

## **Climatology of marine shallow-cloud-top radiative cooling**

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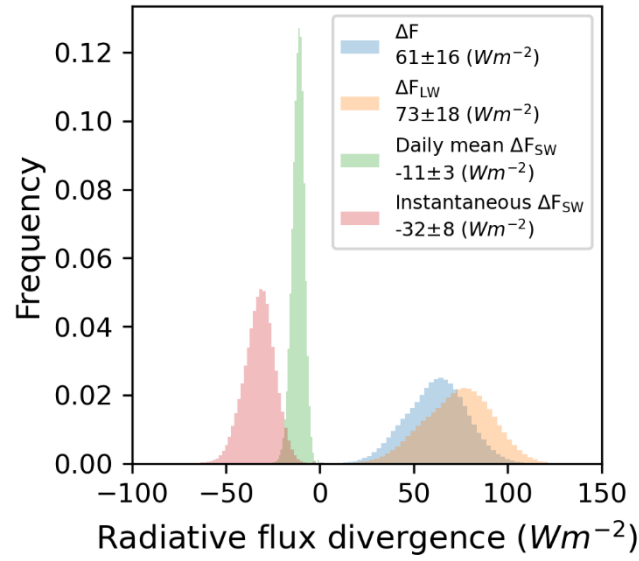
3. Table S1

## **Text S1: Radiative transfer model**

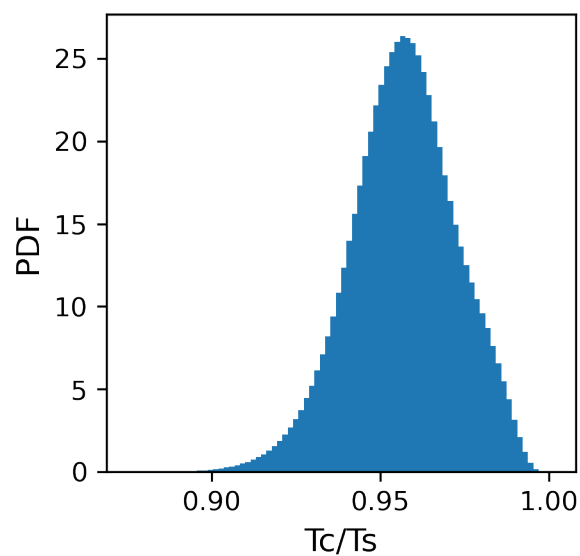
The radiative transfer model we use is the Santa Barbara DISORT Atmospheric Radiative Transfer model (Ricchiazzi et al., 1998). We specify the vertical grids with resolutions of 50 m from the surface to 2.25 km and the grid spacing increases with the altitude until the top of the atmosphere, leading to a total of  $\sim 60$  grids in the vertical. The ozone profile and greenhouse gas concentrations are set to default values. The cloud optical depth is uniformly distributed throughout the cloud layer. The wavelength ranges of longwave and shortwave are set as  $5 \sim 40 \mu\text{m}$  and  $0.1 \sim 5 \mu\text{m}$ , respectively. The wavelength increment is  $0.1 \mu\text{m}$  for shortwave and  $0.2 \mu\text{m}$  for longwave.

## **Text S2: Configuration of the neural network model**

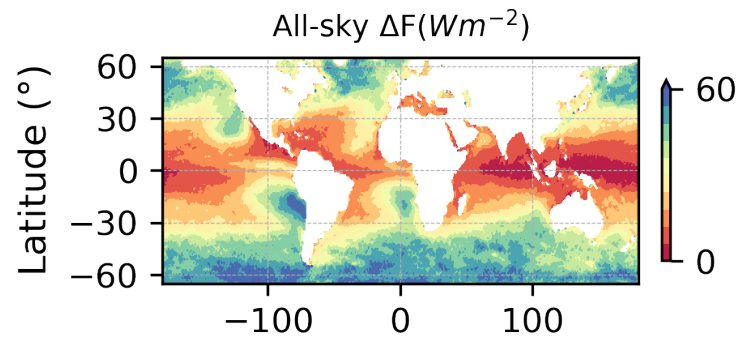
Our NN has a total of four layers. The input and output layers have 25 and 5 nodes respectively, which matches the number of input and output variables. Between them are two fully connected hidden layers with 256 nodes. This adds up to a total of 73733 learnable parameters. We use the Rectified Linear Unit (ReLU) for activation function and the Adam optimizer with a mean squared error loss function. Given the large number of training samples, the specific choices of the hyper-parameters make little difference to the performance. The input data are normalized and shuffled before the training. The total training time was about 7 minutes on a single graphics processing unit.



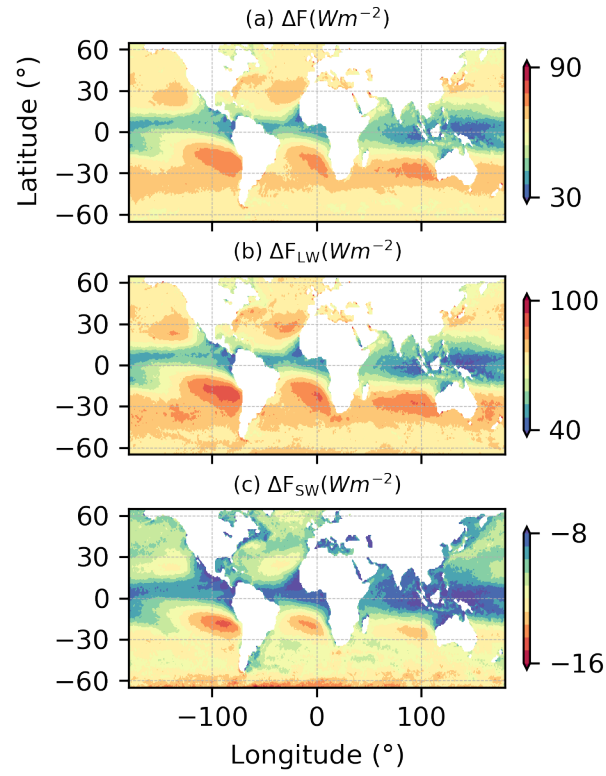
**Figure S1:** Probability density functions of cloud-top radiative flux divergences.



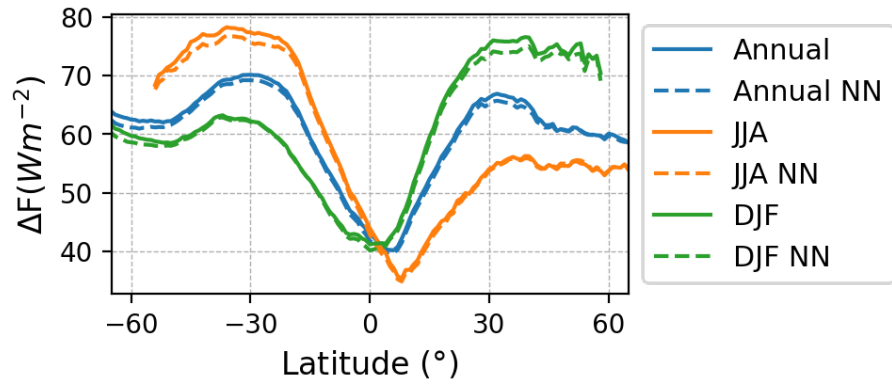
**Figure S2:** Probability density functions of the ratio between cloud-top temperature and sea surface temperature.



**Figure S3:** Annual-mean all-sky cloud-top radiative cooling.



**Figure S4:** Neural-network-estimated global distribution of annually-averaged cloud-top radiative cooling (a), its LW (b) and SW components (c).



**Figure S5:** Zonal-mean meridional variations of cloud-top radiative cooling computed from the radiative transfer model (solid) and the neural network (dashed) for boreal summer (orange) and winter (green).

Input variables	Unit	Output variables	Unit
Cloud optical depth	Unitless	<b>Cloud top radiative cooling, <math>\Delta F</math></b>	$\text{W m}^{-2}$
Cloud droplet effective radius	$\mu\text{m}$	<b>Cloud top longwave cooling, <math>\Delta F_{LW}</math></b>	$\text{W m}^{-2}$
Cloud top temperature	K	<b>Cloud top shortwave heating, <math>\Delta F_{SW}</math></b>	$\text{W m}^{-2}$
Solar zenith angle	degree	Cloud base longwave heating	$\text{W m}^{-2}$
Sea surface temperature	K	Cloud longwave radiative effect	$\text{W m}^{-2}$
Absolute temperature from 1000 hPa to 100 hPa with 100 hPa interval	K		
Relative humidity from 1000 hPa to 100 hPa with 100 hPa interval	%		

**Table 1:** Input and output variables for the Neural Network. The CTRC variables used in this study are highlighted in bold.



148 **Reference**

149 Ricchiazzi, P., Yang, S., Gautier, C., & Soble, D. (1998). SBDART: A research and teaching software tool  
150 for plane-parallel radiative transfer in the Earth's atmosphere. *Bulletin of the American*  
151 *Meteorological Society*, 79(10), 2101-2114.

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