

Chapter 1: Progress in Understanding and Parameterizing Fast Physics in Large-Scale Atmospheric Models

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Abstract

This introductory chapter discusses the atmospheric subgrid processes - collectively called “fast 11 physics” or “fast processes”, and their parameterizations in large scale atmospheric models. It 12 presents a brief historical progression of the parameterization of fast processes in numerical 13 models. Despite great efforts and notable advances in understanding, progress in improving fast 14 physics parameterizations has been frustratingly slow, the underlying reasons for which are 15 explored. To guide readers, this chapter describes the main objectives and scope of this book and 16 summarizes each chapter.

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2 **in Large- Scale Atmospheric Models**

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10 **Abstract**

11 This introductory chapter discusses the atmospheric subgrid processes — collectively called “fast
12 physics” or “fast processes”, and their parameterizations in large scale atmospheric models. It
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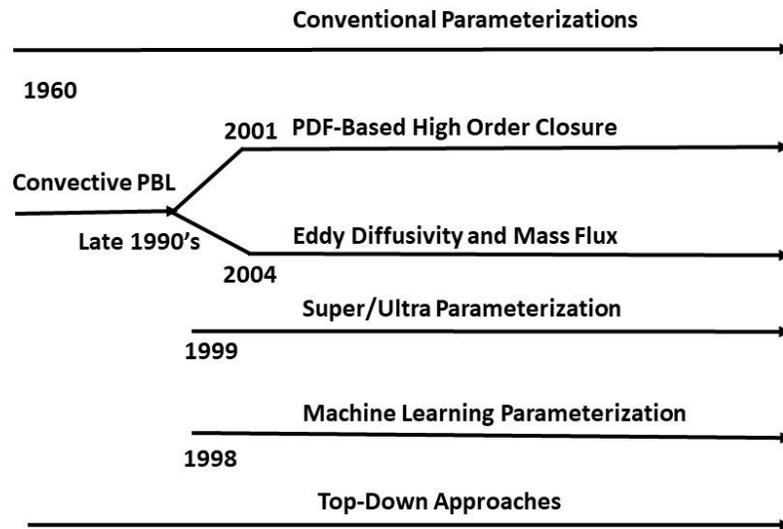
18 **1.1 Fast physics and progress of parameterization development**

19 Large scale atmospheric models are integral components of weather and climate models.
20 Ongoing developments in high-resolution modeling (i.e., global storm-resolving models (GSRMs,
21 Stevens et al., 2019); Energy Exascale Earth System Model (E3SM, Rasch et al., 2019); large-
22 eddy simulations (LES, Gustafson et al., 2020) have resulted in ultra-high resolution numerical
23 simulations of atmospheric systems. Despite these advancements, coarser resolution large scale
24 models remain our main modeling capability for future climate predictions. Many atmospheric
25 processes and phenomena that influence Earth’s weather and climate occur at spatiotemporal
26 scales that are too small to be resolved in these large-scale atmospheric models and must be
27 parameterized — approximately represented by the variables that can be resolved by the model
28 grids. In this book, we refer to this array of parameterized subgrid processes and phenomena
29 collectively as “fast physics” or “fast processes” for convenience, including radiative transfer,
30 aerosol/cloud physics, convection, boundary layer processes, gravity wave, and land-atmosphere
31 interactions.

32 While early parameterizations of fast physics used simple and often empirical or ad hoc
33 relationships (e.g., the Kessler bulk parameterization for representing cloud microphysical
34 processes, Kessler, 1969), later parameterization development has become concerned with
35 building conceptual models with increasingly detailed physical processes by leveraging theoretical
36 analysis, observations, and/or detailed process modeling studies.

37 Furthermore, parallel to the continuing improvement/development of parameterizations for
38 individual fast processes, there has been growing interest in and studies on understanding
39 interactions/couplings among different processes. Significant progress has been made and several
40 promising approaches have emerged since late 1900s and early 2000s. Figure 1 illustrates the
41 approximate timelines in developing fast physics parameterizations in context of the conventional
42 parameterizations that target individual processes and unifying efforts that addresses process
43 interactions.

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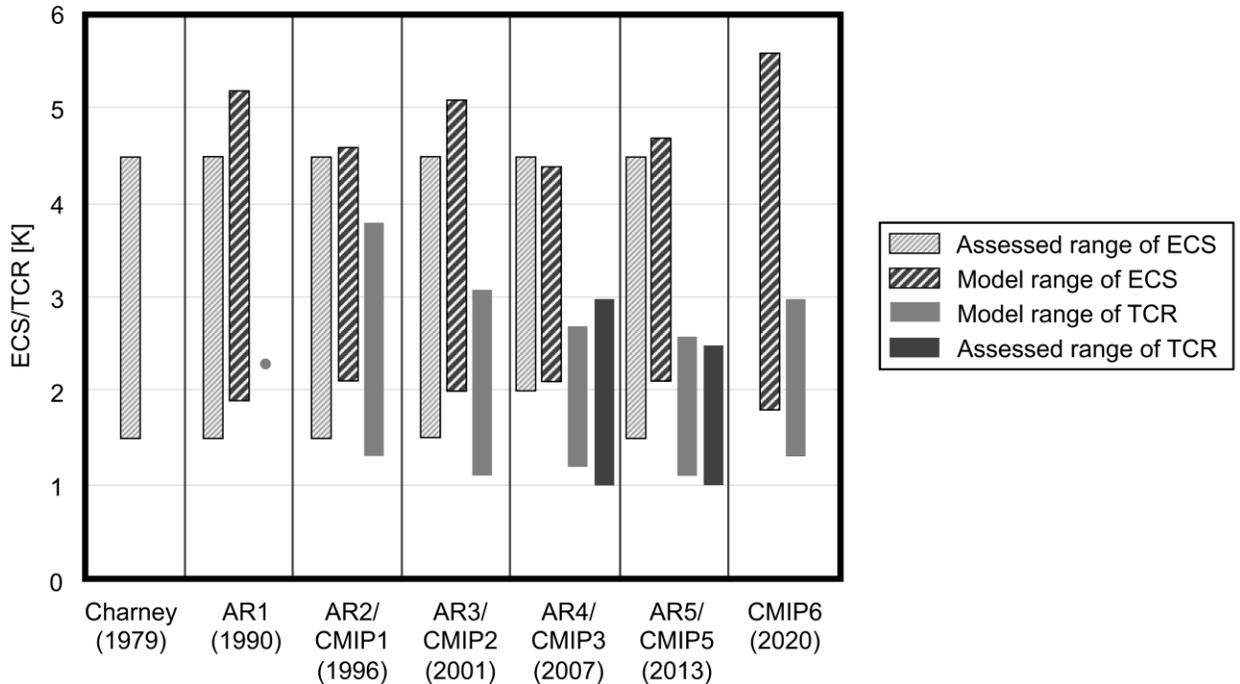


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46 Figure 1. Schematic of the approximate timelines of development of fast physics
 47 parameterizations. Conventional parameterizations are focused on individual fast processes. The
 48 four lines of unifying efforts (PDF-based High Order Closure, Eddy Diffusivity and Mass Flux,
 49 Super/Ultra Parameterization, and Machine Learning Parameterization) aim to unify the
 50 representation of more than two physical processes. Top-Down Approaches borrow holistic ideas
 51 that have been scattered in various disciplines (e.g., nonlinear systems dynamics, statistical
 52 physics, information theory, self-organization, networks, and pattern formation).

53 Despite remarkable efforts and increasing recognition of the importance of these fast
 54 processes over the past few decades; progress remains frustratingly slow in improving their
 55 representation in models. As a result, their impact on future climate predictions remain poorly
 56 understood and highly uncertain. The slow progress is perhaps best attested by the historical lack
 57 of change in the ranges of climate sensitivity across models from the celebrated 1979 U.S. National
 58 Research Council report (Charney et al. 1979) to the latest (6th) Coupled Model Intercomparison
 59 Project (CMIP6) results used in the Intergovernmental Panel on Climate Change (IPCC) report
 60 (Fig. 2). Deficient fast physics parameterizations, and especially those related to clouds, have been
 61 thought to be primarily responsible for the stubborn large spread of model climate sensitivity
 62 (Meahle et al., 2020; Zelinka et al., 2020). Aerosol climate forcing in climate models has been
 63 fraught with similarly unchanged uncertainty (see more in Chapter 3 of this book).

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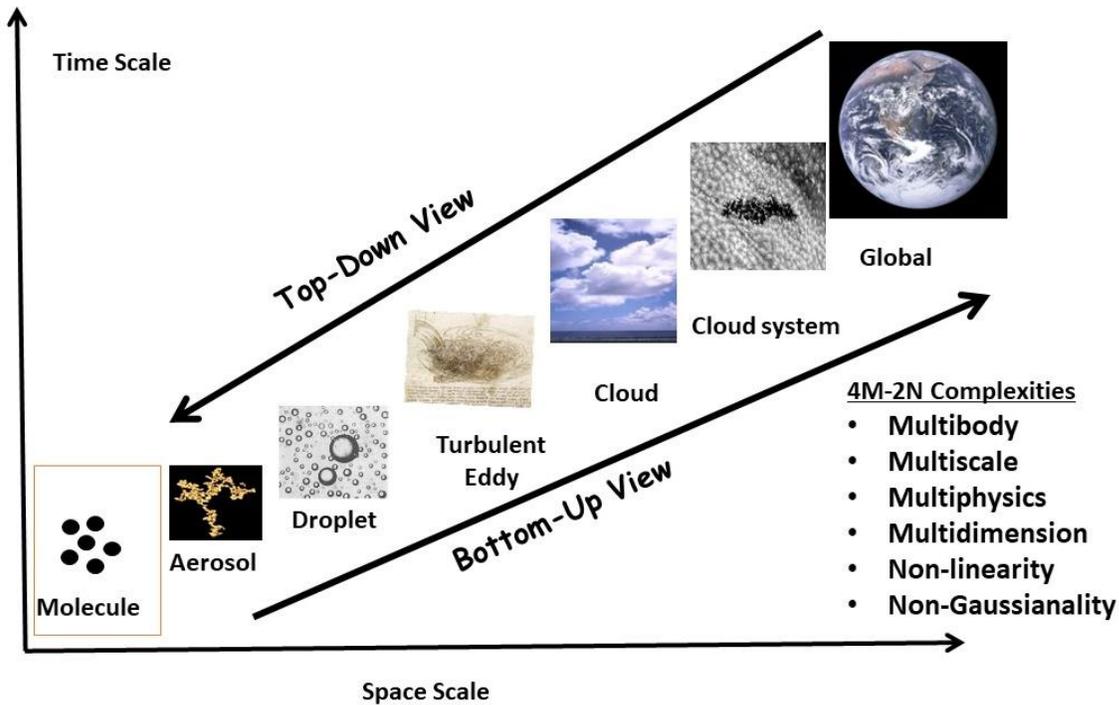


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66 Figure 2. Historical values of equilibrium climate sensitivity (ECS) and transient climate
 67 response(TCR). Figure is modified from Fig. 1 of Meahle et al. (2020) which can be consulted
 68 for details on the data sources and definitions.

69 The slow progress can be attributed to two overarching types of complexities (also see
 70 Randall, 2013; Jakob, 2010). The first lies in the “4M-2N complexities” inherently accompanying
 71 the atmosphere and associated physical processes (Fig. 3). Briefly, fast processes and especially
 72 those cloud-related ones involve *multibody* (sub)systems with numerous particles of different sizes
 73 and shapes, in which multiple physical processes (*multiphysics*) occur over a wide range of
 74 spatiotemporal scales (*multiscale*) and interact with one another, and manifest themselves in a
 75 variety of cloud types such as cumulus and stratiform clouds (*multitype*). The equations describing
 76 these processes are often highly *non-linear*, and exhibit *non-Gaussian* statistics (Lovejoy and
 77 Schertzer, 2010).

78 The other inherent complexity lies with that model development involves an iterative cycle
 79 of developing parameterizations, implementing and evaluating parameterizations against
 80 observations to identify potential parameterization deficiencies and further improvement. This
 81 iterative procedure calls for an organic integration of the key components involved ranging from
 82 modeling to measurements, which in turn demands effective coordination of expertise in distinct
 83 areas that have proven to be extremely challenging. However, effective coordination and
 84 collaboration across different disciplines and institutions are not trivial, and such an “operational
 85 complexity” adds another layer of technological and social challenges in virtually every step of
 86 model development. The issue will become more acute as the field is moving toward more
 87 emphasis on processes interactions with ever increasing data volumes and model resolutions. To
 88 echo Jacob (2010), “... acceleration in model development can only be achieved by significantly
 89 strengthening these weak links through additional research and better coordination across existing
 90 programs.”



91

92 Figure 3. Schematic to illustrate the atmospheric scale hierarchy and involved “4M-2N
 93 complexities”. Together with the “operational complexity” discussed in the text, these science
 94 complexities have posed and will continue to pose challenges to model development in general
 95 and fast physics parameterizations in particular.

96

97 **1.2 Objectives and scope of the book**

98 The objectives of this book are three-fold. First, to survey advances in understanding of
 99 key fast processes and their parameterization developments (Part I). In particular, Part II of this
 100 book is uniquely devoted to unifying efforts. Second, unlike most review articles or the book by
 101 Stensrud (2007) on fast physics parameterizations, this book includes discussions on measurement
 102 techniques and studies that use observations for model evaluation and thus covers approaches to
 103 addressing the weak link in the model development loop. Third, by surveying the recent advances
 104 in key areas, we hope to reveal new challenges, opportunities, and directions for future research.

105 It is worth noting that the related literature is enormous and that the selection of the
 106 material in this text is non-exhaustive and likely biased to the authors’ own research interests. On
 107 the other hand, books focusing on fast physics parameterizations are rare; the only one we are
 108 aware of is Stensrud (2007), which is primarily on conventional parameterizations of individual
 109 fast processes in numerical weather prediction (NWP) models. Bringing together modeling and
 110 measurements with a common goal of parameterization development and evaluation, and
 111 including unifying efforts are unique to this book.

112 **1.3 Book structure and summary of chapters**

113 Fast physics in large scale atmospheric models involves multiple processes that occur over
114 a wide range of spatiotemporal scales. Progress has been made on many fronts and new promising
115 directions of research are emerging. To reflect and synthesize the multiple facets involved, this
116 book is divided into three parts. Part I deals with the major subgrid processes, with eleven chapters
117 (Chapters 2 to 12) covering different fast processes. Beyond conventional treatments, some
118 promising approaches have recently emerged to unify the treatment of (some) processes and thus
119 allows for consideration of processes interactions. Part II is devoted to such unifying efforts, with
120 four chapters (Chapters 13 to 16) that each cover a different endeavor. Part III is devoted to
121 measurements, model evaluation, and model-measurement integration, with four chapters
122 (Chapters 14 to 17) that focus on satellite and airborne remote sensing measurements, surface-
123 based remote sensing measurements, in-situ and laboratory measurements, and model evaluation
124 and model-measurement integration, respectively.

125 **1.3.1 Process studies and parameterizations**

126 Primary to the Earth's climate and weather and understanding climate change is the
127 understanding and representation of the solar (shortwave) and terrestrial (longwave) radiation and
128 of radiative transfer processes such as absorption, scattering, and transmission. In *Chapter 2, Gu*
129 *and Liou* present the fundamentals of radiative transfer and its interactions with the atmosphere,
130 and summarizes the commonly used radiative transfer parameterization schemes in atmospheric
131 models. Also discussed are several more advanced topics in the study of the atmospheric radiation,
132 including cloud vertical overlapping, cloud horizontal inhomogeneity, and 3D radiative transfer in
133 both the cloudy atmosphere and over complex rugged land surfaces such as mountainous terrains.
134 In particular, the chapter highlights that the current commonly used radiation schemes normally
135 represent 1D transport in the vertical direction, although radiative transfer in 3D atmosphere and
136 surfaces could play an important role in determining the radiation budget and radiative heating at
137 the top of the atmosphere, at the surface, and within the atmosphere. Both horizontal and vertical
138 subgrid scale inhomogeneities, and 3D radiative transfer may substantially influence the radiative
139 transfer within clouds and cloud-radiation interactions, suggesting the need for further
140 investigation and for improving their representations in models.

141 Atmospheric aerosols are suspensions of solid particles or liquid droplets in the air.
142 Aerosols contain multiple compositions, exhibiting various morphologies and span a few orders
143 of magnitude in sizes from a few nanometers to tens of micrometers. Aerosol radiative effects
144 constitute one of the largest uncertainties in climate projection, and the large spread of simulated
145 values among general circulation models (GCMs) can be traced to different representations of
146 aerosol processes, including emissions, transport, formation and removal, and aerosol-cloud
147 interactions. In *Chapter 3, Liu* provides an overview of atmospheric aerosols and their climatic
148 impacts through both aerosol direct effects on radiation (aerosol-radiation interactions) and
149 aerosol indirect effects (aerosol-cloud interactions). The authors focus on addressing topics
150 related to three aerosol-related questions: 1) How are aerosol properties and processes as well as
151 aerosol-cloud interactions represented and compared in current GCMs? 2) What are the major
152 assumptions, simplifications and weaknesses of the current representations? 3) Why are there

153 large uncertainties in the aerosol climate effects from GCMs? Several future directions are
154 highlighted.

155 Although entrainment of surrounding dry air into clouds, subsequent turbulent mixing
156 processes, and their microphysical influences haven been known to be essential in determining
157 cloud microphysical and related properties for some time, theoretical understanding of these
158 processes is still far from complete, and their parameterizations in atmospheric models are in
159 their infancy. In *Chapter 4*, the authors (Lu, Liu,) discuss these issues in shallow clouds
160 (cumulus and stratocumulus clouds), focusing on two critical yet understudied aspects:
161 entrainment-mixing mechanisms and entrainment rate. Different conceptual models of
162 entrainment-mixing mechanisms are reviewed, and latest studies on unifying microphysical
163 measures to quantify different entrainment-mixing mechanisms are presented. Approaches for
164 estimating fractional entrainment rate in cumulus clouds are summarized; relationships of
165 entrainment rate to internal cloud properties (e.g., vertical velocity) or external properties (e.g.,
166 relative humidity in environment) are discussed as plausible parameterizations. Three approaches
167 for estimating entrainment velocity in stratocumulus clouds are also discussed. Several topics are
168 highlighted for future research, e.g., the connection between entrainment rate, entrainment-
169 mixing mechanisms, and relationships to other factors (e.g., rain initiation, detrainment, spectral
170 shape of cloud droplet size distributions, entrained aerosols, and environmental relative
171 humidity).

172 Following the discussion on entrainment in shallow cumulus clouds and its role in
173 shallow convection parameterization, *Donner* turns to deep convection from the perspective of
174 large-scale flows in *Chapter 5*, which, as a practical matter, comprises the problem of (deep)
175 cumulus parameterization. The chapter begins with discussing the effects of convection on large-
176 scale flows in which it is embedded, follows with strategies for solving the problem of cumulus
177 parameterization, and concludes with a brief overview of interactions between convection and
178 momentum, chemistry, tracers, cloud microphysics, and aerosols. Emphasized are the roles of
179 convective vertical velocities in treating aerosol-cloud interactions and cloud microphysics
180 related to cloud feedbacks. Major deficiencies in existing parameterizations are discussed,
181 including interactions between deep convection and aerosols, convection-chemistry interactions,
182 understanding and representation of convective organization, and knowledge of convective-scale
183 pressure-gradient forces in treating effects of convection on momentum fluxes. Limitations of
184 mean-state perspectives and the widely used quasi-equilibrium assumption are discussed. Also
185 touched on are connections with other topics (e.g., scale awareness, higher-order closure, multi-
186 scale modeling frameworks and high-resolution models without conventional deep convection
187 parameterizations, shallow convection, boundary-layer processes and gravity waves) detailed in
188 other chapters.

189 Besides convective clouds, stratiform clouds including stratus and stratocumulus clouds
190 constitutes another critical component of the atmospheric system that significantly affects climate
191 and has long been the subject of active research from many perspectives. In *Chapter 6*, *Dong and*
192 *Minnus* provide an overview of such clouds, with a focus on what we have learned from
193 observational studies in terms of improving their parameterization in atmospheric models.
194 Stratus and stratocumulus cloud properties and their importance are discussed based on
195 measurements from trained surface observers, satellite and ground-based remote sensors, and
196 aircraft field campaigns. The processes that determine the variations in stratocumulus properties

197 and govern where and when they occur are discussed, along with such factors as aerosols,
198 radiation, and humidity. Retrieval methods used for extracting information about stratus and
199 stratocumulus clouds from satellite- and ground-based sensors are also briefly reviewed, with an
200 emphasis on the knowledge learned for improving understanding and parameterizations of such
201 clouds in large scale models. Unique consistency between the early trained observers and the
202 state-of-art technologies are demonstrated; synergy of different observational platforms is
203 highlighted for future investigation. Emerging but understudied phenomena are summarized,
204 including the impact of low-level temperature advection, veil clouds developing at the top of the
205 marine boundary layer in areas of open-cell and unorganized cellular convection, the role of
206 gravity waves in the subtropical jet stream in initiating Pocket of Cells (POCs) in some closed-
207 cell stratocumulus over the southeast Pacific, and effects of land-sea breezes. Outstanding issues
208 in profiling marine boundary layer cloud and drizzle microphysical properties are highlighted,
209 including the need for incorporating cloud-top entrainment, drizzle, and vertical and horizontal
210 inhomogeneities to address the issue of nonadiabatic multispectral retrievals.

211 As a layer between the ground surface and the free troposphere, the planetary boundary
212 layer (PBL) is often turbulent, and particularly important because the majority of biota (including
213 humans) and climatically important low clouds like stratocumulus and shallow cumulus reside.
214 Even deep convection is highly related to the properties of the plumes or thermals originating in
215 the PBL. In *Chapter 7, Ghate and Mechem* introduce the PBL structure and the commonly used
216 theoretical approaches for investigating the PBL. A hierarchy of models for representing the
217 boundary layer is presented, including mixed-layer models, first-order closure, 1.5-order TKE
218 closure, and higher-order closure approaches. Challenges for evaluating the emerging advanced
219 schemes (high order, PDF-based, or EDMF) are also discussed in context of the inherent needs
220 for observations of joint PDFs of vertical air motion and thermodynamic variables. The
221 discussion emphasizes the buoyancy-driven convective boundary layer but briefly mentions
222 impacts of shear and clouds. The chapter concludes with a brief historical context and future
223 outlook for representing the boundary layer in large scale atmospheric models. To some extent,
224 this chapter can be viewed as an introduction to Chapters 13 and 14 where the PDF-based and
225 EDMF schemes are detailed.

226 Although the focus of this book is on atmospheric processes, the weather and climate
227 system consists of other sub-systems that strongly interact with the atmosphere over a wide range
228 of spatiotemporal scales. In particular, various surface processes are fundamental to the exchange
229 of heat, water and momentum between the surface and the atmosphere through PBL. As such
230 modeling land-surface processes has been an integral component of atmospheric models. In
231 *Chapter 8, Barlage and Chen* and focus on recent progress in understanding and modeling the
232 biophysical effects of the human dimension, especially urbanization and agriculture, on surface
233 water and energy budgets, and their cascading effects on weather and climate including clouds,
234 aerosols, convection and precipitation. Well-known phenomena are discussed, including the
235 Urban Heat Island (UHI) and urban impacts on precipitation through both cloud microphysical
236 and/or dynamical effects. Also discussed are the unique roles of rough vegetated or urban canopy
237 in determining turbulent fluxes (e.g., evapotranspiration over vegetated regions can exceed
238 evaporative flux from the oceans because larger surface roughness and stronger turbulence
239 whereas the moisture flux can be effectively shut off when the land is dry). Land-surface models
240 (parameterizations) of 3D subgrid structures within urban or vegetation canopies are presented,
241 including the most sophisticated multi-layer scheme — BEP (Building Effect Parameterization).

242 Challenges for evaluating and applying such a comprehensive land-surface model are discussed,
 243 including existence and specification of the large number of tunable parameters used in urban
 244 canopy models.

245 Atmospheric gravity waves (GW) have horizontal wavelengths ranging from 1 to 1000's
 246 of kilometers. Current climate models, and even numerical weather prediction models cannot
 247 resolve significant portions of their momentum flux and parameterizations are necessary to
 248 represent their under- or unresolved effects in atmospheric models. In *Chapter 9, the authors*
 249 *(Kruse, Richter, Alexander, Bacmeister, and Wei)* discuss GWs that are important at nearly all
 250 levels of the atmosphere, especially for the general circulation of the middle and upper
 251 atmosphere. GWs in the tropical stratosphere contribute significantly to the driving of the quasi-
 252 biennial oscillation and the stratospheric and mesospheric semi-annual oscillation, both primary
 253 modes of variability up there. In the extratropics, GWs contribute to the driving of the
 254 stratospheric Brewer-Dobson circulation and significantly influence the strength of the polar
 255 night jet and the corresponding polar temperatures. Additionally, GWs are responsible for the
 256 cold summer mesopause and the reversal of extratropical zonal mean winds in the mesosphere.
 257 Also discussed are both primary sources of atmospheric GWs (i.e., flows over mountains, moist
 258 convection, and imbalances in jets and frontal systems) and secondary GWs generated as a result
 259 of dissipation of primary GWs. The basic theory of GW generation, propagation, and dissipation
 260 and commonly used GW parameterizations are presented. Uncertainties, parameter tuning, and
 261 known missing processes in current parameterizations are explored as well. The importance of
 262 gravity wave in shaping clouds (*Chapter 7*) and in determining cloud microphysical properties
 263 (*Chapter 3*) are gradually recognized as well.

264 **1.3.2: Unifying efforts**

265 *Chapter 10* is the first of the four chapters that introduce the emerging efforts to unify the
 266 parameterizations of different processes, with a focus on higher-order equations closed by
 267 assuming the shape of the probability density function (PDF) of fields on the subgrid scale (PDF-
 268 based method for short). In this chapter, *Larson* presents the general equations involved.
 269 Theoretical analysis of the higher-order equations reveals that they contain the flux-of-flux terms
 270 that lead to non-local cumulus transport, along with a detailed representation of buoyant
 271 generation of turbulence, which is the essential source term of convection and can be closed by a
 272 multivariate PDF. Instead, traditional low-order closure omits the flux-of-flux terms that are
 273 crucial for representing nonlocal cumulus transport. The popular Cloud Layers Unified By
 274 Binormals (CLUBB) is detailed as an example of such PDF-based methods. Other higher-order
 275 closure models are also briefly discussed, including the Intermediately Prognostic Higher-Order
 276 Closure (IPHOC) parameterization (Cheng et al., 2010), which prognoses all the moments
 277 prognosed by CLUBB, plus two additional third-order moments of water vapor and potential
 278 temperature, the Turbulence Kinetic Energy-Scalar Variance (TKESV) parameterization of
 279 Mironov and Machulskaya (2017), which prognoses TKE and scalar variances, and optionally a
 280 third-order moment related to cloud liquid water. The connections of the PDF-based method to
 281 the conventional the mass-flux method for convection and low-order closure for turbulence are
 282 also discussed.

283 Another approach that seeks to unify the treatment of convection and turbulent processes
 284 in PBL is conceptually more direct, combining the widely used eddy diffusivity approach for

285 local turbulent transport with the mass-flux scheme for convection. In *Chapter 11*, the authors
286 (*Teixeira, Suselj and Kurowski*) discuss the EDMF approach. After briefly reflecting on the early
287 development in the late 1990s, this chapter is focused on the new stochastic multi-plume EDMF
288 scheme that can realistically represent the dry boundary layer, stratocumulus, shallow and even
289 deep cumulus convection within a single framework. The surface variability of updraft properties
290 is parameterized using joint PDFs of thermodynamic properties to initialize multiple updrafts.
291 Lateral entrainment is parameterized as a stochastic process. Furthermore, the unified EDMF
292 parameterization explicitly considers the horizontal resolution of the model, paving the way to a
293 scale-aware extension of the scheme. Both the fundamentals and latest results of using the new
294 EDMF scheme are introduced. The multi-plume framework allows for the coexistence of
295 different convective regimes (i.e., dry plumes, shallow moist convection, and even deep
296 convection) within a single grid-box, without any artificial separations between them and with
297 scale-adaptive capabilities for use in next-generation weather and climate models with high and
298 variable horizontal resolutions.

299 The PDF-based and EDMF approaches both aim to unify the parameterizations of
300 turbulence, PBL and convection (especially that of shallow convection). Further coupling with
301 other processes such as cloud microphysics remains an area of active research for both
302 approaches. Around similar times in the late 1990's and early 2000's, ideas of super-
303 parameterization - that embed cloud-resolving models (CRM) in climate model grid column -
304 were proposed and developed as a way to replace all the subgrid processes that the embedded
305 CRM model represents, including turbulence, PBL, convection, cloud microphysics and
306 radiation (Grabowski, 2001; Randall, 2013). Recently similar ideas were extended to using high
307 resolution large eddy simulation (LES) models instead of CRMs in so-called ultra-
308 parameterization (Parishani et al., 2017). Obviously, the benefits of such multiscale modeling
309 approaches come at high computational cost and call for more computationally effective
310 approaches that can be used as alternative to represent multiple fast processes together. In
311 *Chapter 12*, *Krasnopolsky and Belochitski* describe applications of machine learning (ML)
312 approaches to emulate existing parameterizations and developing new ML surrogate models as
313 new parameterizations. The authors first argue that a parameterization can be formulated as a
314 generic problem of mathematical mapping, and then argue that ML tools can be used to emulate
315 and/or approximate the involved mathematical mappings. Four mapping complexities (physical
316 complexity, mathematical complexity, numerical/computational complexity, and functional
317 complexity) are discussed. Further discussed are ML applications to emulate existing
318 parameterizations, to develop new parameterizations, to ensure physical constraints, and control
319 the accuracy of developed applications. Some ML approaches that allow developers to go
320 beyond the standard parameterization paradigm are discussed as well. Limitations of ML models
321 are also discussed, including inability to provide a meaningful physical interpretation of
322 underlying processes, requirements of large data for training and testing purposes, and their
323 limited generalizability to out-of-sample scenarios. Given that neither an ML-only nor a
324 physically based-only approach can be considered sufficient for complex scientific and
325 engineering applications, the research community has been exploring the hybridization of
326 physically-based and ML-based models, where both scientific knowledge and data are integrated
327 in a synergistic manner. It is reasoned that this hybrid paradigm is fundamentally different from
328 the ML mainstream where domain-specific knowledge is often considered secondary, and
329 several differences are discussed. The concept of compound parameterization (CP) that combines

330 an ML parameterization, the original physically-based parameterization, and a quality control
 331 procedure is introduced.

332 Despite the tremendous advances and different extents in dealing with the number of fast
 333 processes and their interactions discussed in the previous chapters, a common theme of those
 334 studies is that they are all essentially bottom-up-based and aim to upscale subgrid scale processes
 335 to grid variables. However, the climate system, including its atmospheric component, is a
 336 multiscale complex system that involves highly nonlinear bottom-up and top-down scale
 337 interactions (recall Fig. 3). Without considering the top-down direction, our understanding would
 338 never be complete, and the physical pictures from the unifying efforts could be as murky as
 339 understanding the output of a full GCM. As another unique addition of this book compared to
 340 existing ones, in *Chapter 13, Feingold and Koren* -summarize innovative ideas that attempt to
 341 consider processes holistically but are scattered in various disciplines including nonlinear
 342 systems dynamics such as chaos theory, statistical physics, information theory, self-organization,
 343 networks, pattern formation, and general systems theory. The “top-down view” is focused on
 344 system-wide behavior and emergent phenomena, distinguishing from the traditional “bottom-up”
 345 view that focuses on individual processes. In particular, a behavior at a larger scale can emerge
 346 from interactions/couplings between detailed processes and between the involved sub-systems at
 347 a finer scale. And this type of order/emergence is not driven by an external force, but instead
 348 grows spontaneously from local interactions, or is ‘self-organized’. Spatiotemporal
 349 communication between components of a system is key to development of synchronization,
 350 patterns, and self-organization. The top-down approach can yield simple holistic models that are
 351 more amenable to interrogation and digestion than complex, detailed models. Concepts and
 352 terminologies that are not that familiar to the atmospheric modeling, esp., the parameterization
 353 community, are introduced and discussed, including fixed points, attractors, limit cycles, chaotic
 354 state, bifurcation points, synchronization, information content, and entropy. In addition to their
 355 distinct foci on local and detailed physical processes vs. process interactions and emergence, this
 356 chapter also provides some intriguing examples to elucidate the conceptual differences between
 357 the bottom-up and top-down approaches from other perspectives: Reductionism vs. Holism;
 358 Basic building blocks vs. an Integrative view; Models representing a Large vs. a Reduced
 359 number of degrees of freedom; Models rooted in mathematical representation of
 360 physical/chemical/biological processes vs. Models that are an abstraction of these processes;
 361 Complexity vs. Simplicity. The authors use aerosol-cloud-precipitation system as a particular
 362 example to demonstrate the great potentials of the top-down view, and the need to integrate the
 363 complementary top-down view and bottom-up thinking in addressing the remaining challenge.

364 **1.3.3 Measurements, model evaluation and model-measurement integration**

365 Reliable observations are always important to improve our understanding of natural
 366 phenomena including atmospheric processes, and serve as the ground truth to verify and evaluate
 367 any theoretical and modeling developments. The synergy between model development and
 368 observations are becoming increasingly important as both fields progress. Earth science
 369 observations in general and atmospheric observations in particular have unique features,
 370 involving different but complementary approaches: surface-based , satellite-based and airborne

371 remote sensing, in-situ field measurements and laboratory studies. This part is devoted to such
372 crucial endeavors, with four chapters focusing on four different topic areas summarized next.

373 *Chapter 14* focuses on surface-based remote sensing for the study of the macro- and
374 micro-physical structure of clouds, precipitation, aerosols, and the clear boundary-layer. In this
375 chapter, the authors (*Lamer, Kollias, Amiridis, Arinou, Loehnert, Schnitt, and McComiskey*)
376 place their emphasis on the unique ability of ground-based system to continuously characterize
377 the atmosphere at high vertical resolution from near the surface to the top of the atmosphere
378 effectively filling observational gaps left by spaceborne and aircraft platforms. Following an
379 overview of the emergence of ground-based observatories details about the measurement
380 principals of their cornerstone instruments (Cloud and Precipitation Radars, Lidars and
381 radiometers) is given. Modern techniques to retrieve cloud and precipitation location,
382 microphysical and dynamical properties as well as planetary boundary layer structure and aerosol
383 properties are discussed including their underlying assumptions and uncertainties. Challenges
384 associated with using ground-based observations for model evaluation are discussed including:
385 1) The fact that most measurements are related to moments of these particle size distributions
386 that differ from those of most interest; the 0th (total number concentration) and 3rd (mass
387 content) moments) being most desirable and the 6th and 2nd being recorded by radars and lidar,
388 respectively. And 2) the fact that high-resolution observations and large-scale models are widely
389 different scales; GCMs having grid resolutions ~50 km while radars have range resolution of
390 ~30m. The growing role of synthetic Observing Systems Simulation Experiments (OSSEs),
391 instrument simulators and sub-column generators in bridging those gaps is emphasized. The
392 chapter closes with an outlook on the next generation of ground-based observatories that should
393 employ scanning systems and distributed networks.

394 *In Chapter 15, -Marshak and Davis* cover remote sensing retrievals of cloud and aerosol
395 properties from overhead instruments, including satellite-based and airborne sensors with
396 standoff distances ranging from NASA's P-3B aircraft at about 3.5 km above cloud top to the
397 Deep Space Climate Observatory (DSCOVR) platform at the Lagrange-1 point, about
398 1,500,000 km toward the Sun. The electromagnetic spectrum covered ranges from the ultra-
399 violet to microwaves, and both traditional passive and relatively recent active (e.g., lidar and
400 radar) and modalities are discussed. The emphasis is on the physics behind the sensors as well as
401 on the retrieval algorithms. The chapter starts with remote sensing of cloud properties,
402 introducing the popular Nakajima–King approach widely used in retrievals of cloud optical depth
403 and particle size, and the Bréon–Goloub approach that is based on the directional signature of the
404 polarized reflectance. Techniques for retrieving (mostly ice) cloud properties with microwave
405 sensors is also briefly described. The chapter then switches to remote sensing of aerosol
406 properties, describing the main aerosol remote sensing approaches used by the major satellite
407 imagers. In addition to passive remote sensing, the active remote sensing methods for aerosol
408 and cloud profiling are also highlighted, with an emphasis on CALIPSO and CloudSat for lidar
409 and radar, respectively. The chapter explores oxygen A- and B-band remote sensing of cloud and
410 aerosol layer height, along with other passive methods for estimating cloud top height. A special
411 section is dedicated to cloud remote sensing at very high spatial resolution either from tasked
412 imaging sensors in space or from airborne platforms deployed above the clouds of interest, such
413 as NASA's ER-2. New studies on the transition zone are reviewed. The chapter closes with a

414 brief discussion of retrieval uncertainty quantification for both cloud and aerosol remote
415 sensing.

416 Laboratory experiments allow more controlled, repeatable measurements of a physical
417 quantity or phenomenon in a well-defined system of interest with known external influences. In
418 *Chapter 16, Chandrakar and Shaw* describe in-situ measurements and laboratory experiments,
419 with a focus on physical processes that are small in spatial or temporal scale such as cloud
420 microphysics and small-scale turbulence. Some illustrative historical examples are given to
421 highlight significant advances and capabilities in three areas of airborne measurement, ground-
422 based measurements, and laboratory experiments, -with a focus on cloud studies. The challenges
423 of operating an aircraft and the inherent sampling limitation of high-speed nature and thus low
424 spatial resolution of most measurement platforms are highlighted, and two developments are
425 introduced to address these challenges: the emergence of un-crewed aerial vehicles (drones) for
426 scientific purposes and the HOLODEC instrument based on digital in-line holography. The
427 HOLODEC provides an estimate of the cloud particle size distribution and particle shape from a
428 single sample volume of cm scales, providing unique opportunities to measure and study
429 outstanding science questions like droplet clustering, particle breakup, and particle shattering. It
430 is pointed out that laboratory measurements have become somewhat less common and lagged
431 field measurement capabilities, although their contributions have been profound.

432 The ultimate test of any models and thus parameterizations is its performance in climate
433 simulations and weather prediction. Efficient and effective evaluation frameworks are needed to
434 test the parameterizations, assess their predictive skills, characterize the model behavior from
435 process level to global scale, and identify sources of potential errors to confidently guide further
436 development. In *Chapter 17, Lin and Xie* describe the approaches and frameworks used for
437 testing and evaluating fast physics parameterizations in climate and weather models. An
438 integrated yet complementary modeling and evaluation framework is advocated that promotes
439 process-oriented evaluation and effectively bridging parameterization development with
440 observations and modeling, with focus on two exemplary such frameworks that have been
441 widely adopted by the modeling centers and the research community. The first is the integrated
442 SCM-CRM-LES framework that has been widely used since it was promoted in the early 1990s.
443 This modeling framework allows for process studies with SCMs, CRMs, and LES models with
444 the scale ranging from a few hundred kilometers to a few tens of meters. With all the models
445 driven by the same large-scale forcing and initial and boundary conditions, the model
446 intercomparison studies this frame have proven useful to identify strengths and weaknesses of
447 model parameterizations; and systematic model deficiencies have been found through these
448 studies. The second framework is based on the idea of running climate models in “weather
449 forecast mode with initial data from NWP analyses to take advantage of the facts that (1) the
450 large-scale state of the atmosphere in the early periods of a forecast is realistic enough that errors
451 may be ascribed to the parameterizations of the atmospheric physical processes; and (2) the
452 atmospheric physical processes (e.g., moist process) are often fast (~hours) and the large-scale
453 state changes slowly (~ days); (3) there is a strong correspondence between the short- and long-
454 term systematic errors in climate models, particularly for those fields that are related to fast
455 physics (e.g., clouds). Further integration of the two evaluation frameworks to better capitalize
456 on their respective advantages are also explored. The metrics and diagnostics designed for model
457 evaluation are also presented, including process-oriented diagnostics and metrics in support of
458 process studies and providing more insights into model errors, and satellite/radar simulator

459 packages that permit direct comparison of model outputs to sensor signals without complicated
460 retrievals.

461 **1.4. How to approach the content in this book**

462

463 The book is targeted at researchers and graduate students working on the relevant areas. Each
464 chapter of this collective volume has its own foci that are closely related to other chapters, and
465 can be read either separately as a stand-alone chapter with its own list of references or together
466 with the other chapters with cross references as needed.

467

468 This book serves a valuable addition to existing literature on fast physics parameterizations in
469 large scale atmospheric models, with several unique features. It would be better read together
470 with: Stensrud (2007), the special fast physics collection in *Journal of Geophysical*
471 *Research:Atmospheres* (Liu, 2019;
472 [https://agupubs.onlinelibrary.wiley.com/doi/toc/10.1002/\(ISSN\)2169-8996.FASTPHYS1](https://agupubs.onlinelibrary.wiley.com/doi/toc/10.1002/(ISSN)2169-8996.FASTPHYS1)), and
473 various topical review articles (e.g., Morrison et al. (2020) on parameterizations of cloud
474 microphysical processes).

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570

571 **Figure caption**

572

573 **Figure 1.** Schematic of the approximate timelines of development of fast physics
574 parameterizations. Conventional parameterizations are focused on individual fast processes. The
575 four lines of unifying efforts (PDF-based High Order Closure, Eddy Diffusivity and Mass Flux,
576 Super/ultra Parameterization, and Machine Learning Parameterization) aim to unify the
577 representation of more than two physical processes. Top-down approaches borrow holistic ideas
578 that have been scattered in various different disciplines (e.g., nonlinear systems dynamics,
579 statistical physics, information theory, self-organization, networks, and pattern formation).

580 **Figure 2.** Historical values of equilibrium climate sensitivity (ECS) and transient climate
581 response(TCR). Figure is modified from Fig. 1 of Meahle et al. (2020) which can be consulted for
582 details on the data sources and definitions.

583 **Figure 3.** Schematic to illustrate the atmospheric scale hierarchy and involved “4M-2N
584 complexities”. Together with the “operational complexity” discussed in the text, these science
585 complexities have posed and will continue to pose challenges to model development in general
586 and fast physics parameterizations in particular.

587