2494(30)	4.0(3)	0.227(17)	1.08(8)	1260(5)	6637(20)	1.002(3)	0.780(1)	0.455(2)	1.012(1)	636.0(5)	643.8(5)
2610	5273(6)	10.35	119.54	2496(30)	4.0(3)	0.226(17)	1.07(8)	1260(5)	6642(20)	1.002(3)	0.779(1)	0.454(2)	1.012(1)	637.9(5)	645.6(5)
2620	5278(6)	10.36	120.10	2498(30)	4.0(3)	0.226(17)	1.07(8)	1260(5)	6648(21)	1.002(3)	0.778(1)	0.452(2)	1.012(1)	639.7(5)	647.4(5)
2630	5283(6)	10.37	120.67	2500(30)	4.0(3)	0.226(17)	1.07(8)	1259(5)	6653(21)	1.002(3)	0.778(1)	0.451(2)	1.012(1)	641.6(5)	649.3(5)
2640	5287(6)	10.38	121.23	2503(30)	4.0(3)	0.225(17)	1.06(8)	1259(5)	6659(21)	1.002(3)	0.777(1)	0.450(2)	1.012(1)	643.4(5)	651.1(5)
2650	5292(6)	10.39	121.79	2505(30)	4.0(3)	0.225(17)	1.06(8)	1259(5)	6664(21)	1.002(3)	0.776(1)	0.449(2)	1.012(1)	645.3(5)	652.9(5)
2660	5296(6)	10.40	122.35	2507(31)	4.0(3)	0.225(17)	1.05(8)	1259(5)	6669(21)	1.002(3)	0.776(1)	0.447(2)	1.012(1)	647.1(5)	654.8(5)
2670	5301(6)	10.41	122.92	2509(31)	3.9(3)	0.224(17)	1.05(8)	1259(5)	6675(21)	1.002(3)	0.775(1)	0.446(2)	1.012(1)	649.0(5)	656.6(5)
2680	5305(6)	10.42	123.49	2512(31)	3.9(3)	0.224(17)	1.05(8)	1259(5)	6680(21)	1.002(3)	0.774(1)	0.445(2)	1.012(1)	650.8(5)	658.5(5)

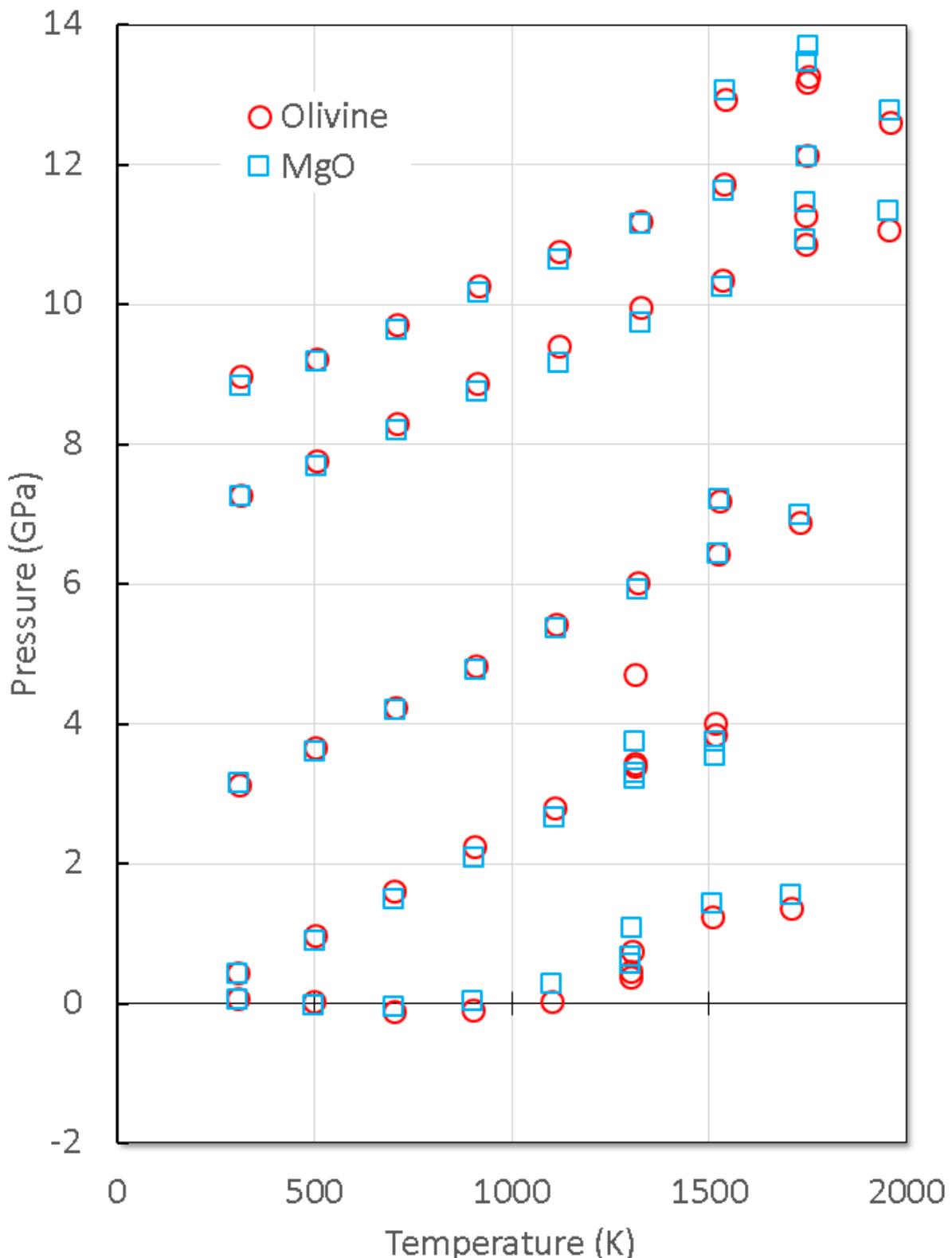
2690	5310(6)	10.43	124.06	2514(31)	3.9(3)	0.224(17)	1.04(8)	1259(5)	6686(21)	1.002(3)	0.774(1)	0.444(2)	1.012(1)	652.7(5)	660.3(5)
2700	5314(6)	10.44	124.63	2516(31)	3.9(3)	0.223(17)	1.04(8)	1259(5)	6691(21)	1.002(3)	0.773(1)	0.443(2)	1.012(1)	654.6(5)	662.2(5)
2710	5319(6)	10.45	125.20	2518(31)	3.9(3)	0.223(17)	1.04(8)	1259(5)	6697(21)	1.002(3)	0.772(1)	0.441(2)	1.012(1)	656.5(5)	664.0(5)
2720	5323(6)	10.46	125.77	2520(31)	3.9(3)	0.223(17)	1.03(8)	1259(5)	6702(21)	1.002(3)	0.772(1)	0.440(2)	1.011(1)	658.4(5)	665.9(5)
2730	5328(6)	10.47	126.35	2523(31)	3.9(3)	0.222(18)	1.03(8)	1259(5)	6707(21)	1.002(3)	0.771(1)	0.439(2)	1.011(1)	660.2(5)	667.8(5)
2740	5332(6)	10.48	126.92	2525(31)	3.9(3)	0.222(18)	1.03(8)	1259(5)	6713(21)	1.002(3)	0.770(1)	0.438(2)	1.011(1)	662.1(5)	669.6(5)
2750	5337(6)	10.49	127.49	2527(32)	3.8(3)	0.222(18)	1.02(8)	1259(5)	6718(21)	1.002(3)	0.770(1)	0.437(2)	1.011(1)	664.0(5)	671.5(5)
2760	5342(6)	10.51	128.07	2529(32)	3.8(3)	0.221(18)	1.02(8)	1259(5)	6724(21)	1.002(3)	0.769(1)	0.436(2)	1.011(1)	665.9(5)	673.4(5)
2770	5346(6)	10.52	128.65	2532(32)	3.8(3)	0.221(18)	1.01(8)	1259(5)	6729(21)	1.002(3)	0.768(1)	0.434(2)	1.011(1)	667.8(5)	675.2(5)
2780	5351(6)	10.53	129.23	2534(32)	3.8(3)	0.221(18)	1.01(8)	1259(5)	6735(21)	1.002(3)	0.768(1)	0.433(2)	1.011(1)	669.7(5)	677.1(5)
2790	5355(6)	10.54	129.82	2536(32)	3.8(3)	0.221(18)	1.01(8)	1259(5)	6740(21)	1.002(3)	0.767(1)	0.432(2)	1.011(1)	671.6(5)	679.0(5)
2800	5360(6)	10.56	130.40	2538(32)	3.8(3)	0.220(18)	1.00(8)	1259(5)	6745(21)	1.002(3)	0.766(1)	0.431(2)	1.011(1)	673.5(5)	680.9(5)

572 #: taken from PREM [Dziewonski and Anderson, 1981].

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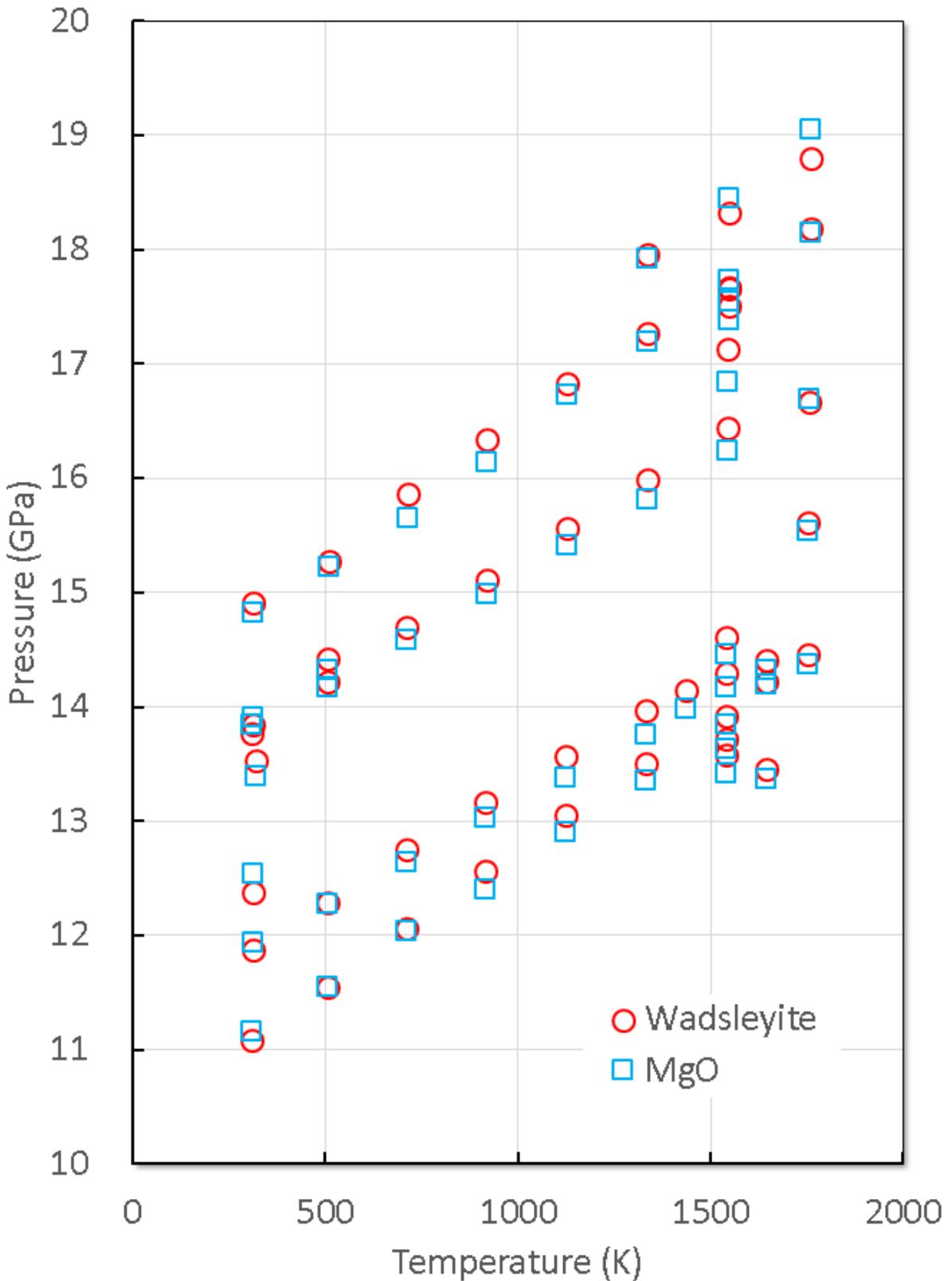


574  
 575 **Figure S1.** Correlation of the endmember transition pressures of the olivine wadsleyite  
 576 transition. Determined by the data from Katsura *et al.* [2004a], corrected using Nishihara *et*  
 577 *al.*'s [2020] thermocouple correction. Each point show the transition pressure obtained in  
 578 each replica data set of the Monte Carlo simulation.



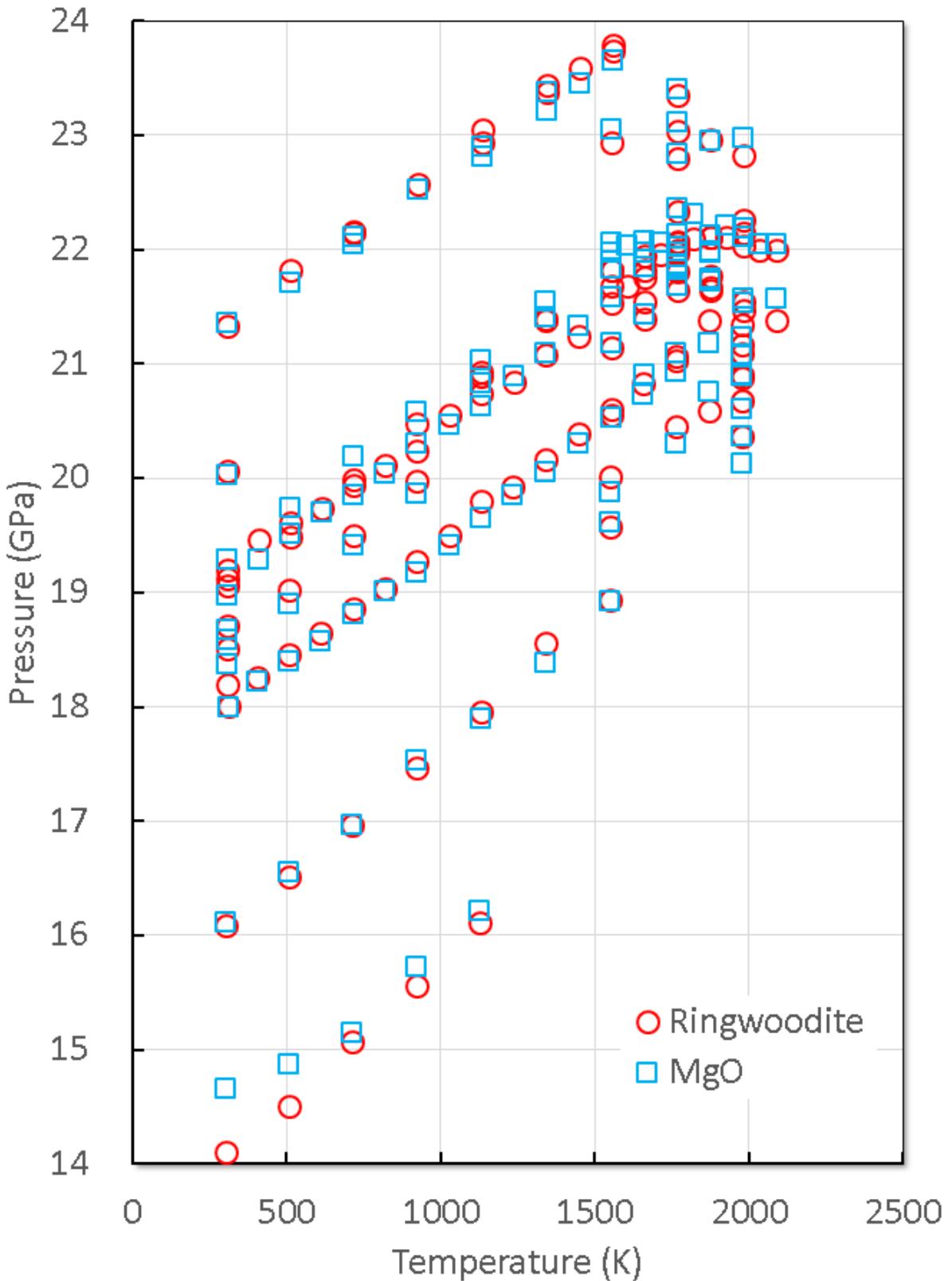
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**Figure S2.** Comparison of the pressures using Tange *et al.*'s [2009] MgO and the current study's olivine EOS's.



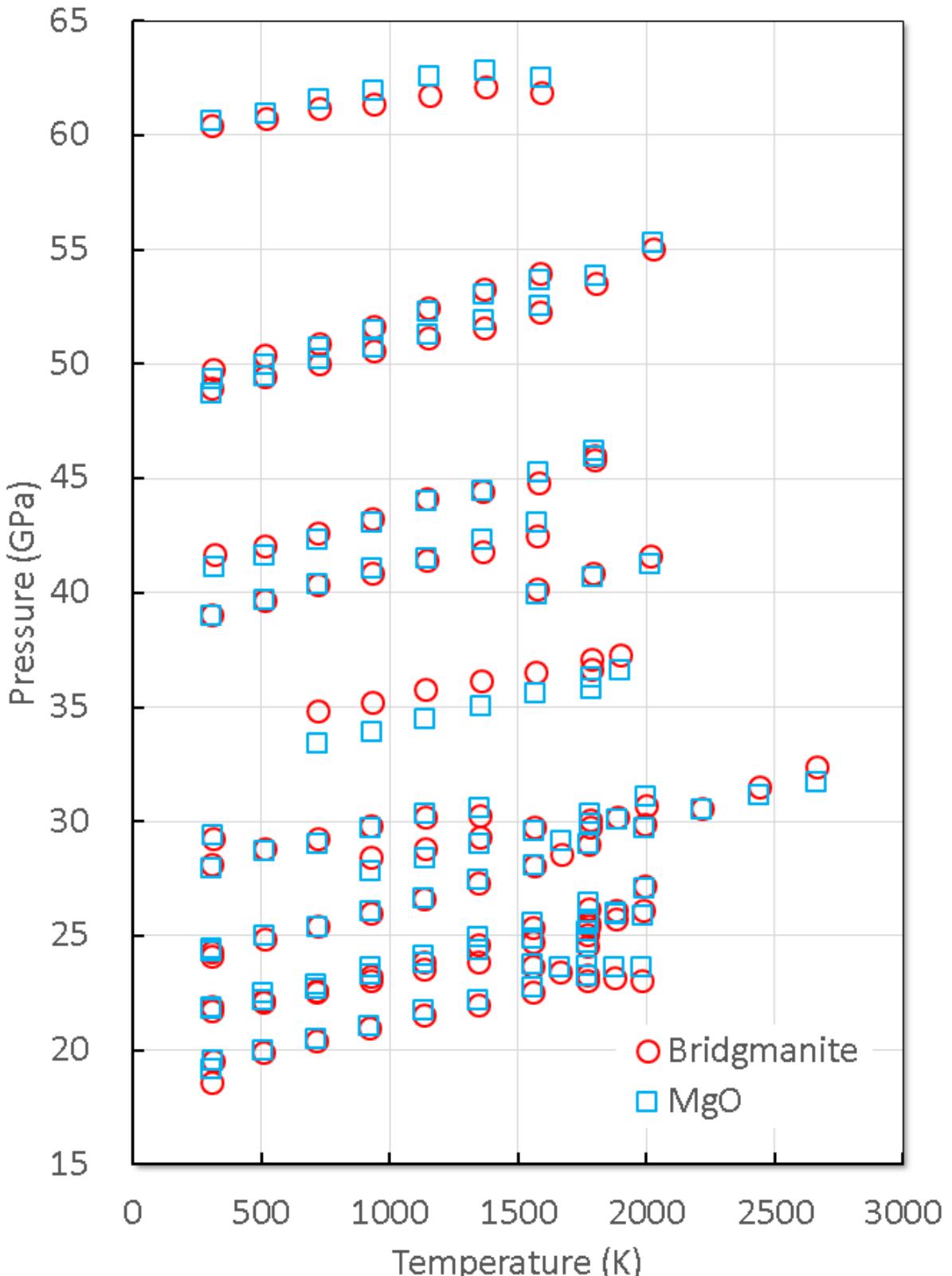
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**Figure S3.** Comparison of the pressures using Tange *et al.*'s [2009] MgO and the current study's wadsleyite EOS's.

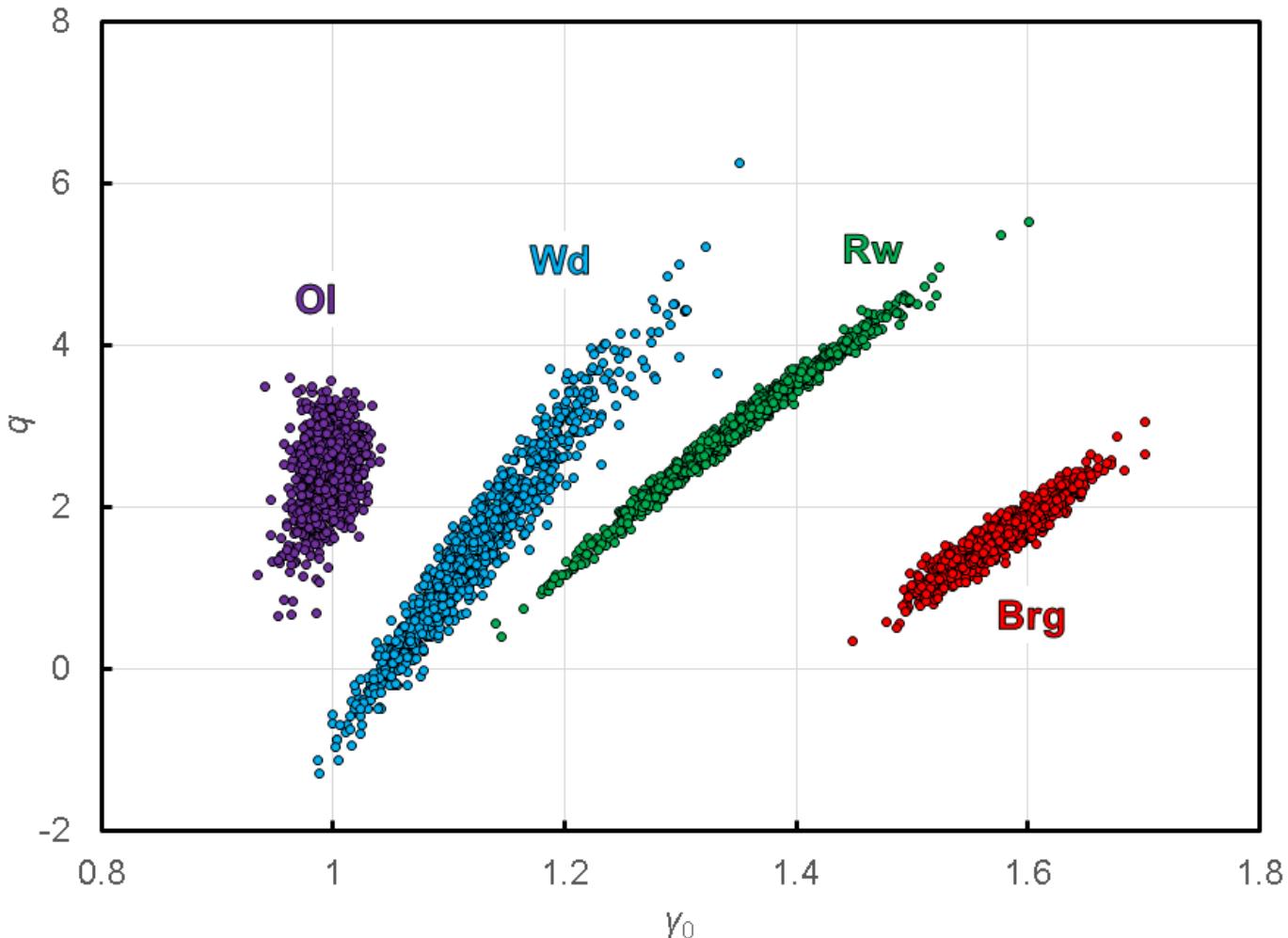


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**Figure S4.** Comparison of the pressures using Tange *et al.*'s [2009] MgO and the current study's ringwoodite EOS's.



588  
589 **Figure S5.** Comparison of the pressures using Tange *et al.*'s [2009] MgO and the current  
590 study's bridgmanite EOS's.  
591 .



592  
593 **Figure S6.** The correlations of  $\gamma_0$  and  $q$ . Violet: olivine, blue: wadsleyite, green: ringwoodite,  
594 red: bridgemanite. Original data from Katsura *et al.* [2004a; 2009a; 2009b; 2009c] are corrected  
595 using Nishihara *et al.*'s [2020] thermocouple correction.  
596