

1 **Occurrence and causes of large dB/dt events in the**
2 **pre-midnight and dawn sectors**

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11 **Key Points:**

- 12 • We investigate the causes of large dB/dt “spikes” in ground magnetometer data
13 in the pre-midnight and dawn sectors
- 14 • Pre-midnight and dawn spikes occur independently; pre-midnight spikes are as-
15 sociated with substorm onsets
- 16 • Dawn spikes are associated with sunward-propagating, azimuthally-spaced auro-
17 ral features embedded in the westward electrojet

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Abstract

We investigate the causes of large dB/dt events observed by SuperMAG, by comparing with the time-series of different types of geomagnetic activity, or “convection state”, for the duration of 2010. Spikes are found to occur predominantly in the pre-midnight and dawn sectors. We find that pre-midnight spikes are associated with substorm onsets. Dawn sector spikes are not directly associated with substorms, but with auroral activity occurring within the westward electrojet region. Azimuthally-spaced auroral features drift sunwards, producing Ps6 (10-20 min period) magnetic perturbations on the ground. The magnitude of dB/dt is determined by the flow speed in the convection return flow region, which in turn is related to the strength of solar wind-magnetospheric coupling. Pre-midnight and dawn sector spikes can occur at the same time, as strong coupling favours both substorms and westward electrojet activity; however, the mechanisms that create them seem somewhat independent. The dawn auroral features share some characteristics with omega bands, but can also appear as north-south aligned auroral streamers. We suggest that these two phenomena share a single underlying cause.

Plain Language Summary

Sudden changes in electrical currents flowing in the auroral ionosphere produce magnetic perturbations on the ground, large dB/dt “spikes”, which can lead to damaging Geomagnetically Induced Currents (GICs) in technological infrastructure. It is known that these spikes occur preferentially in the pre-midnight local time sector and near dawn, though their exact cause is still poorly understood. In this study we show that pre-midnight spikes are associated with a well-known auroral phenomenon known as a substorm which produces sudden southward-directed perturbations in the magnetic field. Dawn spikes are associated with east-west spaced auroral features which drift eastwards and produce wave-like fluctuations in the magnetic field in the east-west direction. The magnitude of dB/dt will depend on the speed at which the forms drift. Pre-midnight and dawn spikes are not directly related, but can occur together when geomagnetic activity is high.

1 Introduction

Geomagnetically Induced Currents (GICs), which negatively impact geographically-extended power-grids and pipelines, are caused by rapid fluctuations in electrical currents flowing in the ionosphere which in turn produce sudden changes in the magnetic field on the ground, also known as large dB/dt events or “spikes”. These ionospheric currents are a component of the electrical circuit which transmits stress from the magnetosphere to the ionosphere as part of the Dungey cycle of convection (Dungey, 1961; Milan et al., 2017), driven by the solar wind-magnetosphere interaction. This study investigates the link between different types of geomagnetic activity, or “convection state”, and the occurrence of large dB/dt events. We identify two main hot-spots of spike occurrence, in the pre-midnight and dawn sectors, which occur for different convection states.

Disturbances of the magnetosphere are produced by the solar wind interaction, so the occurrence of large dB/dt displays dependencies on solar wind conditions and hence the solar cycle (e.g., Milan, Imber, et al., 2023). Magnetospheric activity encompasses a range of behaviours including geomagnetic storms (e.g., Gonzalez et al., 1994), substorms (e.g., Akasofu, 1964), periods of “steady magnetospheric convection” (e.g., Sergeev et al., 1996), intervals of prolonged, repeated auroral activity, termed “multiple intensifications” (Milan et al., 2021), and “high-intensity long-duration continuous AE activity” (HILDCAA) which occurs during high-speed solar wind streams (e.g., Tsurutani & Gonzalez, 1987; Milan, Mooney, et al., 2023). Most of these phenomena have been implicated in the occurrence of dB/dt (e.g., Juusola et al., 2015; Schillings et al., 2022; Milan, Imber, et al., 2023). Different current systems will be responsible for the generation of spikes at different local times during different types of activity. Although there is still

68 debate over exactly which phenomena are responsible, candidates include the substorm
 69 current wedge in the pre-midnight sector, a dawn-side current wedge, omega bands in
 70 the dawn sector, and Kelvin-Helmholtz wave activity in the pre-noon sector (e.g. Ap-
 71 atenkov et al., 2020; Engebretson et al., 2020; Juusola et al., 2015; Kataoka & Pulkki-
 72 nen, 2008; Milan, Imber, et al., 2023; Ngwira et al., 2018; Pulkkinen & Kataoka, 2006;
 73 Sorathia et al., 2023; Weigel et al., 2002, 2003; Zou et al., 2022).

74 We investigate the causal mechanisms of GICs statistically by cross-referencing the
 75 occurrence of dB/dt spikes with the nature of the on-going geomagnetic activity. Milan
 76 et al. (2021) identified the sequence of different types of activity during the year 2010,
 77 and in this study we compare this sequence with the occurrence of spikes determined from
 78 SuperMAG ground magnetometer measurements (Gjerloev, 2012). Although 2010 was
 79 a year of low solar activity, and hence relatively weak geomagnetic activity, the ability
 80 to compare spike occurrence with convection state provides useful insights into spike for-
 81 mation, highlighting the difference between the causes of spikes in the pre-midnight and
 82 dawn sectors. We then further study the difference between pre-midnight and dawn spikes
 83 using two case studies from 2001 and 2010.

84 2 Observations

85 2.1 Convection states

86 As described by Milan et al. (2021), the variation in geomagnetic activity type —
 87 or magnetospheric “convection state” — for the duration of 2010 was determined by con-
 88 sidering the behaviour of the orientation and magnitude of the interplanetary magnetic
 89 field (IMF), the auroral upper and lower electrojet indices AU and AL (Davis & Sug-
 90 iura, 1966), the polar cap or PC index (Troshichev et al., 2006), and the latitude of the
 91 region 1 and 2 field-aligned currents (FACs) measured by the Active Magnetosphere and
 92 Planetary Electrodynamics Response Experiment or AMPERE (Anderson et al., 2000;
 93 Waters et al., 2001). The latitude of the FACs was determined by fitting a circle to the
 94 boundary between the region 1 and region 2 currents (Milan et al., 2015; Milan, 2019),
 95 of which the radius, Λ , was used as a proxy for the open or polar cap flux content of the
 96 magnetosphere, F_{PC} . Convection state is largely determined by the occurrence of day-
 97 side (magnetopause) reconnection and nightside (magnetotail) reconnection, quantified
 98 by Φ_D and Φ_N , which can occur together or independently (Cowley & Lockwood, 1992;
 99 Milan et al., 2007), as shown in Figure 1 of Milan, Mooney, et al. (2023). Reconnection
 100 rates are difficult to measure directly (Chisham et al., 2008; Hubert et al., 2006), so a
 101 proxy for Φ_D based on upstream solar wind speed and the interplanetary magnetic field
 102 (IMF), Φ_D^* , is used (Milan et al., 2012). The following convection states were identified
 103 by Milan et al. (2021) (more details can be found in that paper):

- 104 • Quiet: little or no geomagnetic activity;
- 105 • Weak Activity: geomagnetic activity but with no substorm or other signatures;
- 106 • Substorm Growth Phase: unbalanced dayside reconnection causing an expansion
 107 of the polar cap;
- 108 • Substorm Expansion Phase: a substorm bay (negative excursion) in the AL in-
 109 dex;
- 110 • Substorm Driven Phase: balanced dayside and nightside reconnection leading to
 111 steady magnetospheric convection;
- 112 • Substorm Recovery Phase: unbalanced nightside reconnection leading to a con-
 113 traction of the polar cap;
- 114 • Recovery Bay: recovery phases accompanied by a bay in AL;
- 115 • Multiple Intensifications: prolonged, rapid quasi-periodic variations in AL, usu-
 116 ally during strong solar wind driving, and associated with an expanded auroral
 117 oval.

118 The time-series of geomagnetic activity type for 2010 is available as Milan (2020).

119 Examples of these states are presented in Figure 1, which shows a period of 42 hours
 120 of observations from 18 UT on 28 May 2010. Panels *b* and *c* show cuts of AMPERE-derived
 121 FAC density along the dawn-dusk meridian in the northern and southern hemispheres,
 122 colour-coded with red and blue as upwards and downwards FACs, respectively. Panel
 123 *d* shows the radius of the FACs, Λ , in the northern and southern hemispheres. Panel *e*
 124 shows the AU and AL indices, f the PC-N index and Φ_D^* , g the components of the IMF,
 125 and h the Sym-H index.

126 Vertical lines show the boundaries between different convection states, and letters
 127 at the top label these intervals. A, H, and K are quiet. B, I, and L are growth phases
 128 ($\Phi_D > 0$, Λ increasing). C and M are expansion phases (substorm bay in AL, often ac-
 129 companied by a decrease in Λ). D, F, and N are driven phases ($\Phi_D > 0$ but steady Λ
 130 indicating $\Phi_N \approx \Phi_D$). J and O are recovery phases ($\Phi_D \approx 0$ but Λ decreasing indi-
 131 cating $\Phi_N > 0$). G is a recovery bay (a recovery phase but with a substorm bay in AL)
 132 and P is a period of weak activity. Vertical red dashed lines during intervals E and N
 133 identify events that were termed driven phase onsets by Milan et al. (2021). Driven phases
 134 are periods of balanced dayside and nightside reconnection, and driven phase onsets are
 135 a new substorm onset — presumably associated with the formation of a new near-Earth
 136 neutral line and development of a new substorm current wedge (Milan et al., 2021) —
 137 during already ongoing magnetotail reconnection. Interval E is the type of activity iden-
 138 tified as multiple intensifications by Milan et al. (2021), due to the rapid fluctuations in
 139 AL, and is of especial interest to the present study. The 10-20 min quasi-periodic fluc-
 140 tuations in AL are manifestly different from the substorm growth/expansion behaviour
 141 of states B and C, where a substorm cycle typically lasts 3 hours, and the driven onsets
 142 in intervals E and N. True substorm bays in AL are highlighted by green arrows in panel
 143 *e*.

144 One-minute cadence ground magnetometer data were accessed via the SuperMAG
 145 repository (Gjerloev, 2012). The data have the background field subtracted and are then
 146 rotated into N (North-South, positive northwards), E (East-West, positive eastwards),
 147 and Z (vertical, positive downwards) components. As described by Schillings et al. (2022)
 148 and Milan, Imber, et al. (2023), dB/dt was defined as the minute-on-minute difference
 149 between successive N , E , and Z values, denoted ΔN , ΔE , and ΔZ . We use ΔB to de-
 150 note a spike in any field component. The occurrence and MLT location of spikes with
 151 $|\Delta B| > 150$ nT (that is, $dB/dt > 150$ nT min⁻¹) during the interval encompassed by
 152 Figure 1 is shown in panel *a*. We return to a discussion of these observations later. First
 153 we investigate the occurrence of spikes for the whole of the year 2010.

154 2.2 Statistical study of spike occurrence for 2010

155 Figure 2 presents the magnetic local time (MLT) and magnetic latitude of spikes
 156 with $|\Delta B| > 250$ nT occurring during the growth, expansion, driven, multiple inten-
 157 sifications, recovery, and recovery bay convection states of 2010. Very few spikes are ob-
 158 served during the quiet and weak activity states, so these are omitted. In each panel,
 159 the name of the state is followed by the percentage of time that the magnetosphere spent
 160 in that state during 2010. The number of spikes identified in each state is shown, along
 161 with the percentage of all spikes that this number represents. Finally, the rate of occur-
 162 rence of spikes during that state is shown as the number of spikes per day. Each occur-
 163 rence of a spike is indicated by a coloured dot. The colour-coding represents whether the
 164 spikes tended to occur near the start, near the end, or in the middle of each convection
 165 state (“elapsed time”). A polar histogram in each panel shows the occurrence distribu-
 166 tion of spikes in 1 hour wide MLT bins, on the logarithmic scale shown in the bottom
 167 right panel.

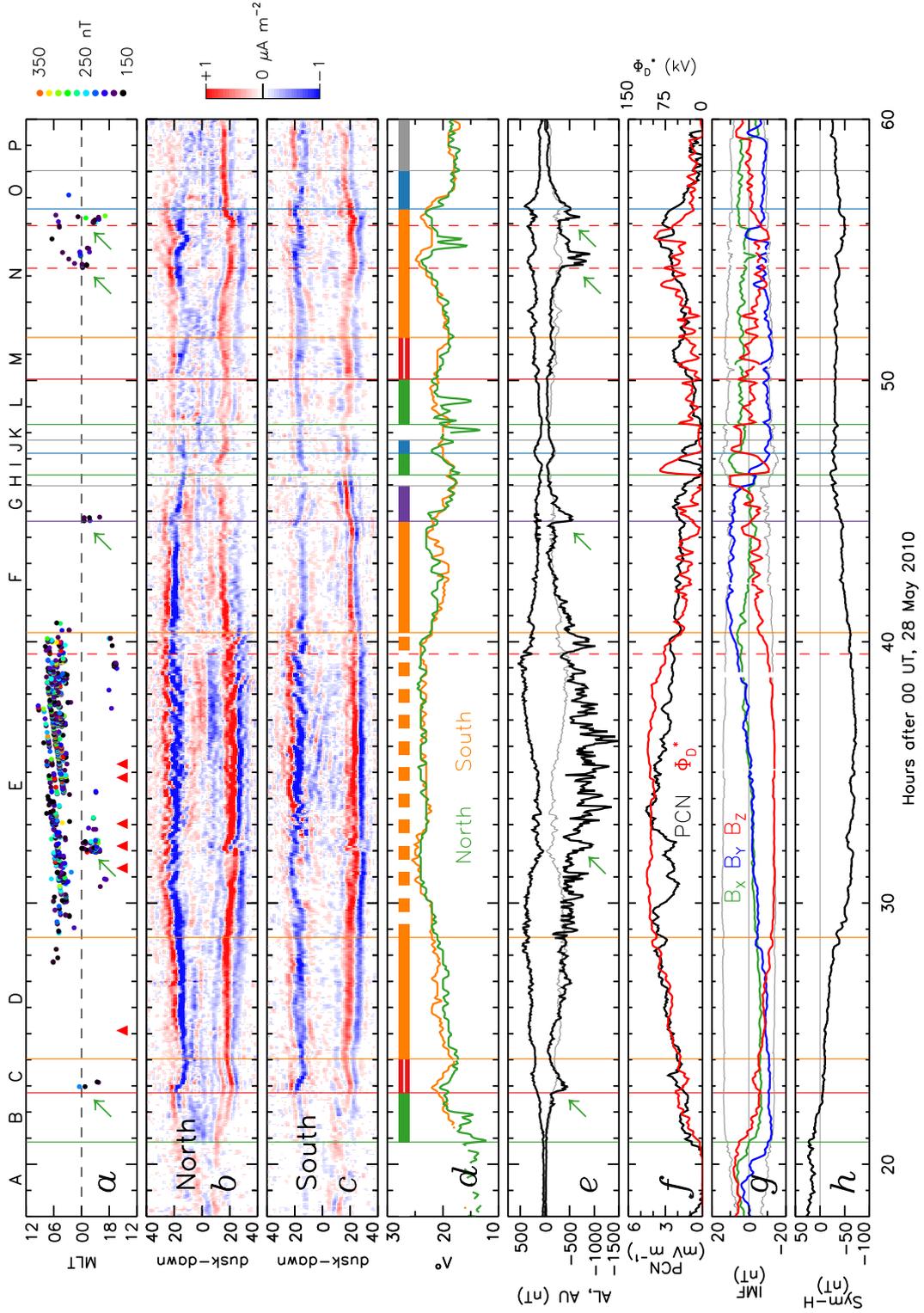


Figure 1. Spike occurrence and context on the 28 to 30 May 2010. Panel *a*: the MLT occurrence of $\frac{dB}{dt} > 150 \text{ nT min}^{-1}$. Panels *b* and *c*: AMPERE observations of field-aligned currents (FACs) along the dusk-down meridians of the northern and southern hemispheres; dawn is towards the top in each case. Panel *d*: the radii, Λ , of the FAC rings in the two hemispheres. Panel *e*: the AU and AL indices. Panel *f*: the PC-N index and Φ_D^* . Panel *g*: the B_x , B_y , and B_z components of the interplanetary magnetic field. Panel *h*: the Sym-H index. Green arrows in panels *a* and *e* highlight the onset of substorms.

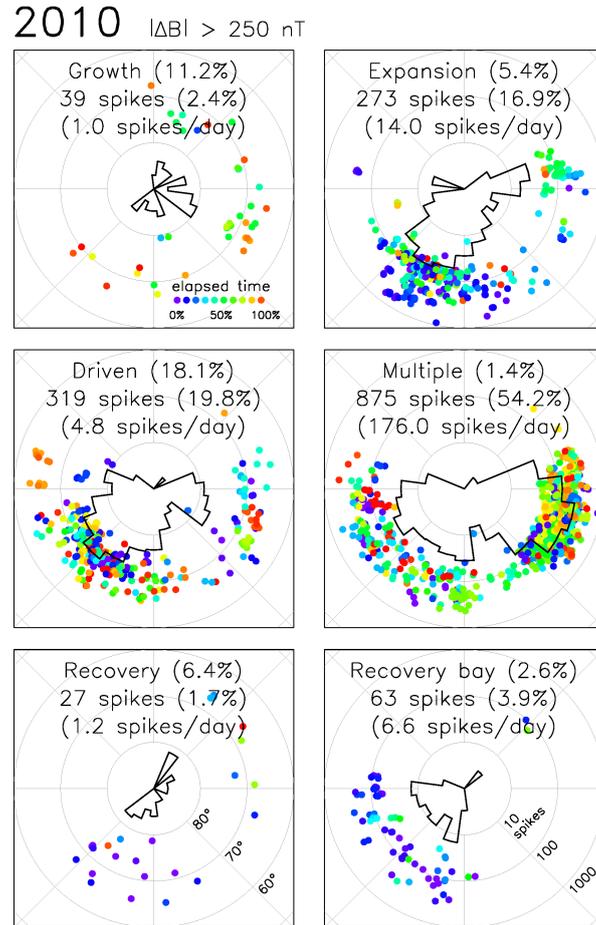


Figure 2. Magnetic local time and magnetic latitude occurrence distribution of $dB/dt > 250$ nT min⁻¹ during different convection states in 2010. Noon is to the top and dawn to the right; grey circles indicate magnetic latitudes in steps of 10°. The events are colour-coded by their time of occurrence within each state. A polar occurrence histogram is shown within each panel, on the logarithmic scale shown in the bottom-right panel.

Overall, spikes occur in two distinct MLT ranges, which we refer to as the pre-midnight (19 to 00 MLT) and dawn (03-07 MLT) sectors. These well-known “hot-spots” have previously been identified by a number of authors (e.g., Juusola et al., 2015; Kataoka & Pulkkinen, 2008; Viljanen et al., 2001; Weigel et al., 2002; Schillings et al., 2022; Milan, Imber, et al., 2023). Here we show that the occurrence of these hot-spots is strongly dependent on convection state. Few spikes are observed during the growth phase (2.4% of the total); those that are observed tend to occur late in the growth phase, so are probably associated with the subsequent expansion phase. A more significant number of spikes (16.9%) is seen in the expansion phase, and usually near the start of this phase, suggesting that they are associated with substorm onset. These are concentrated in the 21 to 00 MLT sector, implying that they are caused by the westward electrojet component of the substorm current wedge. More spikes (19.8%) are seen during the driven phase, with no preference for the elapsed time. These are mainly observed in the pre-midnight sector, but with a small number at dawn. Most of the pre-midnight spikes will be associated with driven phase onsets. The majority of spikes (54.2%) are observed during multiple intensifications, in both the pre-midnight and dawn sectors, but concentrated in the dawn sector (76% in the 01-07 MLT sector as opposed to 23% in the 17-00 MLT sector). Few spikes are seen during recovery phases (1.7%), but there is a small number of spikes associated with recovery bays (3.9%), almost exclusively in the pre-midnight sector. Our main conclusion is that pre-midnight spikes are associated with substorms while dawn spikes are strongly correlated with the multiple intensifications convection state identified by Milan et al. (2021). Although negative excursions in AL are usually assumed to be associated with substorm activity in the pre-midnight sector, in these cases they are caused by westward electrojet activity in the dawn sector. Moreover, while multiple intensifications occupy less than 2% of the time, they produce over 54% of spikes.

We now focus on spikes with $\Delta B > 250$ nT occurring during expansion phase, driven phase, multiple intensifications, and recovery bays. The distributions of spikes during these states are presented in Figure 3, this time colour-coded with respect to spike magnitude (ΔB), the spike component (N, E, or Z), dayside reconnection rate (Φ_D^*), solar wind speed (V_{SW}), IMF B_Z , and the Sym-H index (Iyemori, 1990).

Spike magnitude does not appear to be strongly dependent on convection state nor local time (panels *a*, *g*, *m*, *s*). Magnitudes range from the threshold, 250 nT, up to 450 nT, though with fewer at higher values. (Bear in mind that 2010 was a quiet year and we would expect larger spikes in more active years.) Spikes in all three components can occur in all states at all MLTs, though *N* spikes are favoured in the pre-midnight sector and *E* spikes at dawn (panels *b*, *h*, *n*, *t*), as previously reported (e.g., Schillings et al., 2022; Milan, Imber, et al., 2023).

During expansion phase most spikes occur in the pre-midnight and midnight sectors, mainly between 21 and 02 MLT. These occur predominantly for non-storm times, typically Sym-H > -20 nT (panel *h*), though they occur at lower latitudes for high V_{SW} and more negative IMF B_Z , and hence also for high Φ_D^* (panels *c*, *d*, *e*). A secondary population at 06 to 08 MLT is seen during expansion phases. The most notable aspect of these is that they occur for high V_{SW} (panel *d*), so they resemble the Kelvin-Helmholtz instability spike population described by Weigel et al. (2003) and Milan, Imber, et al. (2023). It is unclear if it is a coincidence that this population appears during expansion phases only (though there is a hint of this high- V_{SW} population during driven phases (panel *j*) also).

Most driven phase spikes occur in the pre-midnight sector, mainly between 19 and 01. This MLT distribution is perhaps shifted to slightly earlier local times than than the expansion phase population. These spikes appear at lower latitudes for elevated Sym-H ≈ -60 nT and $\Phi_D^* \approx 100$ kV (panels *l* and *i*, respectively).

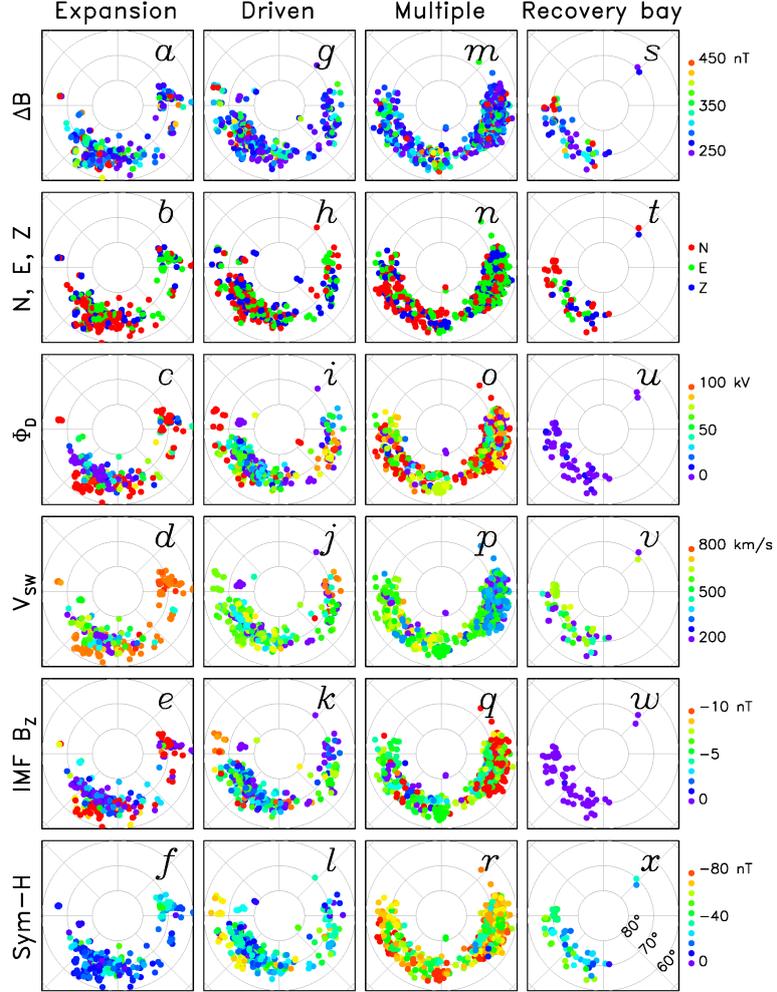


Figure 3. Occurrence distribution of $dB/dt > 250 \text{ nT min}^{-1}$ in a similar format to Fig. 1. The colour-coding refers to the magnitude of dB/dt , ΔB , the component in which ΔB occurs, reconnection rate Φ_D^* , solar wind speed V_{SW} , IMF B_z , and the Sym-H index.

219 The multiple intensification spikes occur at pre-midnight and dawn, though mainly
 220 dawn as described above. This convection state occurs during strong driving (Milan et
 221 al., 2021), and as a consequence Sym-H tends to be elevated, $\text{Sym-H} < -40$ nT (panel
 222 *r*). However, there are distinct differences in the solar wind conditions for the pre-midnight
 223 and dawn populations: pre-midnight spikes are favoured by faster solar wind but lower
 224 values of B_Z , whereas dawn spikes occur for moderate V_{SW} but more negative B_Z (pan-
 225 els *p* and *q*).

226 Finally, recovery bay spikes, being a recovery phase phenomenon, occur for low val-
 227 ues of Φ_D^* , moderate solar wind speed (≈ 500 km s $^{-1}$) and slightly elevated Sym-H \approx
 228 -40 nT (panels *u*, *v*, *x*). These spikes occur between 18 and 00 MLT, a somewhat dif-
 229 ferent distribution than the substorm onset spikes. It is still unclear what leads to re-
 230 covery bays (Milan et al., 2021).

231 2.3 Two case studies of pre-midnight and dawn spikes

232 2.3.1 28 to 30 May 2010

233 We now refer back to Figure 1, in which panel *a* shows the MLT location of spikes
 234 with $\Delta B > 150$ nT. In this example, highlighted by green arrows, weak spikes are seen
 235 in the pre-midnight sector associated with the expansion phase C, with the recovery bay
 236 G, and with the two driven phase onsets indicated by vertical red dashed lines during
 237 interval N. However, the majority of spikes are seen in the dawn sector during state E.
 238 A cluster of spikes is also seen in the pre-midnight sector at 32 hours, highlighted by a
 239 further green arrow, which is associated with a substorm onset occurring during the on-
 240 going period of dawn activity. Green arrows in panel *e* highlight the substorm bays in
 241 AL that are associated with the pre-midnight activity; other fluctuations in AL are pro-
 242 duced by dawn sector activity.

243 Panel *c* shows that the southern hemisphere duskside region 1 and 2 FAC system
 244 (Iijima & Potemra, 1976) is steady during interval E, except for the substorm onset at
 245 32 hours, whereas the dawnside R1/R2 system fluctuates in intensity and location. Panel
 246 *b* shows the same for the northern hemisphere, but with even more structure apparent
 247 at dawn, and some variability at dusk as well after 37 hours. We note that some weak
 248 spike activity is seen between 37 and 40 hours in the dusk sector, recorded in the north-
 249 ern hemisphere, which may be associated with the duskside R1/R2 fluctuations.

250 We now investigate the auroral and FAC signatures associated with the spikes seen
 251 during intervals D and E of Figure 1. Auroral observations are provided by the Special
 252 Sensor Ultraviolet Spectrographic Imager (SSUSI) experiment (Paxton et al., 1992) on-
 253 board the Defense Meteorological Satellite Program (DMSP) F16 and F18 satellites. The
 254 DMSP satellites are in sun-synchronous orbits roughly aligned with the dawn-dusk merid-
 255 ian near an altitude of 850 km. SSUSI recorded a swath of auroral luminosity, extend-
 256 ing sunwards and antisunwards from the orbit, in five wavelength bands. We use obser-
 257 vations at 121.6 nm, which measures Lyman- α emissions associated with proton pre-
 258 cipitation, and the Lyman-Birge-Hopfield short (LBHs) band, 140 to 152 nm, sensitive
 259 to emissions produced by soft electron precipitation.

260 Figure 4 presents six passes of the DMSP F16 or F18 spacecraft over the north-
 261 ern or southern hemispheres, the times of which are indicated by red arrows in Figure 1a.
 262 Most passes are from the southern hemisphere to optimise viewing of the dawn sector
 263 auroral oval, though the coverage of SuperMAG is better in the northern hemisphere.
 264 For each pass we also show the concurrent FAC distribution from AMPERE in both hemi-
 265 spheres. Superimposed on the AMPERE panels, cyan circles show the occurrences of spikes
 266 within ± 7 minutes of the midpoint of the pass (covering the approximate duration of the
 267 pass), where the circle radii indicate the magnitudes of the spikes.

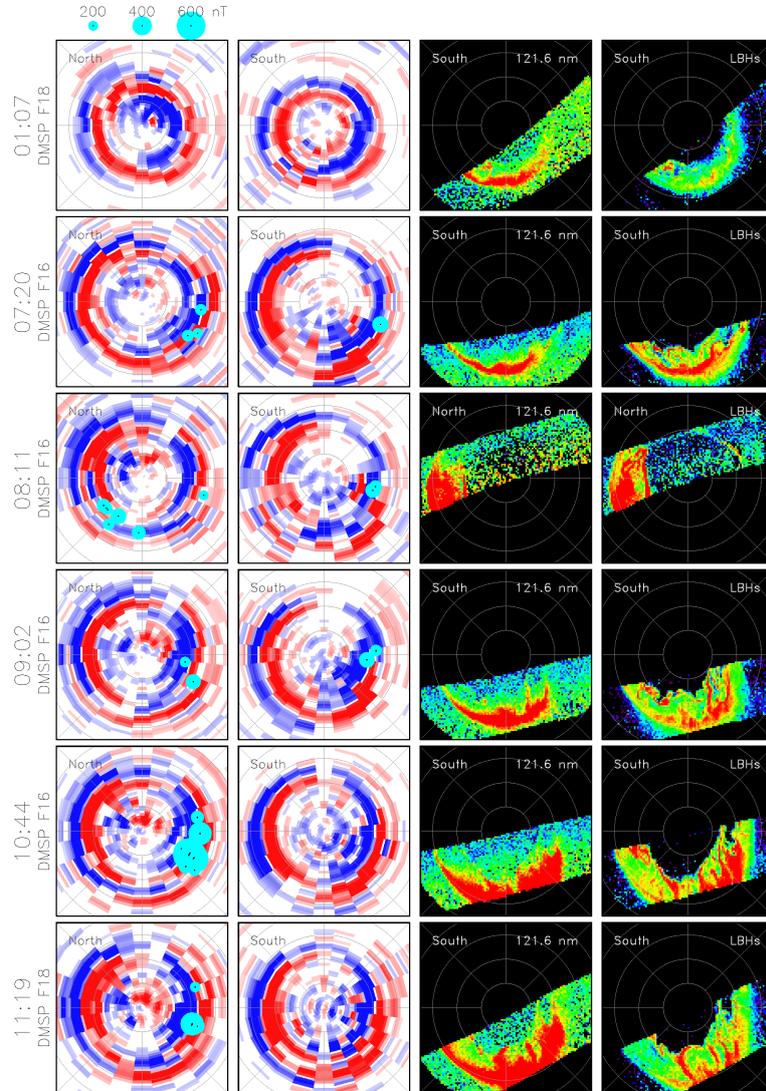


Figure 4. Left panels: AMPERE FAC configuration during six overpasses of DMSP F16 and F18 on 29 May 2010, in the northern and southern hemispheres. Noon is to the top and dawn to the right of each panel. Cyan circles show the occurrence of $dB/dt > 200 \text{ nT min}^{-1}$ in each hemisphere during the duration of each overpass. Right panels: DMSP/SSUSI observations of 121.6 nm and LBHs emissions in each overpass.

268 The first pass occurs during driven phase D. The auroras are relatively featureless,
 269 and the FACs have the characteristic R1/R2 structure (Iijima & Potemra, 1976) asso-
 270 ciated with standard twin-cell ionospheric convection. No spikes are observed at this time.
 271 The other passes occur during the period of multiple intensifications, interval E. The first
 272 is prior to the substorm onset at 32 hours. The auroras have brightened and the FACs
 273 have gained a more complicated structure, especially at dawn. Spikes in excess of 200
 274 nT are seen in the dawn sector of both hemispheres at this time. The third pass occurs
 275 during the onset at 32 hours. The auroras have expanded to lower latitudes and are much
 276 brighter, with a prominent westward-travelling surge seen at dusk, associated with a strong
 277 upwards FAC region. Spikes are seen both at dawn and in the pre-midnight substorm
 278 onset location. In the fourth pass, streamer-like auroral forms have developed in the dusk
 279 sector oval, which has considerably broadened in latitudinal width. These streamers cover
 280 between 15° and 20° of latitude from the poleward edge of the oval to the equatorward,
 281 most bright edge. Spikes are observed in the dawn sectors of both hemispheres at this
 282 time. The streamer-like activity is even more pronounced in the fifth and sixth passes,
 283 with many intense spikes seen especially in the northern hemisphere during the fifth pass.
 284 The lack of spikes in the southern hemisphere at these times is due to a gap in the lo-
 285 cation of magnetometers.

286 Despite the coverage with magnetometers being better in the northern hemisphere
 287 and the DMSP orbits favouring viewing of the dawn sector auroras in the southern hemi-
 288 sphere, we assume that there is interhemispheric conjugacy in the streamer/spike phe-
 289 nomena. The streamers are associated with complicated FAC structures, which are be-
 290 yond the temporal and spatial resolution of AMPERE to resolve accurately. However,
 291 there is evidence for quasi-periodic behaviour in the FACs on timescales of 20 to 30 mins
 292 (see Figure 1*b* and *c*).

293 **2.3.2 20 March 2001**

294 To investigate the dynamics of the auroral features with higher temporal resolu-
 295 tion than available with DMSP, we searched for similar observations during the mission
 296 lifetime (2000 to 2005) of the Imager for Magnetopause-to-Aurora Global Explorer (IM-
 297 AGE), and many examples were found. IMAGE carried the FUV Wide-band Imaging
 298 Camera (WIC) instrument (Mende et al., 2000b), which mainly measured LBH auro-
 299 ral emissions between 140 and 190 nm produced by precipitating electrons, at 2 min ca-
 300 dence. One such event, from 20 March 2001, is presented in Figure 5. Six snapshots from
 301 the WIC camera of the auroras in the northern hemisphere are shown in panels *i* to *vi*,
 302 with noon to the left, midnight to the right, and dawn to the top. Panels *a* and *b* show
 303 the auroral intensity as a function of MLT at magnetic latitudes of 68° and 63° , respec-
 304 tively; dayglow obscures the auroras near noon. Panel *c* shows the occurrence of spikes
 305 $\Delta B > 150$ nT, panel *d* shows AU and AL, and panel *e* shows the dayside reconnection
 306 rate.

307 Auroral intensifications are seen in the pre-midnight sector around 01, 09, and 14
 308 UT (panel *a*, green arrows), associated with substorm onsets seen in AL (panel *d*, green
 309 arrows). Spikes are seen in the pre-midnight sector associated with the substorms at 09
 310 and 14 UT (panel *c*, green arrows). A small number of spikes are seen in the midnight
 311 sector near 01 UT associated with substorm onset at that time. Another group of spikes
 312 after 06 UT suggests an additional substorm onset at this time, when auroral observa-
 313 tions were not available, which was also not captured by AL.

314 Dawn sector auroral structures, periodically-spaced in MLT, are associated with
 315 the substorm at 01 UT, highlighted by a red arrow in panel *a* and visible in panel *i* be-
 316 tween 04 and 07 MLT. The features are separated by close to 1 h of MLT and propa-
 317 gate sunwards at a speed of approximately 250 m s^{-1} for 90 mins after onset. Bursts of
 318 similar features are seen in the dawn sector almost continuously between 09 and 19 UT.

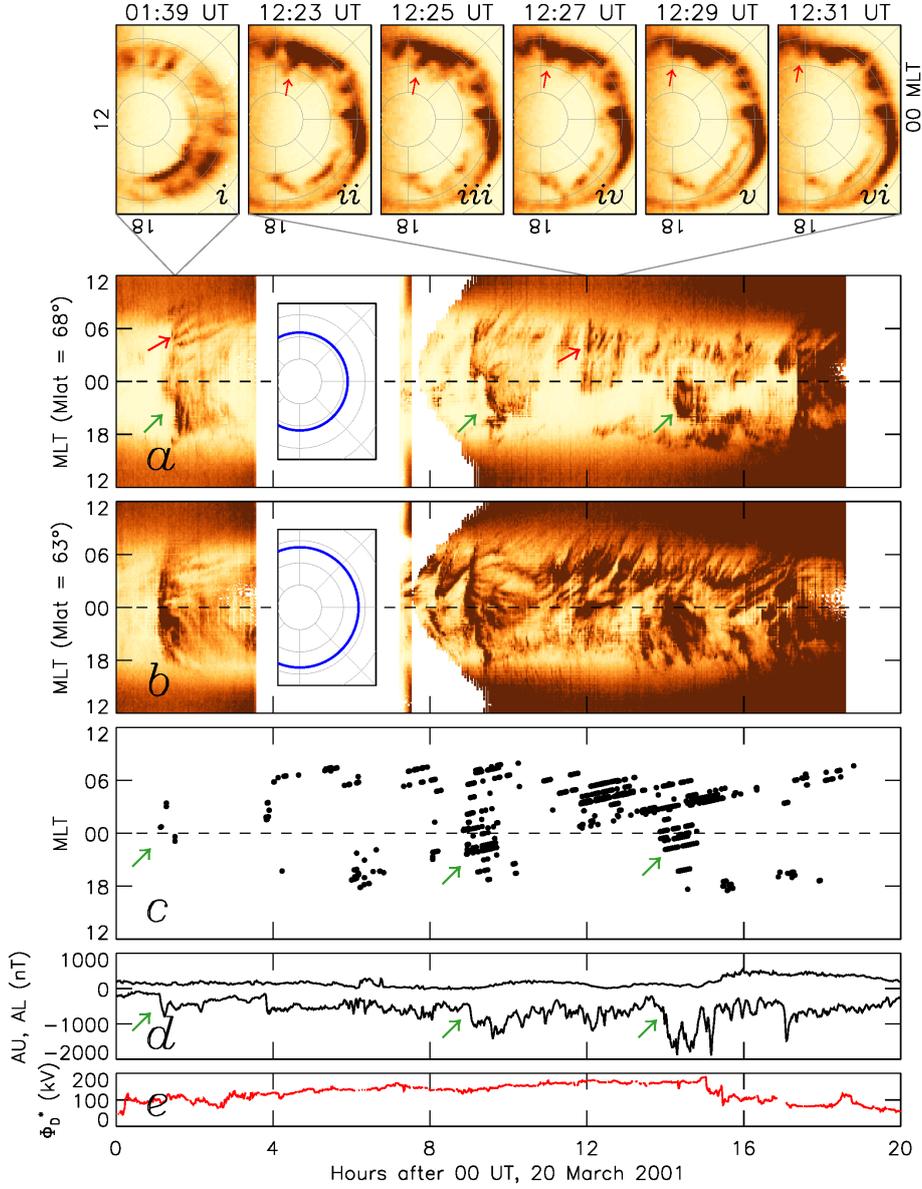


Figure 5. IMAGE/WIC auroral observations in the northern hemisphere during 20 March 2001. Panels *i* to *vi*: individual snapshots of the WIC auroral emissions, with noon to the left, dawn to the top, and midnight to the right. Red arrows track the motion of a single dawnside auroral feature. Panels *a* and *b*: the MLT distribution of auroral emissions at 68° and 63° magnetic latitude; auroras in the noon sector are obscured by dayglow. The inset panels show the latitude from which the MLT distribution is formed. Red arrows indicate the times of panels *i* and *ii* – *vi*. Panel *c*: the occurrence of $dB/dt > 150 \text{ nT min}^{-1}$. Panel *d*: the AU and AL indices. Panel *e*: the reconnection rate Φ_D^* . Green arrows in panels *a*, *c*, and *d* highlight the times of substorm onsets.

319 Although substorms occur at 09 and 14 UT, the dawn sector activity continues long af-
 320 ter each substorm breakup has ceased, such that the two phenomena may be largely in-
 321 dependent. Between 09 and 19 UT the features propagate sunwards much more rapidly
 322 than at 01 UT. One burst is highlighted by a red arrow near 12 UT in panel *a*; the cor-
 323 responding images are shown in panels *ii* to *vi* at 2 min cadence, with one auroral fea-
 324 ture identified by arrows to indicate the motion. The features are again separated by 1
 325 h of MLT but propagate with a speed near 1 km s^{-1} , giving a periodicity of approximately
 326 10-15 mins at a fixed MLT.

327 Figure 5c shows that significant spike activity is seen in the dawn sector, correspond-
 328 ing to the rapid sunward propagating auroral forms between 04 and 19 UT. Spikes are
 329 not seen at 01 to 02 UT associated with the slowly propagating features. Similar to the
 330 interval presented in Figure 1a, dawn spikes and substorms occur somewhat independ-
 331 ently of each other. Similarly, during the interval of dawn spikes, the AL index has the
 332 characteristic variability associated with the multiple intensifications convection state.
 333 During this time $V_{SW} \approx 400 \text{ km s}^{-1}$ and $B_Z \approx -17 \text{ nT}$, similar again to the condi-
 334 tions in Figure 1, and consistent with the statistics presented in Figure 3.

335 Figure 6 reproduces the IMAGE/WIC snapshot from 12:23 UT in panels *i* to *iv*.
 336 On panel *ii* is superimposed the location of SuperMAG stations which observe spikes at
 337 this time, and these are colocated with the propagating auroral features. Panels *iii* and
 338 *iv* have superimposed cross-track convection flows measured by the Ion Driftmeter (IDM)
 339 component of the Special Sensors–Ions, Electrons, and Scintillation thermal plasma anal-
 340 ysis package or SSIES (Rich & Hairston, 1994) on DMSP-F15 and -F14, respectively. These
 341 two passes bracket the time sequence shown in panels *ii* to *vi* in Figure 5, showing a con-
 342 sistent convection pattern. Although the two passes are in the southern hemisphere, they
 343 clearly show that the auroral features are in the convection return flow region, in which
 344 the flow speed is approximately 1 km s^{-1} sunwards. This speed is commensurate with
 345 the eastwards drift of the auroral features.

346 The lower panels of Figure 6 show magnetic measurements from a subset of Super-
 347 MAG stations colocated with the auroral features. Quasi-periodic fluctuations with a
 348 periodicity close to 10 mins (Ps6) and peak-to-peak variations up to 1000 nT are seen
 349 during the passage of the features, especially in the E and Z components. The spikes iden-
 350 tified are associated with the dB/dt of these quasi-periodic variations. Although dawn
 351 sector sunward-propagating auroral features are seen over many hours (Figure 5a and
 352 b), these particular magnetometer stations rotate out of the correct local time sector af-
 353 ter 13 UT.

354 3 Discussion

355 We have conducted a statistical study of the association of large dB/dt events with
 356 convection state during 2010, confirming that spikes occur preferentially in two distinct
 357 hot-spots in the pre-midnight (19 to 00 MLT) and dawn sectors (03-07 MLT). Although
 358 2010 was a geomagnetically quiet year, and the distribution of spikes is known to depend
 359 on phase of the solar cycle (e.g., Milan, Imber, et al., 2023), the 2010 distribution is con-
 360 sistent with several previous studies (e.g., Juusola et al., 2015; Kataoka & Pulkkinen,
 361 2008; Viljanen et al., 2001; Weigel et al., 2002; Schillings et al., 2022; Milan, Imber, et
 362 al., 2023).

363 The pre-midnight hot-spot is expected to be associated with substorm onsets and
 364 the formation of the substorm current wedge, producing negative excursions in the north-
 365 south component of the magnetic field; examples are shown in Figure 6 of Milan, Im-
 366 ber, et al. (2023). Our statistics show that pre-midnight spikes are indeed associated with
 367 substorm onsets, either isolated onsets (our expansion phase state) or onsets that occur
 368 during ongoing activity (driven phase onsets occurring during driven phases). Such on-

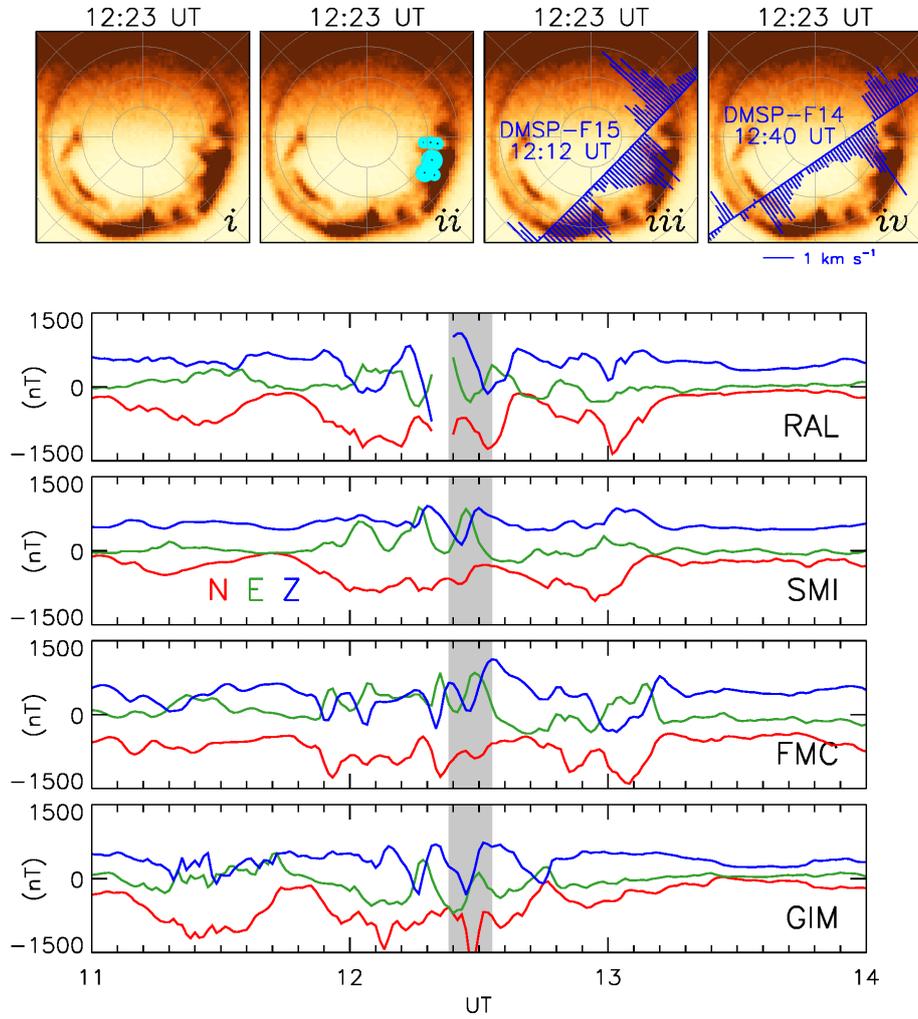


Figure 6. Panels *i* to *iv*: The WIC snapshot from 12:23 UT, with noon at the top and dawn to the right. Panel *ii*: The location of spikes observed by SuperMAG stations during this 2-min exposure. Panels *iii* and *iv*: Cross-track convection measurements from DMSP/IDM onboard F15 and F14, respectively. Lower panels: SuperMAG magnetometer measurements from stations in the dawn sector. The grey region highlights the period presented in Fig. 5 panels *ii* to *vi*.

sets can also occur during intervals classed as multiple intensifications. Substorms recur with timescales of several hours, and this modulates the occurrence of pre-midnight spikes. Pre-midnight spikes have also been observed during HILDCAAs (Milan, Mooney, et al., 2023), which are periods of quasi-periodic substorms produced by high-speed solar wind streams (Tsurutani & Gonzalez, 1987).

Our statistics show that the dawn sector hot-spot is associated with geomagnetic activity classed as multiple intensifications in AL. These spikes have previously been attributed to Ps6 (10-20 min period) magnetic perturbations in the east-west component produced by omega bands (e.g., Apatenkov et al., 2020); examples are shown in Figure 6 of this study and Figure 7 of Milan, Imber, et al. (2023). Omega bands are azimuthally-spaced auroral intensifications which drift eastwards within the sunwards convection return flow, producing quasi-periodic magnetic fluctuations on the ground, mainly in the east-west component. Dawn spikes are most prevalent during strong solar wind driving, as will be discussed in more detail below. This dawnside activity can continue for many hours, with quasi-continuous fluctuations on the ground. Strong driving is also conducive to causing substorms, so pre-midnight spikes can occur during ongoing dawn activity. However, there does not appear to be a direct link between the two phenomena. Schillings et al. (2022) surmised that during geomagnetic storms dawn sector spikes were preceded by pre-midnight spikes, that is omega band activity occurred subsequent to substorm activity. However, this does not appear to be the case in our observations.

Although dawn sector spikes have been attributed to omega bands, there is some discrepancy in the reported characteristics of the two phenomena. Omega band occurrence peaks at 02 to 03 MLT (e.g., Partamies et al., 2017; Vokhmyanin et al., 2021) whereas dawn spikes peak between 03 and 07 MLT. Indeed, Figure 2 and the study of Milan, Imber, et al. (2023) show a distinct dearth of spikes in the 01 to 02 MLT sector. Omega bands are thought to be associated with substorm onset (Wild et al., 2000) and substorm recovery phase (Vanhamäki et al., 2009), but we have shown that dawn spikes can occur independently of substorms. Typically, the omega band Ps6 magnetic signature has peak-to-peak values of a few 10s nT (Sato et al., 2015) and average $dB/dt \approx 50 \text{ nT min}^{-1}$ (Vokhmyanin et al., 2021), though extreme examples have been reported with dB/dt of several 100s nT min^{-1} (Apatenkov et al., 2020). Most reported omega bands have a latitudinal extent of a few 100s km (e.g., Vokhmyanin et al., 2021), consistent with the auroral emissions presented in Figure 5. However, we have also shown that spikes can be formed by features more resembling north-south aligned auroral streamers which can stretch over 15° of latitude, as presented in Figure 4. Indeed, a range of auroral activities occur in this local time sector and it is possible that omega bands and streamers are related (e.g., Henderson, 2022). Both phenomena occur in the same local time sector, both are azimuthally-spaced auroral intensifications, and both drift eastwards. A recent numerical simulation by Sorathia et al. (2023) has shown that streamers can be the auroral signature of bursty bulk flow (BBF) activity in the dawn sector associated with a “dawnside current wedge”; it was suggested in that study that these streamers could produce dawn sector spikes. Hence, it seems that the cause(s) of dawn sector spikes is broader than just omega bands, and that the relationship between omega bands and streamers should be investigated in more detail. Sorathia et al. (2023) suggested that such dawn activity could be caused by a dawn-dusk asymmetry in the ring current distribution, and this will be investigated in a subsequent study.

Omega bands propagate at speeds up to 1500 m s^{-1} , though 500 m s^{-1} is more typical (Vokhmyanin et al., 2021). We expect that dawn sector auroral features (omega bands or streamers) are embedded within the convection return flow so drift at, or close to, the convection return flow speed. If they represent regions of high conductance within relatively laminar return flow then they will be associated with enhanced Hall current, divergence of which at the edges of the features will lead to up/down pairs of FACs. This could explain the dawnside modulated FACs seen in Figure 1. Modulated Hall currents

will also lead to ground magnetic perturbations. The dawnside auroral features have an azimuthal spacing of approximately 1 hour of MLT (700 km) at both 01 and 12 UT in Figure 5. The periodicity of perturbations seen at a fixed MLT will then depend on the convection speed: a period of 10 min at a drift speed of 1 km s^{-1} and 20 mins at 500 m s^{-1} , consistent with the Ps6 pulsations reported to be associated with omega bands (Rostoker et al., 1980) and with our observations (Figure 6). Higher convection speeds will lead to higher frequencies in the ground magnetic perturbations, which in turn will lead to higher dB/dt for quasi-periodic pulsations of fixed amplitude (see also Vokhmyanin et al. (2021); Milan, Imber, et al. (2023)). The features highlighted by arrows in Figure 5a at 01 and 12 UT propagate at different speeds and only the fast features at 12 UT produce spikes. Convection strength is driven by dayside reconnection: Φ_D^* is close to 200 kV at 12 UT when the features propagate rapidly and spikes are observed but only 100 kV at 01 UT when spikes are not observed and features propagate slowly. This could explain the preference of dawnside spikes for stronger solar wind driving, as seen in Figure 3 and reported by Milan, Imber, et al. (2023).

Dawn spikes were the most common type during 2010, 54%, although the multiple intensifications state occurred for less than 2% of the time. Expansion and driven phases, which accounted for the majority of the other 46% of spikes occupied 25% of the time. The rapid negative excursions seen in AL during multiple intensifications are not produced by the substorm electrojet in the pre-midnight sector but by features in the westward electrojet between 03 and 07 MLT. These are not substorm onsets, and are not necessarily directly associated with substorm onsets. However, they will be misidentified as substorm onsets by any substorm identification algorithm based solely on AL (e.g., Newell & Gjerloev, 2011; Forsyth et al., 2015), as shown in Figure 7 of Milan et al. (2019). Milan et al. (2021) coined the name “multiple intensifications” based on the misunderstanding that they represented reintensifications of on-going substorm activity. We should rename “multiple intensifications” as “westjet fluctuations” or something similar.

Our statistical survey hints at a third, minor, population of spikes near 07 MLT, which occur during high V_{SW} (Figure 3d). These spikes are presumably driven by the Kelvin-Helmholtz instability occurring on the dawn magnetopause flank (see Weigel et al., 2003; Milan, Imber, et al., 2023). The year 2010 was close to solar minimum and had low average V_{SW} , explaining the relative lack of such spikes in this study.

4 Conclusions

We have studied the occurrence of dB/dt spikes in SuperMAG observations. We conclude:

- In 2010, spikes occurred preferentially in two local time sectors, pre-midnight and dawn.
- Pre-midnight spikes occurred simultaneously with substorm onsets observed in the AL index and identified in the convection state study of Milan et al. (2021).
- Dawn spikes are accompanied by strong, rapid variations in the AL index produced by activity in the dawn sector westward electrojet – these are not substorms, though are often misidentified as such.
- Dawn spikes and multiple intensifications in AL occurred during periods of strong solar wind driving with moderate solar wind speed but strongly-southward IMF.
- Pre-midnight and dawn spikes can occur independently, though they often occur together as the driving conditions necessary for dawn spikes also favour the excitation of substorms.
- Dawn spikes were associated with azimuthally-periodic auroral forms, sharing characteristics of omega bands and streamers, in the dawn sector which could extend over many degrees of latitude.

- 472 • These auroral forms propagate sunwards in the convection return flow; the speed
473 of propagation presumably depends on the convection speed, which in turn will
474 be related to the rate of dayside reconnection. The passage of these forms produces
475 Ps6 magnetic pulsations on the ground, predominantly in the E and Z components.
- 476 • Faster propagation will lead to large dB/dt for fixed amplitude quasi-periodic mag-
477 netic perturbations, favouring spikes during strong convection.
- 478 • The auroral forms, their propagation, and the magnetic perturbations they pro-
479 duce are reminiscent of omega bands. However, they are larger than most reported
480 omega bands, occur several hours of MLT eastwards of typical omega bands, and
481 their magnetic perturbations are considerably larger.
- 482 • A simulation by Sorathia et al. (2023) suggests that such auroral features are as-
483 sociated with the formation of a dawn sector current wedge and bursty bulk flow
484 activity. However, the similarity to omega bands suggests some unified mechanism
485 should be able to explain both phenomena (e.g., Henderson, 2022). The possible
486 relation of these features to a dawn-dusk asymmetry in the ring current distribu-
487 tion should be investigated.

488 5 Open Research

489 The high resolution (1-min) OMNI data used in this study were obtained from the
490 NASA Goddard Space Flight Center (GSFC) Space Physics Data Facility OMNIWeb
491 portal at https://omniweb.gsfc.nasa.gov/form/om_filt_min.html. The 1-min ca-
492 dence (“low fidelity”) SuperMAG data were obtained from NASA GSFC through the
493 SuperMAG portal at <https://supermag.jhuapl.edu/mag/?fidelity=low>. The IM-
494 AGE WIC data were obtained from CDAWeb (<https://cdaweb.gsfc.nasa.gov>). The
495 DMSP/SSUSI file type EDR-AUR data were obtained from <http://ssusi.jhuapl.edu>
496 (data version 0106, software version 7.0.0, calibration period version E0018). AMPERE
497 data were obtained from <http://ampere.jhuapl.edu>. The DMSP/SSIES data were down-
498 loaded from the Madrigal Database at Millstone Hill ([http://millstonehill.haystack](http://millstonehill.haystack.mit.edu)
499 [.mit.edu](http://millstonehill.haystack.mit.edu)).

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529 References

- 530 Akasofu, S.-I. (1964). The development of the auroral substorm. *Planetary and*
 531 *Space Science*, *12*(4), 273–282.
- 532 Anderson, B., Takahashi, K., & Toth, B. (2000). Sensing global Birkeland currents
 533 with Iridium® engineering magnetometer data. *Geophysical Research Letters*,
 534 *27*(24), 4045–4048. doi: <https://doi.org/10.1029/2000GL000094>
- 535 Apatenkov, S., Pilipenko, V., Gordeev, E., Viljanen, A., Juusola, L., Belakhovsky,
 536 V., . . . Selivanov, V. (2020). Auroral omega bands are a significant cause of
 537 large geomagnetically induced currents. *Geophysical Research Letters*, *47*(6),
 538 e2019GL086677. doi: <https://doi.org/10.1029/2019GL086677>
- 539 Chisham, G., Freeman, M., Abel, G., Lam, M., Pinnock, M., Coleman, I., . . . Vil-
 540 lain, J.-P. (2008). Remote sensing of the spatial and temporal structure of
 541 magnetopause and magnetotail reconnection from the ionosphere. *Reviews of*
 542 *Geophysics*, *46*(1). doi: <https://doi.org/10.1029/2007RG000223>
- 543 Cowley, S., & Lockwood, M. (1992). Excitation and decay of solar wind-driven flows
 544 in the magnetosphere-ionosphere system. *Annales Geophysicae*, *10*, 103–115.
- 545 Davis, T. N., & Sugiura, M. (1966). Auroral electrojet activity index AE and its
 546 universal time variations. *Journal of Geophysical Research*, *71*(3), 785–801.
- 547 Dungey, J. (1961). Interplanetary magnetic field and the auroral zones. *Physical Re-*
 548 *view Letters*, *6*(2), 47. doi: <https://doi.org/10.1103/PhysRevLett.6.47>
- 549 Engebretson, M., Kirkevold, K., Steinmetz, E., Pilipenko, V. A., Moldwin, M., Mc-
 550 Cuen, B., . . . others (2020). Interhemispheric comparisons of large nighttime
 551 magnetic perturbation events relevant to GICs. *Journal of Geophysical Re-*
 552 *search: Space Physics*, *125*(8), e2020JA028128. doi: [https://doi.org/10.1029/](https://doi.org/10.1029/2020JA028128)
 553 [2020JA028128](https://doi.org/10.1029/2020JA028128)
- 554 Forsyth, C., Rae, I., Coxon, J., Freeman, M., Jackman, C., Gjerloev, J., & Fazaker-
 555 ley, A. (2015). A new technique for determining Substorm Onsets and Phases
 556 from Indices of the Electrojet (SOPHIE). *Journal of Geophysical Research:*
 557 *Space Physics*, *120*(12), 10–592.
- 558 Gjerloev, J. (2012). The SuperMAG data processing technique. *Journal of Geo-*
 559 *physical Research: Space Physics*, *117*(A9). doi: [https://doi.org/10.1029/](https://doi.org/10.1029/2012JA017683)
 560 [2012JA017683](https://doi.org/10.1029/2012JA017683)
- 561 Gonzalez, W., Joselyn, J.-A., Kamide, Y., Kroehl, H., Rostoker, G., Tsurutani,
 562 B., & Vasyliunas, V. (1994). What is a geomagnetic storm? *Jour-*
 563 *nal of Geophysical Research: Space Physics*, *99*(A4), 5771–5792. doi:
 564 <https://doi.org/10.1029/93JA02867>
- 565 Henderson, M. G. (2022). Association of mesoscale auroral structures and breakups
 566 with energetic particle injections at geosynchronous orbit. *Frontiers in Astron-*
 567 *omy and Space Sciences*, *9*, 742246. doi: [https://doi.org/10.3389/fspas.2022](https://doi.org/10.3389/fspas.2022.742246)
 568 [.742246](https://doi.org/10.3389/fspas.2022.742246)
- 569 Hubert, B., Milan, S., Grocott, A., Blockx, C., Cowley, S., & Gérard, J.-C. (2006).
 570 Dayside and nightside reconnection rates inferred from IMAGE FUV and Su-
 571 per Dual Auroral Radar Network data. *Journal of Geophysical Research: Space*
 572 *Physics*, *111*(A3). doi: <https://doi.org/10.1029/2005JA011140>
- 573 Iijima, T., & Potemra, T. (1976). The amplitude distribution of field-aligned cur-
 574 rents at northern high latitudes observed by Triad. *Journal of Geophysical Re-*

- 575 search, 81(13), 2165–2174. doi: <https://doi.org/10.1029/JA081i013p02165>
- 576 Iyemori, T. (1990). Storm-time magnetospheric currents inferred from mid-latitude
577 geomagnetic field variations. *Journal of Geomagnetism and Geoelectricity*,
578 42(11), 1249–1265. doi: <https://doi.org/10.5636/jgg.42.1249>
- 579 Juusola, L., Viljanen, A., Van De Kamp, M., Tanskanen, E., Vanhamäki, H., Par-
580 tamies, N., & Kauristie, K. (2015). High-latitude ionospheric equivalent
581 currents during strong space storms: Regional perspective. *Space Weather*,
582 13(1), 49–60. doi: <https://doi.org/10.1002/2014SW001139>
- 583 Kataoka, R., & Pulkkinen, A. (2008). Geomagnetically induced currents during
584 intense storms driven by coronal mass ejections and corotating interacting
585 regions. *Journal of Geophysical Research: Space Physics*, 113(A3). doi:
586 <https://doi.org/10.1029/2007JA012487>
- 587 Mende, S., Heeterds, H., Frey, H., Lampton, M., Geller, S., Abiad, R., ... others
588 (2000b). Far ultraviolet imaging from the IMAGE spacecraft. 2. Wideband
589 FUV imaging. *Space Science Reviews*, 271–285.
- 590 Milan, S. (2019). *AMPERE R1/R2 FAC radii. figshare. Dataset.* [https://doi](https://doi.org/10.25392/leicester.data.11294861.v1)
591 [.org/10.25392/leicester.data.11294861.v1](https://doi.org/10.25392/leicester.data.11294861.v1). doi: 10.25392/leicester.data
592 .11294861.v1
- 593 Milan, S. (2020). *Magnetospheric Geonome Project 2010 University of Leicester.*
594 *Dataset.* <https://doi.org/10.25392/leicester.data.12571307.v1>. doi: 10
595 .25392/leicester.data.12571307.v1
- 596 Milan, S., Carter, J., Korth, H., & Anderson, B. (2015). Principal component anal-
597 ysis of Birkeland currents determined by the Active Magnetosphere and Plane-
598 tary Electrodynamics Response Experiment. *Journal of Geophysical Research:*
599 *Space Physics*, 120(12), 10–415. doi: <https://doi.org/10.1002/2015JA021680>
- 600 Milan, S., Carter, J., Sangha, H., Bower, G., & Anderson, B. (2021). Magne-
601 toospheric flux throughput in the Dungey cycle: Identification of convection
602 state during 2010. *Journal of Geophysical Research: Space Physics*, 126(2),
603 e2020JA028437. doi: <https://doi.org/10.1029/2020JA028437>
- 604 Milan, S., Clausen, L., Coxon, J., Carter, J., Walach, M.-T., Laundal, K., ... others
605 (2017). Overview of solar wind–magnetosphere–ionosphere–atmosphere cou-
606 pling and the generation of magnetospheric currents. *Space Science Reviews*,
607 206(1-4), 547–573. doi: <https://doi.org/10.1007/s11214-017-0333-0>
- 608 Milan, S., Gosling, J., & Hubert, B. (2012). Relationship between interplanetary pa-
609 rameters and the magnetopause reconnection rate quantified from observations
610 of the expanding polar cap. *Journal of Geophysical Research: Space Physics*,
611 117(A3). doi: <https://doi.org/10.1029/2011JA017082>
- 612 Milan, S., Imber, S., Fleetham, A., & Gjerloev, J. (2023). Solar cycle and solar wind
613 dependence of the occurrence of large dB/dt events at high latitudes. *Journal*
614 *of Geophysical Research: Space Physics*, 128(4), e2022JA030953. doi: [https://](https://doi.org/10.1029/2022JA030953)
615 doi.org/10.1029/2022JA030953
- 616 Milan, S., Mooney, M., Bower, G., Fleetham, A., Vines, S., & Gjerloev, J. (2023).
617 Solar wind-magnetosphere coupling during High-Intensity Long-Duration
618 Continuous AE Activity (HILDCAA). *Journal of Geophysical Research:*
619 *Space Physics*, 128(11), e2023JA032027. doi: [https://doi.org/10.1029/](https://doi.org/10.1029/2023JA032027)
620 [2023JA032027](https://doi.org/10.1029/2023JA032027)
- 621 Milan, S., Provan, G., & Hubert, B. (2007). Magnetic flux transport in the Dungey
622 cycle: A survey of dayside and nightside reconnection rates. *Journal of Geo-*
623 *physical Research: Space Physics*, 112(A1). doi: [https://doi.org/10.1029/](https://doi.org/10.1029/2006JA011642)
624 [2006JA011642](https://doi.org/10.1029/2006JA011642)
- 625 Milan, S., Walach, M.-T., Carter, J., Sangha, H., & Anderson, B. (2019). Sub-
626 storm onset latitude and the steadiness of magnetospheric convection. *Journal*
627 *of Geophysical Research: Space Physics*, 124(3), 1738–1752. doi: [https://doi](https://doi.org/10.1029/2018JA025969)
628 [.org/10.1029/2018JA025969](https://doi.org/10.1029/2018JA025969)
- 629 Newell, P., & Gjerloev, J. (2011). Evaluation of SuperMAG auroral electrojet indices

- 630 as indicators of substorms and auroral power. *Journal of Geophysical Research:*
631 *Space Physics*, 116(A12).
- 632 Ngwira, C. M., Sibeck, D., Silveira, M. V., Georgiou, M., Weygand, J. M.,
633 Nishimura, Y., & Hampton, D. (2018). A study of intense local dB/dt vari-
634 ations during two geomagnetic storms. *Space Weather*, 16(6), 676–693. doi:
635 <https://doi.org/10.1029/2018SW001911>
- 636 Partamies, N., Weygand, J., & Juusola, L. (2017). Statistical study of auroral omega
637 bands. *Annales Geophysicae*, 35(5), 1069–1083. doi: [https://doi.org/10.5194/](https://doi.org/10.5194/angeo-35-1069-2017)
638 [angeo-35-1069-2017](https://doi.org/10.5194/angeo-35-1069-2017)
- 639 Paxton, L., Meng, C.-I., Fountain, G., Ogorzalek, B., Darlington, E., Gary, S., ...
640 others (1992). Special sensor ultraviolet spectrographic imager: An instru-
641 ment description. In *Instrumentation for planetary and terrestrial atmospheric*
642 *remote sensing* (Vol. 1745, pp. 2–15). doi: <https://doi.org/10.1117/12.60595>
- 643 Pulkkinen, A., & Kataoka, R. (2006). S-transform view of geomagnetically in-
644 duced currents during geomagnetic superstorms. *Geophysical Research Letters*,
645 33(12). doi: <https://doi.org/10.1029/2006GL025822>
- 646 Rich, F., & Hairston, M. (1994). Large-scale convection patterns observed by
647 DMSP. *Journal of Geophysical Research: Space Physics*, 99(A3), 3827–3844.
648 doi: <https://doi.org/10.1029/93JA03296>
- 649 Rostoker, G., Akasofu, S.-I., Foster, J., Greenwald, R., Kamide, Y., Kawasaki, K.,
650 ... Russell, C. (1980). Magnetospheric substorms—definition and signatures.
651 *Journal of Geophysical Research: Space Physics*, 85(A4), 1663–1668.
- 652 Sato, N., Kadokura, A., Tanaka, Y., Nishiyama, T., Hori, T., & Yukimatu, A. S.
653 (2015). Omega band pulsating auroras observed onboard themis spacecraft
654 and on the ground. *Journal of Geophysical Research: Space Physics*, 120(7),
655 5524–5544. doi: <https://doi.org/10.1002/2015JA021382>
- 656 Schillings, A., Palin, L., Opgenoorth, H., Hamrin, M., Rosenqvist, L., Gjerloev, J.,
657 ... Barnes, R. (2022). Distribution and occurrence frequency of dB/dt spikes
658 during magnetic storms 1980–2020. *Space Weather*, 20, e2021SW002953. doi:
659 <https://doi.org/10.1029/2021SW002953>
- 660 Sergeev, V., Pellinen, R. J., & Pulkkinen, T. (1996). Steady magnetospheric convec-
661 tion: A review of recent results. *Space Science Reviews*, 75(3-4), 551–604. doi:
662 <https://doi.org/10.1007/BF00833344>
- 663 Sorathia, K., Michael, A., Merkin, V., Ohtani, S., Keesee, A., Sciola, A., ...
664 Pulkkinen, A. (2023). Multiscale magnetosphere-ionosphere coupling
665 during stormtime: A case study of the dawnside current wedge. *Jour-*
666 *nal of Geophysical Research: Space Physics*, 128, e2023JA031594. doi:
667 <https://doi.org/10.1029/2023JA031594>
- 668 Troshichev, O., Janzhura, A., & Stauning, P. (2006). Unified PCN and PCS indices:
669 Method of calculation, physical sense, and dependence on the IMF azimuthal
670 and northward components. *Journal of Geophysical Research: Space Physics*,
671 111(A5). doi: <https://doi.org/10.1029/2005JA011402>
- 672 Tsurutani, B., & Gonzalez, W. (1987). The cause of high-intensity long-duration
673 continuous AE activity (HILDCAAs): Interplanetary Alfvén wave trains.
674 *Planetary and Space Science*, 35(4), 405–412. doi: [https://doi.org/10.1016/](https://doi.org/10.1016/0032-0633(87)90097-3)
675 [0032-0633\(87\)90097-3](https://doi.org/10.1016/0032-0633(87)90097-3)
- 676 Vanhamäki, H., Kauristie, K., Amm, O., Senior, A., Lummerzheim, D., & Mi-
677 lan, S. (2009). Electrodynamics of an omega-band as deduced from opti-
678 cal and magnetometer data. *Annales Geophysicae*, 27(9), 3367–3385. doi:
679 <https://doi.org/10.5194/angeo-27-3367-2009>,
- 680 Viljanen, A., Nevanlinna, H., Pajunpää, K., & Pulkkinen, A. (2001). Time derivative
681 of the horizontal geomagnetic field as an activity indicator. *Annales Geophysi-*
682 *ca*, 19, 1107–1118. doi: <https://doi.org/10.5194/angeo-19-1107-2001>
- 683 Vokhmyanin, M., Apatenkov, S., Gordeev, E., Andreeva, V., Partamies, N., Kau-
684 ristie, K., & Juusola, L. (2021). Statistics on omega band properties and re-

- 685 lated geomagnetic variations. *Journal of Geophysical Research: Space Physics*,
686 126(7), e2021JA029468. doi: <https://doi.org/10.1029/2021JA029468>
- 687 Waters, C., Anderson, B., & Liou, K. (2001). Estimation of global field aligned
688 currents using the Iridium® system magnetometer data. *Geophysical Research*
689 *Letters*, 28(11), 2165–2168. doi: <https://doi.org/10.1029/2000GL012725>
- 690 Weigel, R., Klimas, A., & Vassiliadis, D. (2003). Solar wind coupling to and pre-
691 dictability of ground magnetic fields and their time derivatives. *Journal of*
692 *Geophysical Research: Space Physics*, 108(A7). doi: [https://doi.org/10.1029/](https://doi.org/10.1029/2002JA009627)
693 2002JA009627
- 694 Weigel, R., Vassiliadis, D., & Klimas, A. (2002). Coupling of the solar wind to
695 temporal fluctuations in ground magnetic fields. *Geophysical Research Letters*,
696 29(19), 21–1. doi: <https://doi.org/10.1029/2002GL014740>
- 697 Wild, J., Yeoman, T., Eglitis, P., & Opgenoorth, H. (2000). Multi-instrument ob-
698 servations of the electric and magnetic field structure of omega bands. *Annales*
699 *Geophysicae*, 18(1), 99–110. doi: <https://doi.org/10.1007/s00585-000-0099-6>
- 700 Zou, Y., Dowell, C., Ferdousi, B., Lyons, L., & Liu, J. (2022). Auroral drivers of
701 large dB/dt during geomagnetic storms. *Space Weather*, e2022SW003121. doi:
702 <https://doi.org/10.1029/2022SW003121>