

1 Climate drivers of malaria transmission seasonality and 2 their relative importance in Sub-Saharan Africa

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13 **Key Points:**

- 14 • Seasonal malaria transmission in Sub-Saharan Africa is sustained at temperatures
15 well above 15°C or below 40°C.
- 16 • Monthly maximum rainfall for seasonal malaria transmission should not exceed 600 mm.
- 17 • Rainfall and temperature are significant drivers of malaria seasonality in all parts of
18 Sub-Saharan Africa except in west Central Africa.
- 19 • Topography has significant influence on which climate variable is an important driver
20 of malaria seasonality in East Africa.
- 21 • Malaria transmission onset lags behind rainfall only at markedly seasonal rainfall
22 areas, otherwise, malaria transmission is year-round.

23 Abstract

24 A new database of the Entomological Inoculation Rate (EIR) is used to directly link the risk
25 of infectious mosquito bites to climate in Sub-Saharan Africa. Applying a statistical mixed
26 model framework to high-quality monthly EIR measurements collected from field campaigns
27 in Sub-Saharan Africa, we analyzed the impact of rainfall and temperature seasonality on
28 EIR seasonality and determined important climate drivers of malaria seasonality across
29 varied climate settings in the region. We observed that seasonal malaria transmission was
30 within a temperature window of 15°C – 40°C and was sustained if average temperature was
31 well above 15°C or below 40°C. Monthly maximum rainfall for seasonal malaria transmission
32 did not exceed 600 mm in west Central Africa, and 400 mm in the Sahel, Guinea Savannah
33 and East Africa. Based on a multi-regression model approach, rainfall and temperature
34 seasonality were found to be significantly associated with malaria seasonality in all parts of
35 Sub-Saharan Africa except in west Central Africa. Topography was found to have significant
36 influence on which climate variable is an important determinant of malaria seasonality in
37 East Africa. Seasonal malaria transmission onset lags behind rainfall only at markedly
38 seasonal rainfall areas such as Sahel and East Africa; elsewhere, malaria transmission is
39 year-round. High-quality EIR measurements can usefully supplement established metrics
40 for seasonal malaria. The study's outcome is important for the improvement and validation
41 of weather-driven dynamical mathematical malaria models that directly simulate EIR. Our
42 results can contribute to the development of malaria models fit-for-purpose to support health
43 decision-making in the fight to control or eliminate malaria in Sub-Saharan Africa.

44 1 Introduction

45 Sub-Saharan Africa remains the world's region with the greatest malaria burden despite
46 massive efforts over the past decades to lower or eliminate malaria (WHO, 2020). Though
47 poor health care systems and low socio-economic status (Degarege et al., 2019; Yadav et al.,
48 2014) are contributing factors, the climate suitability of the region for malaria transmis-
49 sion has a major influence (Caminade et al., 2014). Generally, climate variables such as
50 temperature, rainfall and relative humidity are known to have significant influence on the
51 development and survival of both the malaria parasites and their vectors. Malaria para-
52 site development is not possible at temperatures below 16°C and temperatures above 40°C
53 have adverse effect on mosquito population turnover (Parham and Michael, 2010; Mordecai
et al., 2013; Blanford et al., 2013; Shapiro et al., 2017). Rainfall provides the environment

for vector breeding (Ermert et al., 2011; Tompkins and Ermert, 2013; Kar et al., 2014) and relative humidity of at least 60% appears necessary for vector survival (Thompson et al., 2005). Rainfall therefore affects the availability, persistence and dimensions of *Anopheles* vectors and their larval habitats (Fournet et al., 2010; Afrane et al., 2012a; Boyce et al., 2016; Asare et al., 2016a). Previous work studying the relationship between sporozoite development and the survival of infectious mosquitoes found optimal temperatures for efficient malaria transmission between 25°C and 27°C (Bayoh, 2001; Lunde et al., 2013a,b). In Sub-Saharan Africa, most countries have annual mean temperatures between 20°C and 28°C (Lunde et al., 2013a). Given Sub-Saharan Africa's warm tropical climate, a plethora of efficient and effective malaria parasite and vectors thrives in this setting (Sinka et al., 2010; Murray et al., 2012). Understanding the relative importance of climate drivers of malaria seasonality is crucial for describing the geographic patterns of the heterogeneous risk and burden of malaria across the sub-region (Gething et al., 2011; Reiner et al., 2015). This could translate to substantial public health gains, taking into account the seasonality in malaria control and prevention interventions, by helping to determine when, where and how to apply vector and parasite control measures.

To our knowledge, there are insufficient field studies using Entomological Inoculation Rate (EIR, defined as the number of infectious mosquito bites person receives per time) data to relate climate to malaria seasonality in Sub-Saharan Africa. Mabaso et al. (2007) assessed the relationship between EIR seasonality and environmental variables in Africa using a rainfall seasonality index (Markham, 1970). The index fails to capture seasonality at areas with bimodal rainfall regimes, however. Furthermore, the study did not take into consideration the impact of diverse climatic conditions on seasonality outcomes but aggregated data from sites of different climate and environmental settings into a single study, which has the potential to skew the results. Other research has examined the link between malaria and climate variables but primarily relied on clinical data or malaria suitability indices (Lowe et al., 2013; Midekisa et al., 2015; Komen et al., 2015). Both malaria indices and case data have drawbacks for studying malaria seasonality.

Malaria indices are derived using statistical relationships between weather and malaria measures and their out-of-sample generalization over space and time for seasonality studies is subject to significant uncertainties. Clinical case data are also subject to significant uncertainties due to the inaccurate diagnostics (often counts of suspected cases, with temporal inconsistency in the use of Rapid Diagnostic Test, RDT or slide analysis) and under-counting

due to varying health-seeking behaviour and health policies (Afrane et al., 2012b). Given that the biology of the malaria parasite and its vector mosquito are temperature and rainfall dependent (Ermert et al., 2011), and that EIR can directly quantify parasite-infected mosquitoes and their propensity to transmit the parasites to humans (MARA, 1998; Shaukat et al., 2010) or estimate the seasonality of the exposure of a population to malaria parasite inoculations (Beier et al., 1999; Takken and Lindsay, 2003), then EIR should be able to usefully relate climate to malaria seasonality better than malaria cases.

In this study, therefore, we investigated the impact of climate variables on EIR seasonality in diverse climate settings across Sub-Saharan Africa with the goal of identifying significant climate determinants of malaria seasonality, their relative importance and variability across the region. To our knowledge, this is the first study to use EIR_m to explore the impact of climatic variables on malaria seasonality in Sub-Saharan Africa on this wider scale. We applied a mixed model statistiacal framework to a high-quality malaria EIR data (Yamba et al., 2018; Yamba et al., 2020) gathered from publicly available field campaigns of sufficient duration and determined the climate effect that explained significant variations in EIR seasonality. Our findings are intended to provide an understanding of geographical heterogeneous malaria risk from climate effect and support future malaria modeling and forecasting efforts. It will contribute to the development of malaria models especially weather-driven dynamical malaria models fit-for-purpose to support health decision-making in the fight to control or eliminate malaria in Sub-Saharan Africa.

2 Data and Methods

2.1 Study Area

The study area includes locations in Sub-Saharan Africa (as shown in Figure 1), where mosquitoes have previously been collected for malariometrics such as Human Biting Rate (HBR), CircumSporozoite Protein Rates (CSPR), and EIR. The geographical coordinates and elevation of each location are detailed in Tables S1 to S4. The study locations are grouped into four distinct climate zones namely Sahel, Guinea, WCA, and EA (see Figure 1). Each zone has a unique climate conditions from others (see Figure S1) and therefore have different climate implications on malaria seasonality (Yamba, 2016). The division into zones is, therefore, to ensure that malaria transmission patterns are consistent across geographical areas with similar climate characteristics. The seasonal distribution of rainfall

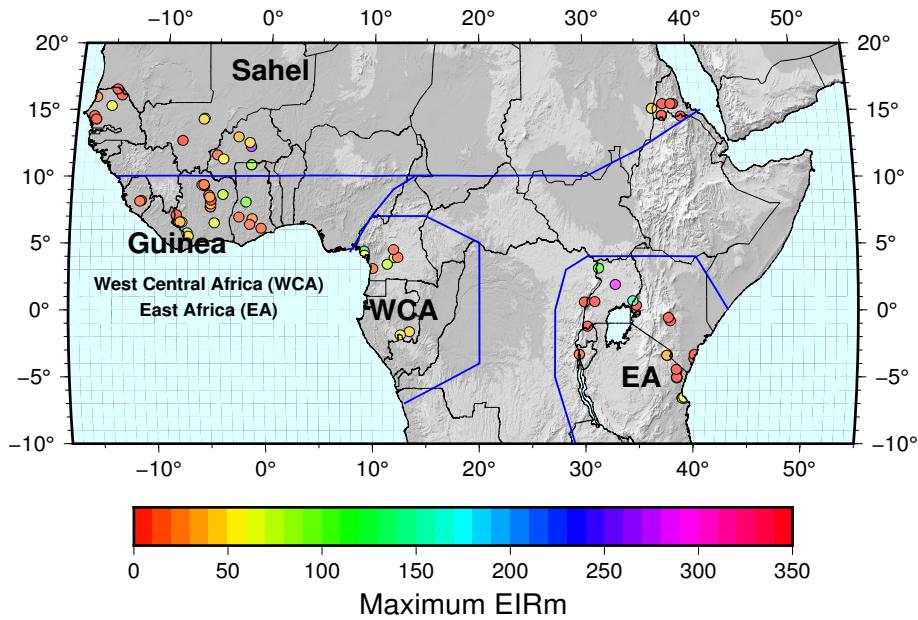


Figure 1: The map of the Sub-Saharan Africa showing field survey sites for EIR. The colour gradient of each site show the maximum EIR available. The blue lines delineate the region into climate zones of Sahel, Guinea, WCA and EA.

and temperature for each zone is shown in Figure S1. In the Sahel, rainfall is markedly seasonal, with a single wet season (usually June to October) and a protracted dry season (November to May). Seasonal temperature ranges between a minimum value of 20 °C during the harmattan season and to a maximum of about 40 °C during the pre-monsoon season. In general, temperatures are higher in the Sahel and colder in EA due to the fact that most areas are characterized by higher altitudes.

2.2 Data

2.2.1 Monthly EIR data

Monthly malaria EIR data (hereafter referred to as EIR_m) were obtained from a newly compiled and published monthly malaria EIR database (Yamba et al., 2018; Yamba et al., 2020) for each study location shown in Figure 1. The years and months for which the EIR_m data were available for each study location is shown in Table S1-S4. Generally, most locations had 12 months of data while other locations had data varying between 24 and 36 months. The data also spanned the period 1983-2013 for all locations. The temporal

duration of the data is mostly limited to one year because sampling mosquitoes for EIR is extremely capital and labour intensive (Kilama et al., 2014; Tusting et al., 2014; Badu et al., 2013). The EIR database from which data were extracted for use in this work is a comprehensive one. It was constructed through an all-inclusive literature review using google scholar and PubMed search facilities. All data in that database was generated from publicly available field campaigns of adequate duration and is freely available for public usage in the PANGAEA repository (Yamba et al., 2018). Details of how this database was constructed including compilation, sources, recording, spatial coverage and temporal resolutions are clearly described in Yamba et al. (2020).

142 2.2.2 Meteorological data

143 Monthly rainfall (RR) and temperature (minimum (T_{\min}), mean (T_{mean}) and maximum
 144 (T_{\max})) data for each study location were gathered. Rainfall data was obtained from the
 145 Global Precipitation Climatology Centre (GPCC) product, version 2018 (Schneider et al.,
 146 2018). The GPCC data is a gridded gauge-analysis products and available globally from
 147 1891-2016 at a spatial resolution of 0.25° . GPCC was chosen because it is a rain gauge-
 148 analysis product built from quality-controlled rainfall data from ground-based weather sta-
 149 tions. Previous validation studies (Manzanas et al., 2014; Atiah et al., 2020) have also
 150 found it to be reliable and consistent with ground-based weather observations. The temper-
 151 ature data was obtained from the European Centre for Medium-Range Weather Forecasts
 152 (ECMWF) Re-Analysis, 5th generation (ERA5) (Hersbach et al., 2020). ERA5 is also a
 153 gridded re-analysis product and available globally on an hourly time scale from 1979 to
 154 present at a high spatial resolution of 0.25° by 0.25° . ERA5 was chosen because previous
 155 evaluation studies of the product (Tarek et al., 2020; Gleixner et al., 2020; Osse et al.,
 156 2020) have widely recommended it for meteorological research. RR, T_{\min} and T_{\max} were
 157 extracted from the respective database for each study location using the nearest grid point
 158 of the location's geographical coordinates. T_{mean} values were estimated by averaging the
 159 T_{\min} and T_{\max} values for the location. The extracted temperature and rainfall data had
 160 to also conform with the exact years and months at which EIR_m data were available for
 161 each location. The study relied on GPCC and ERA5 because, ground-based local weather
 162 stations from which these data could be gathered were mostly not available at the EIR sites
 163 or, if present, often have sparse data.

164 2.3 Data analysis

165 The analysis was conducted for each classified zone as shown in Fig 1. EIR data from locations
 166 characterized with the presence of permanent water bodies and/or irrigation activities
 167 were exempted. Irrigation and permanent water bodies (such as damps, rivers, streams,
 168 swamps etc) have significant influence on the intensity and length of seasonal malaria trans-
 169 mission (Ermert et al., 2011; Tompkins and Ermert, 2013; Asare et al., 2016b; Asare and
 170 Amekudzi, 2017). Their exclusion was, therefore, a means to dissociate the influence of
 171 these hydrological parameters on malaria seasonality and reducing the impact to climate
 172 factors alone.

173 2.3.1 pair-wise comparison

174 The study examined the ranges of RR, T_{\min} , T_{mean} and T_{\max} at which EIR_m occurred using
 175 a simple pair-wise comparison approach. This was done by first aggregating the EIR_m data
 176 from all locations within each zone into a single time series of 12 months irrespective of the
 177 year of availability. Similarly, the corresponding RR, T_{\min} , T_{mean} and T_{\max} data were also
 178 aggregated. The aggregated monthly timeseries of RR, T_{\min} , T_{mean} , T_{\max} and EIR_m were
 179 then matched head-to-head as shown in Figure 2. The ranges of RR, T_{\min} , T_{mean} and T_{\max}
 180 at which EIR occurred were then determined for each zone .

181 2.3.2 Relative importance of climate predictors

The relative importance of RR, T_{\min} , T_{mean} and T_{\max} in predicting EIR_m for each climate zone was analysed using a multiple regression model of the form:

$$EIR_m \sim RR + T_{\max} + T_{\min} + T_{\text{mean}} \quad (1)$$

where EIR_m is the response variable and RR, T_{\min} , T_{mean} and T_{\max} are the predictors. The contribution of each individual predictor to EIR_m outcome was then quantified (see Table 1 and 2). Each regressor's contribution was considered as the R^2 from univariate regression, and all univariate R^2 values add up to the full model R^2 (Grömping, 2007). The R package "relaimpo" (Grömping, 2007) was utilized for the calculation of the contribution of the regressors in the model. It implements six different metrics for assessing relative importance of regressors namely: first, last, pratt, betasg, lmg and pmvd. Among these, lmg and pmvd are computer intensive and has advantage over others in the sense that they decompose R^2 into non-negative contributions that automatically sum to the total R^2

(Grömping, 2007). In this study, lmg was invoked since pmvd is patent protected. The lmg calculates the relative contribution of each predictor to the R^2 with the consideration of the sequence of predictors appearing in the model. It intuitively decomposes the total R^2 by adding the predictors to the regression model sequentially. Then, the increased R^2 is considered as the contribution by the predictor just added. The following are mathematical descriptions of lmg metric referenced from Grömping (2007):

For a model with regressors in set S, the R^2 is given as:

$$R^2(S) = \frac{\text{ModelSS(modelwithregressorsin}S\text{)}}{\text{TotalSS}} \quad (2)$$

To add regressors in set M to a model with the regressors in set S, the additional R^2 is given as:

$$\text{seq}R^2(M|S) = R^2(MUS) - R^2(S) \quad (3)$$

where the order of the regressors is a permutation of the available regressors x_1, \dots, x_p denoted by the tuple of indices $r = (r_1, \dots, r_p)$. Let $S_k(r)$ denote the set of regressors entered into the model before regressor x_k in the order r . Then the portion of R^2 allocated to regressor x_k in the order r can be written as:

$$\text{seq}R^2(\{x_k\}|S_k(r)) = R^2(\{x_k\}US_k(r)) - R^2(S_k(r)) \quad (4)$$

With eq. 4, the metric lmg (in formulae denoted as LMG) can be written as:

$$\text{LMG}(x_k) = \frac{1}{P!} \sum_{r \text{permutation}} \text{seq}R^2(\{x_k\}|r) \quad (5)$$

Orders with the same $S_k(r) = S$ can be summarized into one summand, which simplifies the formula into:

$$\text{LMG}(x_k) = \frac{1}{p!} \sum_{S \subseteq \{x_1, \dots, x_p\} \setminus \{x_k\}} \text{seq}R^2(\{x_k\}|S) \quad (6)$$

The analysis also assessed the relative importance of each regressor (in eqn 1) by looking at what each regressor alone is able to explain (i.e., comparing the R^2 value of regression model with one regressor only without considering the dependence of others as is the case of the metric *lmg*). The metric *first* in the "relaimpo" package was invoked for this purpose because, unlike *lmg*, it is completely ignorant of the other regressors in the model and so no adjustment takes place (Grömping, 2007). Since *first* does not decompose R^2 into contributions like *lmg*), the contribution of the individual regressors alone do not naturally add up to the overall R^2 . The sum of these individual contributions is often far higher than the overall R^2 of the model with all regressors together.

191 Whether *lmg* or *first*, each metric's outcome were bootstrapped to ensure that the relative
 192 importance of each regressor was clearly defined (i.e. those different and those that are
 193 similar in terms of relative importance). Bootstrapping in "relaimpo" was done using the
 194 function *boot* in the package. Prior to calculating the *lmg* and *first* metrics, all data series
 195 (i.e. EIR_m , RR, T_{min} , T_{mean} and T_{max} timeseries) were log transformed. The essence of
 196 the log transformation was to decrease the variabilities in the data pairs and make them
 197 conform more closely to normal distribution with similar variance and standard deviation
 198 (Curran-Everett, 2018).

199 **2.3.3 EIR lag behind rainfall**

200 Seasonal malaria transmission onset lags behind rainfall season onset because of the time
 201 taken for mosquito breeding and vector population growth after rainfall season onset (Tomp-
 202 kins and Ermert, 2013; Badu et al., 2013; Asare and Amekudzi, 2017). This lag time as
 203 influenced by climate and whether it varies from one climate zone to another is not known.
 204 In this analysis, we quantified this lag time for each climate zone using a cross-correlation
 205 statistics performed between RR and EIR_m data pairs. In this statistics, RR was treated
 206 as the predictor variable and the corresponding EIR_m as the response variable. The pairs
 207 were then cross-correlated at lags of -5 to 0 months and the correlation co-efficient at each
 208 lag was calculated. The lag with the strongest positive correlation coefficients was identified
 209 as the optimum period of delay between rainfall onset and the EIR season for the zone.

210 **3 Results**

211 **3.1 pair-wise comparison**

212 Figure 2 shows the EIR_m response ranges of pairs of rainfall (RR) and temperature (T_{min} ,
 213 T_{mean} and T_{max}). In the Sahel, maximum rainfall (RR) ranges were about 400 mm per
 214 month. Temperature ranges generally varied between $20^{\circ}\text{C} - 40^{\circ}\text{C}$ in this zone. T_{max}
 215 ranges were clustered between $25^{\circ}\text{C} - 40^{\circ}\text{C}$, T_{min} within $20^{\circ}\text{C} - 30^{\circ}\text{C}$ and T_{mean} observed
 216 within $25^{\circ}\text{C} - 35^{\circ}\text{C}$. In Guinea, RR ranges were also centered around 400 mm per month.
 217 Temperature response ranges were mostly observed within $25^{\circ}\text{C} - 35^{\circ}\text{C}$ for T_{max} , $20^{\circ}\text{C} -$
 218 25°C for T_{min} and $24^{\circ}\text{C} - 30^{\circ}\text{C}$ for T_{mean} . In WCA, maximum RR ranges were centered at
 219 about 600 mm per month, which is higher compared to ranges observed in the Sahel, Guinea
 220 and EA. Temperature response ranges in this zone were slightly lower than observed in the

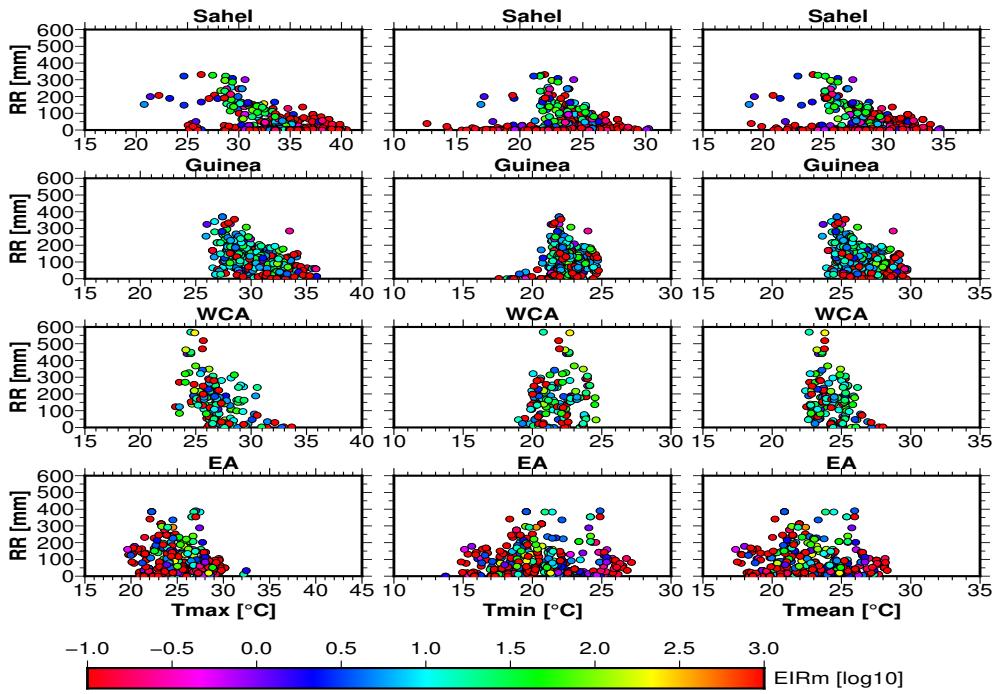


Figure 2: A pair-wise comparison showing the ranges of RR , T_{min} , T_{mean} and T_{max} at which EIR_m occurs. The coloured circles shows log transformed EIR_m values.

Sahel and Guinea. These include $24^{\circ}\text{C} - 32^{\circ}\text{C}$ for T_{max} , $20^{\circ}\text{C} - 25^{\circ}\text{C}$ for T_{min} and $22^{\circ}\text{C} - 27^{\circ}\text{C}$ for T_{mean} . The EA maximum RR ranges were also about 400mm. Temperature ranges of $20^{\circ}\text{C} - 30^{\circ}\text{C}$ for T_{max} , $15^{\circ}\text{C} - 27^{\circ}\text{C}$ for T_{min} and $18^{\circ}\text{C} - 29^{\circ}\text{C}$ for T_{mean} were observed.

3.2 Relative importance of climate predictors

In Table 1 and 2, the relative importance of climate variables in predicting EIR_m is presented for locations with elevations ≤ 500 m and > 1000 m respectively. The predictors with p-value ≤ 0.05 were considered significant and interpreted that the respective climate variable significantly predicted the EIR seasonality in that zone. At lower elevations (≤ 500 m) in Sahel, rainfall and temperature were all significant drivers of EIR seasonality cumulatively contributing about 30.72% of the variations in EIR seasonality. At these lower elevation areas, important predictors of EIR_m seasonality were RR and T_{max} . At higher elevations (> 1000 m), rainfall and temperature are together responsible for about 40% of the variations in EIR_m with insignificant contribution from T_{min} . Like the Sahel, temperature and rainfall

Table 1: **The relative contribution of RR, T_{\min} , T_{mean} and T_{\max} in predicting EIR_m** bootstrapped at confidence interval of 95% for locations with elevations ≤ 500 m. Variables with significant p-values contributions are boldfaced. R^2 represents the total proportion of variance in EIR explained by all the climate predictors. lmg values show the individual contribution of each predictor to R^2 relative to others. *First* is the contribution of each predictor alone to R^2 with complete ignorance of the others.

Zone	R^2 [%]	Variable	lmg [%]	First [%]	Coefficient (R)	P-value
Sahel	30.72	RR	7.73	15.76	0.3497	0.0000
		T_{\max}	12.03	17.54	-15.7276	0.0000
		T_{\min}	4.89	1.79	3.6380	0.0138
		T_{mean}	6.07	3.20	-6.8674	0.0033
Guinea	13.59	RR	5.85	10.22	0.4848	0.0000
		T_{\max}	4.09	9.65	-13.3808	0.0000
		T_{\min}	0.64	0.19	2.7410	0.3760
		T_{mean}	3.01	6.38	-15.7753	0.0000
WCA	1.69	RR	0.34	0.23	0.0974	0.5550
		T_{\max}	0.60	0.95	5.4640	0.3770
		T_{\min}	0.42	0.49	-4.5700	0.5810
		T_{mean}	0.33	0.35	6.8450	0.5280
EA	31.83	RR	0.62	0.00	-0.0141	0.9360
		T_{\max}	10.23	26.50	-23.2210	0.0000
		T_{\min}	8.84	20.10	-12.1120	0.0000
		T_{mean}	12.14	26.04	-18.2160	0.0000

were also significant determinants of EIR_m at lower elevations (≤ 500 m) in Guinea just that their contribution to EIR_m variations is small (about 13.59%) compared to that of Sahel (about 30.72%). In Guinea, also, EIR_m data were unavailable for locations > 1000 m for further analysis in this regard. In WCA, rainfall and temperature were insignificantly associated with EIR_m seasonality whether a lower or higher elevations. Their percentage explanation of the variation in EIR_m were also low (extremely low at lower elevation areas and slightly higher for higher elevation areas) compared to other climate zones. In EA,

Table 2: **Same as Table 1 but for locations with elevations > 1000 m.** In Guinea, EIR data were unavailable for locations at this elevation hence represented as dashed lines.

Zone	R ² [%]	Variable	lmg [%]	First [%]	Coefficient (R)	P-value
Sahel	40.47	RR	7.43	14.83	0.3745	0.0780
		T_{max}	17.66	35.91	-10.8070	0.0036
		T _{min}	3.69	4.17	-1.4543	0.5950
		T_{mean}	11.69	23.40	-7.7660	0.0513
Guinea	-	RR	-	-	-	-
		T _{max}	-	-	-	-
		T _{min}	-	-	-	-
		T _{mean}	-	-	-	-
WCA	16.55	RR	1.41	1.25	-0.0844	0.7653
		T _{max}	6.70	0.23	6.1700	0.6620
		T _{min}	1.53	0.39	14.1000	0.5970
		T _{mean}	6.91	1.32	12.2800	0.5300
EA	18.22	RR	10.44	13.37	0.5510	0.0000
		T_{max}	1.82	3.94	8.1570	0.0289
		T_{min}	2.88	7.77	10.2080	0.0011
		T_{mean}	3.08	6.95	11.0460	0.0026

temperature variables (T_{min} , T_{mean} and T_{max}) were the significant drivers of EIR seasonality at locations ≤ 500 m. It explained about 31% of the seasonality in EIR_m in these areas with extremely insignificant contribution from rainfall. But at areas > 1000 m, all the climate variables were significant contributors with rainfall showing higher contribution to EIR_m variation than temperature.

3.3 EIR lag behind rainfall

Figure 3 shows the seasonal relationship between rainfall and EIR_m (see Figure 3a) and the lag between rainfall onset and EIR onset (see Figure 3b). It is observed in Figure 3a that EIR_m is positively correlated with rainfall. A lag period of 1 month is observed in Sahel

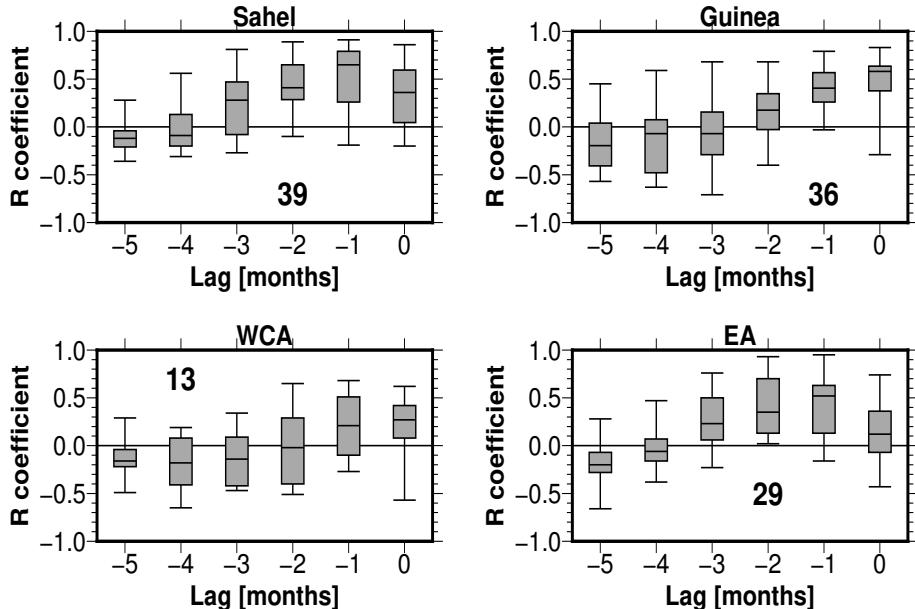
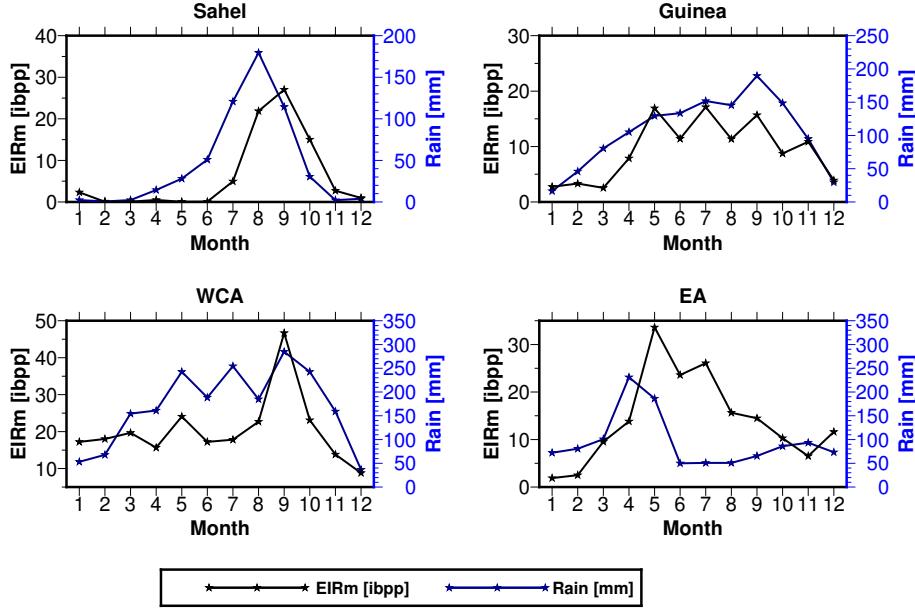


Figure 3: The correlation between rainfall and EIR_m . The numbers 39, 36, 13 and 29 shows the number of location observations contributing to the average timeseries (a) and the box-and-whisker plots (b).

and EA but zero month in Guinea and WCA. Similarly, the cross-correlation statistics determining the lag between the onset of rainy season and the start of the EIR season are shown in Figure 3b. Again, it is observed that the lag at which EIR_m seasonality strongly and positively correlated with rainfall was 1 month in the Sahel and EA but zero month in Guinea and WCA.

4 Discussion

Our study first examined the seasonal ranges of rainfall and temperature at which EIR_m occurred in a pair-wise comparison study. In general, temperature ranges of EIR_m response were mostly clustered between a minimum of 15°C and a maximum of 40°C. This outcome suggests that seasonal malaria transmission is barely impossible below 15°C or above 40°C. Previous studies (Shapiro et al., 2017; Parham and Michael, 2010; Lunde et al., 2013a; Mordecai et al., 2013) have indicated that malaria parasite development is not possible at temperatures below 16°C and that temperatures above 40°C have adverse effect on mosquito population turnover. The outcome of our study using EIR_m corroborates these previous findings. It provides an additional justification that the number of infectious mosquito bites a person receives per time are associated with changes in temperature. While T_{min} may be below 16°C as observed in the Sahel and EA (see Figure 2), the daily T_{mean} must be greater than 16°C particularly for the anopheles mosquitoes for transmission to occur. It should also be significantly less than 40°C for anopheles mosquitoes to survive thermal stress and possible death if seasonal transmission has to take place. Similarly, maximum monthly rainfall values for EIR_m occurrence was 600 mm in WCA but 400 m in the Sahel, Guinea and EA. The higher monthly maximum rainfall in WCA is due to the fact that annual total rainfall is mostly higher in this region than others (Nicholson, 2013; Froidurot and Diedhiou, 2017). Previous works (Craig et al., 1999; Ermert et al., 2011) have demonstrated that the least monthly amount of rainfall required for malaria transmission is about 80 mm. Our findings suggest that the monthly maximum limit required for seasonal malaria transmission should be about 600 mm in WCA but 400 mm in Sahel, Guinea and EA. Excess of these thresholds could result in flooding of breeding grounds and flushing out and killing the water-bound stage vectors (Paaijmans et al., 2010; Ermert et al., 2011).

The evaluation of the relative importance of RR, T_{min} , T_{mean} and T_{max} in predicting EIR seasonality (see details in Table 1 and 2) revealed climate variables that were significantly associated with EIR seasonality in Sub-Saharan Africa. These climate variables are ob-

served as the drivers of malaria seasonality in those zones of the sub-region. The climate variables with highest contribution to EIR variance in each zone are attributed as the most significant drivers. This means that any changes in these significant drivers can result in a substantial changes in malaria seasonality in those areas. Elevation or topography was also observed to play a significant role in determining the important climate drivers of seasonal malaria transmission. In EA for instance, temperature was the important determinant of EIR seasonality at lower elevated areas (≤ 500 m). On the contrary, both rainfall and temperature significantly influenced EIR_m seasonality at higher elevated areas (>1000 m). Though temperature and rainfall are important factors in malaria transmission, our study does not find them to have any significant association with EIR seasonality in WCA. This suggest that malaria seasonality in this zone is importantly driven by other factors other than climate. This require additional studies to unravel these factors driving malaria seasonality in this zone. Mabaso et al. (2007) predicted EIR seasonality from environmental data and found that seasonality in rainfall, minimum temperature, and irrigation were important determinants of seasonality in EIR in Sub-Saharan Africa. Though this study outcome is important, it is not climate specific as it does not justify the implications of diverse climate conditions on EIR seasonality as demonstrated in this study. Other studies (Mabaso et al., 2006; Simple et al., 2018) have used malaria case records from hospitals and found significant correlation between rainfall and temperature. As stated in the introduction, malaria case records have drawbacks for studying malaria seasonality as they are subject to significant uncertainties due to the inaccurate diagnostics and under counting due to varying health-seeking behaviour and health policies (Afrane et al., 2012b).

The cross-correlation statistics showed the lag(s) at which rainfall strongly correlated with EIR_m in each zone. The lag period suggest the time taken for malaria season to start after rainfall season has started. The lag of 1 month in Sahel and EA signifies that malaria transmission season delays 1 month after the start of rainfall season at these zones. In Guinea and WCA, this lag period was zero month suggesting that there is no delay between rainfall season onset and the start of malaria season. Hence malaria transmission in these zones is year-round. In markedly seasonal rainfall zones such as the Sahel and EA, the delay between rainfall onset and the start of the malaria season is expected. Rainfall in the Sahel is markedly seasonal, lasting from June to October, followed by about six to eight months of dry period (Nicholson, 2013; Froidurot and Diedhiou, 2017). Hence, mosquitoes are barely present during the dry and long hot season. Even if present, they are inactive

316 due to low humidity and high temperature and only recover within the rainy season when
 317 rainfall and temperature requirements are suitable. The absence of delay between rainfall
 318 season onset and the start of malaria season at Guinea and WCA is also expected. These
 319 zones are highly humid with shorter dry seasons (Nicholson, 2013; Froidurot and Diedhiou,
 320 2017). For this reason, vectors are able to persist all year round at these zones resulting
 321 in year-round transmission at these areas. Previous studies (Simple et al., 2018; Tompkins
 322 and Di Giuseppe, 2015; Reiner et al., 2015; Ikeda et al., 2017) have reported malaria lagging
 323 behind rainfall at about 1 to 2 months but our study has further demonstrated that malaria
 324 season onset may lag behind rainfall only at markedly seasonal rainfall areas in Sub-Saharan
 325 Africa.

326 5 Conclusion

327 Clinical malaria case data is commonly utilized as a malariometric in examining the rela-
 328 tionship between climate and seasonal malaria transmission in Sub-Saharan Africa. This
 329 data, on the other hand, is fraught with uncertainty due to out-of-sample generalization
 330 over geography and time, erroneous diagnosis, and under-counting due to varying health-
 331 seeking behavior and policy. As a result, in this work, we explored the applicability of
 332 high-quality EIR measurements to link rainfall and temperature seasonality to seasonal
 333 malaria outcomes in Sub-Saharan Africa. The main goal was to determine the climate
 334 variables that significantly drives malaria seasonality and their relative importance in the
 335 sub-region. Sub-Saharan Africa was first divided into four distinct climate zones namely
 336 Sahel, Guinea, WCA, and EA. The division was necessary because each zone has a unique
 337 climate conditions and therefore will have different climate implications on malaria sea-
 338 sonality. Applying a multi-regression statistics, pair-wise comparison and cross-correlation
 339 approaches to a EIR_m database gathered from publicly available field campaigns for each
 340 zone, the climate variables that explained significant variations in EIR seasonality were
 341 determined.

342 Findings in this study affirmed previous understanding that seasonal malaria transmission
 343 is barely impossible below 16°C or above 40°C temperature threshold (Shapiro et al., 2017;
 344 Mordecai et al., 2013). Hence, for seasonal malaria transmission to be sustained, average
 345 temperature should be well above the minimum or well below maximum threshold. While
 346 previous studies (Craig et al., 1999; Ermert et al., 2011) suggest that the monthly minimum
 347 rainfall requirement for seasonal transmission is about 80 mm, our study observed monthly

maximum rainfall limit should be about 600 mm in WCA, and 400 mm in the Sahel, Guinea and EA. While rainfall and temperature were found to be significantly associated with EIR_m seasonality in the Sahel, Guinea and EA, they were not important drivers of malaria seasonality in WCA. Important drivers of malaria seasonality in WCA may be due to other factors other than climate variables. In zones characterized by elevations such as EA, topography has a significant influence on which variable is an important determinant of malaria seasonality. At markedly seasonal rainfall areas such as Sahel and EA, malaria seasonal starts one month later after the rainfall season has started. However, for zones where rainfall season is bimodal such as Guinea and WCA, there is no delay between rainfall season onset and malaria season onset.

In this study, therefore, we showed that high quality EIR_m measurements can usefully supplement established metrics for seasonal malaria by demonstrating evidence for the use of EIR to directly link the risk of humans to infectious mosquito bites to climate. The study informs our understanding of the connection between climate variables and both the malaria vector and parasite biology and how that translates into malaria seasonality in Sub-Saharan Africa. This information is key for the improvement and validation of weather-driven dynamical mathematical malaria models that directly simulate EIR. Our findings provide an understanding of geographical heterogeneous malaria risk from climate effect and support future malaria modeling and forecasting efforts. The study also supplements previous works describing clinical patterns of malaria infection and morbidity. Taking into account the seasonality of malaria management, findings in this study could lead to significant public health advantages by assisting in determining when, where, and how to use vector and parasite control strategies. It can, therefore, help stakeholders establish a robust framework for monitoring, forecasting and control of malaria.

This study does not claim to have identified all the EIR_m data available across sub-Saharan Africa. It relied on EIR_m data available in repository (Yamba et al., 2018) with details explained in (Yamba et al., 2020). The study also acknowledges that the observed EIR_m data were both spatially and temporally limited and thus unavailable for many settings (as shown if Figure 1). This limitation was unavoidable because sampling mosquitoes for the determination of EIR is both labour and cost intensive. Hence, it is very difficult to have EIR_m data available for many locations and for a long period of time. Future mosquito sampling should, therefore, focus on areas of unavailable data in order to consolidate the spatial homogeneity of available EIR_m data distribution. However, an important strength

381 of this study is its restricted geographic and climate relevance. To our knowledge, this
 382 study is the first of its kind and also that EIR_m data has not been explored on such a wider
 383 scale in Sub-Saharan Africa. With the amount of EIR_m utilized for each climate zone, it is
 384 not anticipated that the inherent limitations may have any major adverse influence on the
 385 outcome of the study.

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389 **Competing interests**

390 The authors declare that they have no competing interests.

391 **Authors' contributions**

392 The work presented here was carried out collaboratively among all the authors. Edmund I.
 393 Yamba compiled the database, conducted the analysis for the figures and tables and drafted
 394 the manuscript. Leonark K. Amekudzi, Andreas H. Fink and Adrian M. Tompkins co-
 395 designed the project, supervised the analysis and co-authored the paper. Ernest O. Asare
 396 and Kingsley Badu contributed to result interpretation and co-authored the paper. All the
 397 authors proofread the paper.

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