

The Great Subterranean Spring of Minneapolis, Minnesota, USA, and the potential impact of subsurface urban heat islands (SUHIs)

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Abstract

Anthropogenic subsurface urban heat islands (SUHIs) in groundwater under cities are known worldwide. SUHIs are potentially threats to springs because much spring fauna, like trout, amphipods, and rare plants, is cold stenothermal. The city of Minneapolis, Minnesota, USA, has a SUHI documented by the temperature of an underground spring, dubbed “Little Minnehaha Falls,” inside Schieks Cave, which is located 23 m below the central core of the city. In 2000 the temperature of that spring was elevated 11°C above regional background groundwater temperatures (8°C) at this latitude (45°N). A thermometric survey of the cave and nearby tunnel seepages in 2007 found that an abandoned drill-hole through the bedrock ceiling of the cave was discharging groundwater with a temperature of 17.9°C. By comparison, groundwater in the deep water-table below the cave was closer to natural background temperatures for the region. The unusually warm groundwater was thereby localized to the strata above the cave. This is the strongest signal of anthropogenic groundwater warming in the state of Minnesota and is attributed to vertical heat conduction from basements and pavements. Minneapolis is unique among SUHIs in that a cave forms a natural collection gallery deep below the city surface, whereas the literature is almost exclusively based on data from observation wells.

1 **The Great Subterranean Spring of Minneapolis, Minnesota, USA,**
2 **and the potential impact of subsurface urban heat islands (SUHIs)**

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5

6 **ABSTRACT**

7

8 Anthropogenic subsurface urban heat islands (SUHIs) in groundwater under cities are
9 known worldwide. SUHIs are potentially threats to springs because much spring fauna,
10 like trout, amphipods, and rare plants, is cold stenothermal. The city of Minneapolis,
11 Minnesota, USA, has a SUHI documented by the temperature of an underground spring,
12 dubbed “Little Minnehaha Falls,” inside Schieks Cave, which is located 23 m below the
13 central core of the city. In 2000 the temperature of that spring was elevated 11°C above
14 regional background groundwater temperatures (8°C) at this latitude (45°N). A
15 thermometric survey of the cave and nearby tunnel seepages in 2007 found that an
16 abandoned drill-hole through the bedrock ceiling of the cave was discharging
17 groundwater with a temperature of 17.9°C. By comparison, groundwater in the deep
18 water-table below the cave was closer to natural background temperatures for the region.
19 The unusually warm groundwater was thereby localized to the strata above the cave. This
20 is the strongest signal of anthropogenic groundwater warming in the state of Minnesota
21 and is attributed to vertical heat conduction from basements and pavements. Minneapolis
22 is unique among SUHIs in that a cave forms a natural collection gallery deep below the

23 city surface, whereas the literature is almost exclusively based on data from observation
24 wells.

25

26 **INTRODUCTION**

27

28 Elevated groundwater temperatures are a potentially important threat to springs because
29 much spring biota is cold stenothermal, examples being trout, amphipods, and rare plants
30 (Brick, 2017b). According to Taniguchi et al. (2007), “The heat island effect due to
31 urbanization on subsurface temperature is an important global groundwater quality issue
32 because it may alter groundwater systems geochemically and microbiologically.” In the
33 “Twin Cities” of Minneapolis-St. Paul, Minnesota, USA (**Fig. 1a**) there are designated
34 trout streams within the metropolitan area, whose springsheds are heavily built over and
35 will be impacted by rising groundwater temperatures (Meersman, 2012).

36

37 According to Taylor and Stefan (2008), “Urban development can influence groundwater
38 temperatures in a number of ways: Paved surfaces become much warmer than sod
39 surfaces on clear, sunny days. Heat is conducted from these surfaces into the soil, and can
40 reach shallow groundwater. Taniguchi and Uemura (2005) provide evidence of
41 conduction-based warming of groundwater due to urbanization. Surface water runoff
42 from warm ground surfaces can infiltrate from ponds, channels, or rain gardens.
43 Percolating warm water may carry its heat, not lost in the soil, into an aquifer.”

44

45 Elevated groundwater temperatures below cities are well known worldwide. In Asia,
46 Taniguchi et al. (2009) have documented anthropogenic thermal effects on groundwater
47 in Osaka, Japan, and Bangkok, Thailand. In Europe, Epting et al. (2017) confirm the
48 same trends at Basel, Switzerland. According to Hemmerle et al. (2019) “This is
49 demonstrated for the city of Paris, where measurements from as early as 1977 reveal the
50 existence of a substantial subsurface urban heat island (SUHI) with a maximum
51 groundwater temperature anomaly of around 7 K.” In North America, Ferguson and
52 Woodbury (2004, 2007) describe the subsurface heat island effect below Winnipeg,
53 Manitoba: “Downward heat flow to depths as great as 130 m has been noted in some
54 areas beneath the city and groundwater temperatures in a regional aquifer have risen by
55 as much as 5°C in some areas.” Yalcin and Yetemen (2009), Zhu et al. (2010), Menberg
56 et al. (2013b), and Bayer et al. (2019), among others, go so far as to consider tapping into
57 SUHIs as a source of geothermal energy. In keeping with conventions in the SUHI
58 literature, from this point on C (Celsius) will be used to indicate a measured temperature,
59 whereas K (Kelvin) indicates a temperature difference.

60

61 The earliest known record of groundwater temperature in the Platteville Limestone in
62 what became Minneapolis is from Nicollet (1845). At Coldwater Spring (a Platteville
63 spring about 10 km from Schieks Cave and its underground spring) Nicollet reported an
64 average of 7.8°C in July 1836 and 7.5°C in January 1837 (Nicollet, 1845: 69).

65

66 Taylor and Stefan (2008, 2009) projected a rise in groundwater temperature for the Twin
67 Cities metropolitan area of 3 K (and even higher in global warming scenarios) in research

68 funded by the Minnesota Pollution Control Agency. However, they were most likely
69 unaware that the first measurement indicating elevated groundwater temperature in
70 Minneapolis had already been reported by an undergraduate geology student years earlier
71 from Chalybeate Springs, a mineral water resort that was popular before the American
72 Civil War. These springs emanate from the Platteville Limestone with a temperature of
73 14°C, 6 K above the expected 8°C at latitude of 45°N (Brick, 1993).

74

75 A time series of temperature measurements at Coldwater Spring, on the outskirts of
76 Minneapolis, lasting nearly two years (2013-2015), recorded fluctuations from 10.7°C to
77 13.1°C, well above those of Nicollet (1845) already mentioned. Kasahara (2016)
78 attributed this to “an anthropogenic source of heat within the spring-shed or spring
79 discharge area.” See also Alexander and Brick (2021: 98).

80

81 However, while these elevated temperatures are notable, they are half the temperature
82 changes in a spring and a free-flowing well located inside a cave under downtown
83 Minneapolis at a depth of 23 m below street level, which is the focus of the remainder of
84 this chapter.

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91 **BACKGROUND**

92

93 **Cave and Spring**

94

95 Schieks Cave is the largest cave under downtown Minneapolis, underlying half a city

96 block (**Fig. 1b**). This maze cave in the St. Peter Sandstone has a ceiling of Platteville

97 Limestone. Its exact origin is rather murky (Brick, 2017a).

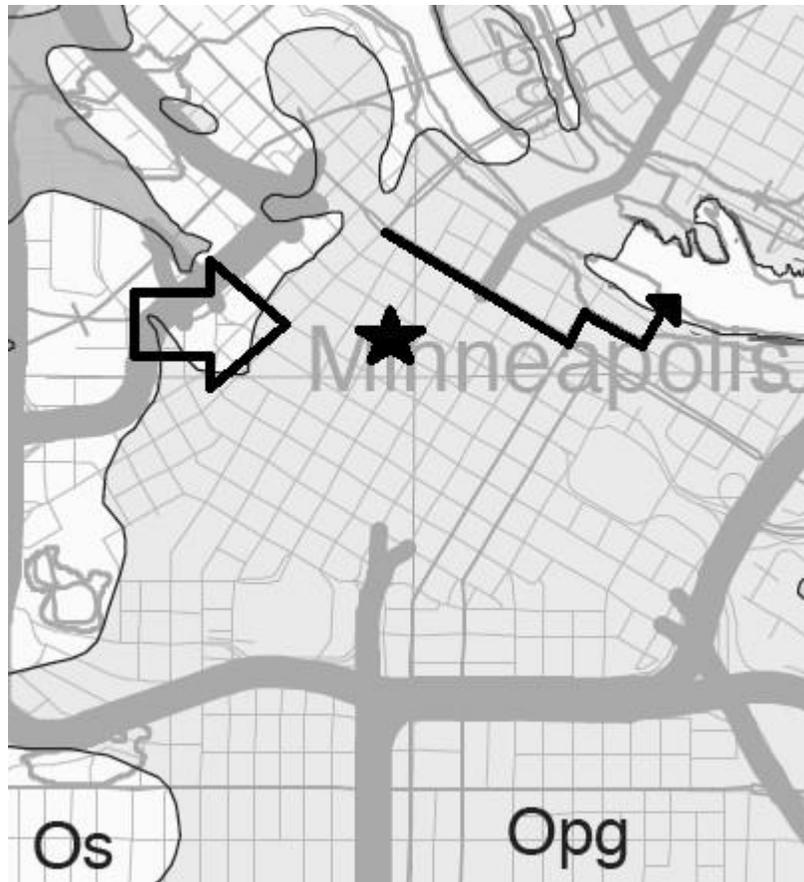


98

99 **Fig. 1a. Location of the Minneapolis-St. Paul metropolitan area (star). Modified**

100 **from Wikimedia/U.S. Geological Survey.**

101



102

103

104 **Fig. 1b. Location of Schieks Cave (star) on Hennepin County bedrock map, north at**
105 **top (Retzler, 2018). Opg=Ordovician Platteville-Glenwood Formation,**
106 **Os=Ordovician St. Peter Sandstone. Big arrow indicates general direction of**
107 **groundwater flow in the shallow Quaternary aquifer based on Kanivetsky (1989).**
108 **Thin arrow represents the Washington Avenue tunnel, where temperature**
109 **measurements were taken in 2007, and its flow direction to Mississippi River (see**
110 **Table 1).**

111

112 The first document regarding Schieks Cave (or its spring) is the 1904 Nic. Lund map.

113 Although rather crude and incomplete, the map is rich in hydrologic details such as

114 “creeks” and “lakes.” At the location of the ceiling spring the map notes “WIDE CRACK
115 IN LEDGE, LARGE BODY OF WATER COMING THROUGH.” Upon first entering
116 the cave, Carl J. Illstrup, city sewer engineer, reported that “Dripping from the ceiling at
117 one place there was a regular curtain of water 30 feet [10 m] in width. The water in the
118 middle [of a pool at the bottom of the curtain] was 20 feet [6 m] deep at one point and
119 tapered down to inches at the shore line. It was a beautiful sight but we had to drain it to
120 remedy the troubles in the Fourth Street tunnel” (Fitzsimmons, 1931).

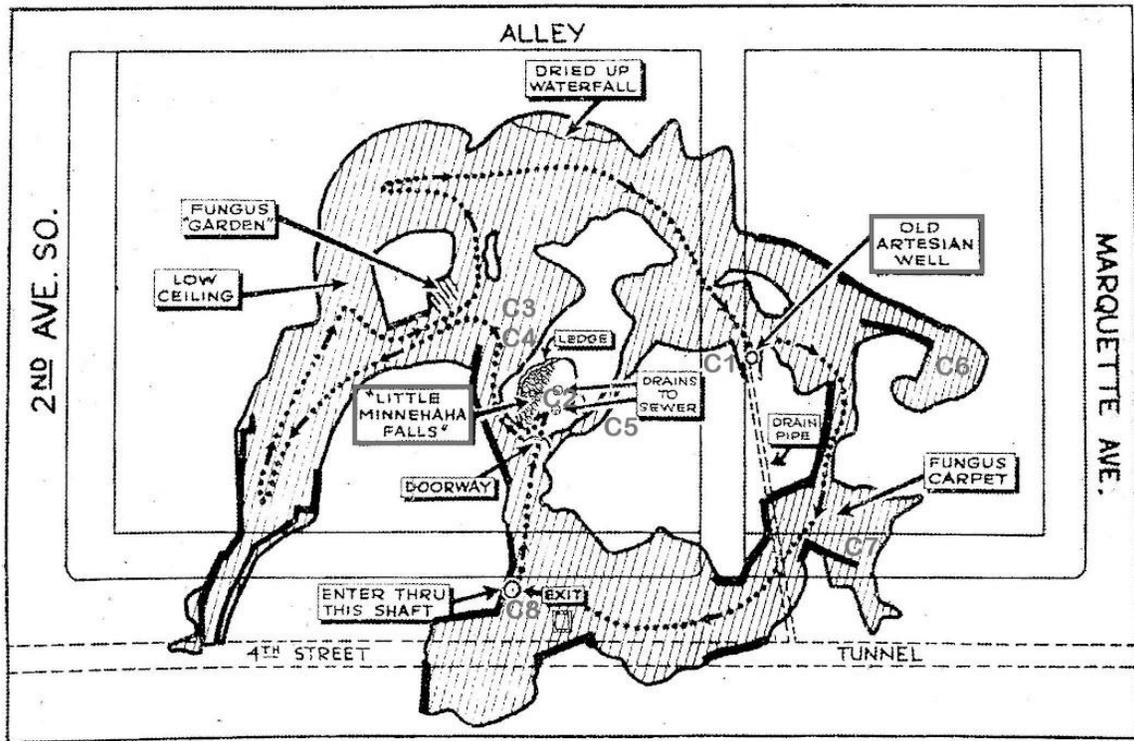
121

122 A second map, by J.E. Lawton in 1929, based on a detailed survey, depicts Schieks Cave
123 extensively modified by the construction of piers, walls, and artificial drainage systems,
124 the latter to prevent further erosion of the soft sandstone. The ceiling spring was shown as
125 “WATER FALLS,” but this time enclosed in a separate concrete chamber provided with
126 floor drains. Both cave maps are reproduced as Plate 3 in Kress and Alexander (1980).

127

128 Longnecker (1907) described the spring and well in print for the first time and Brick
129 (2021) gives an extended analysis of his narrative from a geological perspective.
130 Longnecker included a photo captioned “The Subterranean Falls,” showing how “the
131 crystal-pure spring water gushes forth between the ledges of limestone and falls into a
132 concrete basin built by the city engineer’s staff.” Dornberg (1939) was the first to refer to
133 the ceiling spring as “Little Minnehaha Falls” (LMF) and label it as such on his map (**Fig.**
134 **2**). This was a jocular reference to Minnehaha Falls, a well-known Minneapolis
135 landmark. Zalusky (1953a) described LMF as “a falls which I estimated in the darkness
136 to be about 10 feet [3 m] wide and a drop of 5 feet [1.5 m].” So Illstrup’s 10-m curtain of

137 water had apparently dwindled to a third of its former length. Kress and Alexander
138 (1980) state that “In view of the almost complete cover of the surface by buildings or
139 pavement and the inevitable disruption of the near-surface groundwater flow by the
140 excavation of building foundations, it is not surprising that ‘Little Minnehaha Falls’ is
141 drying up.”
142



143
144
145 **Fig. 2. Dornberg’s 1939 map of Schieks Cave, a maze cave in the St. Peter**
146 **Sandstone, showing the ceiling spring, dubbed “Little Minnehaha Falls” and the**
147 **“Old Artesian Well,” a drill-hole through the Platteville Limestone forming the**
148 **ceiling of the cave. C1 to C8 are temperature measurement points (see Table 1).**

149
150

151 **Discovery of Thermal Anomaly**

152

153 No substantive new information about the spring was forthcoming until my own visit of
154 28 May 2000. An extended account of the trip is found in Brick (2009: 191-203). Schieks
155 Cave is normally accessed by a 23-m shaft from Fourth Street and traffic must be
156 diverted to open the manhole, which is why the Minneapolis Sewer Department rarely
157 visits the cave, so it's difficult to acquire additional data.

158

159 Spring water pours from a bedding plane in the Platteville Limestone, depositing
160 vertically striped flowstone on the walls of the concrete chamber built to contain the
161 spring (**Fig. 3a**). Judging from Dornberg's 1939 photos, the discharge did not appear to
162 have diminished much since then, contrary to other reports. The water fell as an extended
163 sheet about 3 m long (matching Zalusky's dimensions of half a century earlier) and the
164 flow rate was visually estimated at 5 gallons per minute (GPM) (=19 liters per minute).
165 Floor drains convey the water to the North Minneapolis Tunnel, a deep-level sanitary
166 sewer.



167

168 **Fig. 3a. A concrete chamber inside Schieks Cave hosts the anthropogenically**
169 **warmed ceiling spring “Little Minnehaha Falls,” which issues from a bedding plane**
170 **in the Platteville Limestone, depositing “zebra” flowstone on the walls. Photo by**
171 **author, 2000.**

172



173

174 **Fig. 3b. Spring issuing from fissure in the Washington Avenue tunnel. Seepage**
175 **entrains whitish St. Peter sand grains from outside the tunnel lining. Photo by**
176 **author.**

177

178 Upon equilibrating a calibrated SPER Scientific® mercury thermometer ($\pm 1\%$) in the
179 spring orifice for several minutes, I was surprised to note that the groundwater
180 temperature was 19°C , higher than the expected 8°C at this latitude (Brick, 2014).

181

182 I also examined a drill-hole in the limestone ceiling of the cave, which discharged an
183 estimated 50 GPM (=190 liters per minute), falling into a shallow concrete basin (**Fig. 4**).

184 However, I mistakenly thought this was stormwater at the time, only later recognizing it
185 as a water well, so I did not measure its temperature during the first trip. There's no

186 manifestation of this abandoned well at street level, the site being entirely covered with
187 buildings and a car park. Dornberg (1939) refers to it as the “Old Artesian Well” (OAW).
188

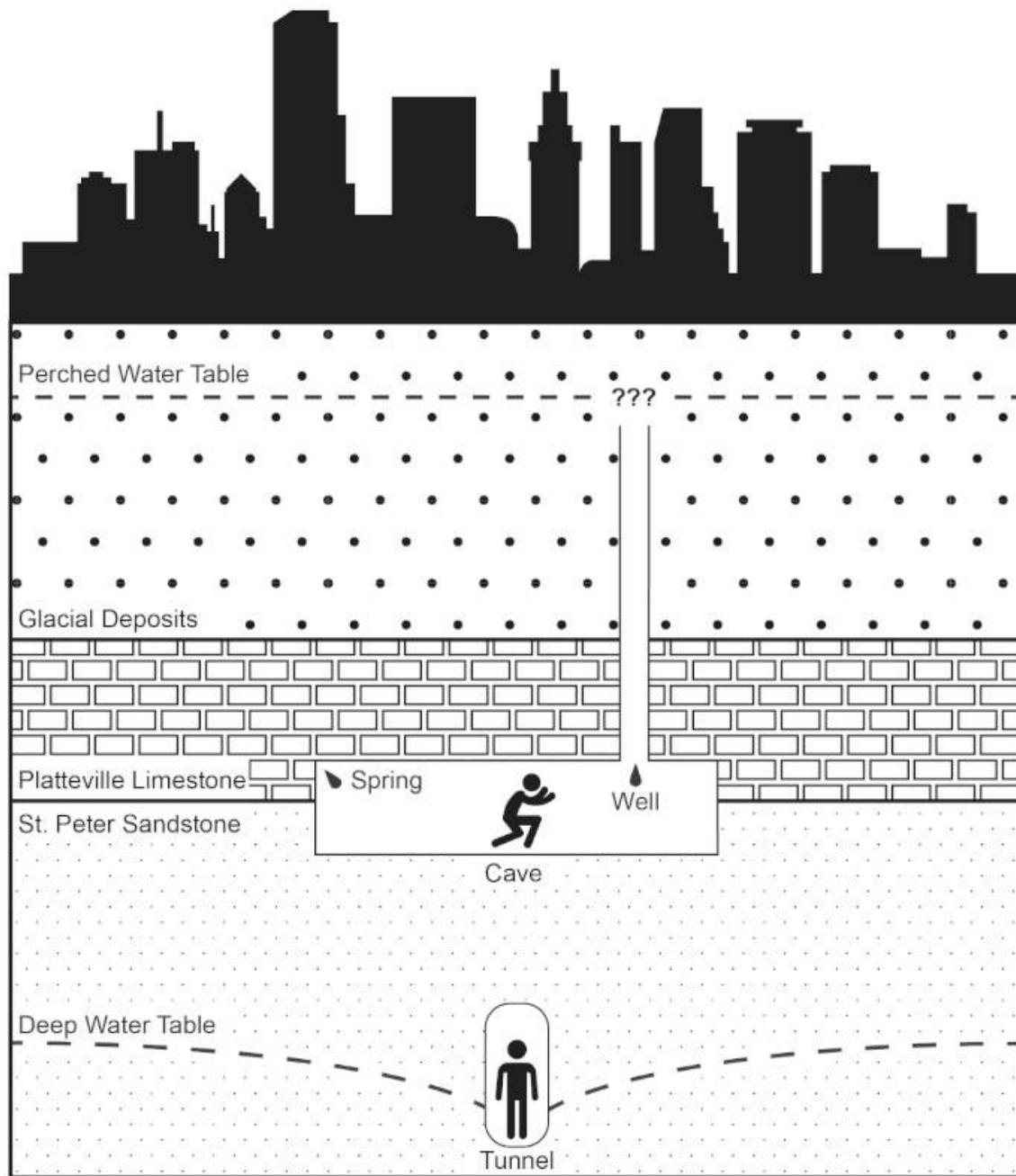


189
190
191 **Fig. 4. The Old Artesian Well flowing in the “wrong” direction, showing bottom of**
192 **drill-hole in the Platteville ceiling of Schieks Cave. Photo by author.**

193
194 **Geology**

195
196 The bedrock layers are Ordovician units making up the Twin Cities basin, overlain by
197 unconsolidated Quaternary glacial sediments (Retzler, 2018). A diagrammatic cross-

198 section of the geology at Schieks Cave is shown as **Fig. 5**. Relative thickness of the layers
199 above the cave are depicted in accordance with Zalusky (1953b), who apparently had
200 access to the driller's fieldnotes from the construction of the Fourth Street entrance shaft.
201



202

203 **Fig. 5. Geological cross-section of Minneapolis at Schieks Cave (not to scale). LMF**
204 **(spring) and OAW (well) are shown inside cave. A one-meter layer of Glenwood**
205 **Shale between the Platteville and St. Peter has been omitted for clarity. Tunnel is 30**
206 **m below ground surface. Artwork by Jessica Rogge.**

207

208 The Platteville Limestone is karstified, acting as aquifer or aquitard depending on the
209 setting (Steenberg et al., 2011). The largest and most abundant springs in Minneapolis
210 emanate from this layer (Brick, 1997). While a meter thick layer of Glenwood Shale,
211 intercalated between the Platteville and St. Peter, has been omitted for clarity, this layer
212 acts as an aquitard within the Twin Cities basin. Inside Schieks Cave, however, LMF
213 issues from a bedding plane separating the two lowermost members of the Platteville
214 Limestone: the thin Pecos (which usually falls away in open voids) and the thicker
215 Mifflin (which remains to form flat ceilings) members. The shale does not appear to play
216 a role here.

217

218 Two water tables are depicted in Fig. 5: a perched water-table in the shallow Quaternary
219 aquifer and Ordovician limestone above the cave, and a deep water-table in the St. Peter
220 Sandstone below the cave. The uppermost sandstone has a separate unsaturated zone.
221 Groundwater infiltrates the tunnels through breaks in the concrete lining.

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223

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225

226 **METHODS**

227

228 Disclaimer: This project reports ad hoc data sets from a site that is difficult to access. The
229 data sets were not obtained in a formally designed academic research project.

230

231 A search for existing well water temperature data was conducted. Two drilled wells
232 penetrate Schieks Cave, one of which is the “Old Artesian Well” described above. The
233 second is the M.L. & T. Co well, with an intact steel casing that passes entirely through
234 the cave to an unknown depth below, about which no further information exists and
235 which will not be further considered here. These wells are not listed in the online
236 Minnesota Well Index (MWI) database, maintained by the Minnesota Department of
237 Health, apparently because they predate the index, which began in 1974. A query of
238 MWI for all wells within a 1.6 km radius of Schieks Cave found no temperature data for
239 wells terminating in the glacial drift, Platteville Limestone, or St. Peter Sandstone (A.J.
240 Retzler, pers. comm., 2021).

241

242 On 15 April 2007, I made another trip to Schieks Cave during which I conducted an
243 extended thermometric survey of groundwater in the cave and underlying tunnel system,
244 using the same thermometer as before.

245

246 LMF was measured in the same spot as before and the groundwater showering from
247 OAW was measured in a concrete basin built into the floor of the cave. The water in this
248 shallow basin undergoes rapid turnover. The same day, seepages in the Washington

249 Avenue tunnel, as well as flowing non-wet weather stormwater where tributary tunnels
250 entered the main tunnel (for comparison purposes) were measured. The tunnel forms a
251 1.4-km transect from northwest to southeast at an average depth of 30 m (**Fig. 1b**). The
252 deep-water table in the St. Peter Sandstone was accessed where it seeped through breaks
253 in the concrete tunnel lining, usually leaving reddish staining on the walls (**Fig. 3b**).

254

255 **RESULTS**

256

257 The water temperatures from the 2000 and 2007 visits are listed in Table 1 along with
258 descriptions of sampling points. The data include temperature measurements from: 1)
259 Platteville Limestone groundwater in Schieks Cave (C1 – C8), which ranged from 16 to
260 17.9°C; 2) St. Peter Sandstone seepage in the Washington Avenue tunnel (T1 - T3), from
261 12.5 to 14°C; and 3) stormwater flow in the same tunnel (T4 – T10), from 13 to 18°C.
262 Employing the terminology of Benz et al. (2017) the anthropogenic heat intensity (AHI)
263 was 9.9 K for this SUHI at the time of the 2007 survey. The AHI reports the difference
264 between the local value and the average rural groundwater background temperature.

265

266 The Schieks Cave features and their temperature sampling points are shown in **Fig. 2**.
267 LMF was 19°C in 2000 and 17.1°C in 2007, while groundwater captured in the OAW
268 concrete basin was 17.9°C. The cave air temperature was 18°C, measured inside the
269 concrete chamber containing LMF. The average surface air temperature for the month of
270 April, 2007, recorded at the Minneapolis-St. Paul International Airport, was 8.44°C
271 (NOAA, 2021).

272

273 Much of the non-wet weather baseflow of the Washington Avenue tunnel is derived from
274 groundwater infiltration and each tributary tunnel (T4 – T8) contributes its own separate
275 temperature. The farthest downstream measuring point, the Chicago Avenue Outfall
276 (T10) was 13.5°C, which presumably represents the mean temperature of all contributing
277 flows in the storm drains mixed together. This temperature falls within the St. Peter
278 seepage range, suggesting its ultimate origin as infiltration water.

279

280 **DISCUSSION**

281

282 The flows from LMF and OAW are documented inside Schieks Cave from the earliest
283 map (1904) until today, more than a century, indicating that the water does not originate
284 from leaking pipes or water mains (cf. Lerner, 1986).

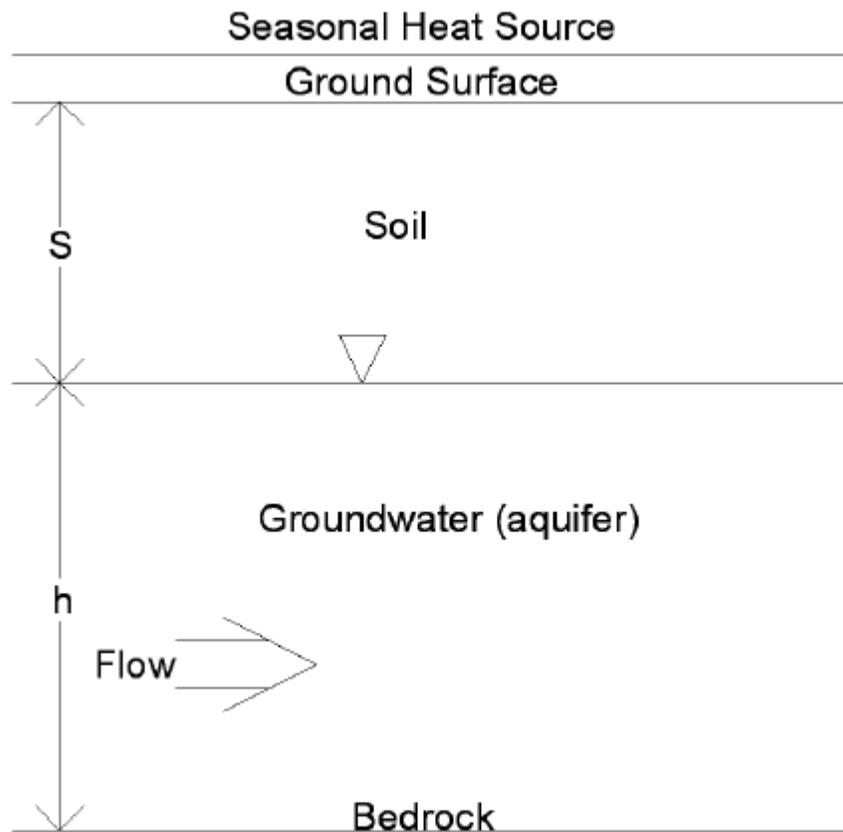
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286 Other possible sources of heat to groundwater can be excluded. Tissen et al. (2019)
287 considered potential natural causes of elevated temperatures for SUHIs, especially hot
288 springs, which are not present in Minnesota. Tissen et al. (2019) also considered Acid
289 Mine Drainage (AMD). While the Platteville Limestone contains pyrites, subject to
290 oxidation—an exothermic reaction—monitoring of Platteville springs has never detected
291 elevated pH among them (Minnesota Department of Natural Resources, 2021).

292

293 Groundwater entering Schieks Cave was elevated above background by 11 K in 2000,
294 larger than the 3 K predicted by Taylor and Stefan (2008, 2009) for the Minneapolis

295 latitude. As a conceptual model, basements and pavements are warming the shallow
 296 Quaternary aquifer by vertical conduction (**Fig. 6**) and this water finds its way via OAW
 297 and fissures down through the Platteville Limestone, into the cave.
 298



299
 300
 301 **Fig. 6. “Schematic of aquifer of thickness h at depth S below the ground surface**
 302 **(from Taylor & Stefan, 2008, p. 21). At Schieks Cave, $S=10.7$ m and $h=5.5$ m. The**
 303 **soil compartment represents the unsaturated zone.**

304
 305 LMF was 19°C in May 2000 and 17.1°C in April 2007, a 1.9 K decrease during this
 306 seven-year interval. According to Taylor and Stefan (2008), seasonal temperature

307 fluctuations penetrate the ground to depths of 10 to 15 m, within the depth range of the
308 shallow Quaternary aquifer at this location. Continuous, multi-year temperature
309 monitoring would be necessary to determine the magnitude of fluctuations at LMF.
310 Kasahara (2016) documented a systematic, seasonal fluctuation in the water temperature
311 in Coldwater Spring of 2.3 K (10.8 to 13.1°C) which lagged the surface air temperature
312 by about four months. The change in temperature at LMF is within the range of
313 documented seasonal fluctuation in the Platteville aquifer in Minneapolis.

314

315 Located within 23 m of the surface, LMF and OAW reveal elevated groundwater
316 temperatures that are also within the range of depth for thermal anomalies measured by
317 Taniguchi et al. (2007) in boreholes in four Asian cities: “The depth of deviation from the
318 regional geothermal gradient was deepest in Tokyo (140 m), followed by Osaka (80 m),
319 Seoul (50 m), and Bangkok (50 m).”

320

321 LMF and OAW, located 27 m apart inside Schieks Cave, had temperatures within 1 K of
322 each other during the only event in which they were simultaneously measured, which
323 would be expected if they are derived from the same aquifer.

324

325 At OAW, the temperature went from feeling “icy” (Longnecker, 1907), and thus
326 presumably normal groundwater temperature of 8°C, to 17.9°C, in the 100 years from
327 1907 to 2007. While the first observation is qualitative and the second is quantitative, the
328 century-long trend is unambiguous. This is strongest signal of anthropogenic
329 groundwater warming in the state of Minnesota.

330

331 Where seepage waters were encountered in the St. Peter Sandstone below the Platteville
332 Limestone, they were several degrees cooler, yet still about 5 K above background
333 groundwater temperatures. The tunnel seepages (T1 – T3) cannot be easily revisited
334 because the Central City Tunnel System project (2020-2023) plans to eliminate
335 infiltration and the proposed construction of relief sewers will have likely disturbed the
336 hydrology of the setting (CDM Smith Inc., 2018).

337

338 The warm groundwater “pool” of the SUHI thus appears to be perched in the strata above
339 Schieks Cave, most likely (given vertical heat conduction considerations) in the shallow
340 Quaternary aquifer but also the Platteville Limestone.

341

342 Attard et al. (2016) defined a thermally affected zone (TAZ) around urban structures as
343 where the groundwater temperature is elevated 0.5 K above expected values. Menberg et
344 al. (2013a) found hotspots below German cities elevated as much as 20 K. Menberg et al.
345 (2013c) concluded that: “By modeling the anthropogenic heat flux into the subsurface of
346 the city of Karlsruhe, Germany, in 1977 and 2011, we evaluate long-term trends in the
347 heat flux processes. It revealed that elevated GST [ground surface temperature] and heat
348 loss from basements are dominant factors in the heat anomalies.”

349

350 In the case of Schieks Cave, the groundwater direction in the shallow Quaternary aquifer
351 indicates flow from the west, through the heavily commercialized Nicollet Mall and the
352 densest skyscraper cluster, towards the cave (Kanivetsky, 1989). The source of the heat is

353 likely due to heated buildings as was shown by Krcmar et al. (2020), who reported
354 groundwater temperatures had risen 3.2 K by flowing past a building in Bratislava,
355 Slovakia.

356

357 Minneapolis is rare among SUHIs in that while the literature is almost exclusively based
358 on data from observation wells, this example involves a cave functioning as a collection
359 gallery for groundwater deep below the city surface. Schieks Cave thus affords a parallel
360 with the historic cellar of the Paris Observatory, which lies at a comparable depth (28 m)
361 and has been used for temperature measurements for centuries (Dettwiller, 1970).

362

363 Scott Alexander (pers. comm., 2021), hydrogeologist at the Earth Sciences Department at
364 the University of Minnesota, stated that several shallow campus wells in the Platteville
365 Limestone (used for teaching purposes) have had a long-term temperature of about 25°C.
366 His opinion is that these elevated temperatures are due to the campus steam tunnel
367 system, which is carved in the underlying St. Peter Sandstone. If true, this campus SUHI
368 would be the inverse of the Schieks thermal anomaly in that the greater heat source is
369 from below, rather than above.

370

371 The “twin” city of Minneapolis is the neighboring city of St. Paul, capital of the state of
372 Minnesota. Unlike its twin, however, St. Paul features a multilevel utility tunnel system
373 carved within the St. Peter Sandstone, on a larger scale than the university campus
374 (Brick, 2009: 179-188). One of the levels contains a district heating system employing
375 hot water. The conjectural St. Paul SUHI would likely involve both the shallow (perched)

376 and deep water-tables. This SUHI could be more intense than that of Minneapolis, even
377 though St. Paul has a smaller population, challenging some fundamental assumptions
378 about urban heat islands based on population modeling (e.g., Oke, 1973).

379

380 **CONCLUSIONS**

381

382 Anthropogenic subsurface urban heat islands (SUHIs) in groundwater under cities are
383 known worldwide. A subsurface urban heat island (SUHI) under the center of the built-up
384 area of Minneapolis, Minnesota, USA, was first detected at a spring inside Schieks Cave
385 located 23 m below street level. SUHIs are potentially threats to springs because much
386 spring fauna, like trout, amphipods, and rare plants, is cold stenothermal.

387

388 Minneapolis is rare among SUHIs in that while the literature is almost exclusively based
389 on data from observation wells, this case involves a cave functioning as a collection
390 gallery for groundwater deep below the city surface.

391

392 A thermometric survey of the cave and tunnel seepages in 2007 revealed a temperature
393 trend extending from 1907 to 2007, showing rising temperature from normal groundwater
394 temperatures for this latitude (8°C) to 17.9°C in one century. This is strongest signal of
395 anthropogenic groundwater warming in the state of Minnesota and is attributed to vertical
396 heat conduction from basements and pavements.

397

398 Future studies of the spring and well in Schieks Cave should involve multilevel sampling,
399 data loggers, lengthy time series spanning several years to detect possible seasonal trends,
400 and MODFLOW visualization of data. Any future trip should take water samples as well
401 as additional temperature readings of this feverish spring.

402

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Table 1. Thermometric Survey of Cave and Tunnels, 15 April 2007*

Origin	Description/location of temperature readings	Temp °C
Platteville Limestone groundwater in Schieks Cave.	C1: Old Artesian Well (OAW) concrete basin (rapid turnover)	17.9
	C2: Little Minnehaha Falls (LMF) ceiling spring at bedding plane*	17.1
	C3: Galvanized drip basin (south)	17.0
	C4: Galvanized drip basin (north)	16.9
	C5: Black Medusa formation, flowing like faucet	16.9
	C6: ML&T Co well hole, from still pool around steel casing	16.5
	C7: Seepage on flowstone in concrete drain	16.5
	C8: Black drip pool at base of ladder, seeping around steel shaft lining	16.0
St. Peter Sandstone seeps in storm drains via gaps in concrete lining.	T1: Red springs near floor, Washington Ave S, half way between 3rd & 4th Ave S	14.0
	T2: Red spring in floor of side passage, Washington Ave S/3rd Ave S	12.5
	T3: Leakage jetting from tunnel lining with red staining, Washington Ave S/Hennepin Ave	13.0

Non-wet weather flows in storm drains measured in tributary just before it joins flow in Washington Avenue tunnel.	T4: Hennepin Ave at Washington Ave S	13.0
	T5: Nicollet Mall at Washington Ave S	15.5
	T6: Marquette Ave at Washington Ave S	14.5
	T7: Side-passage half way between Marquette Ave & 2nd Ave S at Washington Ave S	13.0
	T8: 2nd Ave S at Washington Ave S	18.0
	T9: Iron Gate pool, Portland Ave/Washington Ave S	15.5
	T10: Chicago Ave Outfall at Mississippi River	13.5

***On 28 May 2000, the temperature of LMF was measured at 19°C. This was the only temperature measurement prior to the thermometric survey of 15 April 2007.**