

1 **Impact of optical imagery and topography data resolution on the measurement of surface**  
2 **fault displacement using sub-pixel image correlation**

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8

9 **Key points:**

10 - We perform a quantitative assessment of the effects of optical imagery and topography data  
11 characteristics, primarily ground resolution, on the measurement of fault surface displacements.

12 - Measured displacements across the fault zone are under-estimated by a factor 0.7-0.8 when using  
13 low-resolution (>10 m) compared to high-resolution ( $\leq 1$  m) imagery.

14 - High-resolution (<0.5 m) stereo optical imagery presents as a good candidate for the future STV  
15 Earth observation system from earthquake hazard perspectives.

16

17 **Abstract**

18 The amount and spatial distribution of surface displacement that occurs during an earthquake are  
19 critical information to our understanding of the earthquake source and rupture processes. However,  
20 the earthquake surface displacement generally occurs over wide regions, includes multiple  
21 components affecting the ground surface at different spatial scales, and is challenging to  
22 characterize. In this study, we assess the sensitivity of optical imagery and topography datasets of  
23 different resolutions to the earthquake surface displacement when using optical image cross-

24 correlation (OIC) techniques. Results show that the average noise in the output displacement maps  
25 linearly increases with decreasing image resolution, leading to greater uncertainty in determining  
26 the geometry of the faults and the associated displacement. Fault displacements are, on average,  
27 under-estimated by a factor  $\sim 0.7-0.8$  when using 10 m compared to 0.5 m resolution imagery. Our  
28 analysis suggests that an optical image resolution of  $\leq 1$  m is necessary to accurately capture the  
29 complexity of the ground displacement. We also demonstrate that sub-meter vertical accuracy of  
30 the digital surface/elevation model (DSM/DEM) is also required for accurate image  
31 orthorectification, and is better achieved using high-resolution stereo optical imagery than existing  
32 global baseline topography data. Together, these results highlight the measurement needs for  
33 improving the observation of earthquake surface displacement towards the development of future  
34 Earth surface topography and topography change observing systems.

35

### 36 **Keywords**

37 Optical imagery, surface topography, cross-correlation, earthquake, fault, displacement

38

### 39 **1. Introduction**

40 Continental earthquakes often generate surface displacements that can be directly observed using  
41 satellite imagery. Measuring the amount and spatial distribution of the earthquake surface  
42 displacement then allows for untangling the shallow Fault Zone (FZ) coseismic deformation  
43 processes (Antoine et al., 2024, 2022, 2021; Barnhart et al., 2020; Li et al., 2023; Milliner et al.,  
44 2021; Zinke et al., 2019), bringing insights into the FZ mechanical behavior. Surface displacement  
45 measurements also enable constraining the earthquake rupture processes at depth using numerical  
46 inversion methods (e.g., Fialko et al., 2001; Jin and Fialko, 2021; Jolivet et al., 2014; Ragon and

47 Simons, 2020; Segall, 2010; Simons et al., 2002; Wang et al., 2020). However, the earthquake  
48 surface displacement can include multiple deformation components that affect the ground surface  
49 across a wide range of spatial scales. The broad scale coseismic deformation, of several tens of  
50 kilometers wide, primarily arises from the elastic response of the crust to the deep earthquake  
51 rupture (Avouac, 2015; Segall, 2010). The deformation that occurred on the faults and in their  
52 proximate vicinity, at a spatial scale of ~1-100 m, often leads to inelastic and permanent  
53 deformation of the crust which can take place through a combination of slip along primary faults,  
54 distributed deformation along secondary faults and/or fractures (e.g., DuRoss et al., 2020; Klinger  
55 et al., 2005; Rockwell et al., 2002; Teran et al., 2015; Yuan et al., 2022), and diffuse deformation  
56 of the surrounding medium (Antoine et al., 2024, 2022, 2021). The regions of surface deformation  
57 on- and near the fault correspond to what is referred to as the fault zone (FZ). Our ability to  
58 document the complete spectrum of earthquake surface deformation processes and understand the  
59 earthquake rupture process then hinges on the capability of the imaging sensors and associated  
60 processing methods to measure displacements over a wide range of spatial scales.

61

62 Observation techniques generally used to measure the earthquake surface displacement field  
63 include Synthetic Aperture Radar (i.e., Interferometric SAR, SAR pixel offsets), Global  
64 Navigation Satellite Systems (GNSS), Optical Image Correlation (OIC), and Light Detection And  
65 Ranging (LiDAR) point cloud difference. All of these techniques provide capabilities to image the  
66 ground movements with different footprints, ground resolutions, and measurement accuracies.  
67 GNSS and InSAR are the most sensitive methods with millimeter to centimeter level of accuracy,  
68 and are widely used for this purpose (e.g., Delouis et al., 2023; Floyd et al., 2020; Liu et al., 2021;  
69 Simons et al., 2002; Tong et al., 2010; Wang and Fialko, 2015). However, GNSS only permits

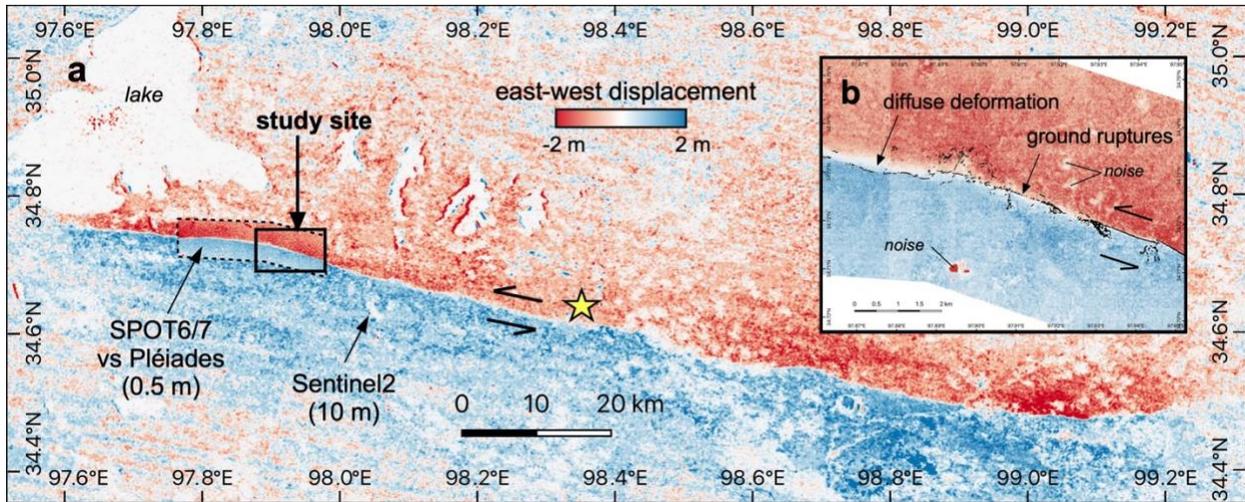
70 measurements at the discrete locations, often at tens to hundreds of kilometers spacing (e.g.,  
71 Fielding et al., 2020; Liu et al., 2021), and InSAR decorrelates in regions where displacement  
72 gradients are greater than the radar phase difference, which is generally along the fault ruptures  
73 (e.g., Fielding et al., 2013; He et al., 2023; Jin and Fialko, 2021; Massonnet et al., 1993; Socquet  
74 et al., 2019). For these reasons, both InSAR and GNSS present low constraints on the surface  
75 displacements that occur within and in the direct vicinity of the FZ. SAR pixel offset methods are  
76 sometimes used to retrieve the displacements closer to the faults (Jolivet et al., 2014; Liu et al.,  
77 2021; Reitman et al., 2023), but these measurements often are limited by the resolution of the  
78 publicly available SAR data which is generally about 10 m.

79

80 OIC (e.g., Aati et al., 2022; Leprince et al., 2007; Rosu et al., 2015), and differential LiDAR (e.g.,  
81 Borsa and Minster, 2012; Nissen et al., 2014; Scott et al., 2018) techniques, on the other hand,  
82 allow for three-dimensional (3-D), spatially continuous, and possibly high-resolution (submeter-  
83 scale) measurement of the earthquake surface displacement field, even in regions of high-  
84 displacement gradients. Between the two techniques, OIC is more commonly used (e.g., Antoine  
85 et al., 2024, 2022, 2021; Delorme et al., 2020; Milliner et al., 2021, 2016; Teran et al., 2015; Zinke  
86 et al., 2019) as satellite optical data cover wide regions ( $> 15^\circ$ ). LiDAR is generally limited to small  
87 study area because of the high-cost associated with LiDAR data acquisitions. Satellite optical  
88 imagery has been widely developed both in the public and private domains, now allowing for dense  
89 data archives in many regions of the globe. Optical imageries with different viewing angles can be  
90 combined to reconstruct the 3D ground surface (digital surface model, DSM) through  
91 photogrammetry methods, and represent a unique dataset to document the ground surface  
92 characteristics, topography, topography change, and horizontal displacement for analyzing the

93 earthquake and FZ deformation processes. Therefore, stereo optical imaging has been proposed as  
94 one of the technology candidates for future Earth Surface Topography and Topography Change  
95 Global Observing systems such as NASA Surface Topography and Vegetation Mission (STV)  
96 (Donnellan et al., 2022, 2017).

97



98

99 **Figure 1. (a) East-west surface displacement along the 2021 Maduo, Tibet, rupture from the**  
100 **OIC of Sentinel-2 optical images at 10 m ground resolution (Antoine et al., 2024). HR (0.5**  
101 **m) OIC result from Antoine et al. (2024), derived from the cross-correlation of SPOT6/7 and**  
102 **Pleiades images, and covering our study site (black rectangle) shown within the dashed black**  
103 **polygon. Epicenter location, from global CMT catalog (GCMT), is shown with a yellow star.**  
104 **(b) East-west surface displacement from the HR OIC results from Antoine (2024) across our**  
105 **study site. Study site area corresponds to the common area covered by all datasets, and is 8.5**  
106 **km-long. Field rupture map from Yuan et al. (2022) is overlaid in black.**

107

108 As of today, low-resolution (LR; 3-10 m) optical imagery (e.g., SPOT1-4, Planet, Landsat,  
109 Sentinel), because it is freely available and presents a regional/global coverage, is widely used to

110 characterize the horizontal earthquake surface displacement field (e.g., Avouac, 2015; Chen et al.,  
111 2020; He et al., 2023; Li et al., 2022; Milliner and Donnellan, 2020; Fig. 1a). High-resolution  
112 imagery (HR; <3 m; e.g., Pleiades, WorldView, SPOT5-7) was though shown to be a more reliable  
113 source of information, especially for mapping the fault ruptures (e.g., Klinger et al., 2005; Reitman  
114 et al., 2023), and assessing the different components of the near-fault surface deformation (e.g.,  
115 Antoine, 2024, 2021; Zhou et al., 2018; Fig. 1b) and the fault zone width (FZW; Ajourlou et al.,  
116 2021) using OIC methods. In the case of stereo acquisitions, HR optical imagery also allows for  
117 modeling the surface topography and measuring topography change (e.g., Antoine et al., 2021,  
118 2022; Barnhart et al., 2020, 2019; Delorme et al., 2020; Zhou et al., 2015). However, HR images  
119 are usually accessible only through purchase or with agreement with satellite agencies (exceptions  
120 of freely available samples exist but only over limited regions, i.e., in the case of disasters  
121 programs), and do not cover the entire globe surface. Moreover, HR images have a smaller spatial  
122 coverage (<20x20 km), and are generally provided as non-orthorectified products, meaning that  
123 they present distortions due to acquisition geometry (viewing angle and satellite attitude  
124 parameters). As a result, HR OIC requires accurate image orthorectification to be performed  
125 beforehand, which consist in a careful modeling of the camera attitude parameters and ground  
126 topography, and projection of the stereo optical images into a common ground geometry (Leprince  
127 et al., 2007; Rupnik et al., 2016; Shean et al., 2016).

128

129 For image orthorectification and vertical displacement measurement, accurate Digital  
130 Elevation/Surface Models (DEMs/DSMs) from pre- and post-earthquake periods are therefore  
131 needed. However, global-coverage DSMs and/or DEMs are generally available at resolutions of  
132 >30 m (e.g., ALOS30, NASA/SRTM, Copernicus , ASTER). HR DSMs and/or DEMs are

133 generally available only locally in some limited regions of interest, usually from previous scientific  
134 publications (e.g., Antoine et al., 2022; Willis et al., 2019), or available for purchase over wider  
135 areas (e.g., ALOS World 3D). Moreover, calculating DSM from HR optical imagery requires  
136 availability and access to stereo data with acquisition parameters compliant with  
137 stereophotogrammetry requirements (e.g., complementary viewing angles, access to the  
138 acquisition parameters) (Hasegawa et al., 2000; Yin et al., 2023). In this regard, publicly available  
139 HR optical and topography data on a global scale are essential for accurate FZ deformation analysis  
140 and feature extraction over different geological and tectonic settings (Donnellan et al., 2017;  
141 Schumann and Bates, 2018).

142

143 Towards developing a global Surface Topography and Vegetation (STV) Earth Observation  
144 System (Donnellan, 2021), in this study, we provide a quantified analysis of the impact of optical  
145 and topography data resolution on the measurement of surface displacement across the FZ from  
146 the Solid Earth fault hazard perspective. The STV mission is in development stage with a primary  
147 goal of global mapping 3-D topography and topography change subject to scientifically defined  
148 resolution and measurement accuracy across broad science and application disciplines including  
149 Solid Earth, vegetation structure, cryosphere, hydrology and coastal geomorphology (e.g., DeLong  
150 et al., 2022; Donnellan et al., 2022). Through this study, we aim to address the following questions:  
151 what is the impact of optical imagery and topography data resolution on the assessment of the FZ  
152 geometry and surface displacement using OIC methods? How does measurement accuracy evolve  
153 with data resolution? To answer these questions, we perform OIC analyses using optical imagery  
154 and DEMs/DSMs datasets of various resolutions and sources to measure the near-fault ground  
155 displacement field, using the 2021  $M_w7.4$  Maduo, Tibet, earthquake rupture as a case study. For

156 different datasets or combination of them, we analyze the signal to noise ratio in the obtained  
157 ground displacement maps, and its impacts on the determination of the FZ geometry, the FZW,  
158 and the associated displacements.

159

## 160 **2. Study site and available observations**

161 The 2021  $M_w$ 7.4 Maduo earthquake ruptured bilaterally and with a strike-slip left-lateral  
162 mechanism the Jiangcuo fault, located within the Bayan Har block of the Eastern Tibetan plateau  
163 (e.g., Fan et al., 2022; L. He et al., 2021; Liu et al., 2022; Wei et al., 2022). This earthquake  
164 generated a 160-km long surface rupture, which has been widely characterized using several  
165 geodetic (e.g., Fan et al., 2022; K. He et al., 2021; Jin and Fialko, 2021; Liu et al., 2021; Tong et  
166 al., 2022; Xiong et al., 2022; Yang et al., 2022; Zhao et al., 2021), field (Pan et al., 2022; Ren et  
167 al., 2022, 2021; Xie et al., 2022; Yuan et al., 2022) and seismic data (e.g., Li et al., 2022; Liu et  
168 al., 2021; Wei et al., 2022; Zhang et al., 2022). Field data report sparse surface ruptures with up to  
169 2.6-2.9 m of horizontal displacement detected locally along primary fault strands (Pan et al., 2022;  
170 Ren et al., 2022, 2021; Xie et al., 2022; Yuan et al., 2022). The total horizontal surface  
171 displacement measured from OIC is 2.27-2.35 m, and occurs over a FZW of 30 m to 2.15 km  
172 (Antoine, 2024; Li et al., 2022; Fig. 1). Thanks to the extensive imagery archives and pre-existing  
173 studies, this event represents a good case to assess the effects of different optical imagery  
174 resolutions (0.5-10 m) on resolving the surface displacements across the FZ. Moreover, the rupture  
175 area is free of vegetation, snow and human activities, which makes it ideal for OIC applications.  
176 We focus on a specific region of the 2021 Maduo rupture, located at 45 km to the north-east of the  
177 epicenter, where both complex fault geometry, and distributed and diffuse deformations were  
178 documented (Antoine et al., 2024; Li et al., 2022; Li et al., 2023). Along this section, a separate

179 analysis using HR optical data revealed an average surface displacement of ~2.81 m and FZW of  
180 ~363 m (Antoine et al., 2024; Fig. 1b).

181

### 182 **3. Data and Methods**

#### 183 **3.1. Optical and topography data**

184 This work uses optical imagery and topography data of various resolutions provided by different  
185 space agencies to perform image orthorectification and OIC (Tab. S1). The results are combined  
186 with pre-existing OIC (Antoine et al., 2024) and SAR results (Liu et al., 2022) for further analysis.  
187 Pre-existing OIC results include pre-earthquake SPOT6/7 and post-earthquake Pleiades OIC at a  
188 common 0.5 m ground resolution, and Sentinel2 OIC at a 10 m ground resolution (Antoine et al.,  
189 2024). Pre-existing SAR results are derived from differential interferometric SAR (InSAR,  
190 DInSAR), pixel offset-tracking (POT), multiple aperture InSAR (MAI), and burst overlap  
191 interferometry (BOI) measurements at a ground resolution of 100 m (Liu et al., 2022). Optical  
192 imagery datasets used in this study include WorldView1/2/3 and Planet images, at resolutions of  
193 0.39-0.67 and 3.125 m, respectively (Tab. S1). Planet images are downloaded as orthorectified  
194 products, whereas WorldView1/2/3 images correspond to non-orthorectified images. We also use  
195 downsampled versions of the WorldView orthoimages we produce (see 3.2. for details on  
196 orthorectification) to analyze separately the effect of image resolution from other effects, including  
197 sensor quality and image orthorectification which vary for different satellite agencies. This  
198 approach also allows us to extend the tested image resolutions (every 1 m from 0.5 to 10 m),  
199 allowing for a more complete analysis of the data resolution effect on the measurements.  
200 Topography datasets include HR pre- and post-earthquake DSMs derived from the HR tri-stereo  
201 WorldView1/2/3 images along with external and publicly available 30-m ground resolution NASA

202 and Copernicus DEMs. Similar to the optical data, we produce downsampled versions of the HR  
203 pre- and post-earthquake WorldView DSMs to test a wide range of topography data resolutions  
204 (from 1 to 30 m) while preserving the native vertical accuracy of the stereo-DSMs (~0.5 m; Rupnik  
205 et al., 2018; Schumann and Bates, 2018; Zhou et al., 2015). The effect of DSM vertical accuracy,  
206 which primarily depend on the sensor type, is assessed by comparing the quality of OIC  
207 displacement measurements derived from orthorectifying the optical images using either the  
208 stereo-DSMs or the Copernicus and NASA DEMs.

209

## 210 **3.2. Method**

### 211 **3.2.1. Optical image cross-correlation (OIC)**

212 Optical image cross-correlation (OIC) is a technique that allows for measuring the continuous  
213 displacement of features between spatially co-registered images (Leprince et al., 2007;  
214 Puymbroeck et al., 2000; Rosu et al., 2015). Using satellite (or airborne) images acquired at  
215 different time periods, one can monitor the ground surface evolution (e.g., Bontemps et al., 2018;  
216 Dehecq et al., 2015). In this study, we use the MicMac OIC method (Rosu et al., 2015; Rupnik et  
217 al., 2017) based on a 2D statistical feature matching exploiting the color information from the  
218 pixels (0-255 in panchromatic images) to measure ground surface displacements between pre- and  
219 post-earthquake orthorectified images. In practice, one defines a correlation window size which  
220 determines the width of the group of pixels that will be used for the statistical matching. At every  
221 pixel position in the pre-earthquake image, the correlator considers the pixel color pattern within  
222 this correlation window, and searches for the best matching pattern in the post-earthquake image  
223 (correlation score from 0 to 1). The search window width in the post-image is set based on a priori  
224 knowledge of the surface displacement, generally from seismic, field or other geodetic

225 observations, to constrain the search area to a reasonable region, and limit the calculation time and  
226 matching errors. The search window moves across the post-earthquake image with step ranging  
227 from 1 to  $1/20^{\text{th}}$  of a pixel through an iterative correlation process, allowing for sub-pixel  
228 displacement detection.

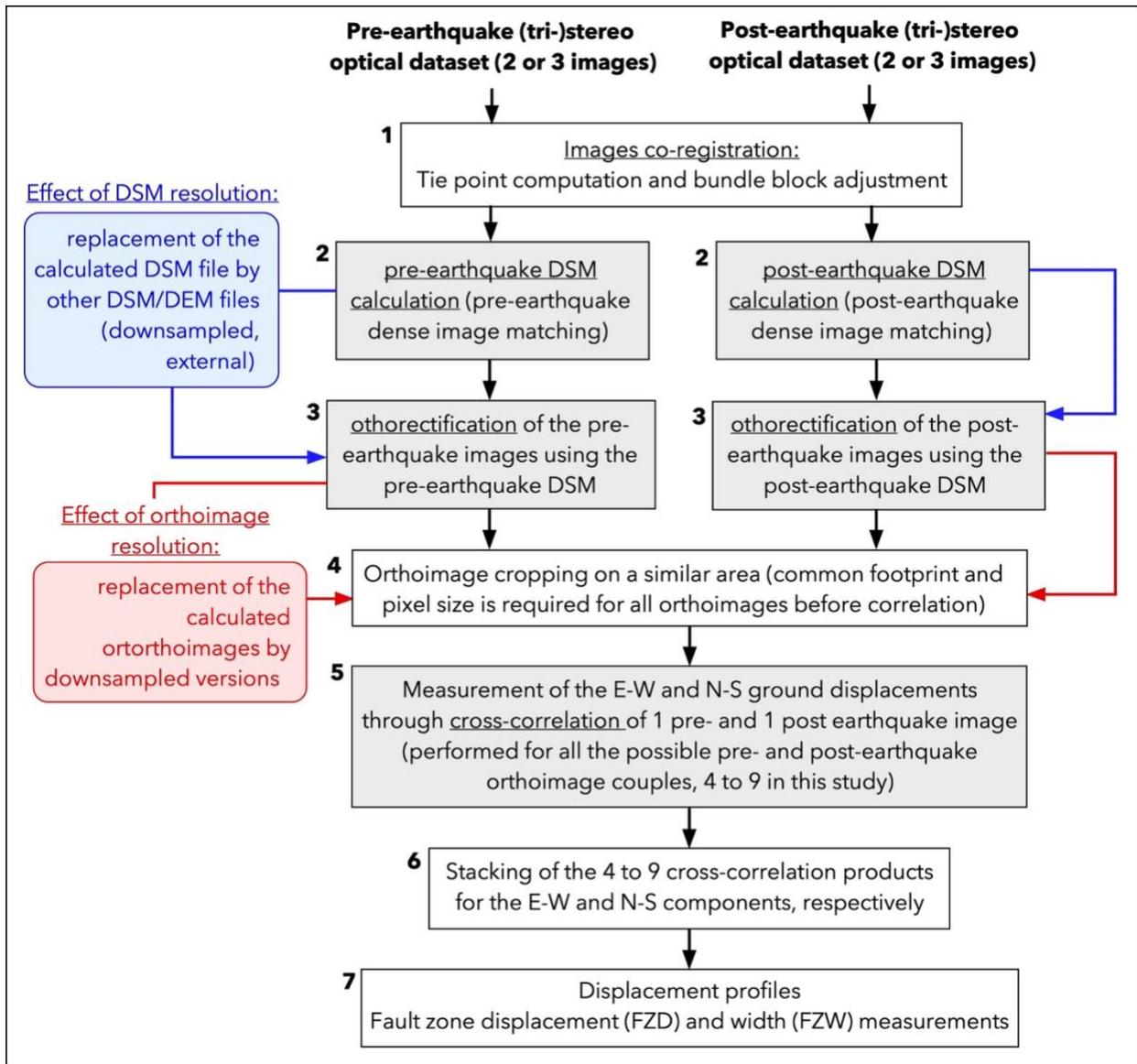
229

230 In this study, we performed OIC for all datasets using the same parameters, including a 5 pixels-  
231 wide correlation window, a search window of  $\pm 2.5$  m, and a regularization of 0.3. These parameters  
232 can be adjusted to improve the OIC results in each case but, in a matter of consistency for the  
233 sensitivity study, we use common parameters throughout the study. The cross-correlation products  
234 consist in pixel displacement maps, in the x and y directions of the images, which correspond to  
235 the east-west and north-south direction of the orthorectified images, along with a correlation score  
236 (0-1) map (Fig. S1). The correlation score map can be used to filter the OIC results, and weight  
237 the displacement measurements (Fig. 2, steps 6 and 7). A wide range correlation parameters (e.g.,  
238 Cofaru et al., 2010; Rosu et al., 2015) and filtering techniques (e.g., Andreuttiova et al., 2022;  
239 Stumpf et al., 2018) can be used to improve the OIC result for each dataset. However, the combined  
240 effects of these techniques and parametrizations are numerous, can vary for each dataset, and are  
241 thus not directly addressed within the scope of this study which primarily focuses on the effect of  
242 the data resolution. However, through simple tests on some of the datasets (Fig. S2) as well as  
243 comparison with other published measurements (Tab. 1), we briefly attempt to assess some  
244 possible effects of different processing techniques and parametrizations.

245

### 246 **3.2.2. Complete processing pipeline**

247 The processing pipeline used in this study to perform OIC of stereo images includes 7 major steps.  
248 Alongside the MicMac software, this work requires the use of GDAL for data cropping and  
249 resampling, Stackprof tool for stacked displacement profile extraction (see also section 3.2.3.), and  
250 python scripts for compilations of the measurements and statistics. This pipeline is similar to those  
251 applied in most photogrammetry and OIC approaches (e.g, Aati et al., 2022; François Ayoub et  
252 al., 2009; Shean et al., 2016), though they all present small differences depending on the type of  
253 data and software used. As mentioned earlier, differences in the correlation method and approach  
254 can lead to some differences in the resulting product (Avouac and Leprince, 2015; Bickel et al.,  
255 2018; Dematteis and Giordan, 2021; Montagnon et al., 2023; Rosu et al., 2015). Based on existing  
256 comparative studies (e.g., Bickel et al., 2018; Dematteis and Giordan, 2021; Montagnon et al.,  
257 2023; Rosu et al., 2015) as well as comparable measurements on similar sites (e.g., Antoine et al.,  
258 2024, 2022, 2021; Barnhart et al., 2020; Cheng and Barnhart, 2021; Li et al., 2023; Milliner et al.,  
259 2021; [Tab. 1](#)), we consider the effect of data resolution to be consistent across OIC approaches.  
260



261

262 **Figure 2. Diagram describing the processing pipeline applied in this study, based on the use**  
 263 **of the MicMac photogrammetry and image cross-correlation software (Rosu et al., 2015;**  
 264 **Rupnik et al., 2016). The pipeline includes 7 steps. Steps 2 and 3 are performed separately**  
 265 **between the pre-earthquake and post-earthquake datasets, whereas other steps involve both**  
 266 **pre- and post-earthquake datasets together. Grey steps correspond to the steps leading to**  
 267 **output products including DSMs (step 2), orthoimages (step 3), and OIC maps (step 5). Tests**

268 **using external and/or resampled datasets, highlighted in color on the side of the diagram,**  
269 **take place at steps 3 and 4, respectively.**

270

271 Step 1 consists in image co-registration through tie point detection (identification of common  
272 features across images), and bundle block adjustment (refinement of camera calibration and  
273 orientation; Rupnik et al., 2016). This step is performed commonly to the pre- and post-earthquake  
274 images to improve their co-registration accuracy. Step 2 calculates pre- and post-earthquake DSMs  
275 separately from the dense matching of the pre- and post-earthquake image pools for which  
276 orientation has been refined earlier (Step 1). Step 3 is the image orthorectification, performed  
277 individually for each image using the DSMs/DEMs of the corresponding period. Orthorectifying  
278 the pre- and post-earthquake images using separate pre- and post-earthquake DSMs prevents  
279 introducing noise from earthquake-related topography changes. Step 4 consists in cropping all  
280 orthoimages onto the same footprint, generating datasets with similar georeferencing, pixel size  
281 and pixel number for the OIC purpose. This step can be done automatically if images are precisely  
282 co-registered, or manually by locating common features in the images. Only then, image sub-pixel  
283 cross-correlation (OIC) can be performed, at Step 5. In the case of stereo (2 images) or tristereo (3  
284 images) datasets, OIC is performed between all the possible pre- and post-earthquake orthoimage  
285 couples, which in this study represent a maximum of 9 combinations. At Step 6, the different OIC  
286 products are stacked, separately for the x (east-west) and y (north-south) components, using a  
287 weighted median method based on the OIC correlation score map (Delorme et al., 2020). Finally,  
288 at Step 7, across-fault displacement profiles are extracted to analyze the across-fault displacement  
289 offsets, and build FZ displacement (FZD) budget and FZ width (FZW) evolution (see section  
290 3.2.3.). The complete processing pipeline is applied to the cases of the non-orthorectified HR

291 optical images (WorldView in this study, similar pipeline applied to SPOT and Pleiades in  
292 Antoine, 2024). However, some data correspond to orthorectified products (Planet, Sentinel) that  
293 were already processed from steps 1 to 3 directly by the space agencies, using a different strategy  
294 generally based on the use of global LR DEMs. In this case, data only require processing from  
295 Step 4 to 7.

296

297 Tests on the effect of DEMs/DSMs resolution on the orthorectification and OIC products are  
298 performed by replacing, at Step 3, the original DSM calculated from the stereo images by the tested  
299 topography dataset, cropped over the same area and resampled with the same pixel size. Tests on  
300 the effect of optical image resolution, apart from those performed on the datasets originating from  
301 different sensors, are performed by using, in Step 4, downsampled versions of the high-resolution  
302 orthoimages originally obtained from the tri-stereo WorldView dataset, in Step 3. Downsampling  
303 of both the DSMs and the orthoimages is performed using GDAL.

304

### 305 **3.2.3. Displacement profiles, and FZ displacement and FZW measurements**

306 This study focuses on the measurement of the fault zone displacements (FZD) and width (FZW)  
307 which represent two main parameters generally considered in earthquake surface displacement  
308 analysis (e.g., Antoine et al., 2024, 2022, 2021; Gold et al., 2021, 2015; Li et al., 2022; Milliner et  
309 al., 2021; Teran, 2015; Zinke et al., 2019). To do so, we use stacked-profiles placed  
310 perpendicularly to the FZ every 200 m along the study area (StackProf tool), representing 44  
311 profiles in total. Profile width is set to be 200 m, which was estimated as a good trade-off between  
312 noise reduction and conservation of the signal complexity (Li et al., 2022). Within the profile box,  
313 displacement measurements are stacked using a weighted median method based on the OIC

314 correlation score map. Displacements initially obtained in the east-west and north-south  
315 components are projected, using the local azimuth of the inferred FZ, onto the fault-parallel and  
316 fault-normal directions. For the case of the 2021 Maduo earthquake, left-lateral displacement is  
317 the primary component of the surface displacement (Fig. 1), hence the measurements are focused  
318 on the fault-parallel component.

319  
320 For each profile, we fit linear regressions to the displacement values outside of the inferred FZ,  
321 and assess the FZD by measuring the displacement offset between these two regressions (e.g.,  
322 Antoine et al., 2024, 2022, 2021; Gold et al., 2015; Li et al., 2022). We also report the width of  
323 this offset, that corresponds to the FZW. In this study, we focus on measuring the maximum offset  
324 between the two regressions, which represents the total surface displacement across the FZ. We  
325 do not quantify the separate contributions of the localized surface slip and that of the more  
326 distributed and/or diffuse deformation, which has been the focus of another study (Antoine et al.,  
327 2024). Uncertainty on each FZD measurement, given by the StackProf tool, considers the error on  
328 the linear regressions on both sides of the FZ. Uncertainty on the FZW, however, is essentially  
329 epistemic and related to the choice of the FZ location by the operator (Reitman et al., 2022). It is  
330 generally poorly documented in the studies, and there exists no unique method to assess such  
331 uncertainty. In this study, we choose to assess the FZW uncertainty by measuring a minimum and  
332 maximum FZW for each profile along with a preferred FZW for which the FZD was assessed,  
333 method previously used by Gold et al. (2015).

334

## 335 **4. Results**

### 336 **4.1. Impact of optical image resolution on the OIC results**

### 337 **4.1.1. Images acquired by different satellite sensors**

338 In this section, we compare ground displacement measurements obtained from the OIC of images  
339 acquired by different satellite sensors and with resolutions ranging from ~0.5 to 10 m (Figs. 3, 4,  
340 S1 and S3, and Tabs. S1 and S2). We also compare these results with the displacement maps  
341 derived from SAR measurements (Liu et al., 2022) at a 100 m resolution.

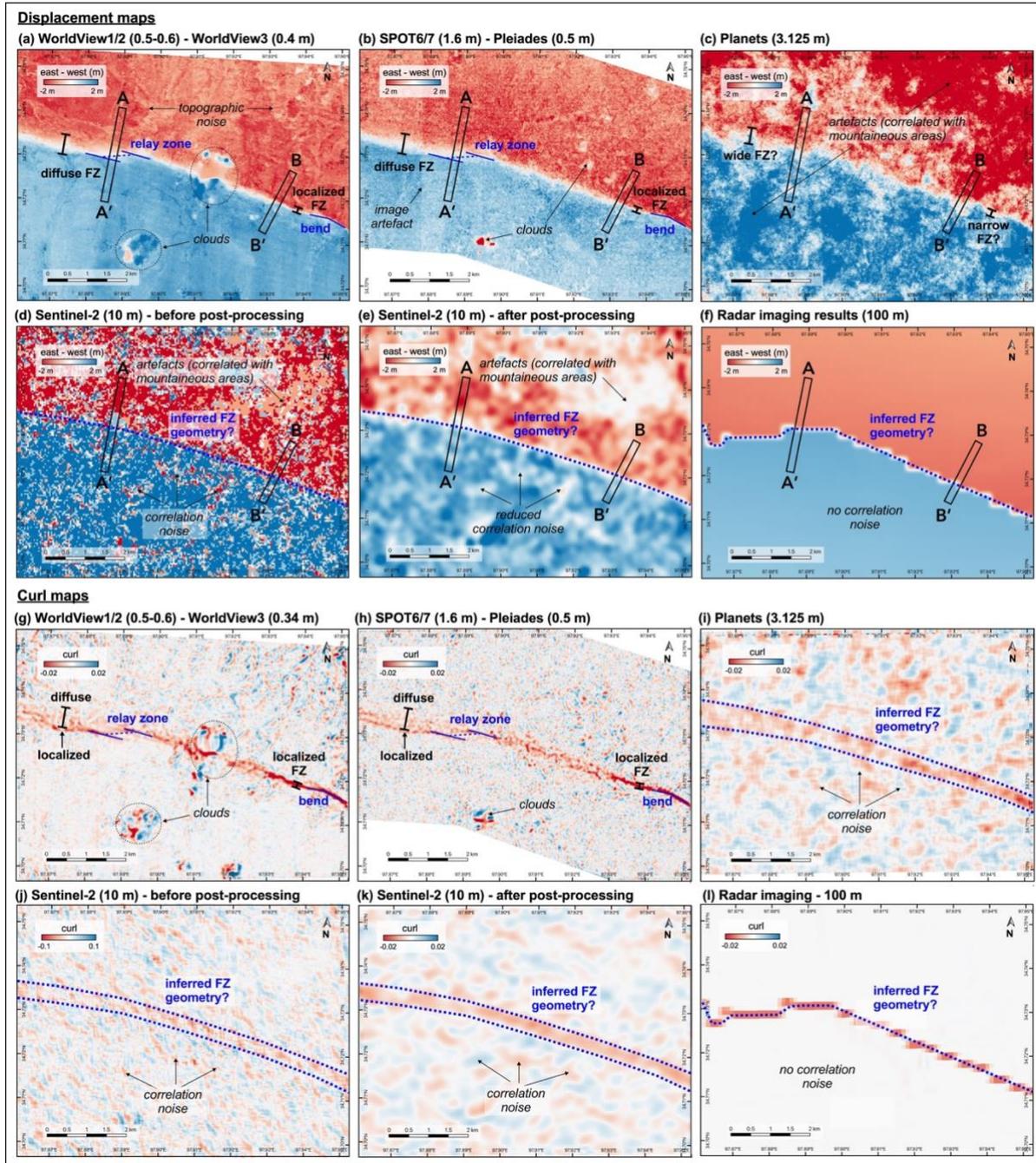
342

#### 343 *4.1.1.1. Surface displacement and curl maps*

344 We present the east-west (E-W) component of the displacement field (north-south (N-S)  
345 component presented in Figure S1) along with curl maps, which were derived from the E-W and  
346 N-S displacements (Figs. 3g,h, and S3). Visual analysis of the displacement maps obtained from  
347 the OIC of the different imagery datasets first highlights a consistent E-W displacement offset of  
348 amplitude  $\pm 2$  m. Transition between the NE and SW regions of the study area, which were  
349 displaced in opposite directions, occurs along an oblique structure lying at the center of the study  
350 area, corresponding to the FZ. Even though, at a first order, consistency is observed between the  
351 different results, we also highlight the differences in the signal to noise ratio, impacting the inferred  
352 FZ geometry and associated FZD and FZW measurements. Spatially coherent noise first arises  
353 from topography residuals due to camera attitude parameter estimations and DEM errors (Fig. 2,  
354 steps 1-3). Such noise is common for data downloaded as orthorectified products, and for which  
355 orthorectification is not always performed using a refined camera model nor a HR and accurate  
356 DEM/DSM (Fig. 2, step 1). Noise also occurs in regions of lesser surface texture change, where  
357 the correlator cannot identify common pixels between the pre- and the post-earthquake images,  
358 with common sources that are vegetation, snow, water, and clouds. Finally, high-frequency noise

359 arises from random cross-correlation error, especially enhanced by image noise (Bornert et al.,  
 360 2018, 2009; Su and Zhang, 2016).

361



362

363 **Figure 3. (a-f) East-west (E-W) surface displacements along the study site calculated from**  
 364 **the OIC of images originating of different sensors and of different resolutions (Tab. S1), and**

365 comparison with SAR-derived measurements from Liu et al. (2022). AA' and BB' across-  
366 fault displacement profiles are presented in Figure 4. (g-l) Map of the curl ( $\omega$ ) calculated  
367 from the E-W ( $u_x$ ) and N-S ( $u_y$ ) downsampled and filtered (Fig. S3) displacement maps using  
368 the following relation  $\omega = \nabla \times \mathbf{U} = \frac{\partial u_y}{x} - \frac{\partial u_x}{y}$  (Zhou et al., 2018). Negative curl (red) corresponds  
369 to anti-clockwise rotation and is consistent with the left-lateral mechanism of the 2021  
370 Maduo rupture.

371

372 Among the different results, those obtained from HR imagery (Fig. 3a,b) allow for describing  
373 spatial variations in the FZ geometry, and separating the diffuse deformation regions, especially  
374 to the NW of the study area, from the localized ones. Fault bends and relay zones can also be  
375 observed along the FZ in these maps (blue lines in Fig. 3a,b). Such complex FZ geometry is  
376 particularly highlighted in the curl maps including, for example, asymmetric and mixed diffuse  
377 and localized patterns across the FZ to the NW of the study area. Across both HR displacement  
378 and curl maps, such complex patterns are consistently observed (Figs. 3a,b,g,h, S1a,b, and S3a,b).  
379 Small differences include greater topography-correlated noise in the WorldView OIC result (as  
380 well as decorrelation due to the presence of a cloud across the FZ), and greater background noise  
381 in the SPOT/Pleiades OIC result. Greater background noise in the SPOT/Pleiades OIC likely arises  
382 from the fact that SPOT6/7 images were upsampled to the 0.5 m native resolution of the Pleiades  
383 images (Antoine et al., 2024), possibly leading greater random correlation, especially among  
384 upsampled pixels.

385

386 Results obtained using LR imagery (Fig. 3c,d,e) present a smaller signal to noise ratio, resulting  
387 in a spatially heterogeneous displacement field both across and outside of the FZ. Noise is

388 particularly visible in the regions outside of the fault zone ( $>1$  km), where one would rather expect  
389 about constant displacements as shown in the HR results. In the curl maps, noise arises as high-  
390 frequency displacement variations, especially visible in the regions outside of the FZ, and with  
391 amplitudes similar to that of the curl signal across the FZ (Fig. 3i,j,k). Noise standard deviation  
392 (std, analyzed to the NE of the inferred FZ within the mountainous areas, Fig. S4) increases with  
393 decreasing image resolution, from a std of 0.30 m in the WorldView OIC results to a std of 1.76  
394 m in the Sentinel-2 OIC results (raw result, no filtering). After filtering and smoothing of the  
395 Sentinel-2 results, background noise reduces with now a std to 0.47 m (Fig. 1e,k; see Antoine et  
396 al., 2021 for methodology). Nevertheless, filtering and smoothing fails to help retrieving the high-  
397 resolution information on the FZ geometry as well as removing the low-frequency noise. As a  
398 result, the FZ primarily shows up as a continuous and curved structure absent of geometrical  
399 complexity in the LR displacement results. The FZ identification and FZD and FZW  
400 measurements in the LR results are then subjected to ad hoc interpretations, and are not always  
401 consistent from one displacement product to another.

402

403 As a case example of comparison between OIC and SAR measurements in the near-fault domain,  
404 we compare our results with the SAR-derived displacement maps from Liu et al. (2022) (Fig. 3f,l).  
405 The SAR results present a 100 m ground resolution and a cm-scale accuracy, which is consistent  
406 with the characteristics of typical multi-look SAR-derived products used to image earthquake  
407 ground displacements (e.g., Fielding et al., 2013; He et al., 2023; Jin and Fialko, 2021; Massonnet  
408 et al., 1993; Socquet et al., 2019; Tong et al., 2022; Zhao et al., 2021). Displacement amplitudes  
409 between the SAR and OIC measurements are consistent at the first order. However, at the scale of  
410 the study area, the FZ location and geometry reported by the SAR measurements is incorrect with

411 regard to both the LR and HR OIC results. This observation supports the previous inference that,  
412 in the near-fault domain, measurements derived from SAR data are generally less accurate than  
413 those derived from optical data.

414

415 *4.1.1.2. Displacement profiles, and fault zone displacement budget, and fault zone width*  
416 *evolution*

417 Displacement profiles (see section 3.2.3 for methodology) are used to assess the evolution of the  
418 FZD and FZW along the study area, for the different displacement products (Figs. 3 and S1). We  
419 first present the analysis of two separate displacement profiles, and then assess the along-strike  
420 FZD budget and associated FZW evolution. The two profiles include one profile across a wide FZ,  
421 characterized primarily by diffuse deformation (AA' in Figs. 3 and 4a), and one profile across a  
422 narrower FZ, characterized primarily by localized deformation (BB' in Figs. 3 and 4a). We first  
423 observe that the two displacement profiles, when taken across the HR results (red and blue profiles  
424 in Figure 4a), report similar displacement patterns. These patterns include a displacement offset in  
425 the middle of the profile, of similar amplitude in both profiles, surrounded by regions of constant  
426 displacement. The position of the displacement offset corresponds to the location of the FZ. The  
427 two same profiles taken across the SAR-derived results (black profile, Figure 4a) display  
428 consistent patterns; however, in this case, the displacement offset is under-estimated, and mis-  
429 located in the case of profile BB', highlighting the lower constraints provided by SAR data in the  
430 near-fault domain.

431

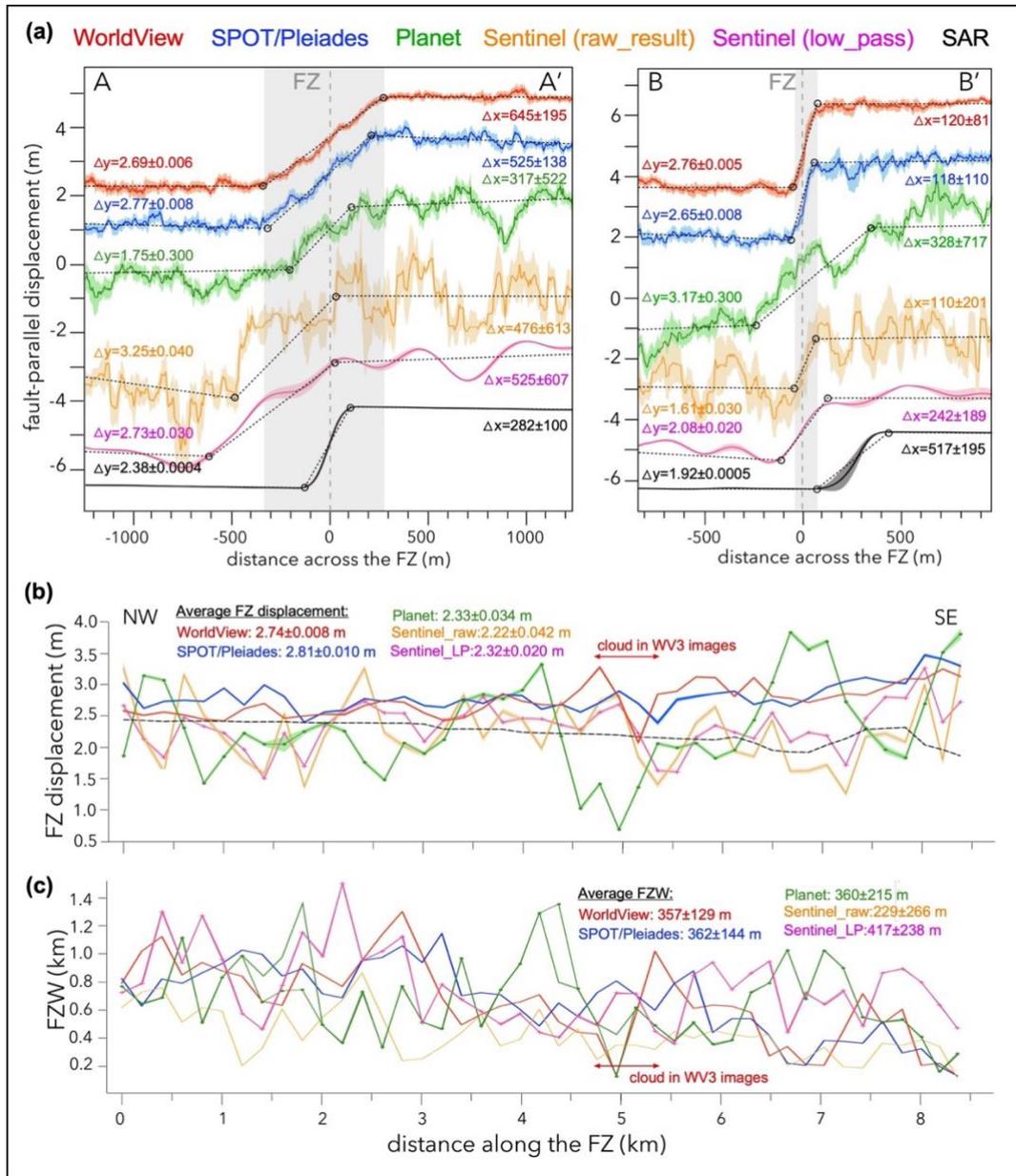
432 Profiles taken across the LR displacement maps (green, orange and pink profiles, Figure 4a) show  
433 displacement variations both at low- and high-frequencies, making the determination of the

434 displacement offset and the associated FZD and FZW more challenging. In fact, these profiles do  
435 not allow for a clear identification of the FZ location. Uncertainties on the measured FZD and  
436 FZW are then on the order of a few tens of centimeters and hundreds of meters in amplitude,  
437 respectively. Uncertainty in the FZW measurement is even larger than the measured value itself,  
438 revealing the arbitrary nature of the FZ location and geometry determination when using the LR  
439 results compared to the HR results. Uncertainty on the FZD and FZW measurements, for the  
440 profiles AA' and BB' respectively, increases by a factor of ~10-100 and of ~2 from the HR to the  
441 LR results (Fig. 4a).

442

443 Analyzing the stacked profiles every 200 m along the study area, 44 profiles in total, we assess the  
444 FZD (Fig. 4b, and Tab. S2) and FZW evolutions (Fig. 4c, and Tab. S2) along the FZ strike.  
445 Measurements derived from the HR results show great consistency (blue and red curves in Fig.  
446 4b,c) with an average difference in the FZD measurements of ~2% between the WorldView and  
447 the SPOT/Pleiades results. Maximum differences along individual profiles reach ~0.5 m, which is  
448 less than 20% of the average FZD ( $2.74 \pm 0.007$  m and  $2.81 \pm 0.009$  m for the WorldView and  
449 SPOT/Pleiades OIC results, respectively). Local differences, in this case, can be explained by  
450 variations in the noise from one dataset to another (e.g., presence of a cloud across the FZ in the  
451 WorldView3; Fig. 3a) which then disrupts the FZD and FZW measurements.

452



453

454 **Figure 4. (a) Cross-fault stacked displacement profiles, in the fault-parallel component,**  
 455 **placed across the FZ in the displacement maps derived from the OIC of different satellite**  
 456 **sensors optical data (Fig. 3a-f). Dashed black lines are the linear regressions, and black**  
 457 **circles are the piercing points used for the displacement offset measurements. FZD ( $\Delta y$ , in**  
 458 **m) and FZW ( $\Delta x$ , in m) are indicated on the left and right side of the profile, respectively.**

459 **(b) FZD budget (m) in the fault-parallel component, derived from the analysis of the 44**

460 profiles, placed every 200 m perpendicularly to the FZ (Tab. S2). (c) FZW (km) measured  
461 for each displacement offset (Tab. S2).

462

463 FZD and FZW curves obtained from the analysis of the LR displacement maps (green, orange, and  
464 pink curves in Fig. 4b,c) show significant differences, sometimes reaching ~2 m for the FZD, and  
465 ~1 km FZW. These differences reflect the increasing uncertainty in assessing the FZ location with  
466 decreasing image resolution, as a result of greater background noise. Uncertainties on the FZD and  
467 FZW measurements increase from 0.8 cm to 42 cm, and from 129 m to 266 m when using the  
468 WorldView and the Sentinel data (raw result), respectively (Fig. 4b,c). In addition, resulting from  
469 the simplified inferred FZ geometry and the un-detected diffuse deformation regions, FZD  
470 measurements derived from the LR OIC results are, on average, under-estimated by >15%  
471 compared to that derived from the HR OIC results. Filtering and smoothing of the Sentinel-2  
472 results allow decreasing the high-frequency noise and improving the recovery of the displacement  
473 by +5% (Fig. 4b,c). However, across the different LR results, including the Sentinel-2 results  
474 before and after filtering and the Planet results, FZW measurements vary considerably around the  
475 HR reference value, obtained from the WorldView results. Therefore, FZW estimated based on  
476 LR imagery are likely unreliable compared to that obtained from HR imagery (Ayorlou et al.,  
477 2021). Still, among the LR results, that obtained from the Planet data at 3.125 m resolution seems  
478 to retrieve better the FZW evolution from narrow to the east to wider to the west of the study area,  
479 suggesting a better sensitivity to FZ complexity compared to the Sentinel-2 data at 10 m resolution.  
480 Finally, average FZW values inferred from the Sentinel-2 measurements before and after filtering  
481 evolve from -25 to +15% around the HR reference, highlighting the influence that post-processing  
482 techniques can have on the FZW assessment in LR results (Fig. 3j,k). Among the datasets tested

483 in this study, HR optical imagery then represents the most reliable source of information for  
484 determining the FZ location, geometry, and associated FZD and FZW.

485

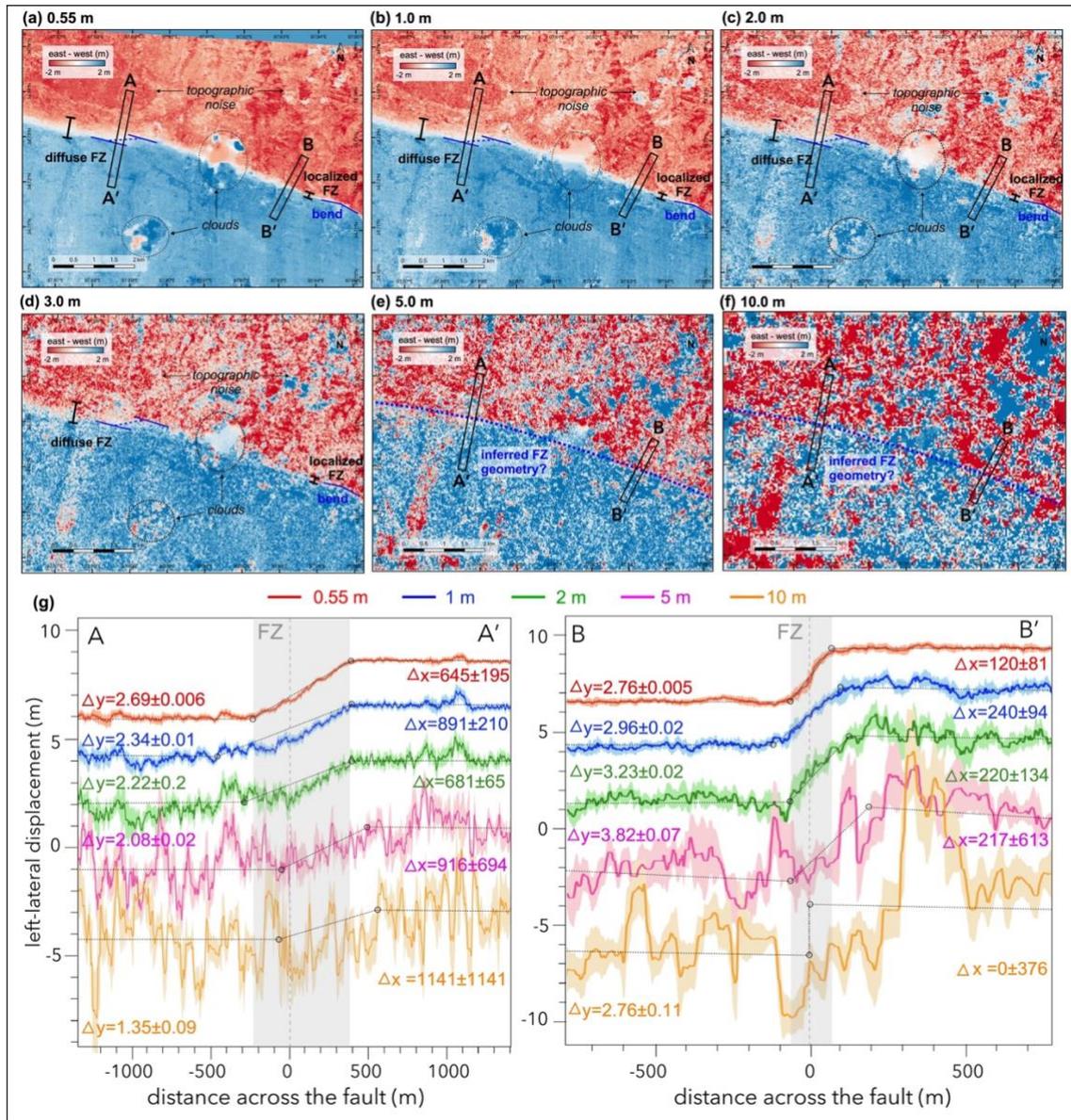
#### 486 **4.1.2. Downsampled pre- and post-earthquake WorldView orthoimages**

487 To separate the effect of data resolution from other data related confounding factors (e.g., sensor  
488 type, acquisition date and geometry, orthorectification quality etc.), we performed OIC on  
489 downsampled versions of the WorldView orthorectified images, at resolutions of 1 to 10 m (Figs.  
490 5, and Tab. S3). This analysis enables cross-examination against our previous observations made  
491 using data from different satellite sensors (Figs. 3 and 4, and Tab. S2).

492

493 Visual inspection of the obtained displacement maps shows an increase in the background noise  
494 with decreasing data resolution (Figs. 5 and 6a), arising both as topography-related noise in the  
495 mountainous area (Fig. 5a-f) and as random high-frequency noise, similar to what was found  
496 previously (Fig. 3). As a result, the identification of the FZ location and geometry is again  
497 challenging in the LR results, especially at resolutions lower than 3 m (Fig. 5). At comparable  
498 resolutions of 3 and 3.125 m, though, OIC results derived from the downsampled WorldView  
499 orthoimages (Fig. 5d) show less noise than those derived from the Planet orthoimages (Fig. 3c).  
500 Such difference most likely relates to the fact that the higher-level Planet data products are  
501 orthorectified using a low-resolution DEM, and without bundle block adjustment (Fig. 2, step 1).  
502 This observations supports the previous inference that sensor and orthorectification quality are  
503 critical parameters for reliable OIC measurements (Antoine et al., 2021, 2022; Leprince et al.,  
504 2007; Shean et al., 2016). Conversely, at a comparable resolution of 10 m, Sentinel-2 OIC results  
505 (Fig. 3d) are less noisy than those obtained using the downsampled WorldView orthoimages (Fig.

506 5f). This difference can arise from i) good Sentinel-2 orthoimage quality, at least compared to the  
 507 Planet data, and ii) possible aliasing in the downsampled products, especially due to large  
 508 downsampling factors applied (García Aranda et al., 2021).  
 509



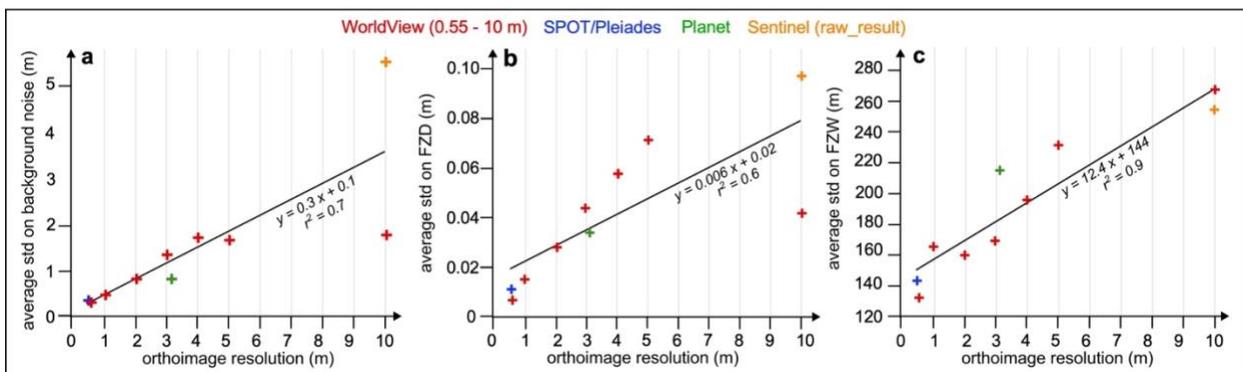
510  
 511 **Figure 5. (a-f) East-west displacement maps calculated from the OIC of (a) WorldView and**  
 512 **(b-f) downsampled WorldView orthoimages. No filtering is applied to the displacement map**  
 513 **outputs, and similar correlation parameters are used for all correlations. FZ and noise**

514 features are highlighted (similar to Fig. 3). (g) AA' and BB' across-fault stacked  
 515 displacement profiles in the fault-parallel component. Similar to Figure 4, FZD and FZW  
 516 values, and associated regression profiles are indicated. Units are meters.

517

518 Profiles taken across the displacement maps show an increase in background noise with decreasing  
 519 orthoimage resolution (Fig. 6a), which is particularly reflected by the std of the FZD and FZW  
 520 measurements. We estimate an increase in the measurement's std with decreasing orthoimage  
 521 resolution by a factor 15 and 20 for the FZD, respectively for AA' and BB', and ~5 for the FZW  
 522 (similar in AA' and BB'). Considering all the measurements together, we find a linear increase of  
 523 the standard deviations (stds) for both the FZD and FZW measurements with decreasing  
 524 orthoimage resolution (Fig. 6b,c). For resolutions greater than 5 m, the signal to noise ratio  
 525 approaches 1 (Fig. 6a), and the estimation of the FZW and associated epistemic uncertainty is  
 526 subjected to the choice from the operator. On average, FZW and FZD estimations are  
 527 underestimated by a factor 0.78 and 0.70, respectively, in the 10-m results as compared to the 0.55  
 528 m results (Fig. S5 and Tab. S3).

529



530

531 **Figure 6.** Evolution of the average standard deviation (std) associated with (a) the  
 532 background noise NE of the FZ (Fig. S4), and the (b) FZD and (c) FZW measurements with

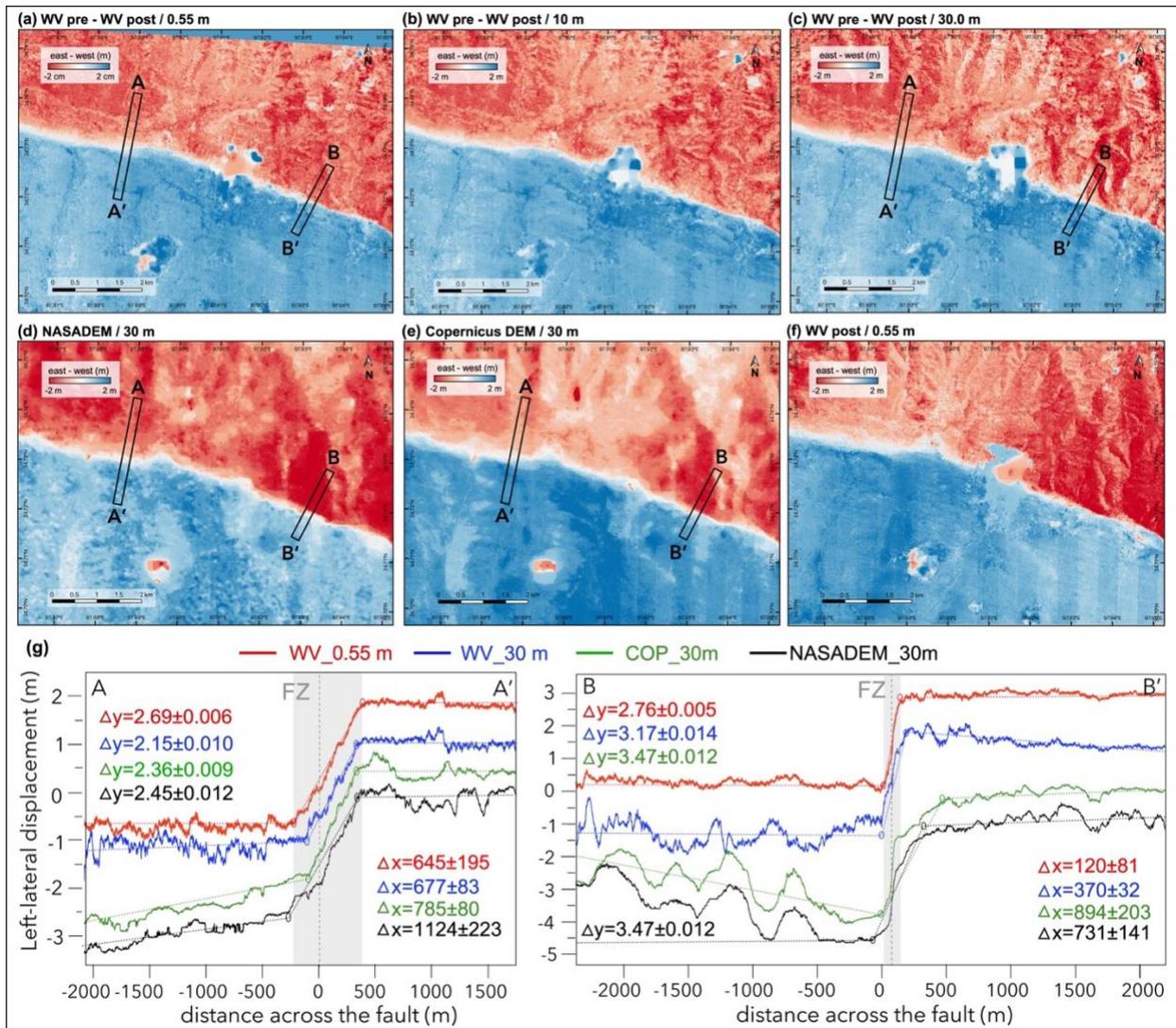
533 **decreasing orthoimage resolution. Points with different colors correspond to the results from**  
534 **the multi-sensor study (Figs. 3 and 4). Red points correspond to the results based on the**  
535 **WorldView orthoimages, both at native and degraded resolutions (Fig. 5). Black lines are**  
536 **linear regressions calculated for all datapoints, with  $r^2$  the associated correlation factor.**

537

#### 538 **4.2. Impact of DSM/DEM resolution on the OIC result**

539 We test the effect of using DEMs/DSMs of different resolutions, acquisition sensors, and dates on  
540 optical image orthorectification and derived OIC quality (Fig. 7). To do so, we performed OIC on  
541 the WorldView1/2/3 image pool orthorectified using different DEMs/DSMs datasets including i)  
542 the high-resolution WorldView DSMs derived from those same images (Fig. 7a), ii) downsampled  
543 versions of these HR DSMs (1-30 m; Figs. 7a-c, and S6b-d) as well as iii) publicly accessible 30-  
544 m-resolution DEMs from the NASA and Copernicus space agencies (Figs. 7d,e, and S6e,f).  
545 Downsampling of the DSMs preserves the native ~0.5 m accuracy of the stereo DSMs across the  
546 range of downsampled resolutions, though it reduces the high-frequency noise (Fig. S6). Thus, we  
547 can test separately the effects of topography data ground resolution (0.55-30 m), vertical accuracy  
548 (~0.5m in the WorldView DSMs, and several meters in the NASA and Copernicus DEMs; see Fig.  
549 S5 and S6), and acquisition time (pre- and/or post-earthquake) on the OIC quality, and FZD and  
550 FZW measurements.

551



552

553 **Figure 7. East-west displacement maps calculated from the OIC of 0.55-m resolution**

554 **WorldView images orthorectified using (a) high-resolution pre- and post-earthquake**

555 **WorldView DSMs, (b-c) downsampled version of the pre- and post-earthquake WorldView**

556 **DSMs, (d-e) the NASA and Copernicus global DEMs (from the pre-earthquake period), and**

557 **(f) the high-resolution post-earthquake WorldView DSM. No filtering is applied to the**

558 **displacement maps, and similar correlation parameters are used for all OIC (see Fig. S2 for**

559 **tests on other parameters). (g) AA' and BB' across-fault stacked profiles in the fault parallel**

560 component across (a) in red, (c) in blue, (d) in black, and (e) in green. Profile annotations are  
561 similar to Figure 4.

562

563 Visual analysis of the displacement maps obtained from the different orthorectified WorldView  
564 products show consistent ground displacement patterns, including the FZ location and geometry  
565 (Fig. 7a-f). This suggests that the effect of the DEM/DSM resolutions (Fig. 2) is less significant  
566 than that of the optical image ground resolution itself though the processing chain. Particularly,  
567 similar patterns are observed across the results based on the downsampled, 1 to 30 m resolution,  
568 stereo DSMs for images orthorectification. These results confirm that the DEM/DSM ground  
569 resolution, as an individual parameter, has minor impact on the orthorectification and OIC quality  
570 (Fig. 7a-c; displacement budget and FZW evolution in Fig. S8). Although larger topography-  
571 related noise appears when using DSMs of resolution >10 m (Fig. 7c), resulting in greater  
572 variability in the FZD and FZW measurement (Fig. 7g). This variation shows no clear correlation  
573 with the DSM ground resolution across the range 10-30 m, and remains on average, within a  
574 confined range of -3 and 0%, and -12 and 0.5% around the HR measurement references for the  
575 FZD and FZW, respectively (Fig. 8, and Tab. S4).

576

577 Displacement maps obtained from the OIC of the WorldView images orthorectified with the  
578 Copernicus (Fig. 7e) and NASA DEMs (Fig. 7f) are in general smoother and with larger  
579 topographic noise than what previously described when using the downsampled WorldView stereo  
580 DSMs. This is particularly visible in the displacement profiles, where low-frequency noise affects  
581 the displacement trends outside of the FZ, resulting in more variable and, on average, over-  
582 estimated of FZD and FZW measurements (Figs. 8 and S9, and Tab. S5). The topography-related

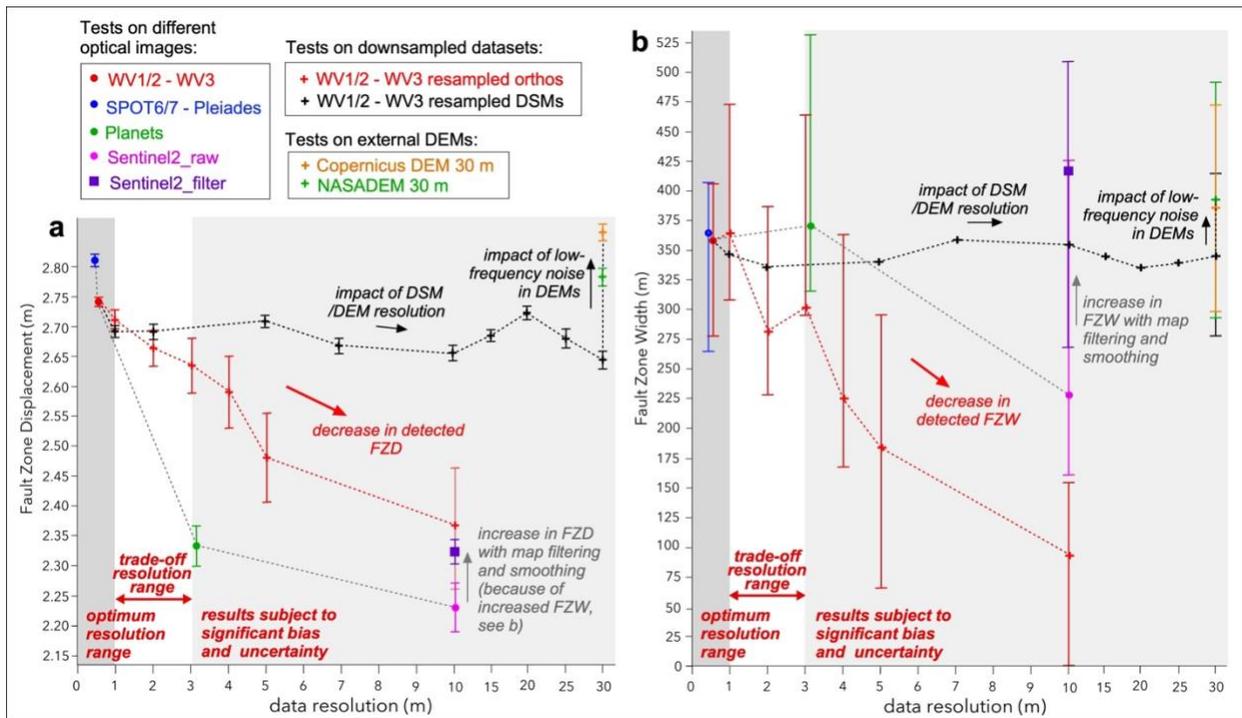
583 noise being consistent with the estimated vertical uncertainty (several meters) of the global DEMs  
584 (Brosens et al., 2022; Florinsky et al., 2018; Uuemaa et al., 2020; Zhou et al., 2018), we suggest  
585 that the DSM vertical accuracy has a direct effect on the OIC result. Noise can also arise from  
586 topography translations on one side of the post-earthquake image compared to the pre-earthquake  
587 image due to the earthquake ground displacement, an effect that was not taken into account when  
588 using the global DEMs acquired before the earthquake event (Fig. 7f). We assess the effect of such  
589 topography translation by performing the orthorectification of the WorldView images using only  
590 the post-earthquake WorldView DSM (Fig. 7f). Results indeed show topography-related noise and  
591 low frequency artefacts in the similar areas as in the results based on the NASA and Copernicus  
592 DEMs, suggesting that part of the noise observed in those three results might arise from the  
593 topography translation due to the earthquake ground displacement. Finally, geographic mis-  
594 alignment between the images and topography datasets, especially when using external source  
595 DEMs that were not processed jointly with the images when refining the camera model (Shean et  
596 al., 2020) can also lead to similar type of noise. Through comparison of the differences between  
597 the pre-earthquake WorldView stereo DSM (derived from the images) and the external DEMs in  
598 the same reference frame (Fig. S7), we assess this effect to be negligible, at least in the case of  
599 the NASADEM (Fig. 7a), although it cannot be fully excluded.

600

## 601 **5. Discussion, and implications towards the development of the Surface Topography and** 602 **Vegetation (STV) Earth Observation System**

603 In this study we investigated the impacts of important characteristics of the optical imagery and  
604 topography data, especially the ground resolution, on accurate imaging of earthquake ground  
605 displacement field. Figure 8 summarizes the average FZD and FZW measurements from the

606 different tests performed in this study. Based on our main observations along with published results  
 607 for other earthquake ruptures (Tab. 1), we highlight important characteristics of optical imageries  
 608 and DEMs/DSMs that should be considered for FZD and FZW measurements using OIC methods,  
 609 and implications for the development of future Earth Surface Topography Observation Systems,  
 610 especially from the fault-related hazard perspective.  
 611



612 **Figure 8. Synthesis of the average measurements of (a) Fault Zone Displacement (FZD) and**  
 613 **(b) Fault Zone Width (FZW) obtained from the different OIC tests performed in this study**  
 614 **(Figs. 3, 5, and 7). Data points include tests using optical imagery originating from different**  
 615 **sensors (colored circles and square, trend in dashed grey), the downsampled WorldView**  
 616 **orthoimages (red crosses, trend in dashed red), the downsampled WorldView DSMs (black**  
 617 **crosses, trend in dash black), and the Copernicus and NASA DEMs (orange and green**  
 618 **crosses, included in the dashed black line). Difference between the WorldView and**  
 619

620 **SPOT/Pleiades FZD measurements at similar resolution can be explained by the presence of**  
621 **a cloud across the FZ in the WorldView3 images (Fig. 4b).**

622

### 623 **5.1. Effect of optical image ground resolution on the measurement of horizontal surface** 624 **displacements using OIC techniques**

625 Using optical images acquired by different satellite sensors along with downsampled versions of  
626 the WorldView dataset, we showed a consistent decrease in the FZD and FZW average  
627 measurements with decreasing image resolution (red and grey curves; Fig. 8). We found that FZD  
628 are underestimated by a factor  $\sim 0.8$  when using  $\geq 10$ -m resolution images (e.g., Sentinel-2)  
629 compared to  $\sim 0.5$ -m resolution images. Under-estimation factor is calculated here as the ratio  
630 between the FZD measured in the LR Sentinel-2 and in the HR ( $\sim 0.5$  m) WorldView results, the  
631 latter considered to be closer to ground truth (Figs. 1b and 3a). Intriguingly, we observe a similar  
632 under-estimation of the FZD, by a factor of  $\sim 0.7$ - $0.8$ , for other earthquake studies that were  
633 documented using both HR and LR optical imagery (Tab. 1). For example, in the case of the 2019  
634 Ridgecrest, California, earthquake sequence, average surface displacement associated with the  
635  $M_w 6.4$  and  $M_w 7.1$  events, respectively, were estimated to be  $0.73 \pm 0.09$  m and  $2.13 \pm 0.06$  m from  
636 WorldView images (Antoine et al., 2021), compared to  $0.55 \pm 0.08$  m and  $1.60 \pm 0.22$  m from  
637 Sentinel-2 images (Chen et al., 2020). Similarly, in the case of the 2013 Baluchistan, Pakistan,  
638 earthquake, average FZD derived from WorldView data is 8.3 m (Gold et al., 2015) compared to  
639 only 6 m using Landsat-8 images at 15 m resolution (Avouac et al., 2014). The fact that similar  
640 under-estimation factor on FZD measurements from LR compared to HR optical imagery is  
641 observed for different earthquakes, where different images, orthorectification strategies and OIC  
642 algorithms were used (e.g., Fig. S2), is suggesting that the optical imagery ground resolution is the

643 primary controlling factor on the resultant OIC quality, and the derived FZD and FZW  
644 measurements accuracy. In the case of Maduo, however, average FZD estimated for the entire  
645 rupture length (~160 km) were similarly estimated from both the LR and HR data. The 2021  
646 Maduo surface deformation was shown to be primarily diffuse and to occur over an average ~600  
647 m FZW (Antoine et al., 2024; C. Li et al., 2022). Therefore, the optical imagery resolution effect  
648 on the FZD and FZW measurement accuracy is especially significant in the case of localized and  
649 narrow surface deformation features.

650

651 **Table 1. Comparison of FZD measurements for the 2021 Maduo rupture with other**  
652 **continental earthquakes.**

| Earthquake  | Sensor                 | Resolution (m) | Average FZD (m) | Reference                       |
|---|------------------------|----------------|-----------------|---------------------------------|
| 2021 Maduo<br>(M <sub>w</sub> 7.4)                      | Sentinel2B             | 10             | 2.27            | Li et al., 2022                 |
|   | SPOT6/7/<br>Pleiades   | 0.5            | 2.35 ± 0.09     | Antoine et al., 2024            |
| 2013<br>Baluchistan<br>(M <sub>w</sub> 7.8)             | Landsat8               | 15             | 6               | Avouac et al., 2014             |
|   | Landsat8               | 15             | 6.7 +0.3/-0.4   | Zinke et al., 2014              |
|   | SPOT5                  | 2.5            | 8               | Vallage et al., 2015            |
|   | WorldView              | 0.5            | 8.3             | Gold et al., 2015               |
| 2019<br>Ridgecrest<br>foreshock<br>(M <sub>w</sub> 6.4) | Sentinel2B             | 10             | 0.55 ± 0.08     | Chen et al., 2020               |
|   | Planet                 | 3.125          | 0.56 ± 0.10     | Milliner and<br>Donnellan, 2020 |
|   | WorldView/<br>Pleiades | 0.5            | 0.73 ± 0.09     | Antoine et al., 2022            |
| 2019<br>Ridgecrest<br>mainshock<br>(M <sub>w</sub> 7.1) | Sentinel2B             | 10             | 1.60 ± 0.22     | Chen et al., 2020               |
|   | Planet                 | 3.125          | 1.68 ± 0.19     | Milliner and<br>Donnellan, 2020 |
|   | WorldView/<br>Pleiades | 0.5            | 2.13 ± 0.06     | Antoine et al., 2022            |

653

654 From these results, we suggest that the optimal image resolution for accurately measuring the FZD  
655 and FZW as well as deciphering the localized versus diffuse nature of surface deformation lies

656 between  $<0.5$  and  $\sim 1$  m, for earthquakes of  $M_w > 6.4$  (Fig. 8 and Tab. 1). In this study, starting from  
657 a resolution of 2 m, background noise approaches  $\pm 1$  m (Fig 6a), making the FZ identification  
658 more challenging (Figs. 3 and 5) and the FZD and FZW measurements less accurate (Figs. 4, 5g,  
659 6b,c, and 7g). For resolution greater than 3 m, FZ geometry cannot be clearly identified anymore.  
660 Provided an accurate image orthorectification, image resolutions of 1-3 meters can represent an  
661 acceptable trade-off range for estimating the FZD and FZW, but not the detailed FZ structure (Fig.  
662 3c versus Fig. 5c). Moreover, this trade-off resolution range would be mostly beneficial only to  
663 medium to large magnitude earthquakes ( $M_w > \sim 6.4$ ) that generate at least 0.5-1 m of ground  
664 displacement (Fig 6a). Detecting surface displacements of amplitudes  $<0.5$  m, for example  
665 associated with distributed and diffuse processes and/or smaller magnitude events, requires sub-  
666 meter resolution imagery. At present, there is a lack of publicly available HR optical imagery on  
667 a global scale, presenting an opportunity that can be addressed by future Earth Surface Observation  
668 Missions such as NASA Surface Topography and Vegetation (STV) (Donnellan et al., 2021).

669

## 670 **5.2. Effect of DEM/DSM ground resolution and vertical accuracy on the measurement of** 671 **earthquake surface displacement**

### 672 *5.2.1. Effect on the measurement of horizontal surface displacements using OIC techniques*

673 In this study, we showed that the effect of the DEM/DSM ground resolution alone was not  
674 significant on the quality of OIC results (Fig.5a-c,g and black curve in Fig. 8) in comparison with  
675 that of the DEM/DSM vertical accuracy, the latter which can relate to or be independent of the  
676 ground resolution depending on the sensor type (Radar, optical, LiDAR). In general LR global  
677 baseline DEMs derived from Radar observations have vertical errors of several meters or tens of  
678 meters, especially in mountainous areas (Brosens et al., 2022; Florinsky et al., 2018; Uuemaa et

679 al., 2020; Zhou et al., 2018). In comparison HR stereo-derived DSMs tend to have an error about  
680 1 meter or below (e.g., Hu et al., 2016; Rupnik et al., 2018; Wang et al., 2019; Zhou et al., 2015),  
681 and usually permit the best possible image orthorectification and OIC result (Figs. 1a and 7a-c).  
682 Therefore, even though global DEMs provide sufficient ground geometry reference for most  
683 applications, the DSMs derived from HR stereo imageries, ideally available for both the pre- and  
684 post-earthquake periods, represent a preferred topography dataset for OIC-based FZD mapping.  
685 Such highly accurate DEM/DSM products can also be obtained using LiDAR data (e.g., Donnellan  
686 et al., 2017; Scott et al., 2020). However, acquiring dense (several points per meter) LiDAR  
687 observations on a global scale is a technical challenge especially regarding the energy supply to  
688 measurement ratio. For these reasons, we suggest HR stereo optical imagery to be a good  
689 candidate for global topography measurements for future NASA Surface Topography and  
690 Vegetation (STV) Earth Observation System (Donnellan et al., 2021).

691

### 692 *5.2.2. Implications on vertical displacement measurements using topography differencing* 693 *methods*

694 Near-fault vertical displacement measurements, similar to the horizontal measurements previously  
695 documented, represent a crucial information to constraining the rupture processes, both at depth  
696 and at the surface (Antoine et al., 2023; Lauer et al., 2020). Such information is especially crucial  
697 for determining possible fault dip angle variations (Teran et al., 2015; Vallage et al., 2015), stress  
698 rotations (Milliner et al., 2022) and slip deficit (Antoine et al., 2023; Fialko et al., 2005) in the  
699 shallow crust. Again, sub-meter resolution stereo optical imagery demonstrated its potential for  
700 measuring sub-meter vertical displacements of from the comparison of the pre- and post-  
701 earthquake stereo derived DSMs (e.g., Antoine et al., 2022, 2021; Delorme et al., 2020; Teran et

702 al., 2015; Zhou et al., 2015). Applied on a global scale, submeter resolution stereo imagery then  
703 would allow for documenting complex and/or small amplitude vertical displacements over a wide  
704 range of geological and tectonic contexts and earthquake magnitudes, allowing to refine  
705 earthquake source models and our understanding of the shallow rupture processes (Antoine et al.,  
706 2023; Marchandon et al., 2021; Xu et al., 2016).

707

708 Vertical accuracy of stereo imagery DSMs and derived vertical displacement maps is proportional  
709 to the image resolution (Rupnik et al., 2018; Schumann and Bates, 2018; Fig. S10). Submeter  
710 resolution optical imagery is thus a prerequisite for accurate measurement of vertical ground  
711 displacements based on topography differencing methods. The acquisition geometry of the stereo  
712 imagery, such as the relative viewing angles between different images, is another limiting factor  
713 of the quality of the stereo DSMs and derived topography change products (Fig. S10), especially  
714 when using archive images that are not acquired with optimal viewing angles. Future stereo  
715 imagery systems then need to address these acquisition geometry requirements, which can vary  
716 depending on the topography amplitude and roughness of the area to be imaged (Hasegawa et al.,  
717 2000; Hu et al., 2016; Loghin et al., 2020). Other confounding factors such as the presence of  
718 vegetation could also limit the quality of ground change measurements. However, ongoing  
719 research starts to show the potential of overcoming these limitations and recovering the ground  
720 surface geometry in sparse vegetated regions (e.g., Yin et al., 2023). The vegetation issue can also  
721 be mitigated by combining optical imagery with dense HR LiDAR data. As mentioned before,  
722 Lidar is able to map 3-D ground surface with high resolution and high vertical accuracy. However,  
723 such data are generally acquired locally from airborne campaigns because of the greater technical  
724 challenge and cost constraints associated with global dense LiDAR acquisitions.

725

726 **Conclusions**

727 This study provides a quantitative assessment of the effects of optical imagery and topography  
728 data characteristics, primarily ground resolution, on the measurement of near-fault earthquake  
729 surface displacements using the 2021  $M_w7.4$  Maduo, Tibet, event as a case study. Our objective is  
730 to provide measurement requirements and viable technology suggestions, through analysis of  
731 existing data and capabilities, to inform the development of future Surface Topography and  
732 Vegetation (STV) Earth observation system from the Solid Earth and earthquake hazard  
733 perspectives. This study uses satellite optical imagery and DEMs acquired by different satellite  
734 agencies and with different resolutions to measure ground surface displacements through sub-pixel  
735 cross-correlation of the orthorectified images (OIC). The main observations from this study are  
736 summarized as follows:

737 - Noise in the ground displacement maps increases linearly with decreasing optical imagery  
738 resolution, resulting in greater uncertainty in the measured displacement amplitudes across the  
739 fault zone.

740 - We infer an under-estimation of the measured displacements across the fault zone by a factor  
741 0.7-0.8 when using low-resolution ( $>10$  m) compared to high-resolution ( $\leq 1$  m) imagery. This  
742 factor is independent of the processing method and has been inferred for multiple satellite  
743 observations on different earthquake case studies.

744 - The use of submeter-accuracy DEM/DSMs in both the pre- and post-earthquake periods allows  
745 for reducing topography related noise in the surface displacement products, and is necessary for  
746 assessing the earthquake vertical displacements through topography differencing.

747 - We suggest that high resolution ( $\leq 1$  m) optical imagery and derived DSM products represents a  
748 preferred dataset for accurately measuring the ground displacements, both in the horizontal and  
749 vertical components, across active fault zones. Especially, the high resolution ( $\leq 1$  m) is necessary  
750 to resolve complex fault geometries, as well as untangle the diffuse, distributed or localized nature  
751 of the surface deformation.

752 - High-resolution ( $< 0.5$  m) stereo optical imagery and associated photogrammetry and OIC  
753 techniques have a unique potential for 3-D analysis of the ground surface characteristics and  
754 change, and present themselves as a good candidate for the future STV Earth observation system  
755 from earthquake hazard perspectives and for other interdisciplinary applications areas such as  
756 volcanos and landslide monitoring, geomorphology and vegetation analysis, and cryosphere.

757

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764

## 765 **Open Research**

766 MicMac (<https://github.com/micmacIGN/micmac>; last accessed on 01/03/24) and StackProf  
767 (<https://github.com/IPGP/stackprof>; last accessed on 01/03/24) are open source. WorldView and  
768 Planet data were accessed through the Commercial Smallsat Data Acquisition ([CSDA](#)) program  
769 of the National Aeronautics and Space Administration

770 (<https://www.earthdata.nasa.gov/esds/csda/commercial-datasets>; last accessed on 01/03/24).

771 Supplementary figures providing details on the methodology and on the results of this study, along  
772 with surface displacement maps and fault displacement and width measurements will be available  
773 for the published version of the manuscript.

774

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