Coastal retracking using along-track echograms and its dependency on coastal topography

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Abstract

Although the Brown mathematical model is the standard model for waveform retracking over open oceans, coastal waveforms usually deviate from open ocean waveform shapes due to inhomogeneous surface reflections within altimeter footprints, and thus cannot be directly interpreted by the Brown model. Generally, the two primary sources of heterogeneous surface reflections are land surfaces and bright targets such as calm surface water. The former reduces echo power, while the latter often produces particularly strong echoes. In previous studies, sub-waveform retrackers, which use waveform samples collected from around leading edges in order to avoid trailing edge noise, have been recommended for coastal waveform retracking. In the present study, the peaky-type noise caused by fixed-point bright targets is explicitly detected and masked using the parabolic signature in the sequential along-track waveforms (or, azimuth-range echograms). Moreover, the power deficit of waveform trailing edges caused by weak land reflections is compensated for by estimating the ratio of sea surface area within each annular footprint in order to produce pseudo-homogeneous reflected waveforms suitable for the Brown model. Using this method, Jason-2 altimeter waveforms are retracked in several coastal areas. Our results show that both the correlation coefficient and root mean square difference between the derived sea surface height anomalies and tide gauge records retain similar values at the open ocean (0.9 and 20 cm) level, even in areas approaching 3 km from coastlines, which is considerably improved from the 10 km correlation coefficient limit of the conventional MLE4 retracker and the 7 km sub-waveform ALES retracker limit. These values, however, depend on the coastal topography of the study areas because the approach distance limit increases (decreases) in areas with complicated (straight) coastlines.

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I. Introduction

- Sequence weaker reflections by lands







2. Data

- The 20 Hz along-track waveform of Jason-2 in Sendor Geophysical Data Record (S-GDR) are used for Cycle 1-252 (2008.7-2015.4; AVISO).
- For comparison, Adaptive Leading Edge Subwaveform (ALES) Jason-2 coastal altimetry product (Passaro et al., 2014) are used (http://www.coastalt.eu/community), in addition to the traditional S-GDR retracking algorithm for open oceans.
- Tracks #36 (across Tsushima Island, Japan) and #164 (across Taiwan) are selected for locations with different topography (Fig. 3). The area south of Tsushima Island (red rectangle in Fig.3) is first discussed in Section 4.1, comparing with tide gauge record at *Izuhara*. Then the northern area (blue rectangle) and the south of Taiwan (yellow rectangle) are compared in Sections 4.2 and 4.3, respectively.

Fig.3 Jason-2 descending tracks #36 and #164 (middele panel; *Google*). The track #36 passes the Tsushima Island located in the Tsushima Strait between Japan and Korea (right panel), and the track #164 passes Taiwan (left panel). Circles with the 10-km radius are plotted for reference.

3. Retracking Algorithms 3.1 Removal of bright parabolas

Since the geometry between a bright reflector on the sea surface and a satellite altimeter (Fig. 4) results in a fixed-shape parabola in the along-track echogram, the parabolic shapes are fitted recursively for all bright echos in along-track echograms. By removing these parabolas, waveforms without contaminations can be obtained, to which the theoretical Brown waveform can be fitted (Fig. 5).

3.2 Compensation for land echo losses

Assuming that the reflected ALT signals from lands are neglegible w.r.t. these from the sea, we enhance the echo intensity by the ratio of the land area in a co-range circle (Fig. 6). This procedure significantly changes the tailing edge, so that estimations of the wind speed (or sigma-0) and Significant Wave Height (SWH) are influenced. Meanwhile, its influence on the SSH is less than 10 cm in the root-mean-squared difference (RMSD).

4. Area Dependency 4.1 South of Tsushima Island

The SSHAs from the 252-cycle mean are calculated for traditional SGDR and ALES products and for the present method. All extreme outliers (larger than 100m or smaller than -130m) are removed.

- **<u>km</u>** away for this study (Fig 7).
- Time series of the SSHA at 34.21N (Fig. 8), *i.e.* 5 km away from the shoreline, shows that ALES retrieves correct SSHA, except several intermittent extreme cycles. Presence of these extreme cycles can be which is used as a simple performance index in the next subsections.

distance d0.

Fig.5 The same along-track echogram in Fig.2 but after removal of the bright parabolas (left). In this case, five parabolas were removed. The waveform at 34.2N (right) contains no "pseudo leading edges" seen in Fig.1.

Fig.6 Witin the given co-range circle, the length ratio of the arcs over lands (red curve) is determined from the high-resolution coastline data (Wessel and Smith, 1996; left panel). The echo intensity is compensated based on this ratio (right panel). The slope of the trailing edge (*i.e.*, antenna mispointing angle) becomes small.

The correlation coefficient (CC) and RMSD with the tide gauge record at *Izuhara* indicate that the SSHA is available at a point **10 km** away from the shoreline for the traditional SGDR, 7 km away for ALES and only 3

identified by the RMS (or, the standard deviation) of the SSHA at a point, \bigcirc 30

Fig.8 Time series of the SSHA at 34.21N for tide gauge (green), ALES (blue) and this study (red).

Wessel and Smith (1996) J.G.R., **101**(B4), 8741-8743.