

Midwinter dry spells amplify post-fire snowpack decline

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Key Points:

- A factor of 9.8 increase in satellite-based fire detections are observed in California's snow zones in 2020-2021 compared with 2001-2019.
- Measured accumulation season snow albedo declined 25-71% with these reductions driving decreased snow covered days and snow cover fraction.
- Compared with a meteorologically similar 2013 dry spell, snow albedo declines led to rapid midwinter melting in post-fire environments.

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Abstract

Increasing wildfire and declining snowpacks in mountain regions threaten water availability. We combine satellite-based fire detection with snow seasonality classifications to examine fire activity in California's seasonal and ephemeral snow areas. We find a nearly tenfold increase in fire activity during 2020 and 2021 compared to 2001-2019 as measured by satellite data. Accumulation season snow albedo declined 17-77% in two burned sites as measured by in-situ data relative to un-burned conditions, with greater declines associated with increased soil burn severity. By enhancing snowpack susceptibility to melt, decreased snow albedo drove mid-winter melt during a multi-week midwinter dry spell in 2022. Despite similar meteorological conditions in 2013 and 2022, which we link to persistent high pressure weather regimes, minimal melt occurred in 2013. Post-fire differences are confirmed with satellite measurements. Our findings suggest larger areas of California's snowpack will be increasingly impacted by the compounding effects of dry spells and wildfire.

Plain Language Summary

Satellite fire detections indicate substantial increases in wildfire activity in California's snow-covered landscapes during 2020 and 2021, suggesting wildfire is increasingly altering mountain hydrology. During 2022, a multi-week mid-winter drought, or dry spell, occurred. A meteorologically-similar dry spell occurred in 2013, and the 2022 event provides a test case to examine how post-fire changes (canopy loss and deposition of burned dark material on snowpack) alter snowmelt patterns. Using field observations, weather station data, and satellite remote sensing of snow, we find large reductions in snow albedo drove rapid melt during the 2022 dry spell in burned areas whereas during 2013, minimal melt occurred. Our findings motivate additional research into assessing and planning for post-fire hydrologic changes in snow-dominated landscapes as both wildfire and dry spells will increase in frequency with climate warming.

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1 Introduction

Communities and ecosystems worldwide rely on snowpack to meet water demands (Immerzeel et al., 2020). A warming climate changes the spatial patterns and timing of snowpack accumulation and melt via alterations to rain-snow partitioning, decreases in cold content, extended dry spells, and extreme precipitation events (Lynn et al., 2020; Gershunov et al., 2019; Siirila-Woodburn et al., 2021). During the dry season, reduced snowpack combined with warming and drying enhances evaporative demand (Abatzoglou & Williams, 2016; Alizadeh et al., 2021) and lowers fuel moisture (McEvoy et al., 2019).

A warming, drying, and disturbance-prone climate combined with fire suppression promotes severe wildfire at higher elevations in the western U.S. (Millar & Stephenson, 2015; Alizadeh et al., 2021). From 1984-2017, a 9% increase per year in area burned in the seasonal snow zone (Gleason et al., 2019) has been accompanied by a 7.6 myr^{-1} up-slope increase in average wildfire elevation (Alizadeh et al., 2021). High burn severity areas also increased during these decades (Parks & Abatzoglou, 2020).

Severe fires alter mountain snowpack processes near and below the treeline in two key ways. First, the burned canopy increases incoming solar radiation and black carbon sourced from burned vegetation reduces snow albedo, which together, accelerate snowmelt rates by up to 57% (Gleason et al., 2013; Kaspari et al., 2015; Gleason & Nolin, 2016; Gleason et al., 2019; Skiles et al., 2018; Aubry-Wake et al., 2022). Second, the decrease in forest canopy reduces interception. Every 20% increase in tree mortality increases below-canopy snow accumulation by 15% (Maxwell & Clair, 2019). In the absence of fire, reduced canopy shifts the timing of peak snowpack later (Cristea et al., 2014).

Additional wildfire impacts on mountain hydrology include changes in soil hydraulic properties (Ebel & Moody, 2017), shifts in surface and subsurface water partitioning and flow pathways that increase water yields (Maina & Siirila-Woodburn, 2020), as well as interactions between forest structure, snow, and fire effects (e.g., Moeser et al., 2020; Wilson et al., 2021). By altering the snow-vegetation-hydrology dynamics, severe fire in montane forests threatens ecosystems and the volume of snowpacks (Stevens, 2017; Gleason et al., 2019; Siirila-Woodburn et al., 2021). While it is well-documented that spring snowmelt rates increase after wildfire (e.g., Gleason et al., 2019; Uecker et al., 2020), the mid-winter impacts remain understudied.

Our work is motivated by two recent phenomena adversely affecting California's snow hydrology: widespread severe wildfires of 2020-2021 reaching into the seasonal snow zone of mountain watersheds (Figure 1) and the multi-week, midwinter dry spell during the winter of 2022. We examine how the post-fire environment during the unusually dry conditions amplified snowmelt rates. We hypothesize fire-impacted regions undergo declines in midwinter snow albedo that drive more rapid and earlier snowmelt compared with pre-fire or unburned conditions.

2 Methods

2.1 Satellite Fire Detection in Seasonal and Ephemeral Snowpacks

In-situ wildfire activity is difficult to quantify. Satellite-based fire detection is a useful proxy for generally assessing wildfire activity by providing consistent overflight return intervals across multiple years (Justice et al., 2002). We acquired daily fire detections at 1 km horizontal resolution from the MODerate resolution Imaging Spectroradiometer (MODIS) via the FIRMS database (<https://firms.modaps.eosdis.nasa.gov/download/>) for the period spanning January 2001 to December 2021. We subset all Californian fire detections into seasonal, ephemeral, and non-snow environments based on the concept of snow seasonality (Petersky & Harpold, 2018; Hatchett, 2021): the duration of time a landscape is continuously snow-covered. To assess seasonality, we applied

103 the snow classifiers to a gridded, 4-km horizontal resolution, daily snow water equiva-
104 lent (SWE) product (Zeng et al., 2018; Broxton et al., 2019) across California. Seasonal
105 snowpacks are defined as gridcells with an annual median of more than 60 days of con-
106 tinuous snow cover spanning 1982–2018. Ephemeral snowpacks are defined as gridcells
107 with intermittent (i.e., > 60 days of continuous) snow cover.

108 2.2 Snow Remote Sensing

109 Daily observations of snow surface properties from the vicinity of the Caldor Fire
110 from Terra MODIS are available from a collaboration between the National Snow and
111 Ice Data Center and the Institute of Arctic and Alpine Research called Snow Today [[https://
112 nsidc.org//snow-today](https://nsidc.org//snow-today)]. The website provides snow cover percent, snow cover days,
113 snow albedo, as well as the reduction of snow albedo from light absorbing particles (LAP).
114 Initial estimates of snow surface properties are based on the MODIS Snow Covered Area
115 and Grain size model (MODSCAG, (Painter et al., 2009)) and the MODIS Dust Radiative
116 Forcing in Snow model (MODDRFS, (Painter et al., 2012)). Data from the two mod-
117 els are combined to create spatially and temporally complete (STC) daily images that
118 account for forest canopy, off-nadir viewing, and cloud mis-identification (Rittger et al.,
119 2020). Snow cover errors from MODSCAG have been shown to be half the size of stan-
120 dard MODIS products (Rittger et al., 2013) and albedo estimates from STC-MODSCAG/MODDRFS
121 show 5% RMSE with no bias (Bair et al., 2019). STC-MODSCAG/MODDRFS data has
122 been previously used for SWE reconstruction (Rittger et al., 2016; Bair et al., 2016), real-
123 time estimates of SWE (Bair et al., 2018), estimating trends in snow cover at regional
124 scales (Ackroyd et al., 2021), understanding snow darkening related to LAP (Sarangi et
125 al., 2019, 2020; Huang et al., 2022) and improving snow albedo modeling (Hao et al., 2022).

126 2.3 Albedo Measurements

127 Our study sites for snow albedo include the 2021 Caldor Fire (measured once in
128 January 2022) and the 2020 Creek Fire (measured once in both February and April 2021;
129 Figure 2a). These are two of the larger fires in California history—the Dixie Fire burned
130 389,900 ha between July and October 2021, and the Caldor Fire burned 89,800 ha be-
131 tween August and October 2021 [[https://www.fire.ca.gov/media/4jandlhh/top20
132 _acres.pdf](https://www.fire.ca.gov/media/4jandlhh/top20_acres.pdf)]
133 and were notable as both crossed the hydrographic crest of the Sierra Nevada. Spectral
134 albedo measurements were made using a Spectral Evolution RS-3500 Portable
135 Spectroradiometer (RS-3500) equipped with a 180° field of view diffuser mounted on an
136 extendable 1.2 m pole (Figure 2b). The RS-3500 has a spectral resolution of 1 nm over
137 the spectral range 350–2500 nm. Measurements were made every 10 m along approxi-
138 mately flat terrain with one 100 m transect for each burn severity class: high, moder-
139 ate, and unburned. Soil burn severity for each fire was determined using maps produced
140 by the United States Department of Agriculture Forest Service Burned Area Emergency
141 Response [<https://burnseverity.cr.usgs.gov/products/baer>] using field-checked,
remotely-sensed pre- and post-fire visible reflectances (Key & Benson, 2006).

142 2.4 Snowpack and Meteorological Observations

143 We used station-based observations of SWE, precipitation and solar radiation to
144 examine the impacts of wildfire in burned and unburned areas. Daily SWE observations
145 (1 October 2011–15 April 2022) spanned the two mid-winter dry spells of interest from
146 four stations in the California Cooperative Snow Survey Network (Rattlesnake, Robin-
147 son Cow Camp, Greek Store, and Alpha) and two stations from the Snowpack Teleme-
148 try Network (SNOTEL; Central Sierra Snow Laboratory and Echo Summit; Figure 2a).
149 Two stations, Rattlesnake and Alpha, were burned in 2021 by the Dixie and Caldor Fires,
150 respectively, but remained functional.

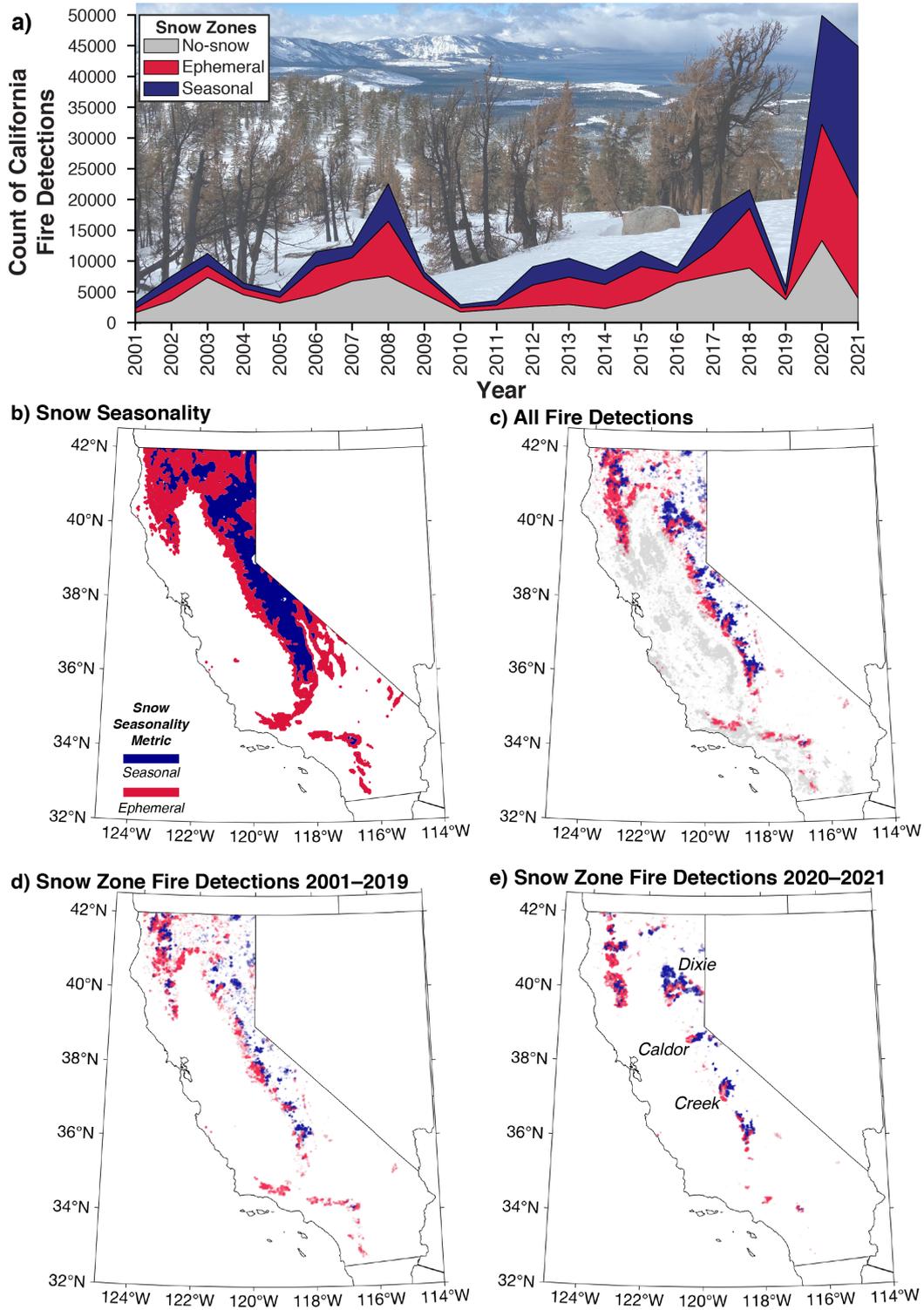


Figure 1. (a) Annual fire detections subset by snow seasonality (snow zone). (b) Snow seasonality classifications for California. (c) All fire detections (2001-2021). Fire detections during (d) 2001-2019 and (e) 2020-2021, noting fires named in the text.

151 To characterize the frequency of midwinter dry spells and place recent dry spells
 152 in a climatological context, we used daily precipitation data spanning 1 October 1917–
 153 15 April 2022 from the Tahoe City National Weather Service Cooperative Observation
 154 Program. Dry spells were defined as consecutive periods of time with no daily precip-
 155 itation exceeding 2.54 mm between October and April.

156 2.5 Weather Regimes

157 To provide a synoptic-planetary perspective and compare atmospheric circulation
 158 patterns during the two dry spell winters, we used the weather regime catalog of Guirguis
 159 et al., which evaluates the daily joint phase relationships between four regionally impor-
 160 tant modes of atmospheric variability (Guirguis et al., 2020). We extended this prod-
 161 uct to include the winter of 2021–2022. We focus on the days of the dry spell period (30
 162 December to 18 February) shared between the two years.

163 3 Results

164 3.1 Fire activity increased in seasonal and ephemeral snow zones

165 Fire detections show peaks during singular years (2008, 2018) or groups of years
 166 (2012–2016, 2020–2021; Figure 1a) across California’s seasonal and ephemeral snow zones
 167 (Figure 1b). In 2020 and 2021, an abrupt increase in fire detections in the snow zones
 168 occurred. $\sim 50\%$ of total 2001–2021 fire detections in seasonal snow zones and $\sim 35\%$ in
 169 ephemeral snow regions occurred in 2020–2021. A factor of 9.8 increase in mean annual
 170 fire detections in the seasonal snow zone occurred in 2020–2021 compared with the 2001–
 171 2019 average. Fire activity in snow zones was widespread throughout 2001–2021 (Fig-
 172 ure 1c), with a broad distribution of fire occurrence prior to 2020 (Figure 1d). However,
 173 very large fires including the Dixie, Caldor, and Creek Fires, as well as fire complexes
 174 elsewhere, during 2020 and 2021 clustered fire detections in snow zones (Figure 1e).

175 3.2 Snow albedo declines following wildfire

176 Snow albedo measurements in both the accumulation season (January and Febru-
 177 ary) and the ablation season (April) show time-dependent decreases in snow albedo (Fig-
 178 ure 2c). Decreases in the visible portion of the spectrum (0.4–0.7 μm) ranged from 17-
 179 31% (moderate severity) and 45–49% (high severity) compared to unburned during the
 180 accumulation season. In April, high burn severity areas showed a snow albedo decrease
 181 of 60% compared to unburned areas. For the NIR (0.7–2.5 μm), accumulation season
 182 declines ranged between 31 to 69% (moderate severity) to 47 to 77% (high severity). When
 183 comparing moderate to high severity burned areas, April albedo decreased by 34% in the
 184 NIR compared to 11% in the visible wavelengths. The opposite occurred in January where
 185 NIR decreased by 10% while the visible decreased by 34%. We note that the decreased
 186 albedo in the visible wavelengths is mainly due to light absorption by black carbon (Wiscombe
 187 & Warren, 1980; Warren, 1982), while decreased albedo in the near-infrared wavelengths
 188 is likely due to a combination of increased grain size and light absorption by black car-
 189 bon (Wiscombe & Warren, 1980; Warren, 1982; Skiles & Painter, 2019). Unlike dust, which
 190 has primary absorption in the visible wavelengths (He et al., 2019), black carbon is a “gray”
 191 absorber and can absorb throughout the solar spectrum. Translating measured snow albedo
 192 changes to net radiation using the Beer-Lambert Law indicates post-fire increases in net
 193 radiation from 0.931 Wm^2 (unburned) to 40.3 Wm^2 (high severity) during January and
 194 from 27.7 Wm^2 (unburned) to 161 Wm^2 (high severity) during April (Koshkin, 2022).

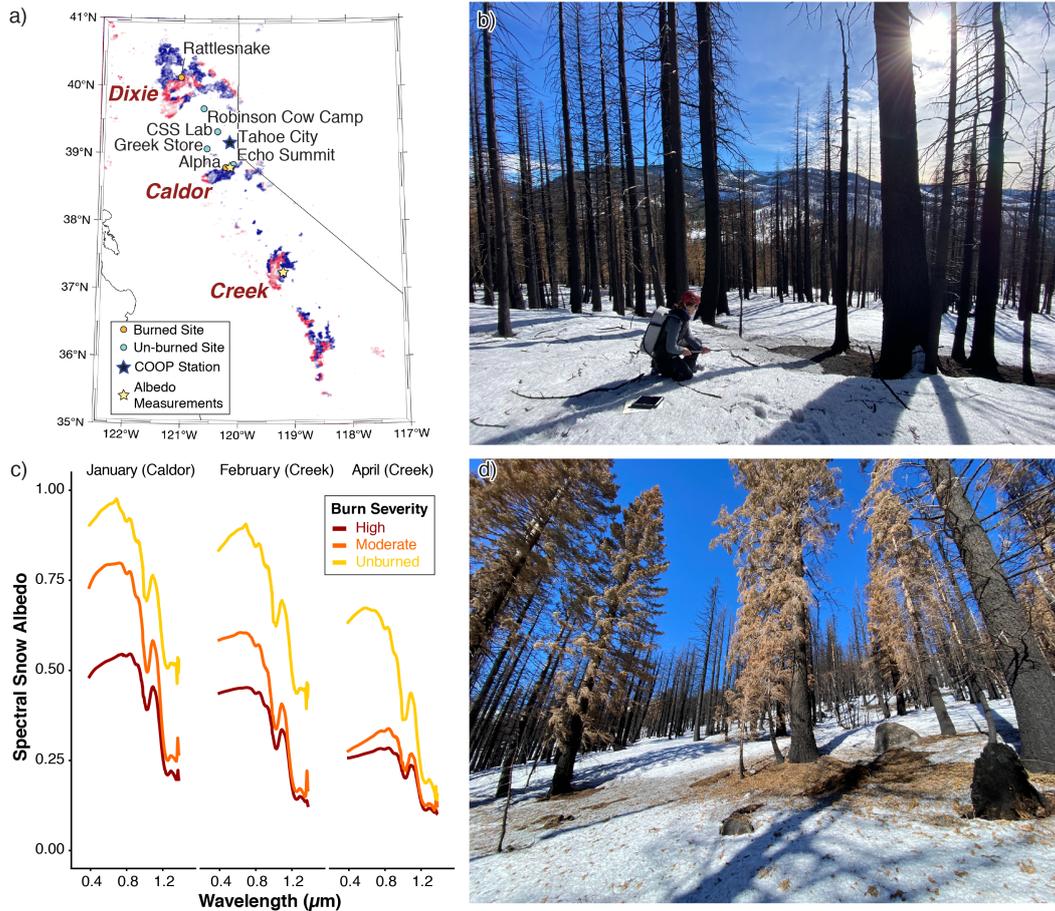


Figure 2. (a) Map of stations and sample locations with fire names. (b) Albedo measurements collection from high burn severity forest during January in the Caldor Fire. (c) Changes in spectral snow albedo for unburned (gold), moderate burn severity (orange) and high burn severity (red) during January (Caldor), February (Creek), and April (Creek). (d) Foreground shows burned debris deposited onto the snow surface.

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3.3 Rapid snowmelt during a dry spell following wildfire

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The long-term median midwinter dry spell at Tahoe City is 22 days. During water year (WY) 2022, Tahoe City experienced its second-longest midwinter dry spell (46 days). With observations beginning in 1917, three of the five longest midwinter dry spells have occurred since WY2011, including WY2015 (third longest with 43 days), WY2022 (second with 46 days), and the record-setting WY2012 (60 days). Although WY2013 (tied for 11th with 36 days) did not experience as prolonged a midwinter dry spell as WY2022, the well-below average precipitation following a wet start to the water year provides an object lesson year for comparison. In both WY2013 and WY2022, heavy precipitation during October–December produced substantial early season snowpacks (338–770 mm SWE), but were followed by persistent well-below average precipitation (Figure 3a) and similar radiation. Compared with WY2013, 5% more accumulated solar radiation occurred during WY2022 between 28 December–18 February period at the Red Baron RAWS (adjacent to the Caldor Fire) but approximately equal radiation between 28 December–1 March (Figure S1).

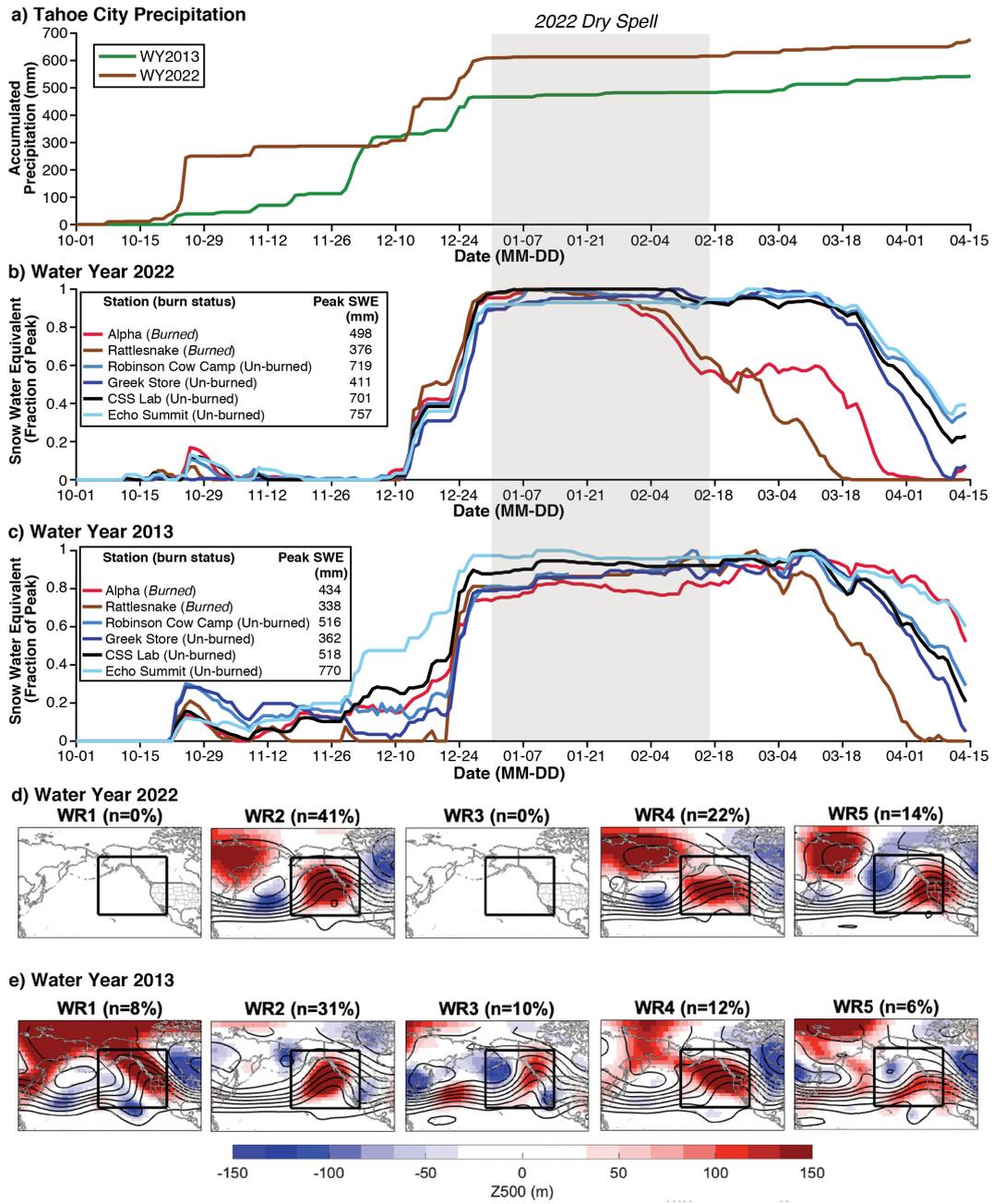


Figure 3. (a) Accumulated precipitation at Tahoe City during WY2013 and WY2022. (b) Snow water equivalent (SWE) as a fraction of peak SWE during water year (WY) 2022. (c) As in (b) but for WY2013. Primary dry weather regimes (WR) and their frequency during the dry spells of December 30-February 18 of (d) WY2022, and (e) WY2013.

SWE declined faster at the two burned sites, Alpha and Rattlesnake, compared to the unburned sites during WY2022's dry spell (Figure 3b). In contrast, all stations behaved similarly during the WY2013 dry spell (Figure 3c), though Rattlesnake began melting in mid-March. Compared to the date of maximum SWE, the SWE at unburned sites declined by 0-4 percentage points in WY2013 and 0-8 percentage points in WY2022 over the course of the dry period. During WY2013, Alpha and Rattlesnake declined by 2-9 percentage points during the dry spell. However, in WY2022 after the wildfire, these stations declined 41-45 percentage points, consistent with enhanced net shortwave radiation loading in burned environments (Figure 2). After a small precipitation event on 18 February 2022, snowpack continued to decline at Rattlesnake but remained consistent at Alpha before declining in late March. Compared to WY2013, snow at both Alpha and Rattlesnake disappeared earlier during WY2022 (Figure 3b-c).

3.4 Midwinter dry spell weather regimes

Analysis of weather regimes (WR; (Guirguis et al., 2022) reveals broad similarities between WY2013 and WY2022 during the midwinter dry spells (Figure 3d-e) but also throughout the accumulation season (Figure S2). The bulk of the snow accumulation in December during both years was associated with wet weather regimes (Figures S2 and S3), which favor anomalously wet conditions over California and anomalous positive snowpack accumulation in the Central Sierra Nevada (Guirguis et al., 2022). Beginning in late December (WY2013) or early January (WY2022) a WR shift occurred bringing atmospheric ridging conditions over/offshore from California (Figure S2). Similar dry-type WRs but with varying frequencies occurred during the respective dry spells (Figure 3d-e). The cessation in SWE accumulation is associated with the onset and persistence of WRs favoring persistent high pressure ridging bringing about mid-winter drought conditions (Figure 3d-e). The ridging patterns were more persistent during WY2022 (Figure 3d). WY2013 was more variable with short-lived weather pattern changes allowing for small snow accumulation events (Figure 4b). These events appear as intermittent breakdowns of the ridging patterns and development of patterns (e.g., WR3) producing weak onshore flow (Figure 3e).

3.5 Snow remote sensing indicates post-fire snowpack decline

Despite similar conditions in snowpack at the beginning of each dry spell (Figure 3b-c) and generally similar meteorological conditions (Figure 3d-e and S2), remote sensing shows widespread rapid post-wildfire snowmelt throughout the accumulation and melt seasons within the Caldor Fire perimeter (Figure 4). Snow covered area declined faster during during January and February WY2022 compared to WY2013 (Figure 4a), with 50% less snow cover at the end of the dry spell in WY2022. Albedo resets following snowfall were more common in WY2013 than WY2022 (Figure 4b), with 2022 demonstrating the lowest basin-average snow albedos on record in early February. Consistent with lower albedo (Figures 2c and 4b), melt occurred faster after storms in WY2022 compared to WY2013 (Figure 4a). Impacts within the fire perimeter are clear with more than 50 days less snow cover days by the end of April (Figure 4c), making WY2022 the year with the lowest snow cover days in the MODIS record. Dry conditions during November (Figure 3a) and melt-out of October snowfall (Figure 3b) contributed to the initially (1 January) low cumulative days of snow cover in WY2022 (Figure 4c). In contrast, WY2013 was near-to-above average in terms of snow cover days until late April (Figure 4c).

Spatial comparisons for February mean snow cover fractions show WY2013 had near-to-slightly below the 2001-2022 mean, whereas WY2022 had well-below mean snow cover fractions within the Caldor Fire perimeter (and Tamarack Fire perimeter; Figure 4d-e). By 1 March, snow cover days were 20-50 days below average only in the lowest elevation (western-most) regions during WY2013 whereas strong correspondence between anoma-

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lous below-mean snow cover days and the Caldor fire perimeter occur during WY2022 (Figure 4f-g). These differences increased as the season progressed (Figure S4).

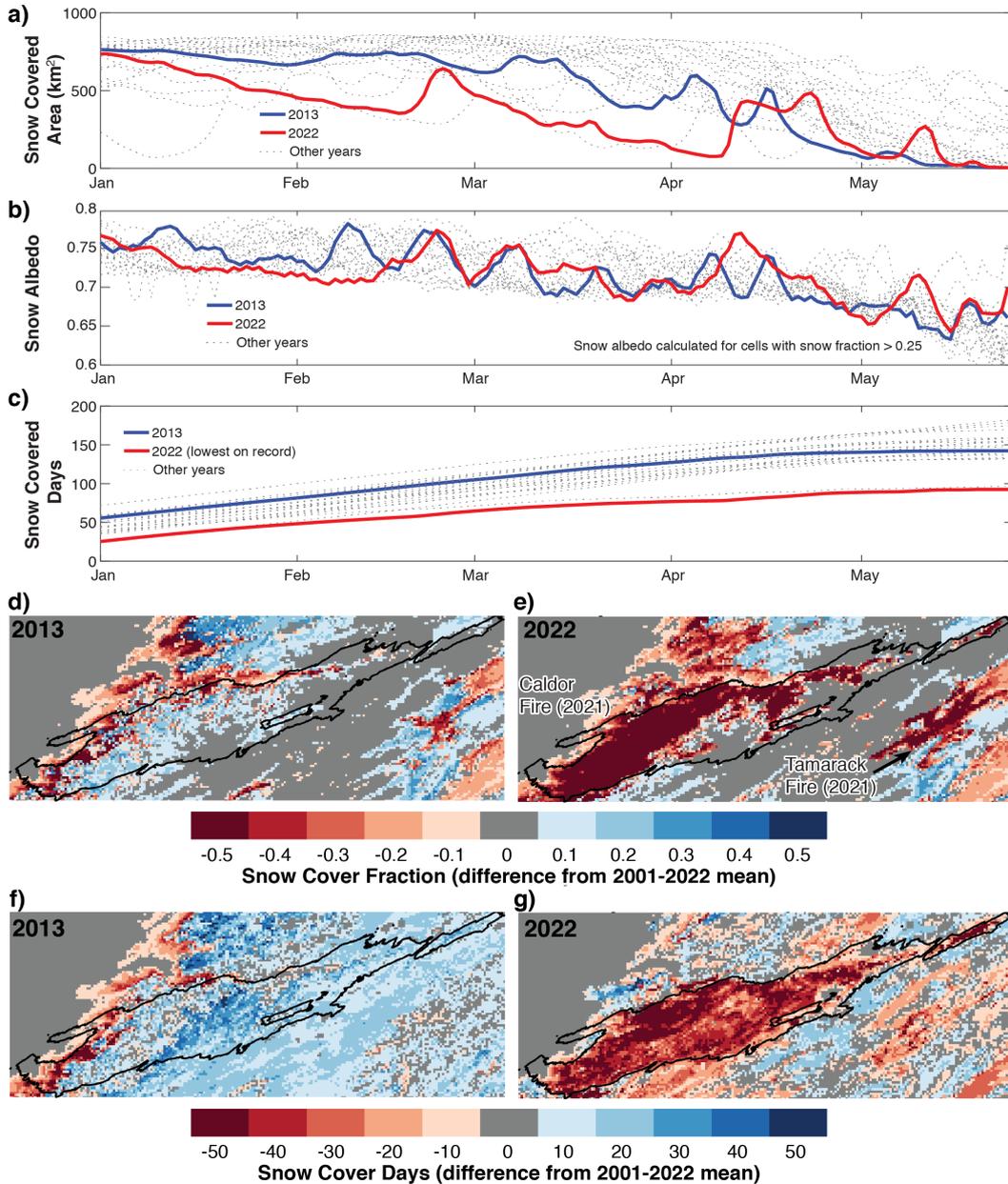


Figure 4. (a) Snow covered area (km^2) for Caldor Fire perimeter between 1 January and 31 May for WYs 2001-2022 (light dashed lines) with WY2013 and WY2022 shown as thick blue and lines respectively. (b) As in (a) but for snow albedo. (c) As in (a) but for snow covered days. February snow cover fractions, as differences from 2001-2022 mean for (d) WY2013 and (e) WY2022. End of February snow cover days (cumulative from 1 October), as differences from 2001-2022 mean for (f) WY2013 and (g) WY2022.

4 Discussion

Wildfires in seasonal and ephemeral snow zones are expected. Our identified abrupt, near-10-fold increase in fire activity during 2020-2021 in California's snow zones relative to the previous 18 years (Figure 1) is embedded in an increasing trend in California wildfire activity (Alizadeh et al., 2021; Gleason et al., 2019). Conditions conducive to large, severe fires will increase as the climate warms (Abatzoglou & Williams, 2016; Gutierrez et al., 2021; Williams et al., 2019) and becomes more volatile (Gershunov et al., 2019). This implies future fire activity in snow zones will more frequently resemble 2020 and 2021.

In moderate to high severity burned areas, the albedo decreases and canopy removal (Figure 2b-d) combined to enhance snowmelt during midwinter dry spells (Figure 3c) leading to reductions in snow covered area (Figure 4c). Similar results associated with local and long-range transport and deposition of fire-generated LAPs occurs in seasonally snow-covered regions (Gleason et al., 2019; Uecker et al., 2020) and glacial environments (Aubry-Wake et al., 2022). However, those studies focused on the ablation season rather than the accumulation season. Our results indicate strong potential for enhanced post-fire midwinter melt under persistent high pressure (Figure 3d).

The wavelength dependence of snow albedo reductions (Figure 2c) suggests impurities are a more dominant control on albedo changes compared to snow grain size during the accumulation season when snowpacks are colder and grain growth is slower (Colbeck, 1982). During the melt season, wet snow causes rapid grain growth (Colbeck, 1982), and these increases in grain size strongly decrease snow albedo in the near-infrared wavelengths (Warren, 1982). Also during the melt season, the concentration of LAPs decreases snow albedo in the visible wavelengths (Warren, 1982; Sterle et al., 2013). The combined effect is a significant overall decrease in snow albedo across the solar spectrum. Similar to dust-on-snow (Skiles & Painter, 2019), in post-fire environments, radiative forcing-induced positive feedbacks likely occur between grain size growth, albedo decline from melt-driven LAP accumulation, and larger-scale land surface albedo decline as the land surface becomes snow-free (Huang et al., 2022; Koshkin et al., 2022).

Guirguis et al. (2022) found increasing frequencies of three midwinter dry patterns that parallel observed declines in California snowpack and water availability (Mote et al., 2018; Siirila-Woodburn et al., 2021). These same atmospheric circulation patterns are associated with the midwinter dry spells in WY2013 and WY2022. While both years demonstrate similar weather patterns to one another, WY2013 had slightly more active weather compared to WY2022 (Figure S2). This implies observed melt resulted predominantly from altered land surface conditions (e.g., burned canopy) rather than meteorological differences. It is also likely that feedbacks between grain size and snow albedo (Koshkin et al., 2022) further accelerated melt, especially at lower elevations (Figure 4j).

Amplified midwinter melt in burned areas during midwinter dry spells raises concern for hydrologic resources and hazards. Enhanced radiation-driven midwinter melt with greater snow accumulation (Maxwell & Clair, 2019) has the potential to elevate soil moisture earlier in the water year and make snowpacks more hydrologically active (Brandt et al., 2022). Additional soil moisture increases runoff efficiency and soil pore water pressures, leading to elevated runoff during rain-on-snow events and higher probabilities for shallow landslides.

Midwinter runoff affects reservoir operations as traditional regulatory frameworks may not allow for additional reservoir storage when flood risk reduction is the primary management concern (Maina & Siirila-Woodburn, 2020; Williams et al., 2022a). Moreover, higher rates of sediment influx from burned areas entering reservoirs (Sankey et al., 2017; Murphy et al., 2018) reduce water quality (Murphy et al., 2012) and can damage infrastructure (Randle et al., 2021).

313 The compounding effects of post-fire impacts on snow and midwinter dry spells pose
 314 challenges for climate projections and operational forecasts. More frequent fire activity
 315 in snow zones and additional dry days are both expectations of a warming climate (Westerling,
 316 2018; Polade et al., 2014; Hatchett et al., 2022). Midwinter snowpack loss may enhance
 317 intraseasonal climate variability towards drought by shifting peak SWE earlier and caus-
 318 ing earlier snow disappearance (Gleason et al., 2019; Uecker et al., 2020; Smoot & Glea-
 319 son, 2021) leading to drier late-season soil and vegetation conditions (Harpold & Molotch,
 320 2015). Less efficient spring runoff in exchange for more efficient midwinter runoff (Maina
 321 & Siirila-Woodburn, 2020) may offsetting slower melt in unburned areas (Musselman et
 322 al., 2017). Skillfully predicting weather regimes associated with high-impact weather (anoma-
 323 lous wet or dry conditions) at subseasonal-to-seasonal scales provides lead-time to im-
 324 plement mitigation measures for altered hydrology (Guirguis et al., 2022). However, mit-
 325 igation hinges upon skillful hydrologic forecasts. If post-fire effects on snow exacerbates
 326 emerging trends towards elevated runoff (Uzun et al., 2021; Williams et al., 2022b), di-
 327 rect updates of snow albedo to operational hydrologic models and improved parameter-
 328 izations of fire-snow relationships in Earth system models is required (Hao et al., 2022).

329 5 Conclusions

330 The societal connection between mountains and humans will be strained as moun-
 331 tains face increasing climate-related stressors (Immerzeel et al., 2020). Midwinter drought,
 332 snow loss, and increasing wildfire are expectations of a warming world. Addressing these
 333 challenges requires innovative water and forest management paradigms (Millar & Stephen-
 334 son, 2015; Sterle et al., 2019; Siirila-Woodburn et al., 2021). We identified abrupt increases
 335 in wildfire activity in California’s snow zones that reduced snow albedo and accelerated
 336 melt during extended midwinter drought. To enhance water supply reliability, reduce
 337 flood hazards, and inform adaptation strategies—aspects impacted by wildfire’s effects
 338 on mountain snowpacks—we recommend improving our process-based representation and
 339 inclusion of wildfire’s impacts in the snow zone in short- and long-term operational hy-
 340 drologic and Earth system models.

341 6 Open Research

342 MODIS fire detections are available from the NASA Fire Information for Resources
 343 Management System: <https://firms.modaps.eosdis.nasa.gov/>. Weather regime data
 344 (Guirguis et al., 2022) is available from the UCSD library digital collections at: [https://](https://doi.org/10.6075/J089161B)
 345 doi.org/10.6075/J089161B. The University of Arizona Snow Water Equivalent Prod-
 346 uct is available from the NASA National Snow and Ice Data Center Distributed Active
 347 Archive Center at: <https://doi.org/10.5067/0GGPB220EX6A>. Station data is publicly
 348 available for SNOTEL stations from the United States Natural Resources Conservation
 349 Agency: <https://wcc.sc.egov.usda.gov/reportGenerator/> with RAWS data avail-
 350 able from the Desert Research Institute: <https://raws.dri.edu>. The MODIS data is
 351 available from the Dryad repository at:

352 [https://datadryad.org/stash/share/xWCdmAowGgAjlyISPY9jirfiYmvm2bqTYNJz77mG](https://datadryad.org/stash/share/xWCdmAowGgAjlyISPY9jirfiYmvm2bqTYNJz77mG-0E)
 353 [-0E](https://datadryad.org/stash/share/xWCdmAowGgAjlyISPY9jirfiYmvm2bqTYNJz77mG-0E); upon acceptance this link will be archived as: [https://doi.org/10.5061/dryad](https://doi.org/10.5061/dryad.7wm37pvx7)
 354 [.7wm37pvx7](https://doi.org/10.5061/dryad.7wm37pvx7). Spectrometer data is available in the Supporting Materials.

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