Probabilistic reconstruction (or forecasting) of distal runouts of large magnitude ignimbrite PDC flows sensitive to topography using mass-dependent inversion models.

Willy Aspinall¹, Andrea Bevilacqua², Antonio Costa², Hirohito Inakura³, Sue Mahony¹, Augusto Neri², and R Sparks¹

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Abstract

We describe a new method for the reconstruction (or forecast) of probabilities that distal geographic locations were inundated by a large pyroclastic density current (PDC) in terms of the flow mass and related uncertainties. Using appropriate model input uncertainty distributions, derived from expert judgements using the equal weights combination rule, we can estimate the mass amount needed to reach a marginal locality at any given confidence level and compare this with ambiguous or inexact peripheral field data. Our analysis relies on different versions of the Huppert and Simpson (1980) integral formulation of axisymmetric gravity-driven particle currents. We focus on models which possess analytical solutions, enabling us to utilize a very fast functional approach for enumerating results and uncertainties. In particular, we adapt the 'energy conoid' approach to generate inundation maps along radial directions, based on comparison of the mass-dependent kinetic energy of the flow with the potential energy control by topography in the direction of flow at distal ranges. We focus on two alternative conceptual models: (i) Model 1 assumes the entire amount of solid material originates from a prescribed height above the volcano and flows as a granular current slowed by constant friction; (ii) Model 2 is a multi-phase formulation and includes, in addition to suspended particles, interstitial gas thermally buoyant with respect to surrounding cold air. In the latter case, the flow stops propagating at the surface when the solid fraction becomes less than a critical value, and there is lift-off of the remaining mixture of gas and small particulates. Our model parameters can be further constrained where there is reliable field data or information from analogue eruptions. Finally, we used a Bayes Belief Network related to each inversion model to evaluate probabilistically the uncertainties on the mass required, estimating correlation coefficients between input variables and the calculated mass. For any major magnitude ignimbrite PDC scenario, our method provides a rational basis for assessing the probability of distal flow inundation at critical peripheral locations when there is major uncertainty about the actual or predicted extent of flow runout. Example case histories are illustrated.

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1. A method to model the minimum volume and mass to inundate an at-risk city from a volcano source site

Our analysis relies on the implementation of several versions of the integral formulation for axisymmetric gravity-driven particle currents, based on the pioneering work of Huppert and Simpson (1980). The theory is detailed in Bonnecaze et al., (1995) and Hallworth et al. (1998). We focus on models which possess analytical solutions, enabling us to utilise a very fast functional approach in the uncertainty quantification process. Further details on the physical equations we adopted, as well as the expression of analytical solutions, can be found in Biagioli et al., 2019.

In particular, we focus on two different models:

- Model 1 [Rock avalanche dynamics with constant stress over the flow basal area] This model for energy dissipation described in Dade and Huppert (1998). assumes the entire amount of solid materia falls from a prescribed height. Constant
- stress dynamics has been further explored in Kelfoun et al. (2009); Kelfoun (2011). **INPUT** parameters:

*one third of the constant stress value

- H Collapse Height *rho c -* Flow density tau – equivalent stress'
- Model 2 [Density current dynamics with particle deposition and buoyancy effects This model is described in Dade and Huppert 1995. It has been developed for the simulation of oceanic turbidity currents and then adapted to the simulation of large-scale ignimbrites (Dade and Huppert, 1996).

INPUT parameters: phi0 – initial solid fraction ws – velocity of settling of the solid particles

rho – density of solid particles *rho a* – density of ambient air *rho i* – density of interstitial gas In our implementation, Model 2 assumes monodispersed solid particles, because the modelling of the full Total Grain Size Distribution (TGSD) does not produce analytical solutions. The Sauter diameter is believed to provide a reasonable approximation to the dynamics of the full TGSD. We assume a fixed volume instantaneous release of collapsing material, as in Neri et al. (2015); Bevilacqua (2016); Esposti Ongaro et al. (2016); Bevilacqua et al. (2017), Biagioli et al., (2019)

- Model 2a. This variant includes interstitial gas, thermally buoyant with respect to surrounding cold air. The flow stops propagating when the solid fraction *phi(t)* becomes lower than a critical value *phi_cr*, and the remaining mixture of gas and particles lifts off, possibly generating a phoenix
- Model 2b. The modelling equations in this variant are equivalent to the previous model, but an alternative input of ws is adopted, expressed as a range of values. This range is based on the law of particle terminal velocity (Armienti et al., (1988); Bonadonna and Phillips, (2003); Dioguardi et al., (2017)) at the scale of the Sauter diameter for analogue flows. These flows are: Mt St. Helens (Costa et al., 2016), Campanian Ignimbrite (Costa et al., 2012; Marti et al., 2016), Youngest Toba Tuff (Costa et al., 2014). Moreover, this variant implements an alternative input range of phiO, based on
- Model 2c. This variant assumes an interstitial gas equivalent to ambient air. Thermal buoyancy effects are absent, and the flow stops when phi(t)=0

Model 2 assumes thermal properties remain constant for the duration of the flow. Thermodynamics modelling of cooling effects (e.g. Bursik and Woods, 1996; Fauria et al., 2016) could be further explored in follow up research.

3. Output ranges Minimum PDC Volume and Equivalent Mass

The following results represent the required volume necessary to inundate an at-risk city at 130 km from the source of the flow, i.e. an hypothetical volcano. The value of the Pth percentile is the volume amount that has probability P to inundate the at-risk city, according to the probability distribution of the model inputs.

Bayes Net calculation results, using Volume

Model	Maximum distance	Minimum PDC Volume [km³]						
Model	130 km	1%ile	5%ile	50%ile	mean	95%ile		
D&H98	(1) Elicited inputs	11.5	26.0	263	450	1474		
D&H95 w/hot gas	(2a) Elicited inputs	5.68	10.2	55.9	77.9	226		
	(2b) Modified inputs*	5.35	7.30	16.7	17.4	30.0		
D&H95	(2c) Elicited inputs	3.34	6.80	43.5	64.5	201		

*modified phiO and ws based on MDR modelling and Sauter diameter of analogues.

To obtain a consistent comparison of the two models requires the calculation of equivalent mass, because the model volumes have a different meaning.

Model 1, based on Dade and Huppert (1998), does not include gas. The volume calculated represents the bulk solid material, at a density of about 1000 kg/m³ (deposit).

Model 2, based on Dade and Huppert (1995), is multi-phase and includes solid particles and interstitial gas. As a model option, this gas can be assumed hot and buoyant with respect to ambient air.

The volume calculated only represents the solid phase, at a density of about 2300 kg/m³ (rhyolite, Bonadonna et al., 2003).

Optional corrections of these values can include: Site shielding effects of local topography: MinVol +104% • Asymmetries in the flow propagation: *MinVol* -15%.

Bayes Net calculation results, using Equivalent Mass

Model	Maximum distance	Minimum PDC Mass [10 ¹² kg]						
Wiodei	130 km	1%ile	5%ile	50%ile	mean	95%ile		
D&H98	(1) Elicited inputs	11.5	26.0	263	450	1474		
D01105 /1 /	(2a) Elicited inputs	13.1	23.5	129	179	520		
D&H95 w/hot gas	(2b) Modified inputs*	12.3	16.8	40.7	40.0	69.0		
D&H95	(2c) Elicited inputs	7.68	15.6	100	148	462		

Model 2 is reported in three variants:

- Model 2a includes hot gas effects, and relies on the elicited input ranges of phi0, ws, rho, rho_a, rho_i.
- Model 2b includes hot gas effects, and relies on modified input ranges based on MDR modelling and Sauter diameter of analogues. These results impose lower phi0 and ws values than in the previous case.
- Model 2c does not include hot gas effects, and relies on the elicited input ranges of phi0, ws, rho, rho_a.

Through a Monte Carlo simulation randomly sampling the models, we also averaged the four probability distributions obtained. We assigned equal weights to the four modelling choices (Model 1, 2a, 2b, 2c). This is equivalent to a linear combination of the models' pdf or cdf (see Cooke, 1991; Bevilacqua, 2016).

		Minimum PDC Mass [10 ¹² kg]						
Model Mixture	1%ile	5%ile	50%ile	mean	95%ile			
	10.9	18.4	86.0	204	735			

2. Input ranges based on Expert Judgment

We based our input range estimation on structured expert judgment (Cooke, 1991; Aspinall, 2006). For judgment aggregation, we implemented the equal weight combination rule; we did not apply performance-based scores because of the relatively small number of experts participating and because the overheads and time demands involved in implementing a formal elicitation protocol were not warranted in this case.

Reported values express the percentiles of the probability distribution obtained by pooling experts' judgments (also called the solution Decision Maker DM).

We assumed the inputs to be formed of an array of independent variables. Further research may explore the effects of possible correlations between them.

|Scale|

Elicitation solution #2

Case name: Model 2

Nr. | Id

Elicitation solution #1

Case name: Model 1

Resulting solution (joint DM distribution of values assessed by experts)

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Nr.		Scale	1	5%	50%	95%	Units	I
			l	I	I	I		
1 0	Collapse H	uni	2	566	5752	9629	m	
2 F	Flow density	uni	68	6.3	992	1511 k	g/m^3	
3 5	Stress	uni	24	4.3	1868	7666	Pa	
4 I	Lambda ⁺	uni	1.	945 3	3.044	3.142	rad	

+in the sequel lambda will be fixed to pi unless otherwise stated.

2.1 Energy cone inference as "Model 0"

Before running the models described above, we used the elicited Collapse Height H to calculate the Minimum Volume (MinVol) required to reach 130 km runout, according to the **Energy Cone linear regression** in Ogburn and Calder (2017).

Minimum PDC Volume [km³]						
1%ile	5%ile	50%ile	mean	95%ile		
13.4	16.9	669	1.92 x 10 ⁵	4.1 x 10 ⁵		

The minimum volume (MinVol) estimate reported here appear to capture the same order of magnitude as results in the other models in terms of the lower percentile values. However, the mean and 95th percentile of MinVol look **infeasibly large**, numerically; Model 0 will not be considered when compiling a combination of the models.

08/01/2019

0.001789| 0.01103| 0.03675| 2.460| 2|Ws 0.04492| 1089| $2357 | kg/m^3 |$

Resulting solution (joint DM distribution of values assessed by experts)

1.023| 1.284| kg/m^3 | 4 | rho a 0.3184| $0.7957 | kg/m^3$ 5|*rho i* Modified input range based on MDR modelling and Sauter diameter of analogues

Nr. Id	Scale	MIN	MAX	Units	
	111	l			_
1* <i>phi0</i>	uni	0.002	0.01		
2* ws	uni	0.04	0.3	m/s	

*these results approximately impose input values in the range [min, median] assessed by joint DM.

of the flow mass and related uncertainties.

5. Conclusions

THIS STUDY SPECIFICALLY FOCUSES ON THE **NEW METHOD**.

WE APPLIED THE METHOD IN A **HYPOTHETICAL EXAMPLE** OF MAJOR IGNIMBRITE SCENARIO.

WE USED A NON-EXISTING TOPOGRAPHY AND ARBITRARY ERUPTIVE **MASS BOUNDS**.

ANY RELATION TO REAL VOLCANOES OR REAL AT-RISK CITIES IS PURELY COINCIDENTIAL.

6. Sensitivity analysis

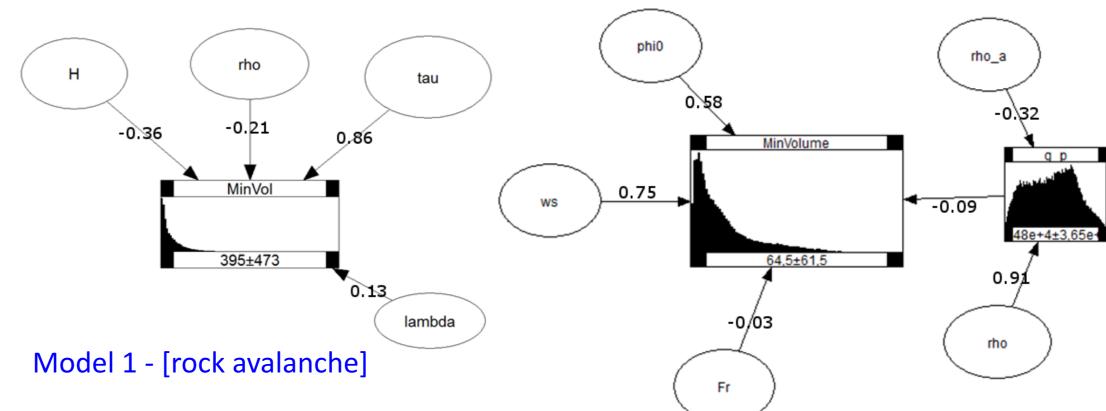
We use a Bayes Belief Network (BBN) related to each model to characterize uncertainties on the MinVol estimate, including in the calculational scheme correlation coefficients between the input variables and the calculated MinVol [Ababei (2016); UNINET BBN: https://lighttwist-software.com/ accessed 25 Nov 2019].

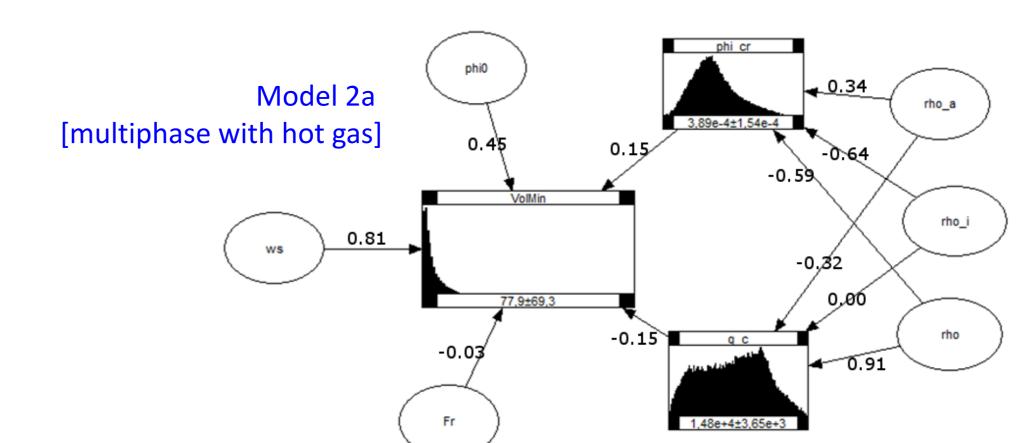
The large rectangular panels in each BNN illustrate the shape of the uncertainty distribution for the minimum volume needed to reach the at-risk city, based on the uncertainties in the relevant input variables, with the mean value and standard deviation shown in the lower frame of the panel.

Larger correlation values (marked on linking arc arrows) highlight the more sensitive variables. Negative correlation means that an increase of that input decreases the MinVol required to reach the at-risk city. Positive correlations produce the opposite effect.

The correlation coefficients are based on the joint DM distribution from elicitation and on the 130 km maximum runout distance. Further research is merited to develop better constraints on the uncertainty distributions of the most influential input parameters.

Model 2c [multiphase with cold gas]





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In particular, we adapted the 'energy conoid' approach to generate inundation maps along

radial directions, based on comparison of the mass-dependent kinetic energy of the flow with the **potential energy control by topography** in the direction of flow at distal ranges.

We described a new method for the **reconstruction** (or forecast) of probabilities that distal

geographic locations were inundated by a giant pyroclastic density current (PDC) in terms

Using appropriate model input uncertainty distributions, derived from **expert judgments**

using the equal weights combination rule, we estimated the mass amount needed to reach

a particular distal locality at any given confidence level and compared this with a range of

formulation of axisymmetric gravity-driven particle currents. We focused on models which

Our analysis relied on different versions of the Huppert and Simpson (1980) integral

possess analytical solutions, enabling us to utilize a very fast functional approach for

plausible eruptive masses. in a hypothetical major ignimbrite scenario

We focused on two different models:

enumerating results and uncertainties.

(i) Model 1 assumes the entire amount of solid material originates from a prescribed height above the volcano and flows as a granular current slowed down by constant

(ii) Model 2 is a multi-phase formulation and includes, in addition to suspended particles, interstitial gas thermally buoyant with respect to surrounding cold air. In the latter case, the flow stops propagating when the solid fraction becomes less than a critical value, and there is **lift-off of the remaining mixture** of gas and small particulates.

Our model parameters can be further constrained where there is reliable field data or with information from analogue eruptions.

Finally, we used a **Bayes Belief Network** related to each inversion model to evaluate probabilistically the uncertainties on the mass required, estimating correlation coefficients between the input variables and the calculated mass.

For any major magnitude ignimbrite PDC scenario, our method provides a rational basis for assessing the **probability of flow inundation** at critical geographic locations within distal areas when there is major uncertainty about the actual or predicted extent of flow runout.

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4. PDC inundation probability at an at-risk city based on eruption size In Figure 1 we show the cumulative distribution of Equivalent Mass required to reach the at-risk city, according to the four models described. We based a

PROBABILITY P OF PDC FLOW REACHING THE AT-RISK CITY,

Model 2b (Sauter-based ws, MDR-model-based phi0) - 100%

170 km RUNOUT [VDRE 200km³, with topo-effects]

• Model 1 (avalanche *tau*-friction) - 48.1%

• Model 2a (box model with hot gas) - 75.1%

• Combined model - 76.0%

first analysis on the Dense Rock Equivalent (DRE) volume of a major ignimbrite eruption, first we assume an example volume of 200 km³. We converted this DRE volumes to a mass of 500 • 10¹² kg units), assuming a density of 2500 kg/m³ (DRE rhyolite).

PROBABILITY P OF PDC FLOW REACHING THE AT-RISK CITY, 130 km RUNOUT [VDRE 200km³]

- Combined model 89.9%
- Model 2a (box model with hot gas) 94.2%

• Model 1 (avalanche *tau*-friction) - 68.9%

CUMULATIVE DISTRIBUTION OF EQUIVALENT MASS

- Model 2b (Sauter-based ws, MDR-model-based phi0) 100%
- Model 2c (cold gas) 96.4%

inundation probability:

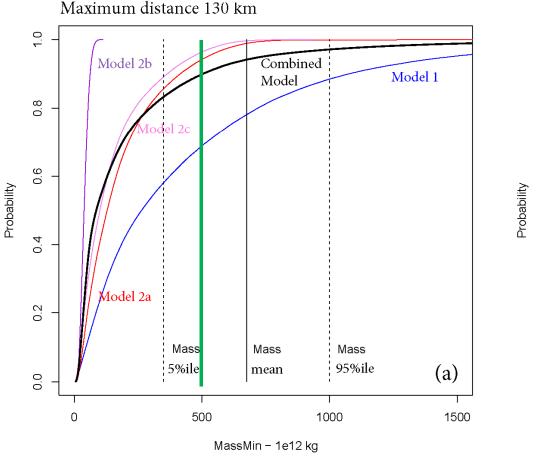
Model 2c (cold gas) - 80.9%

Fig. 1 also implements more detailed estimates: i.e. [140, 270, 400] km³, 5th, mean, and 95th percentile values

respectively. According this uncertain erupted volume range, and the combined runout model results, we have

An optional correction for flow direction asymmetries produces the following adjustment: P = [71.6%, 85.4%, 92.1%].

P = [67.5%, 82.2%, 89.5%]



Maximum distance 170 km

runout distance of 130 km, i.e. between volcano and at-risk city; (b) is related to a maximum runout distance of 170 km, allowing for the flow to overcome possible topography shielding effects near the site. The **black bold** line is the

Model 2b, Pink line is Model 2c A green vertical line marks the mass estimate related to the VDRE of 200 km³. **Black** vertical lines mark the mean and

distribution of the minimum mass Equivalent Mass required to reach the at-risk city.

(a) is related to a maximum

above. Blue line is Model 1, Red line is Model 2, Purple line is

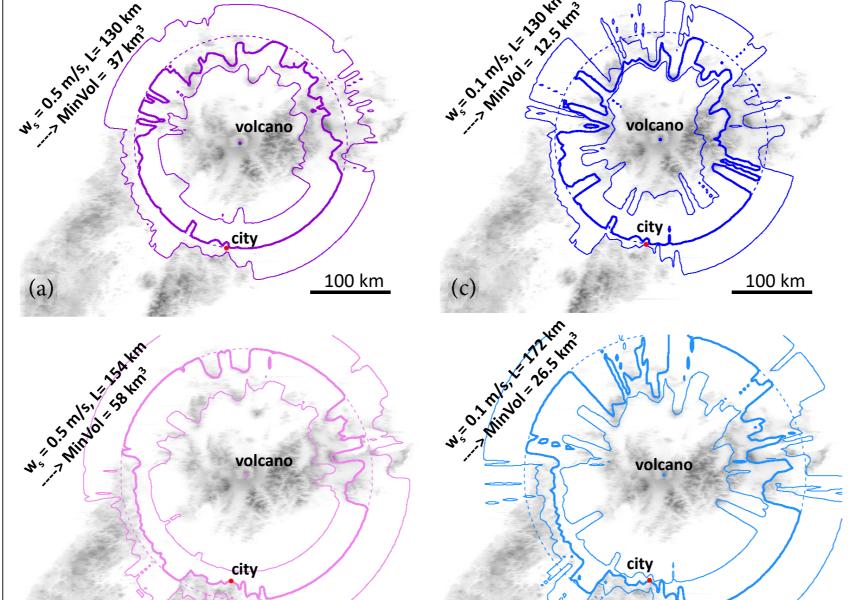
combined model, described

uncertainty percentiles of an estimate of plausible erupted mass, i.e. obtained from VDRE [140, 270, 400] km³.

5. Example of the effects of distal and medial topography

The realization of 130 km runout L in absence of topography is **not a sufficient condition** to impact a distant important city and any evacuation decision. We report examples of inundated regions as a function of runout distance L, according to the energy conoid technique.

We follow the 'energy conoid' approach adopted in Neri et al., (2015) and Bevilacqua et al., (2017). Figure 2 shows inundation maps based on the comparison of kinetic energy available to the flow and local topography along radial direction. Energy calculation is based on the equations of Model 2c, with illustrative values of phi0=1% and rho=2000 kg/m³.



topography. Shades of black display topographic obstacles (i.e. mountain ranges). (a),(b) are based on $w_s = 0.5$ m/s, and (c),(d) assume $w_s=0.1$ m/s. All the pictures assume phi0=1% and rho=2000 kg/m³, and the minimum

volume estimate V is displayed.

FIGURE 2. Panels (a),(c) show the

inundated region assuming L=130

km³; (b),(d) show the inundated

region assuming the minimum L

required to affect the at-risk city

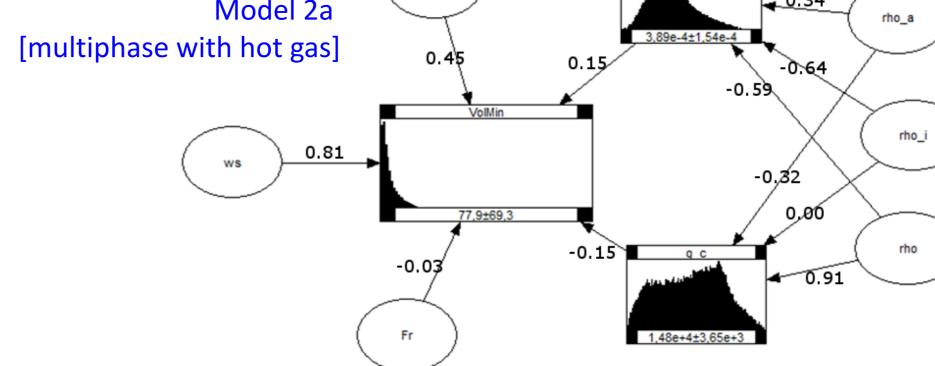
considering shielding effects of local

A coloured dot marks the volcano, and a red dot marks the at-risk city. A bold coloured line marks the **coundary of the inundated** region based on volume V. A thin dashed line marks a circle of radius L. Thin coloured lines mark the boundaries of the inundated region based on V/2 and 2V, for comparison. Topography, volcano, and at risk city are

According to the analytic expression of the *MinVol* variable, a more conservative assumption of requiring a runout of 170 km

an hypothetical example.

would double the required erupted volume in order to inundate the at-risk city. We remark that our energy conoid approach is only sensitive to the shielding effect of topography close to the at-risk city, and not on the large-scale topography around the source site



other experts participating in the elicitation sessions.