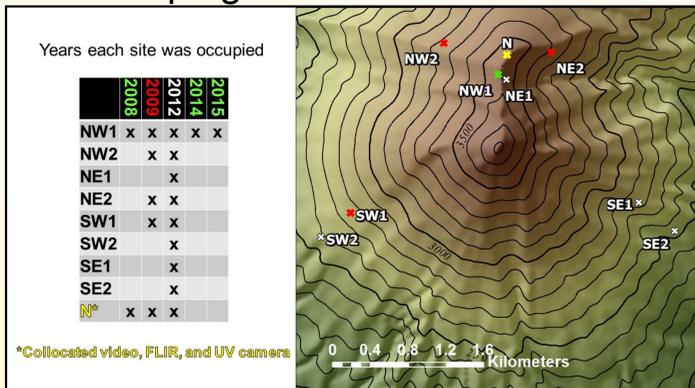


Long Term Stability of Conduit Dynamics at Fuego Volcano, Guatemala, 2008-2015

Background

Field Campaigns from 2008 - 2015



Overall energy of system varied, but examples of persistent behavior:

Vent location constant within 200 m radius

- Often two vents



Very Long Period (60 sec-10 sec) Seismic Events

- Sub-vertical dike feeding a dipping sill¹
- Different emission styles with different waveforms²
- Large range of amplitudes
- SO₂ emissions linked to seismicity³

Long Period (0.5 – 2 Hz) Seismic Events

- Series of cracks around conduit margin⁴
- No visible emissions associated

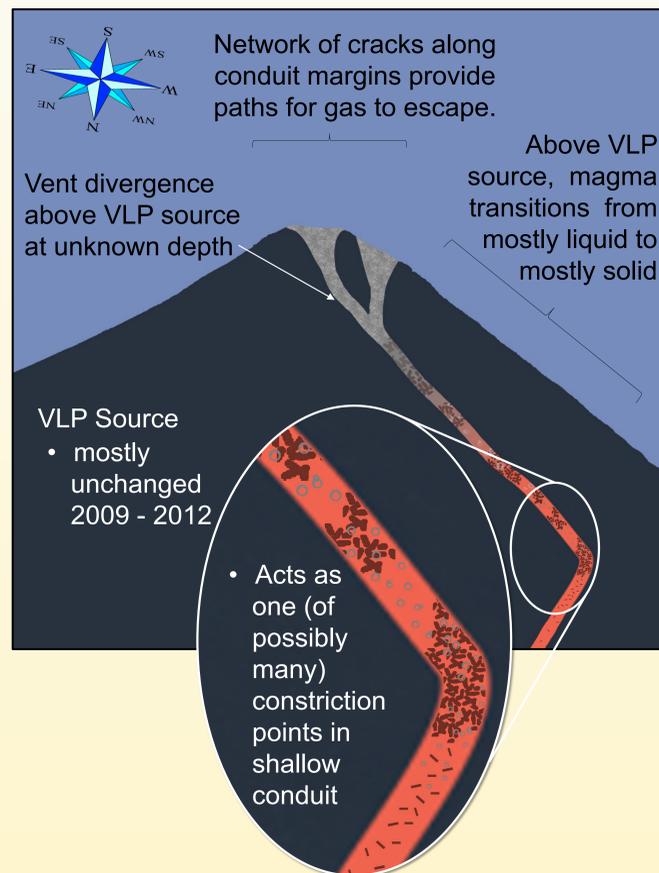
Consistent Average Annual Gas Emission Rates

- Harmonic and broadband tremor with visible plumes⁵
- Also aseismic interevent degassing



Clustered ($C_v > 1$) Repose Periods for Seismicity

Model



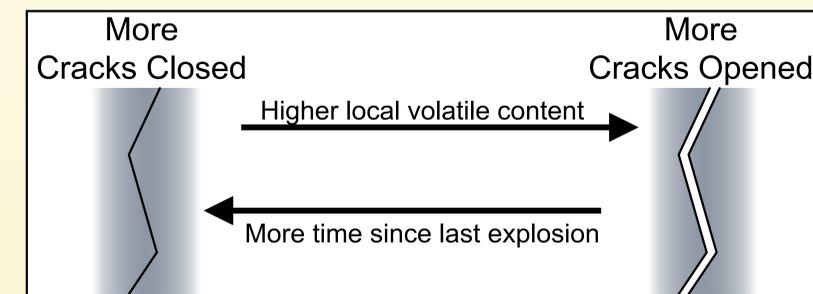
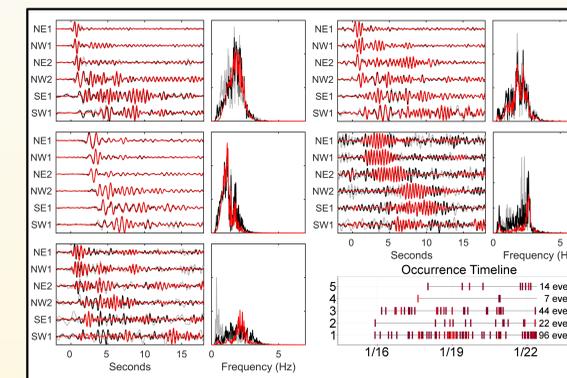
Two styles of cracks

I. Along Conduit Margins

- More stable (LP sources)
- Provide paths for gas to leak

II. Within Cap

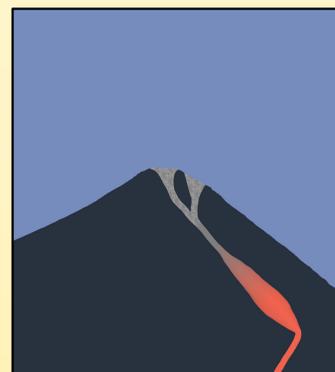
- Random distribution
- As cap properties vary, so do cracks
- Cap forms from undercooling



Event size correlates with repose time when $t_r > 2$ hrs

Broadband and harmonic tremor along with aseismic degassing.

When cap seals:

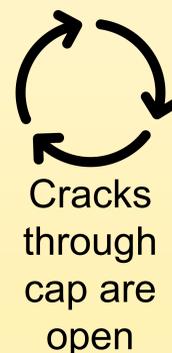


- Pre-event inflation of sill (exaggerated)
- Evident in tilt data
- Cap fractures near surface to relieve pressure.

When cap ruptures:



- Emission size depends on depth fragmentation front penetrates into conduit.
- Largest events have VLPs
- Emissions take path of least resistance.



Questions

Magma fragmentation rates

- What process governs the different speeds of magma fragmentation and the different depths that the front reaches?
- Are local aggregations of crystals and bubbles enough to account for different VLP event styles and changes in event times?

Longer term (week-month) magma supply rate changes

- What drives changes in liquid-solid transition depth?
- What drives behavior regime changes: lava flows vs. strombolian vs. paroxysms?

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